

Abstract

Multidisciplinary field projects can be very useful to a more fundamental understanding of the world around us, though these projects are not as common as they should be. In particular, the combination of archeology and geology combines our understanding of human behavior and human use of the landscape with an intimate knowledge of geologic processes and the materials available for human use in order to gain a broader understanding of human-Earth interaction. Here we present data from a cross-disciplinary project that uses a common dataset, archeological artifacts, to explore the anthropological and geologic implications of useage patterns. Archeological excavations and surveys conducted by the Office of Contract Archeology in 2007 along the route of the proposed international border fence reveal patterns of use of geologic materials by Archaic, Formative and Protohistoric Native Americans in the Boot Heel of southwestern New Mexico. Thousands of artifacts were recorded in multiple sites from Guadalupe Pass in the southern Peloncillo Mountains to the Carrizalillo Hills west of Columbus. We identified the lithologies of artifacts, ranging from projectile points to groundstones, and then constructed material movement maps based on either known procurement sites (“quarries”) or outcrops identified as the closest source to a given site for each lithology. Not unexpectedly, the majority of the rock types utilized by native peoples are local siliceous volcanic materials. However, several artifacts constructed from obsidian were transported into the region from northern Mexico and eastern Arizona, indicating long-distance travel and/or trade routes. We also examine useage pattern difference between Archaic, Formative and Protohistoric sites. Additionally, a dramatic change in distribution of sources for geologic materials occurs between one pre-Spanish site and one post-Spanish site that are adjacent to one another.

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1 Introduction

Lithic procurement strategy in the Boot Heel region of southwestern New Mexico has largely gone untreated in the archaeological literature. This is primarily because there has been relatively little archaeology done in this area in recent years, but is also due to a lack of documentation of the lithic raw material sources in the area. Here we provide a description of the lithic resources in and around the Boot Heel and use this information to examine patterns of lithic procurement among the Archaic, Formative and Protohistoric periods of the area. Data were collected from sites recently recorded by the Office of Contract Archeology (OCA) for the proposed tactical barrier fence to be constructed on the US-Mexico international border by the Department of Homeland Security. Lithic material types that were utilized at these sites are identified and described in detail and possible source localities for these materials are identified. The relative proportions that these sources were utilized on the sites assigned to the temporal periods in question are then analyzed to determine similarities and differences among them.

The artifacts under analysis in this chapter were compiled from three separate project areas related to the border fence project. These projects ranged from the Dog Mountains in the southeastern Boot Heel to the Carrizalillo Hills in the east and included the eastern border of the Boot Heel (Figs. 1, 2). The broad project areas are referred to here, from west to east, as HV 5-8, Three Sites and Nineteen Canyon.

Nearly all of the lithic materials on the sites were likely procured from the major geologic features of the region. The Boot Heel of New Mexico, as well as southeastern Arizona and northern Mexico are part of the Basin and Range Province of western North America, which is characterized by extensional faulting, producing uplifted blocks of Paleozoic and Mesozoic strata and intervening basins with predominantly Quaternary sediments. In addition, the San Juan volcanic field extends along the Arizona-New Mexico border from northern Mexico into southwestern Colorado and is comprised of Tertiary-aged silicic igneous flows and domes.

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2 Methods

2.1 Material recording

The artifacts and raw materials used here were recorded by several individuals in two different settings. All of HV 5-8 were recorded in the field during pedestrian survey. Approximately 10 recorders (including CH) made designations of material type for lithic artifacts based on their accumulated knowledge, a common practice during cultural resource survey. The remaining artifacts were recovered during the data recovery phase and were recorded in the OCA laboratory. Artifacts were washed with water and toothbrushes upon their return to lab. Material recording was then conducted by three individuals with some level of lithic analysis training (including CH), occasionally with the aid of a low-powered microscope.

A representative sample of specimens was selected and examined by a professional geologist (KZ) and described in detail. These included recovered artifacts and source samples that were analyzed in-hand and specimens in high-resolution photographs from survey. Munsell colors for recovered specimens are from the 1991 Geological Society of America Rock-Color Chart. Items that are tentatively identified from high resolution photographs are described only by basic lithotype and Munsell colors for these specimens are estimated. Inferences made from these examinations were selectively applied to the recorded artifact data mentioned in the previous two paragraphs in order to fix observed errors in the database.

The geologic base map for the Boot Heel region was constructed by scanning in geologic maps of Arizona (Reynolds, 1988), New Mexico (Wilks, 2005) and northern Mexico and redrafting the units into a consistent map. High resolution geologic maps (1:24 000 scale) were used to modify the base map where needed (Zeller, 1959, 1962, 1966, 1970; Soule, 1972; Seager and Clemmons, 1988). Many units have been simplified, especially for sedimentary packages that include multiple rock types.

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2.2 Temporal period designations

The temporal periods used in this analysis follow guidelines established for the New Mexico Cultural Resource Information System (NMCRIIS) in 1993. This is a standard system used by most Cultural Resource Management (CRM) firms throughout New Mexico. Only sites with components belonging to a single temporal period are considered in this study. Sites with multiple recognizable Native American components belonging to more than one temporal period and sites of unknown temporal affiliation are not included. It is understood that there is always potential for multiple components of different temporal periods to contribute to any site assemblage regardless of recognizable components. However, for the purpose of making this study feasible, all artifacts from a site assigned to a given period are considered to belong to that period also.

2.2.1 Archaic

The Archaic period spanned from 5500 BC to 200 AD. This period is typically defined as being a time during which people were primarily mobile hunters and gatherers that were slowly transitioning to a more sedentary lifestyle. It is distinguished from the earlier Paleoindian period by an increased emphasis on gathering as opposed to a primarily hunting-based subsistence. This is reflected by increased utilization of groundstone tools. The Archaic is generally distinguished from the later Formative period by a lack of ceramics, permanent village structures and intensive agriculture. However, there is often substantial overlap between the periods regarding these characteristics because the transition to sedentism was a very gradual process.

Archaic period designations were primarily based on the presence of diagnostic projectile points but were also based on groundstone assemblage analysis and radiocarbon dates obtained from thermal features.

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2.2.2 Formative

The Formative period for southwestern New Mexico is regarded as being occupied by the Mimbres Mogollon tradition between AD 200 and AD 1400. This period is typically characterized by sedentism at the village level and higher, intensive agriculture and the presence of ceramics. This period ends when these sedentary people exit the archaeological record. Around AD 1400 large villages and agricultural plots ceased to be constructed and traditional ceramics were no longer manufactured. This signals the transition to the Protohistoric period.

Formative period designations were primarily based on the presence of diagnostic ceramics but were also based on diagnostic projectile points and radiocarbon dates obtained from thermal features.

2.2.3 Protohistoric

The NMCRIS guidelines do not provide dates for a distinctive Protohistoric period. For the purposes of this study, it is assigned to cultural groups that existed after the Formative period that had not yet been assimilated into cultural groups with written histories (i.e. Europeans). The maximum extent of this period spanned from AD 1400 to AD 1850 with different groups appearing, disappearing and becoming assimilated at different times. These groups included the Apache, Suma and Manso. People belonging to these groups were typically highly nomadic, built largely impermanent dwellings and used agriculture and ceramics far less than Formative period people.

Protohistoric period designations were primarily based on radiocarbon dates obtained from thermal features and lithic analysis results but were also based on diagnostic projectile points.

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2.3 Site cluster localities

The sites used in this study are located over a broad area that is interspersed with lithic source materials. Therefore, an analysis of the combined data set by period alone would not be able to roughly model the direction and distance that materials travel from source to destination. Breaking down the assemblage into the individual sites would create data sets that are too small for comparisons and the number of potential comparisons would be unfeasibly large. Instead, sites are divided into clusters that consist of aerially associated sites and are distinctly separated from one another. This allows for larger data sets while retaining some association with location. These clusters are defined below from west to east.

