Spatial variability of soil properties and soil erodibility in the Alqueva dam watershed, Portugal

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Abstract

The aim of this work is to investigate how the spatial variability of soil properties and soil erodibility (K factor) were affected by the changes in land use allowed by irrigation with water from a reservoir in a semiarid area. To this, three areas representative of different land uses (agroforestry grassland, Lucerne crop and olive orchard) were studied within a 900ha farm. The interrelationships between variables were analyzed by multivariate techniques and extrapolated using geostatistics. The results confirmed differences between land uses for all properties analyzed, which was explained mainly by the existence of diverse management practices (tillage, fertilization and irrigation), vegetation cover and local soil characteristics. Soil organic matter, clay and nitrogen content decreased significantly, while K factor increased with intensive cultivation. The HJ-biplot methodology was used to represent the variation of soil erodibility properties grouped in land uses. Native grassland was the least correlated with the other land uses. K factor demonstrated high correlation mainly with very fine sand and silt. The maps produced with geostatistics were crucial to understand the current spatial variability in the Alqueva region. Facing the intensification of land-use conversion, a sustainable management is needed to introduce protective measures to control soil erosion.

Keywords: Geostatistics, HJ-Biplot, land use management, spatial variability, soil erodibility.

Introduction

Soil erosion is a significant economic and environmental problem worldwide as a driving force affecting landscape (Zhao *et al.*, 2013). It is a very dynamic and complex process, characterized by the decline of soil quality and productivity, as it causes the loss of topsoil and increases runoff (Lal, 2001; Yang *et al.*, 2003). Furthermore, soil erosion often causes negative downstream impacts, such as the sedimentation in rivers and reservoirs decreasing their storage volume as well as lifespan (Pandey *et al.*, 2007; Haregeweyn *et al.*, 2013).

One of the main cause of soil loss intensification around the world is associated with land-use change (Leh *et al.*, 2013). The relationship between different land use and soil susceptibility to erosion has attracted the interest of a variety of researchers (Yang *et al.*, 2003; Cerdà and Doerr, 2007; Blavet *et al.*, 2009; Biro et al., 2013; Wang & Shao, 2013), which have shown the impact of changes on vegetation cover and agricultural practices on soil properties and therefore in overland flow. Generally, cultivated lands experience the highest erosion yield (Cerdà *et al.*, 2009; Mandal & Sharda, 2013). In the Mediterranean regions, in combination with these anthropogenic factors, the climate change has amplified the concerning about soil erosion since it is expected the increase of dry periods followed by heavy storms with concentrated rainfall (Nunes *et al.*, 2009).

Some models have been developed to predict soil loss and sediment delivery. The Revised Universal Soil Loss Equation (RUSLE) is the most used empirical equation for modeling annual soil loss from agricultural watersheds (Renard *et al.*, 1997). The susceptibility of soil erosion and land degradation depends largely on various inherent soil properties, namely chemical, physical, biological and mineralogical properties (Cambardella *et al.*, 1994; Pérez-Rodríguez *et al.*, 2007). However, according to the RUSLE model only some of the soil's properties define soil erodibility (K factor), such as particle-size composition, the content of organic matter, soil structure and permeability. Therefore, the K factor is the most used and is an important index to measure soil susceptibility to erosion (Panagopoulos and Antunes, 2008).

Spatial variability in soils occurs naturally as a result of complex interactions between geology, topography and climate. Moreover the spatial variability of soil properties, which influence soil susceptibility to erosion, is highly related with anthropogenic factors particularly in cultivated lands (Paz-González *et al.*, 2000; Wang & Shao, 2013). Then, information on the spatial variability and the interactions between soil properties is essential for understanding the ecosystem processes and planning sustainable soil management alternatives for specific land-uses (Pérez-Rodríguez *et al.* 2007; Ziadat & Tamimeh, 2013).

