1 SHORT COMMUNICATION

2 Grassland fire effect on soil organic carbon reservoirs in a semiarid environment

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10 Abstract

11 The aim of this work was to investigate the effect of an experimental fire, used for grassland management, 12 on soil organic carbon (SOC) stocks. The study was carried out on Hyparrhenia hirta (L.) Stapf (Hh) 13 grassland and Ampelodesmos mauritanicus (Desf.) T. Durand & Schinz (Am) grasslands, located in the north 14 of Sicily. Soil samples were collected at 0-5 cm before and after experimental fire and SOC was measured. 15 During grassland fire soil surface temperature was monitored. Biomass of both grasses was analyzed in order 16 to determine dry weight and its chemical composition. The results showed that SOC varied significantly with 17 vegetation type, while it is not affected in the short period by grassland fire. Am grassland stored more SOC 18 compared with *Hh* grassland thanks to lower content in biomass of labile carbon pool. No significant 19 difference was observed in SOC before and after fire which could be caused by several factors: first, in both 20 grassland types the measured soil temperature during fire was low due to thin litter layers; second, in 21 semiarid environment higher mineralization rate results in lower soil carbon labile pool; and third, the SOC 22 stored in the finest soil fractions, physically protected, is not affected by fire.

23 Keywords: SOC, experimental fire, grassland, Mediterranean environment

25 Introduction

26 Fire is part of the Earth System and has been for millennia a tool for many societies (Pyne, 2001) and this is well-known in grasslands (Dickie and Parson, 2012). Fire is regarded as an active ecological agent able to 27 28 mobilize nutrients and restore soil fertility (Snyman, 2003) but, also as a primary cause of soil degradation 29 due to nutrients loss for volatilization, leaching and erosion, especially in severe wildfires. It is, in fact, 30 considered a major disturbance in many ecosystems, which lead to important shifts in the soil properties and 31 vegetation (Certini, 2005; Granged et al., 2011a). One of the most common effects of fire is the alteration in 32 composition and amount of soil organic matter (Knicker, 2007, Terefe et al., 2008). Several studies recorded 33 a decrease (Fernández et al., 1997; Novara et al., 2011) in soil organic carbon (SOC) after fire, while results 34 of other studies showed no significantly changing or even increase of previous SOC content (Kavdir et al., 35 2005). These discrepancies occur due to the large amount of controlling factors and therefore the effect of 36 fire is highly variable in space and time. Among these factors, fire intensity, fire severity, fire regimen, type 37 of burned vegetation, connectivity, distribution of fuel on soil surface, type of ash produced and dispersion, 38 topography, soil properties, aspect, regional climate and meteorological conditions in the immediate period 39 after the fire play a key role to determine SOC alteration and accumulation in soils (Certini, 2005; Pereira et 40 al. 2010; Pereira et al., 2013).

41 In semiarid areas fire is one of the common management tools used by shepherds to enhance pasture regrowth and this is found in many other grass covered soil in the Earth of any latitude (Wang et al., 2005; Li 42 43 et al., 2012; Busso et al., 2012). In fact, the recovery of vegetation canopy after fire in the Mediterranean 44 area can be quite rapid due to adaption of plant communities to the disturbances caused by fire as observed in 45 several studies (Trabaud, 1981; Oba, 1990; Woube, 1998; Barberis et al., 2003; Pausas and Verdú 2005). It is 46 known, moreover, that fire is considered an important factor for arid and semiarid grasslands because it 47 avoids invasion of trees and shrubs with implications on soil carbon storage (Briggs et al., 2005). Despite the 48 importance of fire on grassland ecosystems (Bond et al., 2005), its impact on SOC is not well understood in 49 the immediate period after the fire in the Mediterranean grasslands (Snyman, 2003). The aim of this work is 50 to quantify SOC stock change as a result of an experimental fire of two of the most widespread types of 51 Mediterranean grasslands (Brullo et al., 2010; Díez-Garretas and Asensi, 1999) and, therefore, to establish if

this practice could be used as a sustainably management tool for grazing recovery (Álvarez-Martínez et al.,
2013).