The Dog Mountains cluster consists of sites located between Eagle Mountain in the west and the eastern foothills of the Dog Mountains in the east. The Little Hat Top Butte cluster consists of four sites located north and northeast of the Alamo Hueco Mountains and south and southeast of the Big Hatchet Mountains. The Archaic sites are located near the foot of Little Hat Top Butte on the ridge that joins the Alamo Hueco and Big Hatchet Mountains. The Formative sites are located immediately next to one another approximately 7 km to the east along an unnamed drainage that leads from Little Hat Top Butte. The Lower Hachita Valley cluster is comprised of sites that are located primarily in the Hachita Valley between the Big Hatchet Mountains and Sierra Rica Mountains. This area is confined to the southeastern part of the valley that is in the United States. The Upper Hatchita Valley site is located in the Hachita Valley approximately 4 km west of the Sierra Rica Mountains and 6 km northeast of the northwestern Big Hatchet Mountains. The Doyle Peak cluster includes four sites that are located along a stretch of access road approximately 10 km long that runs between the Sierra Rica Mountains and the Apache Hills. Doyle Peak is located near the middle of this stretch.

The Three Sites cluster consists of three sites on the international border approximately 3 km west of the Carrizalillo Hills and 4 km south of the foothills of the Cedar

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Mountains and the Nineteen Canyon site located along drainage on the eastern slopes of the Carrizalillo Hills approximately 200 m north of the international border.

2.4 Procurement strategy analysis

In this contribution, we utilize the information gathered from the material descriptions, local geology and geology maps to examine potential procurement patterns in different areas of the Boot Heel and during different time periods. Distances to source areas and assemblage proportions are the primary data that are used to conduct this analysis.

For the purposes of creating calculable distances for material travel, site clusters are simplified to single points on the maps. These are calculated as the geographic centers of all sites within a cluster belonging to the same temporal period. The Universal Transverse Mercator (UTM) coordinates for the site datums, obtained with GPS units, were averaged to create a single set of coordinates that represents the geographic center of each cluster. These coordinates were then converted to latitude and longitude and plotted on the map to the nearest minute (Figs. 4–6).

Exact provenances for most of the materials in this study are largely unknown. Therefore, sources for materials were selected with the assumption that the most parsimonious options (i.e. the closest sources) were utilized by inhabitants of the site clusters under study. It is understood that this is highly unrealistic given the nature of human mobility patterns, but it allows for a basis of comparison. Sources for materials are defined as either the geologic units identified during the construction of the map or known source localities identified during survey. The source for each material type in each cluster is then defined as the potential source that has the shortest radial distance to the geographic center of the cluster. Distances were estimated by measuring with calipers to the nearest hundredth of a millimeter on the map and converting to kilometers. A few of the longer distances that could not be accurately estimated on the map were calculated with Google Earth. These distances are represented with arrows on the maps (Figs. 4–6).

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Proportions of material types were analyzed statistically among the assemblages using Homogeneity of Proportion tests based on the chi-square probability distribution. Further analysis was conducted on the individual material types using Fisher's Exact Two-Proportion tests with Bonferroni error rates. Significance was determined at $p = 0.05$. These analyses were carried out with MINITAB and SPSS statistical software programs. Proportions are represented on the map with shaded arrows indicating material type contributions to the assemblage in 25% intervals.

The distance and proportion data were also combined to form the Average Minimum Distance (AMD) of transport for all materials to the individual site clusters. This figure is calculated with the following formula:

$$i = s$$

$$AMD = \sum p_i d_i$$

$$i = 1$$

where p_i is the proportion and d_i is the distance in kilometers to nearest potential source for the i -th material type in the assemblage containing s types. This provides a figure that accounts for the contribution of each artifact to the distance that materials travelled from source to destination (site cluster). Comparisons of this figure were conducted among time periods both within site clusters and for the entire assemblage.

3 Material descriptions

Samples characteristic of each area of the Boot Heel are described in detail below and are divided into appropriate rock type categories.

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3.1 Igneous rocks: extrusive

3.1.1 Rhyolite

Rhyolite specimens in the HV 5-8 area were examined in photographs and include large cobbles commonly found in the centers of roasting pits and some projectile points. Many of the projectile point specimens are pale pink to pinkish gray in color and are porphyritic with large euhedral, transparent quartz crystals and lesser lathe-shaped, transparent crystals tentatively identified as sanidine. Some projectile points from the area were constructed from pyroclastic rhyolite with fragments of other rock types (primarily pumice) included in a fine-grained groundmass. The roasting pit cobbles are more difficult to identify because most of them are covered with thick layers of char. However, some that were chipped and show fresh surfaces appear to be primarily pyroclastic rhyolite fragments.

Samples of rhyolite from the Three Sites area are pale red in color and show less color variation than rhyolites from the other areas. Some of the samples are pyroclastic and included rounded pumice and fragments of what have been tentatively identified as siltstone. Rhyolite samples are porphyritic aphanitic, sometimes faintly foliated, with small phenocrysts of euhedral quartz, sanidine and plagioclase feldspar. Samples of rhyolite recovered from the Nineteen Canyon area are pale red to moderate red in color, with little color variation. Samples are porphyritic aphanitic, sometimes faintly foliated, with small phenocrysts of euhedral quartz, sanidine and plagioclase feldspar. One sample contains plagioclase feldspar and small amphibole needles.

3.1.2 Andesite

Andesite specimens examined in photographs from the HV 5-8 area are aphanitic to fine-grained porphyritic. Porphyritic specimens exhibit small, white, opaque feldspar crystals. Because these materials were examined in photographs, it is not possible to determine which type of feldspar is present.

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Specimens of andesite from the Three Sites excavations, examined in the laboratory and ranging in color from grayish red to light brownish gray, are porphyritic aphanitic and have a distinctive mottled appearance. Phenocrysts include pyroxene, amphibole (commonly twinned) and quartz. No biotite was identified in hand sample. Pyroxene crystals tend to be weathered out of the host rock, leaving small holes that show two cleavage planes at 90°. One sample of andesite recovered from the area may be pervasively thermally altered as all phenocrysts and the groundmass are the same grayish red color.

Andesite specimens from the Nineteen Canyon excavations are medium gray to pale brown and are porphyritic aphanitic, with mottled appearances. Phenocrysts include pyroxene, feldspar, and rare biotite. No quartz was identified in hand sample. Individual crystals of biotite (and possibly very small magnetite) have red oxidation halos.

3.1.3 Basalt

Basalt from the HV 5-8 area was examined in photographs and several projectile points were constructed from aphanitic, nonvesicular basalt that is black in color and often weathered to an orange-brown. Manos were constructed from vesicular basalt that is also black in color and weathered to a pale gray. These specimens were all examined in photographs, so the presence or composition of any small phenocrysts could not be positively identified.

Specimens of basalt in the Three Sites area are dark gray and occasionally weathered to a pale yellowish brown. They are aphanitic and locally contain abundant vesicles, some of which are long and thin, lending a “burrowed” appearance to the rock. One sample has rare, very small feldspar lathes, although phenocrysts are usually not visible. Specimens from the Nineteen Canyon area are dark gray, aphanitic and locally contain abundant vesicles that are aligned along flow bands. As with the Three Sites basalt, phenocrysts are not visible.

3.1.4 Dacite

In the HV 5-8 area, a dark gray aphanitic dacite with rare sanidine and biotite crystals has been collected from a possible quarry site. The groundmass is generally too fine grained to identify its composition, but has local areas of larger crystals with either biotite or very small pyroxene crystals. While this material does not contain quartz phenocrysts, as would be expected in a dacite or a rhyolite, more mafic minerals (such as pyroxene) are not present in the mineral assemblage, such that this rock should not be termed an andesite, in spite of its darker color. One specimen collected displays flow banding, with at least one band defined by a siliceous layer.

3.1.5 Ash-flow tuff

Ash-flow tuff specimens from the HV 5-8 area were identified in photographs and are pale gray in color and appear ash rich. Most of the tuffaceous materials used to construct lithic artifacts were semi-welded tuffs that are harder and less obviously ash-rich.

3.1.6 Obsidian

Obsidian objects from the HV 5-8 area were examined in photograph and consist primarily of projectile points that are either opaque black or translucent olive green to black.

A single obsidian fragment was recovered from the Three Sites area that is grayish black in color, opaque with a waxy luster and with radial fracture lines. Fragments of obsidian recovered from the Nineteen Canyon are black and opaque. Near the edges, where the rock is thin enough to be translucent, the material exhibits an olive green tinge, similar to obsidian flakes from the Antelope Wells obsidian field.