Classical statistics and geostatistics methods have been widely applied on studies about spatial distribution of soil properties (Pérez-Rodríguez *et al.*, 2007, Tesfahunegn *et al.*, 2011). Geostatistical techniques based on predictions and simulations have been used to describe areas where predicted information is established by a limited number of samples (Goovaerts, 1997). Geostatistics provides tools for analyzing spatial variability structure and distribution of soil properties and evaluating their dependence (Panagopoulos et al., 2014).

The Biplot methodology provides an added value for analyzing spatial variability of soil properties. This multivariate statistical technique allows the graphical representation of a large data matrix (Gabriel, 1971), whereby it is possible to interpret the relations between individuals (samples) and between variables, as well as between both. Biplot can also indicate clustering of units with close characteristics, showing inter-unit distances as well as displaying variances and correlations of the variables (Gallego-Álvarez *et al.*, 2013). The HJ-Biplot permits not only the analysis of the behavior by sample but also the determination of which variable is responsible for such behavior (Garcia-Talegon *et al.*, 1999), allowing a visual appraisal to establish relations between soil properties and land uses.

The construction of the Alqueva dam in a semiarid area of South Portugal created one of the largest artificial lakes in Europe. Taking advantage of water availability from the reservoir, this Mediterranean region has been subjected to land-use conversion from the native Montado grassland to intensive agricultural uses. Land-use conversion from the native ecosystem to agriculture may alter physical, chemical and biological soil properties which consequently may increase soil erosion and siltation in the reservoir. Soil erosion in the area has to be carefully evaluated in order to take sustainable soil management measures. Therefore, the aim of this study was to evaluate the effects of cultivation practices on some chemical and physical soil properties and on soil erodibility (K factor on RUSLE), and to characterize their spatial variability using geostatistics and HJ-Biplot methodology.

Material and methods

1.1. Study Area

Localized in the semiarid Alentejo region of Portugal, at the Guadiana River, the Alqueva reservoir (8º30' W, 38º30' N) covers an area of 250 km^2, and the capacity of the reservoir is 4.15 km^3. The main arguments for the implementation of what is considered the largest artificial lake in Europe were based on the need to combat the growing effects of desertification and to prevent the annual and monthly fluctuations in precipitation. One of the main goal of the Alqueva Multipurpose Project was the implementation of 120,000 hectares of new irrigated land in the Alentejo. The Alentejo region, covering an area of 27,000 km^2 is considered one of the most depressed regions of the European Union and characterized by a Mediterranean climate with very hot and dry summers and mild winters. The average temperature ranges from 24 to 28°C in hot months (July/August) and from 8 to 11°C in cold months (December/January). The average annual precipitation at the nearest meteorological station, for the last 30 years, is 517.2 mm. The region is affected by intense dry periods followed by heavy, erosive rains concentrated in the autumn season.

The study experimental site (farm “Herdade dos Gregos”), located in the surrounding area of the reservoir (Figure 1), is a private property with 900 ha. The landscape is characterized by its hilly topography with significant altitude variations (mainly between 100 and 250 meters). The bedrock of the study area is rocky and according to World Reference Base for Soil Resources (FAO, 2006), the two types of soil in this area are: Haplic luvisols (LVha) and Lithic leptosols (LPli). This farm was selected to include a diversity of land uses, including native Montado grassland and more intensive land-uses, with irrigation, namely Olive tree orchard and Lucerne cultivation. Direct pumping from Alqueva reservoir is done in this private property since it is near the reservoir.

The typical landscape in the Alentejo region is the Montado native grassland, an agrosilvopastoral system characterized by savannah-like, low density woodlands with evergreen holm oaks (*Quercus ilex*). For that reason, an area of the Montado grassland (20.7 ha), used as a permanent pasture for the cattle, was selected for this study. This small area is located in the high altitudes of the “Herdade dos Gregos” (from 200 to 240 m) with a slope that varies from 1.4 to 20.9 %. Tillage (at about 15 cm depths) was done only once every 10 years to decrease shrub competition (the most recent one was four years before the study implementation), and the soil is not subjected to any fertilizer. Four years before the study implementation, there was a fire on this agrosilvopastoral area of the farm.