54 Materials and Methods

The field studies were carried out in the province of Palermo, Sicily (Italy) (350 m a.s.l.) (Fig. 1). Local soil type is an *Eutric Cambisol* according to WRB (WRB, 2006) with sand and clay contents of 18% and 46%, respectively. The climate is Mediterranean, with mean annual rainfall of 580 mm and yearly average temperature of 16 °C.

59 An experimental fire was conducted on July and September 2009 on five (replicas) delimited square areas 60 (50×50 cm) in two different grassland types, dominated by Hyparrhenia hirta (L.) Stapf (Hh) and 61 Ampelodesmos mauritanicus (Desf.) T. Durand & Schinz (Am). Each sampling square was about 2m distant 62 from the neighbor square. In order to simulate a natural wildfire, burning was allowed to take its natural 63 course until it extinguished itself. The fire was generated with a match, starting from leeward in each plot. 64 Soil surface temperature during the burning was measured using a thermocouple system (type K Inconel 600 65 insulated). In each selected area three soil samples were collected at 0-5 cm depth before and immediately after fire. On three one meter square in both grasslands (dominated by Hyparrhenia hirta or Ampelodesmos 66 67 mauritanicus) all plants were cut, oven dried for 3-4 days at 60-65 °C, and weighted. SOC content was 68 measured using a CHN-Elemental Analyzer. For the δ 13C analysis, an EA-IRMS (elemental analyzer 69 isotope ratio mass spectrometry) was used. The International Atomic Energy Agency (IAEA), Vienna, 70 distribute IAEA-CH-6 as a reference standard material. The results of the isotope analysis are expressed as a 71 δ value (‰) relative to the international Pee Dee Belemnite standard as follows:

72

$$\delta(\%_0) = \frac{Rs - Rst}{Rst} * 1000$$

73

74 where $\delta = \delta 13C$, R = 13C/12C, s = sample, and st = standard.

Dry biomass weight and its chemical composition (ADF acid detergent fiber, NDF neutral detergent fiber,
 Cellulose, Hemicellulose, Lignin, Ash) were determined on three 0.5 m² square area subsamples for each
 grassland types.

Data analysis was conducted using the SAS statistical package (SAS Inst., 2002). Normal distribution of data was verified previously to statistical data comparisons and analysis of variance (ANOVA) was conducted. Significant differences were considered at a p < 0.05.

81

82 **Results and discussion**

83

SOC ranged from 20.3 to 37.0g kg⁻¹ and from 15.4 to 32.5 g kg⁻¹ before and after experimental fire, 84 respectively, in soil covered by *Hh*, and from 32.5 to 38.2 g kg⁻¹ and from 38.3 to 49.1 g kg⁻¹ before and after 85 experimental fire, respectively, in soil covered by Am. The experimental fire did not have significant 86 differences in SOC in both grassland types (Fig. 2). Similarly to SOC results, δ^{13} C was not affected 87 significantly by fire. The average by time of δ^{13} C values measured in *Hh* grassland were -25.418 ± 0.25‰ 88 89 and $-25.161 \pm 0.40\%$ in soil sampled before and after fire, respectively; while in Am grassland were -26.87390 \pm 0.16‰ and -26.98 \pm 0.31‰ before and after fire, respectively. Our results are in agreement with similar 91 observations reported by other authors (Granged et al., 2011b) who found no change in SOC content before 92 and after prescribed fire. The experimental fire has a moderate fire severity, similar to prescribed fire 93 described by Granged et al. (2011b). The time of combustion was 12 ± 2 minutes and 7 ± 1 minutes for *Hh* and 94 Am, respectively (Fig. 3). The maximum temperature measured at soil surface was around 480 °C in both 95 grasslands. Temperatures over 200 °C persisted for 5 minutes and 3 minutes for Hh and Am, respectively. 96 The burning time and intensity was low due to low amount of fuel in both grasslands. Mediterranean 97 environmental conditions involve high organic matter mineralization rates and, thus, negligible amounts of 98 litter biomass stock. The low temperatures registered during low severity fires does not have important 99 effects on SOC stock (Úbeda et al., 2005). The loss of organic carbon by burning can occur even at relatively 100 low temperatures such as 200 °C, but total combustion is only observed at high temperatures 450-500°C (De 101 Bano et al., 1998). When comparing the two grasslands, the SOC amount and the effect of fire on SOC stock was contrasted. The lower SOC content was measured under *Hh* grassland, which also recorded the lower 102 biomass yield. The above ground biomass estimated is 4.76 Mg ha⁻¹ and 11.60 Mg ha⁻¹ of dry matter for *Hh* 103 104 and Am grassland, respectively.