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3.2 Sedimentary rocks

Sandstone and siltstone

HV 5-8 sandstone samples are dark brown to purple brown volcanoclastics that are medium to coarse grained, poorly sorted and with angular to subangular grains. Siltstone samples from the Three Sites area are pale red or grayish orange to light brown in color. Samples are too fine grained to identify mineral composition and may have undergone very low-grade metamorphism. Sedimentary structures identified include relict cross-laminations and planar laminations. Sandstone samples from the Nineteen Canyon area are arkoses that are grayish red to grayish orange in color. These rocks are medium to coarse grained, moderately well sorted and the grains themselves are subangular. They contain up to 80% feldspar and approximately 20% quartz with rare lithic grains. Hematite cement coats the grains, giving these rocks a mottled red and pink color. No sedimentary structures were observed in hand-sample. Siltstone samples range in color from moderate red to pale olive. Few of the samples are metamorphosed fully to slate, but show signs of very low temperature heating and/or slightly higher pressures. Most are quartzose with relict cross-laminations or tabular bedding. Some samples are micaceous and others have been metamorphosed to such a degree that they are very silicic and almost chert-like in appearance. One sample has a distinctive red and yellow mottling that probably reflects diagenetic mottling prior to alteration.

3.3 Metamorphic

Quartzite

Quartzite from the HV 5-8 area are buff to purple in color and range from coarse grained to very fine grained. Some of the very fine-grained quartzites can be mistaken for coarsely crystalline chert without examination under a binocular microscope. Usually

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relict sedimentary structures and grains can be identified in the quartzites and include laminations, cross-laminations and occasional pebbles of other lithologies.

Specimens of quartzite from the Three Sites area range in color from pale purple to moderate orange. All are fine grained with occasional medium-grained quartz crystals that are transparent. These larger, very clear quartz crystals may indicate higher degrees of metamorphism than samples without these crystals were subjected to. Two samples include pebbles of quartzite of a different texture.

Quartzite samples from the Nineteen Canyon area are pale red to grayish orange and very fine to fine-grained. Several samples include 10 to 15 per cent lithic grains, although all samples are predominantly quartzose. Many display relict cross-laminations and some samples have been metamorphosed enough that grain boundaries have become diffuse and silicic rinds have developed on the outer surfaces. Four samples are arkosic to subarkosic with pyrite crystals or pyrite replaced with botryoidal goethite. The pyrite and goethite are enclosed in red oxidation halos.

3.4 Chert and chalcedony

3.4.1 Chert

Chert from the HV 5-8 area was examined in photographs and includes a variety of colors and textures, including the distinctive golden-brown color of Eagle Mountain chert, gray, red, white, pink, black and yellow. Textures range from brecciated to massive.

Chert from the Three Sites area are brown, black, red and ivory in color and seams of pale blue or white-blue are common. Textures observed in the chert from this area include lacy, massive, coarsely crystalline, mottled, brecciated, distorted laminated and dendritic. All specimens are opaque with a waxy luster, though some are duller than others. Weathering rinds (cortex) are common as are pores filled with prismatic quartz crystals. Two samples are red-orange to red-brown in color with brecciated texture. Chert samples from the Nineteen Canyon area include black, red blue, white, pink and gray varieties. Textures observed in these specimens include lacy, massive, coarsely

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crystalline, mottled, distorted laminated, banded (faint) and concentrically zoned. Veins of more translucent material are common. Weathering rinds (cortex) are common as are pores filled with prismatic quartz crystals. Two samples are orange mottled chert.

3.4.2 Chalcedony

Chalcedony found in the HV 5-8 area were examined in photographs, as well as in the field and is a translucent to near-opaque white and bluish brown material with some pink and pale red coloring. The weathered exterior or rind is usually ivory or brown in color.

Chalcedony fragments from the Three Sites excavations are a variety of colors, including yellow, white, pale blue, gray, and pale brown, pink and orange, and are translucent with semi-waxy to dull luster. Textures identified include concentrically zoned mottles (2–3 cm in diameter), streaking and massive. Some specimens still have a weathering rind (cortex) present that is nodular. Chalcedony fragments from the Nineteen Canyon excavations are light gray, pale orange, light yellow and bluish gray, and are massive, with few variations in texture. Cortex is common on most specimens, as is the occurrence of small clusters of prismatic quartz crystals.

4 Local geology

4.1 HV 5-8

In the eastern Boot Heel, geologic source areas include the Alamo Hueco and Dog Mountains, the Big and Little Hatchet Mountains, the Sierra Rica and Apache Hills, outcrops in northern Mexico and possibly minor sources in the southern and eastern Animas Mountains. The southern and eastern Animas Mountains are dominated by rhyolitic units that are Eocene to Oligocene in age and these flows are related to both eruptive suites found throughout the Boot Heel (35.2–32.7 Ma, 27.6–26.8 Ma; McIntosh

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and Bryan, 2000). The OK-Bar Conglomerate, which covers large areas of the southern Animas Mountains, is comprised of boulders and cobbles of locally derived volcanic materials and includes significant amounts of obsidian weathering from the Park Tuff (Zeller, 1962; Shackley, 1995). The Alamo Hueco and Dog Mountains are comprised of mostly Tertiary volcanics, including rhyolites, tuffs and andesitic basalts (Zeller, 1959; Wilks, 2005). In some places, local faults are excellent sources of hydrothermally deposited chalcedony: LA 162029 is a local procurement area for chalcedony that is near a faulted outcrop. There are local outcrops of Cretaceous limestone and sandstone to the northwest of and in the northern Alamo Huecos that describe the hinge of a syncline plunging to the southeast. U-Bar Ridge, the area for the type section of the U-Bar Formation, forms the hinge of the syncline. Sites LA 161106 and LA 161996, Eagle Mountain Chert procurement sites, lie on small remnants of this syncline.

The Big Hatchet Mountains are primarily Paleozoic limestone of Pennsylvanian and Permian age, with smaller outcrops of the Cambrian-Ordovician Bliss Sandstone and Mississippian to Devonian carbonates along the margins of the range (Zeller, 1965). Paleozoic limestone is a potential source area for chert of a variety of colors (white, gray, lavender, red, brown and black). Specifically, the Permian Concha Limestone is the only unit that has been documented that produces red and lavender chert nodules (Zeller, 1965). Other colors of chert cannot be specifically linked to any Paleozoic limestone unit. The Bliss Sandstone contains arkosic sandstones, conglomerates that contain some Precambrian quartzite boulders, a white orthoquartzite and a sequence of thick sandy dolomite (Zeller, 1965).

The Little Hatchet Mountains are comprised of fault blocks of a variety of lithologies. The southern Little Hatchet Mountains include outcrops of Precambrian and Tertiary granite, upper Cretaceous to lower Tertiary diorite and monzonite intrusions, Pennsylvanian carbonates, Lower Cretaceous limestone and sandstone of the U-Bar and Mojado Formations and Upper Cretaceous basalt (Ringbone Formation) (Zeller, 1970). The central Little Hatchets are primarily Cretaceous sedimentary rocks, including the Lower Cretaceous Hell-to-Finish, Mojado, U-Bar Formations and Upper Cretaceous

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being procured from the Sierra Fresnal secondary field in northernmost Mexico. The Carrizalillo Hills are a Lower Oligocene rhyolite and AFT complex (Wilks, 2005). The Apache Hills and Sierra Rica are a source for quartzite and Paleozoic outcrops in northern Mexico may be a procurement area for a variety of cherts. The Cedar Mountains include upper Oligocene to lower Miocene basalts and basaltic andesites, Miocene to Pliocene basalts to andesites, upper Oligocene rhyolites and lower Oligocene silicic domes, flows and rhyolite AFTs. Sedimentary strata present in the Cedar Mountains include minor outcrops of Lower Cretaceous strata (Wilks, 2005). The Klondike Hills are probably not an immediately useful procurement source but they do contain relatively well-preserved sequences of Paleozoic limestone, dolostone and shale that may have served as a chert source (Armstrong, 1970).

The Tres Hermanas are a Tertiary monzonite and Upper Eocene to Lower Oligocene andesite complex with minor outcrops of Pennsylvanian and Permian strata (Seager and Clemons, 1988; McLemore, 2000a; Wilks, 2005). Granite identified in the Nineteen Canyon excavations is most likely to have come from the Tres Hermanas monzonite outcrops (as opposed to the granite present in the southernmost Little Hatchets). The northern Burro Mountains, north of Lordsburg, contain outcrops of the Beartooth Quartzite, which contains very small cubes of pyrite (McLemore, 2000b). There is a possibility that quartzites from the Lower Cretaceous Mojado Formation may contain pyrite, though the occurrence of pyrite in these quartzites has not been documented. Disseminated pyrite has been found in the Tres Hermanas in fissure veins, though it has not been reported as being found in a quartzite (McLemore, 2000a).