Taking advantage of the water availability, another land use (with 33.5 ha) is an irrigation area (Pivot Sprinkler Irrigation System) on which Lucerne (*Medicago sativa*) is sown four times a year. Lucerne, once dried, is nutritional for cattle, and it incorporates nitrogen in the soil. In this area, conventional tillage is used, involving multiple aspects: plough (about 20 cm depth) in fall, fallowing cultivator (about 15 cm depths) and disc harrow (about 10 cm depths) subsequent to soil tillage. Inorganic fertilizers were applied to the cultivated field at a rate of 100 kg ha^{-1} NPK. This land use is placed in the midland (194-220 m), and the slope varies from 0 to 9%.

Other irrigated land use consists of an Olive tree plantation (57.5 ha), which is done in strips. This cultivation has a drip irrigation system, is fertilized once every two years and is ploughed once a year to decrease weed competition. The Olive orchard is located in the low elevations of the farm (150-186 m), and it is on the side of the reservoir (Figure 1). The slope varies from 0 to 14.2%.

1.2. Soil sampling and laboratory analysis

Since the objective was to study the relation between soil properties and K factor from RUSLE, the soil samples were collected from 0 to 20 cm depth, according to Renard *et al.* (1997). In order to predict variations in short distances, 25, 27 and 52 soil samples were randomly collected respectively in Montado, Lucerne and the Olive orchard (see Figure 1). Samples were air-dried and then dried for about 6 hours at 40ºC on a ventilated oven, and they were passed through a 2 mm sieve to remove rocks and gravels. The particle-size distribution was determined by the Bouyoucos hydrometer method (Bouyoucos, 1936). Soil organic matter content was determined using the Walkley & Black (1934) method, a wet oxidation procedure. The soil's total nitrogen content was determined according to Kjeldhal digestion, distillation and the titration method (Bremmer & Mulvaney, 1982). Soil pH and electrical conductivity were measured with glass electrode in a 1:2.5 soil/water suspension (Watson & Brown, 2011).

1.3. Soil erodibility factor

Soil erodibility factor (*K*) (Mg ha h ha^{-1} MJ^{-1} mm^{-1}) was estimated using soil property values, such as particle-size composition, content of organic matter, soil structure and permeability, in the 104 samples points described above. This factor represents the soil-loss rate per erosion index unit for a specified soil as measured on a standard plot (Renard *et al.*, 1997). An algebraic approximation of the nomograph (Equation 1) was used to estimate *K* factor (Renard *et al.*, 1997):

$$K = [2.1 \times 10^{-4}(12 - OM) \times M^{1.14} + 3.25(s - 2) + 2.5(p - 3)]/759 \quad \textbf{(1)}$$

where *OM* is the percentage of organic matter, *s* is soil structure class, *p* is permeability class, and *M* is the product of the percentage of modified silt (silt particles and very fine sand) or the 0.002–0.1 mm size fraction and the sum of the percentage of silt and percentage of sand. *K* is expressed with SI units of Mg ha h ha^{-1} $MJ^{-1}mm^{-1}$. To estimate the permeability the field-saturated hydraulic conductivity was measured in the field using a double-ring infiltrometer (6 site-measurements per land-use, each one with 5 repetitions). Permeability class and soil structure class were defined in accordance with Renard *et al.* (1997).

1.4. Statistical and Geostatistical Analysis

Data were subjected to classical analysis using SPSS 17.0 software to obtain descriptive statistics, namely the mean, minimum and maximum, standard deviation (SD), coefficient of variation (CV) and skewness of each parameter.

Soil data were introduced in the ArcGIS environment and geostatistical analysis were performed using Geostatistical Analyst Tool, in other to examine spatial distribution of soil properties. Prior to geostatistics to obtain prediction maps, a preliminary analysis of data were done to check data normality and global directional trends. Skewness is the most common statistic parameter to identify a normal distribution that is confirmed with skewness values varying form − 1 to + 1. Data transformation to normal distribution was necessary for some soil properties and geostatistical analyst tools were used (log or box-cox method). Trend analysis was performed to examine the presence of any global directional trend in our data, an overriding process that affects all measurements in a deterministic way (nonrandom). So, when necessary, the trend removal was done using Geostatistical Analyst tools to more accurately model the variation (Panagopoulos et al., 2006).

