



## Supplement of

## Anatomy of a high-silica eruption as observed by a local seismic network: the June 2011 Puyehue–Cordón Caulle event (southern Andes, Chile)

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## SUPPLEMENTARY MATERIAL

Here we present a number of figures and tables that compliment part of the results discussed in the main text.



**Figure S1.** Continuous seismic record at PHU station seismogram (vertical component) 24 hours/30 minutes registered on June 04<sup>th</sup>, 2011.Two classes of earthquakes are highlighted given that represent the seismic onset process 4 June eruption and the subsequent posteruptive seismic change. The first was classified as VLP recorded before the eruption (17:54 GMT, yellow dot). The second earthquake highlighted was generated after the eruption (23:00 GMT; VT; red dot).



**Figure S2.** Hybrid (HB) seismic inversion  $M_L$ :4.7 registered on April 27<sup>th</sup>, 2011 filtered 0.06-0.15 Hz. Red seismogram shows the synthetic signal, while the black seismogram represents the signal recorded by different stations. The blue numbers represent the reduction of variance. The amplitudes of the signals are normalized and also are shown in meters. The name POC station was change by name QIR\* after the eruption owing to some improvement in the received seismic signal in OVDAS.



**Figure S3.** Volcano-tectonic (VT) seismic inversion  $M_L$ :4.1 registered on July 11<sup>th</sup>, 2011 filtered 0.04-0.17 Hz. Red seismogram shows the synthetic signal, while the black seismogram represents the signal recorded by different stations. The blue numbers represent the reduction of variance. The amplitudes of the signals are shown in meters. Ensuring the proper functioning of ISOLA software, we tested the quality of the solution using Full MT inversion. The results from the large VT earthquake were compared using waveform inversion v/s fist polarities focal mechanism, obtaining similar results (Fig. 6, focal mechanism 11/7 in red rectangle).



1D Local Velocity Model Puyehue Cordón Caulle Volcanic Complex (PCCVC)

**Figure S4.** One-dimensional seismic velocity model (Vp red; Vs blue) obtained in this work (bold lines) compared with the one of Bohm et al. (2002, segmented line). The level 0 Km is related to the average height of the Cordón Caulle graben (-1.5 km above the sea level). The Bohm et al. (2002) velocity model started in -2 km and the 0 km is the sea level. Local seismic velocities resulting during the seismicity relocation process, are 1-2 km/s lower than the regional model of Bohm et al. (2002) at depth shallower than 6 km, with values of Vp near 5 km/s and Vs around 2.8 km/s. This large difference gradually decreases between 6 and 10 km however, below 6 km depth, both models get closer (Fig. S9). This is not surprising, since the model of Bohm et al. (2002) is for the entire the crust between 38° and 40°S, and our model has a local significance beneath the PCCVC, located into the intra-arc volcanic domain.



**Figure S5.** Quasi-harmonic tremor polarization analysis from Puyehue (PHU), Quirralco (QUI) and Antillanca (ANT) stations. The left side shows the seismic record and the frequency spectrum. Below, it shows the three component (Z: vertical; N: north-south and E: east-west) seismic record filtered between 0.9-1.1 Hz. The right side show the signal normalized and below then, the squares (2 sec. window) show different orientations of the particle motion trajectory polarization (L: longitudinal; Q: normal; T: transversal). The light blue rectangle selected to the analysis show the P wave onset. The NE box (north-east) is a plan view that show the azimuth (arrival) of the P wave. The blue arrows shows the 2011 volcanic vent azimuth related to the station location, and the RZ box represent the incident angle. All the waves arrival from quasi-harmonic tremor are consistent with a P wave, however, two station (PHU and ANT) suggest a shallow source (RZ $\approx$ 90°) coming from the 2011 vent.

| Data<br>(time GMT) | Deep (km)<br>Magnitud | Stations          | Filter<br>inversion | % DC v/s<br>%CLVD+ISO        | Variance<br>Reduction | Focal<br>Mechanism |
|--------------------|-----------------------|-------------------|---------------------|------------------------------|-----------------------|--------------------|
|                    | e                     |                   |                     |                              | & Mw                  |                    |
| April 27<br>06:56  | 4.3<br>Ml:4.7         | QIR<br>OSO<br>CAL | 0.07-0.15           | DC=16.6<br>CLVD+ISO=83.<br>4 | 0.7<br>Mw:4.8         |                    |
| May 05<br>02:09    | 4.6<br>Ml:4.2         | QIR<br>PHU<br>RAN | 0.09-0.18           | DC=39.6<br>CLVD+ISO=60.<br>4 | 0.6<br>Mw:4.4         |                    |
| May 17<br>14:31    | 4.1<br>Ml:3.7         | QIR<br>PHU<br>RAN | 0.1-0.21            | DC=34.5<br>CLVD+ISO=65.<br>5 | 0.62<br>Mw:3.9        |                    |
| May 27<br>08:30    | 4.2<br>Ml:4.1         | QIR<br>PHU<br>RAN | 0.07-0.17           | DC=9<br>CLVD+ISO=91          | 0.55<br>Mw:4.2        |                    |
| May 30<br>19:36    | 4.0<br>Ml:3.6         | QIR<br>PHU<br>RAN | 0.12-0.2            | DC=28.6<br>CLVD+ISO=71.<br>4 | 0.56<br>Mw:3.7        |                    |

**Table S1.**-Focal mechanisms inversions of large HB seismic events of the Enhanced Unrest Phase. Main characteristics of the earthquakes before the eruption made with ISOLA software. It highlights its depth, Local Magnitude (MI), the stations used in the inversions, the filter range, the percentages of the mechanism involved in the seismic source (DC-CLVD-ISO), the reduction of variance, Mw of the earthquake and finally a graph to the focal mechanism solution. The locations of these earthquakes are showed in the figure 3A. DC=double couple, ISO=isotropic, CLVD=compensated linearly dipole vector. Full wave inversion of some of these events confirmed the dominance of the NDC component. The Mw calculated by ISOLA coming from scalar moment Mo and it is equal to scalar product between observed and synthetic data. Therefore, Mo is proportional to a match value between synthetic and observed data. This could be the reason of the overestimated Mw value related with a  $M_L$ .