5.1.2 Dog Mountains procurement site descriptions: Eagle Mountain A chert (LA 161106)

The Eagle Mountain A chert procurement site is located along a very small outcrop of the Lower Cretaceous U-Bar Formation that had not previously been mapped (e.g., Zeller, 1959; Fig. 3). The chert, a distinctive orange-brown in color with red flecks and massive, brecciated or mottled textures, occurs as nodules and laterally discontinuous irregularly shaped beds within a limestone that is a laminated to hummocky laminated gray wackestone to packstone. The laminations often exhibit soft-sediment deformation and fossils have been replaced chert. In field samples and photographs, none of the fossil material could be positively identified. We interpret the Eagle Mountain A limestone outcrop to be Lower Cretaceous U-Bar Formation because the outcrop lies on the trace of the southwestern limb of a southeastward plunging syncline described by mapped outcrops of U-Bar Formation to the north and east (U-Bar Ridge describes the hinge of this syncline). This outcrop is a primary source for Eagle Mountain chert.

5.1.3 Dog Mountains procurement site descriptions: Eagle Mountain dacite (LA 161105)

The Eagle Mountain dacite procurement site occurs on a low rise immediately north of the international border (Fig. 3). Low, narrow ridges of dacite trend northwest-southeast across the rise and hundreds of artifacts have been observed in the immediate vicinity. The dacite is aphanitic and pale gray with rare phenocrysts of sanidine and biotite. The site occurs along the western edge of a band of rhyolite outcrop and is a localized occurrence of dacite that is a primary procurement source. Dacite is not common in the Boot Heel and this procurement locality is the only are where this material was worked extensively.

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5.1.4 Dog Mountains procurement site descriptions: Middle Mountain A chalcedony (LA 161108)

The Middle Mountain A chalcedony procurement site, in the western part of the pass between the Dog Mountains and the Alamo Hueco Mountains, is located in a broad valley on a low hill of exposed welded tuff/rhyolite bedrock (Fig. 3). The chalcedony is ivory, pale blue or pale red in color, is often laminated to nodular in texture and occurs in very dense concentrations on the surface of the site. Chalcedony is a common hydrothermally deposited mineral phase in areas with silicic igneous rocks (e.g. White and Corwin, 1961). The Middle Mountain A site is a secondary deposit of chalcedony that probably originated in fractures within the andesite and rhyolite outcrops immediately to the north and has weathered out onto the modern land surface.

5.1.5 Dog Mountains procurement site descriptions: Middle Mountain B chalcedony (LA 162029)

The Middle Mountain B procurement site is in the eastern part of the pass between the Dog Mountains and Alamo Hueco (Fig. 3). The site is on a south-facing slope where the toes of two hills meet. These hills are two different igneous units with the contact between the two occurring in the saddle between the two hills. The Middle Mountain B chalcedony is probably weathering out of the contact zone or fractures within either igneous unit and is a primary chalcedony source. The Middle Mountain B chalcedony is pale blue or pale brown in color with primarily massive or nodular textures. The chalcedony found loose on the surface of the site occurs as much larger blocks and nodules than at Middle Mountain A.

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5.1.6 Dog Mountains procurement site descriptions: Dog Springs chalcedony (LA 89049)

The Dog Springs chalcedony site is located on the eastern flank of the pass between the northern Dog Mountains and the Alamo Hueco Mountains (Fig. 3). It occurs in a small valley between two ridge lines of rhyolite and is on or near two contacts between flow units and is immediately north of a thrust fault. The surface of the site is primarily rhyolite outcrop and the chalcedony is probably weathering out of the contacts between flow units or from fractures within the units that may be related to the fault to the south and is therefore a primary source. Chalcedony from this procurement locality is similar in colors and textures to that of other chalcedony procurement sites.

5.1.7 Dog Mountains procurement site descriptions: Eagle Mountain B chert (LA 161996)

The Eagle Mountain B procurement site contains chert nodules similar in color and texture to the Eagle Mountain A chert. The Eagle Mountain B chert is golden-brown to reddish brown. Golden-brown pieces tend to have lacey to massive textures, whereas red specimens have lacey to brecciated textures. Blue translucent chalcedony is present as discrete blebs or disseminated through the red varieties. The range of textures seen in the Eagle Mountain B chert is not as diverse and there seems to be a stronger association between color and texture type than is observed in the Eagle Mountain A chert. The Eagle Mountain B chert occurs as nodules littering the land surface on a north-facing slope near Dog Springs Ranch (Fig. 3). There are no documented outcrops of limestone near the site, indicating that the chert nodules are a secondary deposit. Tertiary fanglomerates containing Paleozoic limestone to the east are down-slope of the site, and are unlikely to be the source for the Eagle Mountain B chert. A plausible explanation is that the chert nodules represent erosional remnants of a local outcrop of U-Bar Formation that has since been eroded away, leaving only the chert nodules.

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5.1.8 Dog Mountains procurement site descriptions: Dog Springs quartzite (LA 85792)

The Dog Springs quartzite procurement area is located on the alluvial plain between the Dog Mountains to the west and Haystack Mountain to the east (Fig. 3). The site occurs on the north side of a low rise formed by a weathered outcrop of bedrock that is a primary source for the artifacts at this site. This site was originally described by Human System Resources in 1991 and they identified the bedrock as quartzite. Site survey undertaken in the summer of 2007 reclassified the site as resting on basalt bedrock. However, geologic maps of the area describe the hills immediately east of the site as Lower Cretaceous Mojado Formation and do not document basalt near this site (Zeller, 1959; Wilks, 2005). Therefore, it is unlikely that the bedrock here is basalt, but is most likely the “brown-weathered sandstone” or quartzite observed by Zeller (1959).

5.2 Hachita Valley and Doyle Peak procurement sites

5.2.1 Hachita Valley and Doyle Peak local geology

The Hachita Valley is bounded by the Big Hatchet Mountains to the southwest, the Little Hatchet Mountains to the west and the Sierra Rica and Apache Hills to the northeast and east. Doyle Peak is located in the valley between the Sierra Rica to the south and the Apache Hills to the north-northeast. The Big Hatchet Mountains are composed primarily of Paleozoic limestone units with extensive outcrops of the Cambrian Bliss Formation at the northern end of the range (Zeller, 1966). The Little Hatchet Mountains contain a variety of lithologies that range from Precambrian and Tertiary granite at the southern end of the mountains to Paleozoic and Cretaceous sedimentary strata and Cenozoic igneous outcrops in the central and northern mountains (Zeller, 1970). The Sierra Rica are composed of the Lower Cretaceous Mojado Formation that has been so silicified that the sandstone units are now quartzite (Zeller, 1966). Minor fault blocks of Paleozoic limestone occur on the western end of the Sierra Rica. The Apache Hills

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are a mixture of silicified Mojado Formation and extensive andesite flows. Doyle Peak is a fault block of Paleozoic limestone (Wilks, 2005).

5.2.2 Hachita Valley and Doyle Peak procurement site descriptions: Hachita Valley banded chert (LA 161991)

5 The Hachita Valley banded chert procurement site is not close to prominent outcrops of chert-bearing rock, and so is tentatively identified as a procurement site on the basis of the high density of raw banded chert nodules found (Fig. 3). The site is on a southwest-northeast trending gravelly terrace that overlooks the Hachita Valley to the southeast. The Big Hatchet Mountains are about two miles to the southwest. Chert utilized at
10 this site is a distinctive lavender and reddish pink banded material that is most likely to have come from the Permian Concha Limestone. The Concha Limestone is exposed in the center and southwestern flank of the Big Hatchet Mountains. The Hachita Valley procurement site is a secondary deposit of this material after the chert was eroded out of outcrops in the Big Hatchet Mountains and redeposited as terrace gravels within the
15 Hachita drainage.