The geostatistical methodology is based on the creation of a semivariogram (SV), a graphical representation (Equation 2) that describe how samples are related to each other in space and it is based on:

$$\gamma(h) = 1/2N(h) \times \sum[Z_i - Z_{(i+h)}]^2 \quad \textbf{(2)}$$

where *γ(h)* is the variance (the most related samples have lower values of variance), *N(h)* is the number of samples that can be grouped using vector *h*, *Zi* represents the value of the sample, and *Zi+h* is the value of another sample located at a distance ||*h*|| from the initial sample *Zi* (Chiles and Delfiner, 1999).

Ordinary Kriging (OK) was selected as geostatistic method. OK is considered one of the most accurate interpolation technique which assumes that variables close in space tend to be more similar than those further away (Goovaerts, 1999).

Using the Geostatistical Analyst Tool (ArcGIS) and selecting the OK methods, a semivariogram was created for each measured property. In the kriging method different semivariogram models can be used (*e.g.* spherical, exponential) and the selection is usually performed by employing the cross-validation technique, which permits the evaluation of the prediction accuracy. Cross-validation was executed to investigate the prediction performances through the statistical values, as the mean error [ME] or root-mean-square standardized error [RMSSE]), which results from comparing the estimated semivariogram values and real observed values. Additional semivariogram parameters were analyzed to better understand the spatial structure and dependence of each variable. Nugget is the variance at distance zero and reflects the sampling error. Sill is the semivariance value at which the semivariogram reaches the upper bound and flattens out after its initial increase; it is the variance in which the samples are no longer spatially related at the study area.

Once cross-validation process was completed, interpolation maps of spatial distribution, for each soil variable, were produced according the semivariogram model selected, in the ArcGIS software.

1.5. HJ-Biplot

HJ-Biplot represents a matrix, without assumptions related to its probabilistic distribution, permitting a graphic representation of the geometric data structure, representing the dataset (samples and variables) variability. The prefix "bi" is due to a simultaneous representation of the matrix rows and columns, searching for the maximum representation quality possible, at the same scale (Martín-Rodríguez *et al.*, 2002; González-Cabrera *et al.*, 2006; Gallego-Álvarez *et al.*, 2013).

A data matrix X suffers a factorization to reduce its dimensionality through single value decomposition, the algebraic base of biplot representation (Eq. 3) (Gabriel, 1971).

$$X_{(n\times p)} = U_{(n\times r)}\, \Lambda_{(r\times r)}\, V'_{(r\times p)} \qquad \textbf{(3)}$$

where $\Lambda_{(r\times r)}$ is a diagonal ($\lambda_1, \lambda_2, \ldots, \lambda_r$) corresponding to the r eigenvalues of XX' or X'X, $U_{(n\times r)}$ is an orthogonal matrix whose columns are the eigenvectors of XX', and $V'_{(r\times p)}$ is an orthogonal matrix whose columns are the eigenvectors of X'X.

With the *MultiBiplot software*, developed by the University of Salamanca (Vicente Villardón, 2014), an HJ-Biplot was used to determine the relation between soil properties, between land uses, and the correlations between both (soil properties and land uses), thereby defining patterns and clustering the samples in groups.

On the HJ-Biplot graphic representation, the points represent individuals (samples), and the vectors represent variables (in this case, chemical and physical soil properties). To interpret and discuss the graphs obtained with this methodology it's essential to be aware of (Gallego-Álvarez *et al.*, 2013):

- The distance between points represents the variability and can be interpreted as similarity or dissimilarity, i.e. the close samples have similar behaviors;
- the angle formed by variable vectors is interpreted as correlation, i.e. small angles between variables represent similar behaviors with high positive correlations, and the

obtuse angles that are almost a straight angle are associated with variables with high negative correlations; i.e. the cosine value of the angles represents the correlation between variables.