105 Even if the SOC change before and after fire was not statistically significant, after fire SOC content 106 decreased of 11.5% in Hh and increased of 27.9% in Am grassland. The increase of SOC after fire could 107 occur due to external inputs of charred material and ash, as commonly is observed in low severity fires due 108 fuel and organic matter incomplete combustion. In particular, the burned material returns to soil as particles 109 smaller than 2 mm in the form of ash, which are mixed in the top horizon, and which cause a net increase of 110 SOC content (González-Pérez et al., 2004). The reason for the slight SOC increase after fire only in Am 111 grassland may depend on different characteristics of the two considered grasses. Firstly, Am biomass 112 contains more lignin and cellulose than *Hh* biomass (Table 1), and, thus, more recalcitrant compounds that 113 under low temperature do not completely volatilize. Secondly, Am has a densely caespitose habit: this feature 114 impedes a complete burning and favors the retention of not completely burnt plant residues. The ash of *Hh* 115 is, instead, lighter and quickly eroded by wind (Cerdà and Doerr, 2007). This is clear evidence that Hh 116 grassland burned at higher severity, despite the similar temperatures observed. Previous studies observed that 117 fire severity is different according the burned specie (Pereira et al., 2011). Thirdly, biomass of Am contains 118 siliceous compounds that obstruct burning.

119 Conclusions

Data here reported confirm that the use of experimental fire to favour plant recovery in *Hh* and *Am* grassland does not affect SOC stock, even if these grasslands did not burn for many years, therefore it is possible to adopt the system of controlled burning to maintain grassland formations.

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124 Acknowledgement

125 This research was financially supported by the MIUR through the PRIN project "CARBOTREES".

126 **References**

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Table caption

- Table 1. Biomass composition (% of dry biomass) of *Hyparrhenia Hirta* (Hh) and *Ampelodesmos mauritanicus* (Am). Values in parenthesis are standard deviations. Abbreviations: ADF = acid detergent fiber, NDF = neutral detergent fiber.

Grassland	ADF	Cellulose	NDF	Hemicellulose	Ash	Aboveground	С
						biomass	Biomass
						$(Mg ha^{-1})$	$(g kg^{-1})$
Am	6.91	37.72	73.03	23.99	4.02	4.76	43.8
	((0.58)	(1.58)	(2.65)	(1.32)	(1.10)		
Hh	5.98	34.00	72.01	28.26	4.34	11.60	45.8
	(0.68)	(1.20)	(1.53)	(1.76)	(1.49)		

212 Figure Caption

- Fig. Localization of the study area (black point) in Sicily (rectangle in inset) and with respect to the next
- urban settlement Cefalù. Grey scale represents altitudinal gradient (1 shade of grey = 200 m).



- 220 Fig. 2. Soil organic carbon before and after fire in Hyparrhenia Hirta (Hh) and Ampelodesmos mauritanicus
- 221 (Am) grassland.



Figure 3- Soil temperature during fire under *Hyparrhenia Hirta* (Hh) (blu line) and *Ampelodesmos mauritanicus* (Am) (red line) grassland