5.2.3 Hachita Valley and Doyle Peak procurement site descriptions: Doyle Peak chalcedony (LA 161989)

The Doyle Peak chalcedony procurement site is on a west-facing slope west of a low ridge that overlooks an incised arroyo to the south. Doyle Peak is approximately two
20 to three miles to the west-southwest (Fig. 3). The chalcedony occurs as two to ten centimeter diameter nodules on the modern land surface and is a secondary deposit. The original source of the chalcedony is probably from flow contacts or within-unit fracture zones in the andesite in the southern Apache Hills.

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6 Analysis results

The results below are visually represented in the maps (Figs. 4–6) and Table 1 below. Statistical results are summarized in the text. Once again, site clusters are presented from west to east.

6.1 Dog Mountains

This cluster consists of Archaic and Formative sites. Each set of sites has a distinct location separate from one another. The Archaic sites are located on the eastern flanks of the Dog Mountains and the Formative sites are west of the Dog Mountains centered an area south of Middle Mountain.

Exactly 50% of the Archaic assemblage is made up of Eagle Mountain chert and chalcedony. The nearest Eagle Mountain chert source is approximately 1 km away at LA 161996. The site cluster lies on a chalcedony producing fault zone and one of the sites in the cluster, LA 89049, is a chalcedony procurement site. Rhyolite and tuff are also immediately available. Basalt is available in the southern Alamo Hueco Mountains, also very nearby. Other cherts are available in the southern Dog Mountains in northern Chihuahua. The most distant source is for obsidian, at the Antelope Wells field.

The Formative assemblage is dominated by chalcedony. This chalcedony is abundant in this region, outcropping at both LA 161108 and LA 162029, which is located exactly at the geographic center for the Formative sites. Rhyolite and tuff are also immediately available. Basalt, once again, arrived from the southern Alamo Hueco Mountains. Eagle Mountain chert came from LA 161996 and other cherts are again available in the southern Dog Mountains. The nearest quartzite is a small Cretaceous outcrop east of the Dog Mountains. Permian chert is coming from the southern end of the Big Hatchet Mountains and the obsidian is coming from the Antelope Wells obsidian field 22 km to the west.

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The Homogeneity of Proportions test between the two assemblages was highly significant ($p < 0.000$). The Fisher's Tests indicate that the Archaic assemblage has a significantly higher proportion of Eagle Mountain chert (24% more of the assemblage) and that the Formative assemblage has a significantly higher proportion of chalcedony (39% more of the assemblage). These differences appear to be a direct result of the locations of the site clusters. The Archaic sites are located very near an Eagle Mountain chert source and the Formative sites are sitting on an expansive chalcedony outcrop. Another interesting difference that is not statistically significant is the higher proportion of obsidian in the Archaic assemblage.

The AMDs for the Archaic and Formative assemblages are 4.00 and 1.73 km, respectively. These are the lowest values in this study. This is because these clusters are immediately near sources of materials that were most utilized in their assemblages. In fact, the only reason the Archaic AMD is more than twice as large as that of the Formative is the higher proportion of obsidian from the Antelope Wells obsidian field.

6.2 Little Hat Top Butte

Similar to the Dog Mountains site cluster, the Archaic and Formative sites for the Little Hat Top Butte cluster group separately in distinct areas. The Archaic sites are grouped around the base of Little Hat Top Butte and the Formative sites are grouped farther to the east, about approximately one mile west of the international border.

The majority of artifacts in the Archaic assemblage are Eagle Mountain chert, other cherts and rhyolite. The nearest known Eagle Mountain chert source is LA 161996, however these sites are located very near a large outcrop of U-Bar Formation, which could potentially yield other sources of Eagle Mountain chert. Quartzite is also immediately available in the U-Bar Formation. Other cherts, including Permian cherts, arrived from nearby in the southern Big Hatchet Mountains. Rhyolite, tuff and chalcedony came from less than a kilometer away in volcanic flows north of the Alamo Hueco Mountains. Sandstone could be procured immediately to the south in the Emory Canyon area.

Basalt is available in the northern and central Alamo Hueco Mountains. By far, the farthest potential source is the Antelope Wells obsidian field 33 km to the southwest.

The Formative assemblage is very similar to the Archaic assemblage in terms of material proportions. Furthermore, the nearest sources of materials are identical. The exception is the presence of limestone in the Formative assemblage, the nearest known source of which is the same as for chert, in the southern Big Hatchet Mountains. In general, the sites are slightly further from the sources than the Archaic sites.

The Homogeneity of Proportions test between the two assemblages was not significant ($p = 0.644$). Fisher's tests were also run to confirm this result and they too showed no significant differences. The proportions among the assemblages are very similar. The only noticeable anomaly is the presence of one piece of obsidian in the Archaic assemblage which had to be procured from a relatively long distance.

The AMDs for the Archaic and Formative assemblages are 7.43 and 9.37 km, respectively. These are relatively low and, like the Dog Mountains figures, are the result of being fairly close to multiple procurement areas. The slight difference of these values is, again, the result of the Formative sites being slightly farther from the source areas.

6.3 Lower Hachita Valley

This cluster contains sites from the Archaic, Formative and Protohistoric periods. Nearly all of the sites are located in a tight group near the center of the Hachita Valley basin. As a result, the geographic centers of the Archaic and Protohistoric sites are very close to one another near the international border. There is, however, one Formative site located closer to the foot of the Big Hatchet Mountains. This site draws the geographic center of the Formative sites a few miles to the west.

The Archaic assemblage is comprised mostly of quartzite, basalt and chert. Quartzite is available in the southern Sierra Rica Mountains, basalt in the northern Alamo Hueco, and chert (and limestone) in the Big Hatchet Mountains. Permian chert is once again available in the central Big Hatchets. The nearest chalcedony source is LA 161989 in the Doyle Peak region. Rhyolite and tuff are available in the volcanic

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substantial contribution to all the assemblages. In the case of the Protohistoric assemblage, basalt accounts for more than three-quarters of the exceptionally high AMD value. If there is a nearer basalt source, these AMDs would be much lower. Rhyolite provides most of the remainder of the Protohistoric figure because of its large contribution to that assemblage. The slight difference between the Formative and Archaic AMDs appears to be the slightly closer proximity of the Formative geographic center to the sources and the higher proportions of quartzite and Permian chert in the Archaic assemblage.

6.4 Upper Hachita Valley and Doyle Peak

The Formative site in the upper Hachita Valley lies approximately 7 km southwest of the geographic center of the Doyle Peak Archaic sites. These sites comprise a single analytical cluster and are only presented with different names to better describe their locations. This need arose because of the large expanse that the Archaic sites are spread across.

The Archaic assemblage is mostly made up of equal parts basalt, Permian chert and other cherts. The nearest source of basalt is at Flat Hill immediately east of the town of Hachita. Permian chert once again is coming from the central Big Hatchet Mountains. Other cherts and limestone are immediately available around the base of Doyle Peak. Quartzite and sandstone are very nearby in the Mojado Formation of the Sierra Rica Mountains. The nearest known chalcedony source is LA 161989, which is located within the Archaic site cluster area. Diorite is available in the central Little Hatchet Mountains. The closest rhyolite is found in the Cedar Mountains. More distant is LA 161996, the nearest known source of Eagle Mountain chert. The Sierra Fresno secondary field is the nearest source of obsidian, but Antelope Wells obsidian was also identified in the assemblage.

The Formative assemblage at LA 161080 was limited to 5 lithics. The two pieces of quartzite likely came from the Mojado Formation of the Sierra Rica Mountains, the

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same as the Archaic. The two pieces of basalt and single piece of rhyolite, however, were likely procured from the central Little Hatchets.

All statistical tests on proportions were not significant. The Formative assemblage is simply too small to derive any inferences about it.

The AMDs for the assemblages are very similar at 13.87 km for the Archaic assemblage and 13.01 km for the Formative. These values are average among all values for the clusters. Basalt and, for the Archaic assemblage, Permian chert were the largest contributors to these values because of their high proportions in the assemblages and the long distances to sources.

6.5 Three Sites and Nineteen Canyon

The Three Sites and Nineteen Canyon assemblages are both Protohistoric and located about 6 km from each other. Radiocarbon dates were obtained from both clusters indicating that Nineteen Canyon, which is the eastern cluster, is slightly older with a proposed occupation range between AD 1420 and 1540. Three Sites, to the west, is suggested to date from AD 1620 to 1800. These assemblages are compared to explore any potential differences between these periods/culture groups.