- The proximity of individual points and variable vectors means high preponderance; in other words the closer a point is to a variable vector, the more important this sample is to explain this variable;
- The length of the vector represents the variable's variability and the longer is the vector the higher is this variability.

2. RESULTS AND DISCUSSION

2.1. Descriptive Statistics

The descriptive statistics of soil properties are given in the first part of Table 1. All measured parameters varied considerably within the areas (different land uses) as indicated by the coefficient of variation (varies from 4.2 to 70.2%). Nitrogen (N) and organic matter (OM) show the highest variation values, especially for cultivated fields (Lucerne cultivation and Olive orchard), that can be explained with the lack of homogeneous fertilization or tillage practices applied to soil in these areas.

The skewness results, which vary from -1.48 to 3.54 in this study, indicated that some soil properties of the different uses were not normally distributed, especially OM and N. The principal reason for some soil properties having non-normally distributions may be related with soil management practices (Tesfahunegn *et al.*, 2011). As it was already mentioned data was transformed to normal distribution when necessary (see Table 1).

These mean results show significant differences between land uses for all the properties analyzed. From the particle size distribution reported in Table 1, the soils are mostly sandy loam, formed mainly of sand, followed by silt and low quantities of clay. However, there are some differences between land use areas that can be explained by soil type. The Lithic leptosols (LPli) soils are characterized by a thin layer (about 10 cm), in that case upon a schist rock, justifying the higher clay content at the Montado grassland. The Haplic luvisols (LVha) soils in the Lucerne cultivation and the Olive orchard are characterized by a loam or sandy loam layer (first 20 cm) with good drainage over clay-enriched subsoil (upon a basic crystalline rock), explaining the lower values of clay and fine sand, especially in the Olive orchard. Despite the same soil type, soil texture is different between Lucerne and Olive orchard that can be justified by land-use. The Lucerne is a more intensive cultivation (intensive irrigation, tillage and continuous cultivation, fertilizers and lime application), conditions that promote changes in the soil weathering and moisture, and consequently on soil texture (Yimer *et al*., 2008). On the other hand the soil between olive trees is kept without vegetation for most of the year and it can explain the clay drainage to a sub-layer.

Montado shows the highest content of OM (5.22%), whereas Lucerne and Olive fields show the lowest values (with 2.08% and 2.10%, respectively). Other studies suggest that OM is higher in no-tillage soils compared to minimum tillage that increases aeration (Celik, 2005). Tillage mixes the subsoil with topsoil; after soil erosion, the nutrients are easily leached and the surface becomes poor in nutrients (Al-Kaisi & Licht, 2005). As for OM, the highest value of N nutrient occurs in the Montado (0.19%) and the lowest values in Lucerne (0.11%) and the Olive orchard (0.10%), which is related to the tillage practice that is frequently employed in these last two land uses, while in the Montado grassland the cattle enriches the soil.

Soil EC values (Table 1) were similar when comparing the Montado grassland (0.100 dS/m) and the Lucerne field (0.107 dS/m); they were slightly higher in the Olive orchard (0.182 dS/m) but not enough to raise salinity problems. Usually, the addition of fertilizers (that happens on Lucerne and the Olive orchard) can cause high EC due to the percentage of the salts, which are leached by water irrigation (higher in the Lucerne field).

The soil pH was significantly higher in the Lucerne cultivated land (7.1) compared to the Montado grassland (5.9) or in the Olive tree orchard (5.5) (Table 1). The soil pH in the Lucerne was greater due to lime application to increment the soil pH in that area. Lucerne's optimum pH for production is between 6.5 and 7.2, and lime application has been found to produce a significant improvement in nodulation of Lucerne (both number and dry weight of nodules per plant) (Grewal & Williams, 2001).