The Three Sites assemblage is primarily made up of chert and rhyolite. Chert is immediately available from a Paleozoic outcrop that extends southward into Chihuahua. Rhyolite and tuff are available just to the east in the Carrizalillo Hills, though there is a possibility that rhyolite was being procured immediately, at LA 85779. Andesite, basalt and chalcedony were procured a little further away from the southern Cedar Mountains and northern Carrizalillo Hills. The nearest siltstone and quartzite outcrops are in the Mojado Formation of the northern Sierra Rica Mountains. The single piece of obsidian was likely to have come from the Sierra Fresnal secondary field. The only known source of Permian chert is in the Big Hatchets. The furthest source is for Eagle Mountain chert at LA 161996.

Nineteen Canyon is dominated by a distinctive grey-to-purple fine-grained quartzite from the northern Sierra Ricas. Many of the other sources are identical to that of

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the Three Sites assemblage. The exceptions are basalt from domes east of the Carizalillo Hills, a piece of granite from the Tres Hermanas, Antelope Wells obsidian and Beartooth quartzite from the Burro Mountains southwest of Silver City.

The Homogeneity of Proportions test between the two assemblages was highly significant ($p < 0.000$). The Fisher's tests indicated that the Nineteen Canyon assemblage has a significantly higher proportion of quartzite and the Three Sites assemblage has significantly higher proportions of rhyolite, chalcedony, chert, Eagle Mountain chert and Permian chert. Clearly the large abundance of quartzite in the Nineteen Canyon assemblage is the largest source of difference between the assemblages. Quartzite makes up 63% more of its assemblage than it does in Three Sites.

The AMDs for Three Sites and Nineteen Canyon are 8.91 and 18.10 km, respectively. The Three Sites figure is relatively low and the Nineteen Canyon figure is one of the highest. The largest proportions of materials for Three Sites could have been procured locally, while the quartzite at Nineteen Canyon had to have traveled at least 22 km.

6.6 Entire assemblage

A final analysis was conducted among the Archaic, Formative and Protohistoric assemblages of the entire data set. The Archaic and Formative assemblages are directly comparable, because the sites were sampled in the same fashion at all site clusters. The Three Sites and Nineteen Canyon assemblages, however, were directly targeted for excavation without sampling any sites around them. These sites represent the bulk of the Protohistoric artifacts and are located somewhat separately from the remainder of the sites. The distinct location of the Protohistoric sites would likely influence any results. It is for this reason that the Protohistoric assemblage was not included in the proportion analysis below. It is, however, included in the AMD analysis because distance to source is a general phenomenon that should be less influenced in this case. It is important to note that there are significant taphonomic processes affecting these assemblages, including (but not limited to) sorting by geologic processes, collection of artifacts by nonarcheologists and "recycling" of material by Apache groups.

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The Homogeneity of Proportions test between the Archaic and Formative assemblages was highly significant ($p < 0.000$). Fisher's tests indicated multiple significant differences. The Archaic assemblage has significantly higher proportions of andesite, obsidian, rhyolite and chert. The only andesite is from western site clusters, suggesting only local use of this material by Archaic populations. Western site clusters are also primarily responsible for the differences in the other three materials because of its large contribution to the overall assemblage. In fact, chert often has a similar or higher proportion in the Formative assemblages of other clusters.

The Formative assemblage has significantly higher proportions of basalt, chalcedony and Eagle Mountain chert. Once again, the western site cluster assemblage plays a large role in these results. The low proportions of basalt and Eagle Mountain chert in that assemblage are the primary reasons for those significant differences. In the case of the chalcedony, however, it is the extraordinarily high proportion in the Dog Mountains Formative assemblage that provides the largest contribution to the difference. Again, this is because those sites sit in a chalcedony procurement "zone".

AMDs among all periods were examined using a standard one-way ANOVA. The values used for this test are found in Table 1. There were no significant differences ($p = 0.403$). This appears to be primarily because of the small number of values. The means for these samples are interesting though. The means for the Archaic and Formative samples are similar at 10.65 and 10.53 km, respectively, while that of the Protohistoric is higher at 15.72 km. One might have expected the mean Archaic and Protohistoric AMDs to be higher than the Formative because of higher mobility. Instead the similar values for the Archaic and Formative suggest similar procurement patterns and, perhaps, similar levels of mobility. The larger mean Protohistoric AMD likely reflects higher mobility during this period.

7 Summary and discussion

The preceding analysis is a first attempt at understanding and quantifying lithic procurement strategies of Native American populations of the Boot Heel area. The questions under study are: (a) *are there noticeable patterns of lithic procurement within and among the different temporal periods?* and (b) *are there substantial differences in lithic procurement patterns between the different temporal periods?* Viewing procurement in terms of material type proportions within assemblages and distances to nearest sources, the analysis provided some intriguing observations.

Obviously, the model being used in this study suffers from some unavoidable inadequacies. For instance, ideally it would be desirable to have the exact locations of lithic sources and have exact identifications and provenances of the artifacts in question. Also, for the most part, this study only recognizes primary sources of raw materials while it is known that many materials for lithic artifacts were procured from secondary deposits including gravel beds. Unfortunately, this type of information was not available at the time of this study. Another potential problem lies in the identification of temporal periods of sites. These identifications rely heavily on observed diagnostic artifacts. The presence or absence of multiple occupations is highly susceptible to being masked by collecting activity. In this area, it is known that local residents frequently collect diagnostic projectile points and ceramics on private land. Should the only diagnostic artifact(s) of a given component on a site be removed, that component will likely not be recognized. Additionally, it is suspected that Protohistoric populations of the region frequently scavenged artifacts of previous temporal periods and deposited them on their own sites, potentially causing Protohistoric components to be misidentified (Seymour, 2010: 12). Recent research on the Three Sites and Nineteen Canyon projects have produced evidence supporting this claim. Several diagnostic Archaic points and a Clovis point were recovered from these sites, but radiocarbon dates and lithic artifact analyses strongly suggest predominantly Protohistoric occupations. As a result of collecting phenomena, sites in this study identified as being of one component could, in fact, be Protohistoric or have mixed components.

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A major assumption that was made going into this analysis was that Archaic and Protohistoric populations were far more mobile than Formative populations in this region and that these differences would be reflected in lithic procurement patterns. The results of this analysis do not entirely support this assumption. While the average minimum distance of transport for the Protohistoric assemblage is relatively high, those of the Archaic and Formative assemblages are similar. This could reflect similar mobility levels in the Archaic and Formative that were much lower than mobility levels during the Protohistoric. Alternatively, it could reflect a disconnect between overall mobility and lithic procurement, or possibly the inherent inadequacies of the model. The observation that only one of the artifacts from very distant sources is in the Formative assemblage and several are in the Archaic assemblage provides potential support for high mobility in the Archaic.

Preferences for individual material types appear to be heavily influenced by site location and proximity to lithic sources. That being said, there are some interesting associations of material types with temporal periods. Archaic populations appear to have preferred to use obsidian slightly more than Formative populations. The contribution of Antelope Wells obsidian to all assemblages decreases with distance to the east and north of the procurement field (as noted by Findlow and Bolognese, 1980). However, the rate of this drop appears to be slightly lower in the Archaic assemblages. This could reflect one or more of three conceivable phenomena: (1) Archaic populations may have been slightly more willing to travel greater distances to acquire obsidian, (2) Archaic populations were indeed more mobile and (3) vectors of mobility for Archaic populations were more focused to the north and east away from the Antelope Wells obsidian field than were the other populations. The large abundance of obsidian on LA 159188 in the Guadalupe Mountains, another Archaic site, could indicate a movement vector away from the Antelope Wells field or that the unsubstantiated obsidian sources in the southern Peloncillo Mountains do, in fact, exist. Shackley (1995) examined multiple outcrops in the southern Peloncillo Mountains, but was unable to locate a point source for obsidian.

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Eagle Mountain chert also shows a pattern of diminishing abundance away from the source area. The large abundances in the Little Hat Top Butte assemblages suggests that there are undiscovered sources of Eagle Mountain chert as far north as the southern margin of the Big Hatchet Mountains, likely in the U-Bar Formation limestones. The abundance of this material drops rapidly north of Little Hat Top Butte and west of the Dog Mountains area. The exception is a high abundance of Eagle Mountain chert in the Formative assemblage at Cloverdale Playa. This could indicate a movement vector for this population heading west from the Dog and Alamo Hueco Mountains.

The distinctive gray-to-purple quartzite identified solely in the Nineteen Canyon assemblage was observed by one of the authors (KZ) at an outcrop in the Mojado Formation of the Sierra Rica Mountains 22 km to the west. The large abundance of this material in the assemblage could indicate a rapid movement of this population to the east. This information is even more interesting when considering the Protohistoric assemblage from the Lower Hachita Valley. This assemblage is dominated by basalt and rhyolite, the nearest sources of which are to the south and west. If these sites are related, the combined picture suggests a group of people moving to the north and east away from the Alamo Hueco Mountains. Further evidence would be necessary to confirm any link between these sites.

A further distinction noted in the analysis is the greater AMD of the Nineteen Canyon assemblage than that of Three Sites. Notably most of the dates for Nineteen Canyon appear to predate Coronado and the Spanish colonization of New Mexico while Three Sites dates to the period during European presence. The change in AMDs could reflect a change from highly mobile populations to populations that had more restricted movement patterns due to the threat of conquest and raids from increasingly hostile neighbors.

As a final note on material types, basalt is found in unusually high abundance for all temporal periods in the Hachita Valley and Doyle Peak areas. There are no observed basalt sources in the area. The nearest source for any of the clusters is 15 km away and the average distance is 20.68 km. The only other recent project in the Boot Heel also

tion on the period of occupation. Hopefully, all of these issues can be improved on in the future allowing for better understanding of lithic procurement, not only in the Boot Heel of New Mexico, but globally.

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Table 1. Count and percentage of assemblage of geologic materials documented in the eastern Boot Heel region. Upper number in cell is the count, lower number is the percentage of the assemblage. A = Archaic, F = Formative, P = Protohistoric.

Material	Dog Mountains		Little Hat Top Butte		Lower Hatchita Valley			Upper Hatchita Valley	Doyle Peak	Three Sites	Nineteen Canyon
	A	F	A	F	A	F	P	F	A	P	P
Rhyolite	3 16.7	30 14.4	6 13.6	7 10.9	8 4.7	2 3.3	13 25.5	1 20.0	9 4.9	228 29.8	107 10.5
Tuff	1 5.6	3 1.4	3 6.8	8 12.5	2 1.2	-	1 2.0	-	-	7 0.9	5 0.5
Andesite	-	80 2.71	-	-	-	-	1 1.96	-	-	6 0.8	18 1.8
Basalt	2 11.1	20 9.6	3 6.8	3 4.7	38 22.2	16 26.2	32 62.8	2 40.0	44 24.0	18 2.4	33 3.2
Diorite	-	-	-	-	-	-	-	-	2 1.1	-	-
Granite	-	-	-	-	1 0.6	-	-	-	-	-	1 0.1
Obsidian	2 11.1	2 1.0	1 2.8	-	8 4.7	2 3.3	-	-	6 3.3	1 0.1	-
Antelope Wells Obsidian	-	1 0.5	-	-	-	-	-	-	2 1.1	-	3 0.3
Sierra Fresnal Obsidian	-	-	-	-	-	-	-	-	-	-	3 0.3
Prismatic quartz	-	-	-	-	-	-	-	-	1 0.6	-	-
Chalcedony	4 22.2	127 61.1	4 9.1	3 4.7	12 7.0	5 8.2	1 2.0	-	11 6.0	69 9.0	31 3.0
Chert	1 5.6	13 6.3	8 18.2	13 20.3	33 19.3	21 34.4	1 2.0	-	39 21.3	298 38.9	101 9.9
Eagle Mtn. Chert	5 27.8	8 3.85	13 29.6	20 31.3	8 4.7	4 6.6	-	-	7 3.8	35 4.6	12 1.2
Permian Chert	-	3 1.4	3 6.8	6 9.4	19 11.1	3 4.9	-	-	43 23.5	86 11.2	12 1.2
Gypsum	-	-	1 2.3	-	-	-	-	-	-	-	-
Limestone	-	-	-	3 4.7	1 0.6	-	-	-	3 1.6	-	-
Sandstone	-	-	1 2.3	-	1 0.6	-	-	-	1 0.6	-	5 0.5
Siltstone	-	-	-	-	-	-	-	-	-	5 0.7	19 1.9
Quartzite	-	1 0.5	1 2.3	1 1.6	40 23.4	8 13.1	2 3.9	2 40.0	15 8.2	13 1.7	662 65.0
Beartooth Quartzite	-	-	-	-	-	-	-	-	-	-	7 0.7
Total Count	18	208	44	64	171	61	51	5	183	766	1019

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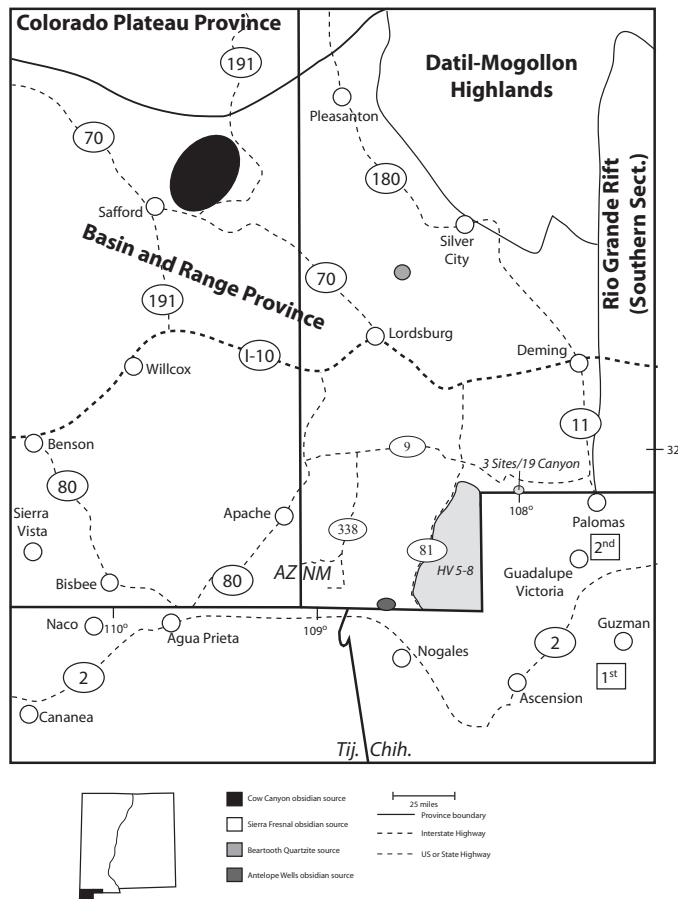


Fig. 1. Location of the study area in regional context. Known locations for distinct geologic materials are indicated by shaded areas.

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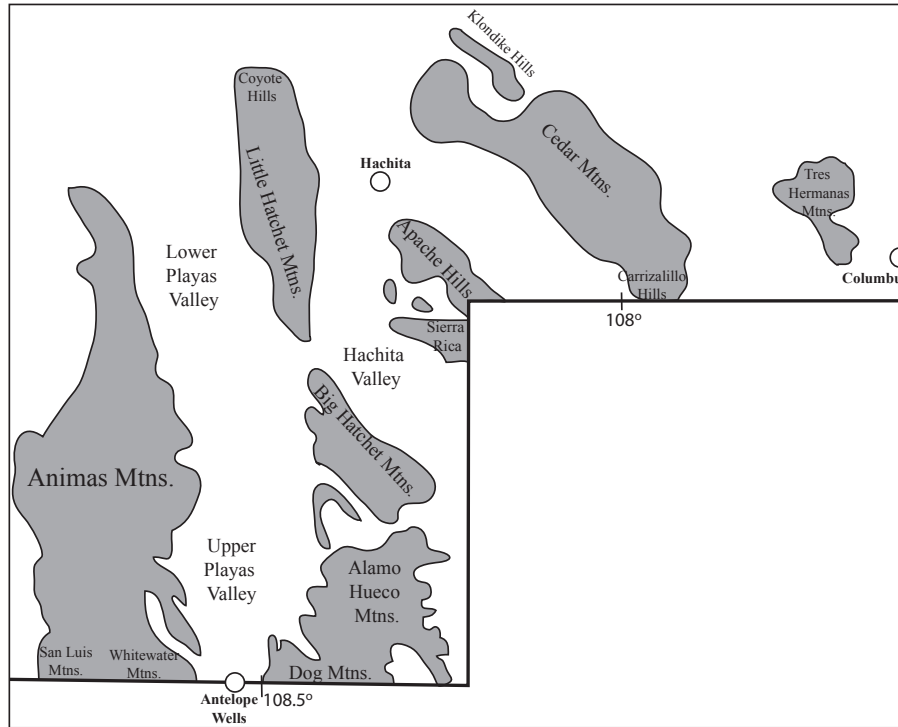


Fig. 2. Physiographic map of southwestern New Mexico.