Saturated Hydraulic Conductivity (HC) values were greater in the Lucerne area (5.95 cm/h), slightly lower in the Montado grassland (4.56 cm/h) and lowest in the Olive orchard (2.60 cm/h). The lower permeability in the Olive orchard can be explained by the clay-enriched subsoil or soil crust problems, and may explain the higher values of EC, i.e. the greater concentration of salts. Also it can be explained by the frequency of tillage in the different land uses because aggregate stability and water infiltration rate are higher in soils subjected to limited tillage systems (Alvarez & Steinbach, 2009).

As a result, K factor was different for the typical land use, Montado grassland, compared to the Lucerne cultivation and the Olive orchard. The values increased with the intensification of the cultivation field, with the lowest values for Montado grassland (0.021 Mg ha h ha^{-1} MJ^{-1} mm^{-1}) and the highest for the Lucerne cultivation (0.039 Mg ha h ha^{-1} MJ^{-1} mm^{-1}) and the Olive orchard (0.038 Mg ha h ha^{-1} MJ^{-1} mm^{-1}). Other studies had similar results, showing that the removal of permanent vegetation, the loss of OM and the reduction of aggregation, caused by intensive cultivation, contribute to decrease K factor (Celik, 2005).

2.2. Spatial dependence of soil properties

Model selection for each soil property was based on the nugget, sill, mean error (ME) and the root-mean-square standardized error (RMSSE) presented in the second part of Table 1 (Geostatistics).

Nugget is low in most soil properties studied, implying strong spatial dependence. The nugget to sill ratio is used to define spatial dependence of soil properties: if the ratio was <0.25, there is strong spatial dependence; if it was 0.25 to 0.75, there is moderate spatially dependence; and if the ratio was >0.75, spatial dependence is weak (Cambardella *et al.*, 1994). As shown in Table 1 the ratio values indicate the presence of high to moderate spatial dependence for all soil parameters (values between 0 and 0.64). In general, there is stronger spatial dependence in Montado (low nugget to sill ratio), which can be explained with the non-existence of extrinsic factors, such as management cultivation practices, that influence soil properties, and soil is left as it is for permanent pasture.

Cross-validation facilitated the selection of the best-fit semivariogram for an interpolation map, which could provide the most accurate predictions. Closer values of the ME to zero, and closer values of the RMSS to 1 suggested that the prediction values were close to measured values (Wackkernagel, 1995). Most of the soil properties were best fitted with an Exponential model, particularly in the Montado area and Olive orchard, whereas in Lucerne the semivariogram models Gaussian, Circular and Stable were used.

2.3. Spatial distribution

The interpolation maps obtained with geostatistics are useful to better understand spatial variability and its influences. The variability of spatial soil properties can be influenced by natural factors (as particle-size composition and topography) and anthropogenic factors (as land cover or management practices) (Tesfahunegn *et al*., 2011). Sometimes, the effect of some factors is at least one order of magnitude greater (as topography or soil type) than the land-use. So, as mentioned trend analysis was performed to study the existence of directional trends caused by these factors with large scale of variation, and it is shown in the Figure 2. Global trend exists if a curve that is not flat (i.e., a polynomial equation) can be fitted to the data (for example for total nitrogen (N) in Montado or very fine sand (VFS) in olive orchard). These trends were identified for part of the soil properties and for different land-uses (Figure 2). The strongest influence of directional trend was identified from southeast to the northwest, which can be associated with the topography (Figure 1) since the altitudes increase according these direction. So, trend removal is crucial to create more accurate prediction maps in order to justify an assumption of normality.

The interpolation maps for some studied soil properties are shown in Figure 3. Looking at the VFS distribution, it was noticed that the higher fractions of these particles (Figure 3) were measured on low altitudes or flat slopes such as the valley (see elevation on Figure 1). This can be explained by erosion-deposition processes because these particles are easily detached and transported by water.

The highest percentages of N and OM were found on Montado, as discussed previously. These two properties present similar distributions for all land uses. The nitrogen existing in the soil is mostly organic, and the inorganic forms (ammonium and nitrate) are easily leached or assimilated by plants. So, when OM breaks down due to mineralization, the N fraction decreases (Varennes, 2003). There were higher values in Montado because the soil is not frequently tilled as it is in the other land uses. In the Lucerne cultivation and the Olive orchard, the variation of OM and N can be explained by inadequate management practices (*e.g.* inadequate fertilization rates, tillage, irrigation rates, seed rates, etc.).

Figure 3 illustrates the interpolation map for K factor which was estimated through the Wischmeier nomograph (Eq. 1). The values vary from 0.006 to 0.061 Mg ha h ha^{-1} MJ^{-1} mm^{-1}, and the prediction map show the highest values for Lucerne and the Olive orchard, especially where the soils have more silt and very fine sand (VFS), along with less OM and N (see HJ-Biplot). In the surrounding area of the reservoir, the types of soil differ with the topography and land use; therefore, the knowledge of soil properties is fundamental when facing the intensification of cultivation that could increase K factor. These intensive practices decrease OM in soils, making them poor and vulnerable to the soil erosion process.

Looking for natural vs anthropogenic impact on the K factor, for each land-use, it's evident that in the Montado the spatial variability is mainly associated with natural (intrinsic) factors (as texture), being soil properties and erodibility distribution more homogenous. In the Lucerne and Olive orchard the spatial variability is more dependent from not homogenous anthropogenic causes such as fertilization and irrigation rates and tillage/plough processes.

2.4. HJ-Biplot

The HJ-Biplot representation matrix of soil properties is showed in Figure 4. It was observed that the dominant axis (axis 1) takes 35.83% of the total inertia (information) of the system. With both dimensions, an accumulative inertia of 61.04% was achieved. Regarding this graphic representation, it was observed that samples were grouped according to the land use. The Montado samples were close to OM, N and Clay vectors, showing their preponderance to be a characterization of these variables. The Lucerne samples were important to describe the pH and Silt content. On the other hand the Olive samples were more disperse but related to EC, Permeability class, Sand, VFS and K.

The variables demonstrating a more positive correlation between them were OM and N, as previously noticed. Clay and Silt were also positively correlated, but negatively correlated with sand as expected, because soils with more sand have less clay and/or silt.

Through the matrix representation it was detected that soils with more sand have higher EC (Olive orchard), although EC normally increases with the percentage of clay. This may be explained by the addition of fertilizers, as previously discussed, that can contribute to an EC increase. These results for EC show low variability between land uses, revealing a low cation exchange capacity (CEC) of these soils. This is frequently caused by intensive soil mobilization (Paz-González *et al.*, 2000).

Permeability class increases as the HC_{sat} decreases, as defined by Renard *et al.* (1997). So, contrary to what was expected, for this study the soils with more sand (occurring in the Olive orchard) have less hydraulic conductivity (high permeability class). It can be explained by a clay-enriched sub-layer under the sandy loam layer or/and by the soil compaction/degradation processes. The soil compaction and degradation can be related to repeated plow operations to reduce shrubs between olive rows and irrigation (Pagliai *et al.*, 2004). This permeability decrease in the Olive orchard was correlated with the increase of K factor.

Nevertheless, the properties more positively correlated with K factor were the very fine sand (VFS) and silt; this is due to the susceptibility of these particles to erosion since they can be easily detached and transported by water (Morgan, 2005). The OM and N content were negatively correlated with K and permeability. The higher OM reduces the susceptibility of the soil to detachment and increases infiltration (Bronick & Lal, 2005). The nitrogen (N) content is not used to estimate K; however, especially for soils without fertilization, the existent N is mostly associated to OM. Nevertheless, nutrients decrease in soils that are more erodible, according to the literature (Tesfahunegn *et al.*, 2011). The clay content also shows a negative correlation with K factor, as expected (Renard *et al.*, 1997).