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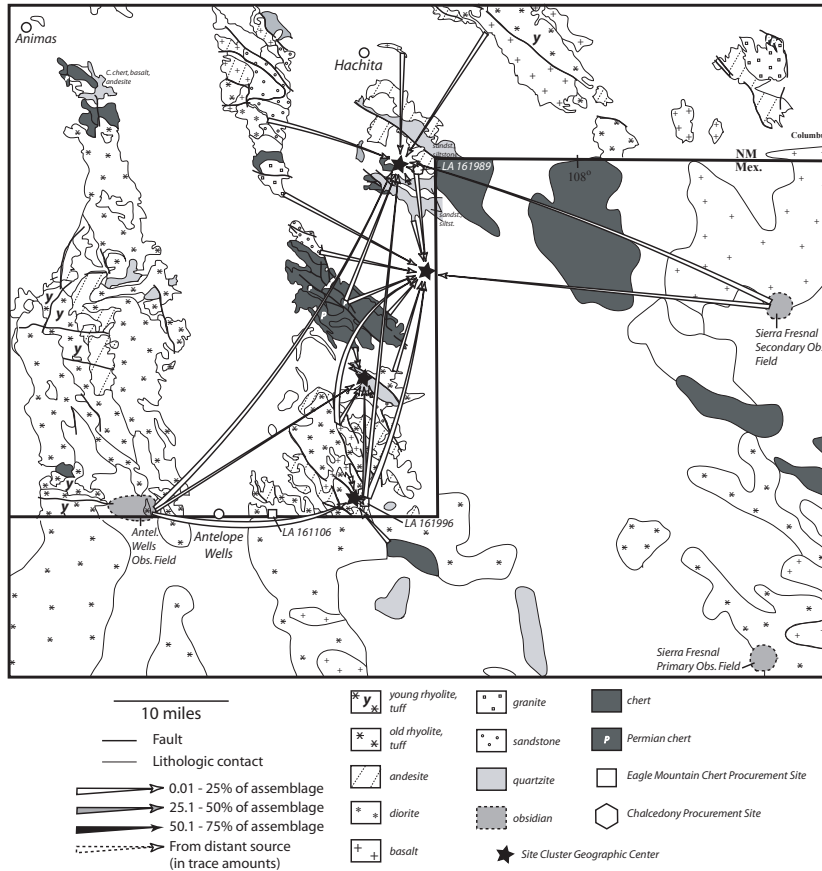


Fig. 3. Geologic map of the Boot Heel of southwestern New Mexico with locations of site clusters and minimum distance movement of geologic materials for sites identified as Archaic.

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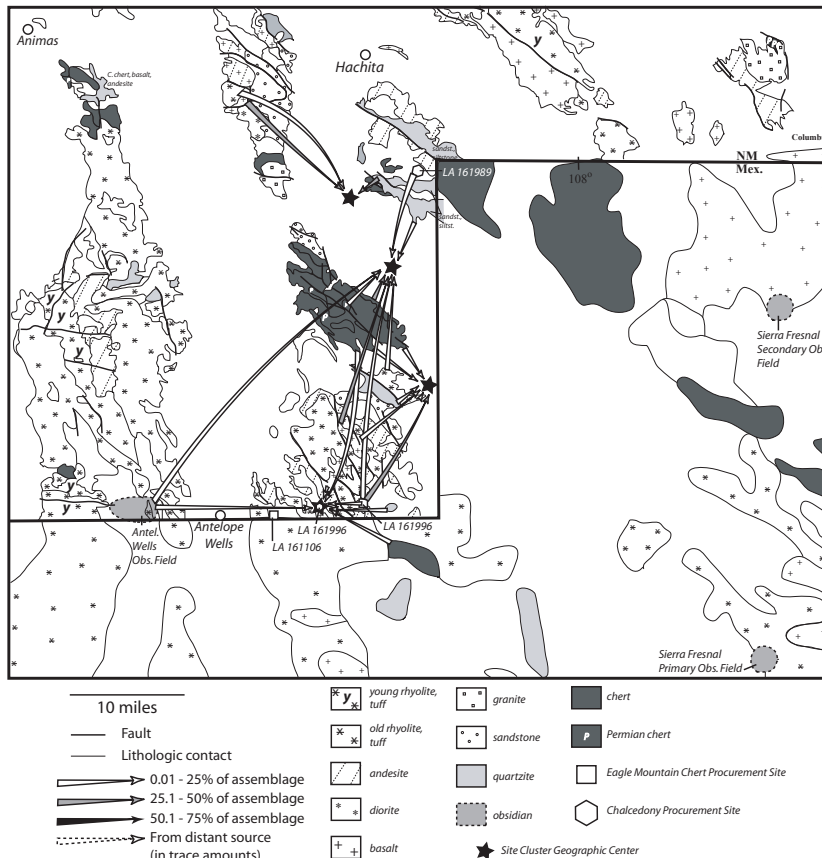


Fig. 4. Geologic map of the Boot Heel of southwestern New Mexico with locations of site clusters and minimum distance movement of geologic materials for sites identified as Formative.

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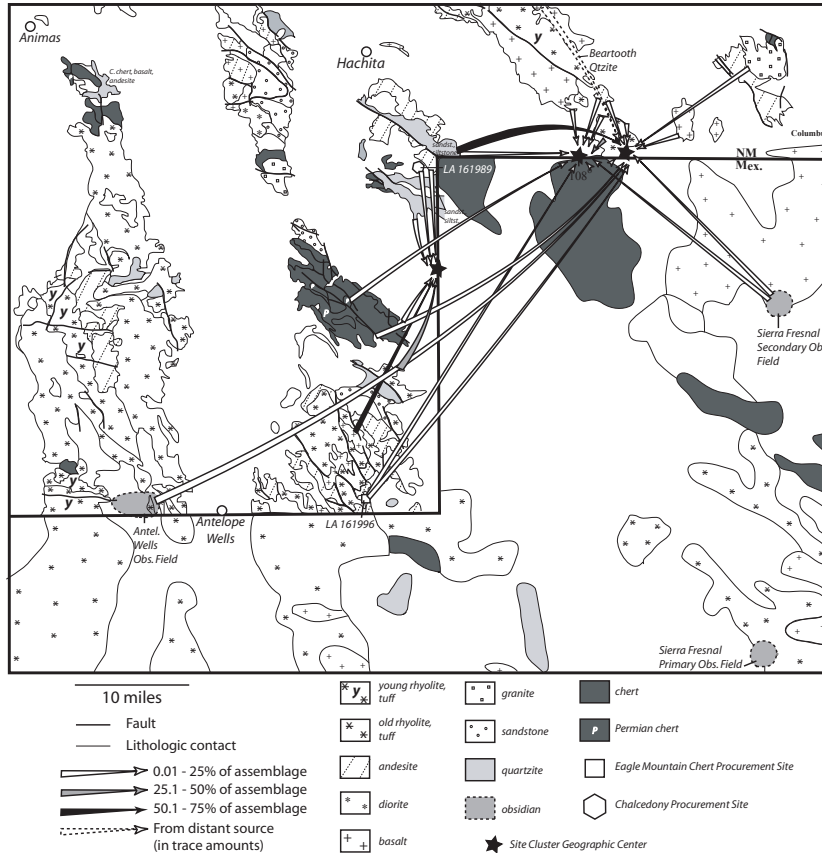


Fig. 5. Geologic map of the Boot Heel of southwestern New Mexico with locations of site clusters and minimum distance movement of geologic materials for sites identified as Protohistoric.

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