Figure 5 shows the hierarchical clusters representation. Using HJ-Biplot methodology and the aggregation tool *ward*, 3 clusters were obtained. The samples were grouped by land uses (that were already detected by the matrix representation, see Figure 4). *Cluster 1* is represented by a majority of samples from Lucerne, *Cluster 2* by samples from Montado and *Cluster 3* by samples from the Olive orchard. This was explained by the effect of different management practices, vegetation cover and local soil characteristics, as discussed. Some samples in each land use had different values (higher or lower than the majority) and were grouped in a different *cluster*. Identifying the location of the sample, the cause of displacement can be studied and can help to improve land management practices.

Therefore, the cluster analysis is convenient to identify the effect of different land-use and management on soil properties and consequently on soil erosion. On the other hand, the cluster analysis could support the delineation of zones according to soil properties, and subsequently according to erosion susceptibility, that could be used for site specific soil management recommendations.

3. Conclusions

This study demonstrated that the variability of soil properties and K factor is associated to land use, cultural practices (tillage type, fertilizer rates, conservation measures, etc.) and local conditions (complex topographic landscape, soil type, etc.). The K factor showed high correlation especially with organic matter, nitrogen, silt and very fine sand. Soils with intensively cultivated land use, and consequently with more tillage and irrigation, had lower organic matter and lower nitrogen content. This translates into a lower cation exchange capacity producing lower aggregate stability and, consequently, an increase of the K factor.

Therefore, in the surrounding area of the Alqueva reservoir, the ongoing change in land use and soil management practices can have a significant effect for chemical and physical soil properties. As a result, this affects the soil erodibility index, intensifying the risk of erosion. The increase of soil loss in the watershed might have a significant impact on a reservoir's ability to storage of water, reducing its lifespan.

Knowledge of soil spatial variability is fundamental for environment management and can help in the sustainable use of the resource soil. The prediction maps produced with geostatistics are an important monitoring tool, showing the exact position in the field of the specific soil properties. The HJ-Biplot methodology was demonstrated to be useful in gaining a better understanding of how soils properties were correlated and allowed not only a determination of the behavior by sample but also a conclusion as to which variable is responsible for such behavior. The simultaneous use of HJ-Biplot with geostatistics allow this information to be found on the map, which has important theoretical and practical significance for precision agriculture. Facing the intensification of cultivation in the surrounding area of the reservoir, site-specific soil management and careful land use planning are needed to take into account the spatial variability of soil properties, delineating management zones, variable fertilization management, irrigation scheduling, conservation practices and other efforts.

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References

Al-Kaisi, M.M., and Lichtm, M.A.: Soil carbon and nitrogen changes as influenced by tillage and cropping systems in some Iowa soils. *Agric. Ecos. Environ.*, **105**, 635-647, 2005.

Alvarez, R. and Steinbach, H.S.: A review of the effects of tillage systems on some soil physical properties, water content, nitrate availability and crops yield in the Argentine Pampas. *Soil Till. Res.*, **104,** 1-15, 2009.

Biro, K., Pradhan, B., Buchroithner, M. and Makeschin, F.: Land use/land cover change analysis and its impact on soil properties in the Northern part of Gadarif region, Sudan. *Land Degrad. Dev.*, 24, 90- 102, 2013.

Blavet, D., De Noni, G., Le Bissonnais, Y., Leonard, M., Maillo, L., Laurent, J.Y., Asseline, J., Leprun, J.C., Arshad. M.A. and Roose, E.: Effect of land use and management on the early stages of soil water erosion in French Mediterranean vineyards. *Soil Till. Res.*, **106,** 124-136, 2009.

Bouyoucos, G.: Directions for making mechanical analysis of soils by the hydrometer method. *Soil Sci.*, **42**, 225-228, 1936.

Bremmer, J. and Mulvaney, C.: *Nitrogen-Total*. In: Page AL, Miller RH and Keeney DR (Eds.), Methods of Soil Analysis Part 2, Chemical and Microbiological Properties. Soil Science Society of America. Madison WI, USA, pp. 595-625, 1982.

Bronick, C.J. and Lal, R.: Soil structure and management: A review. *Geoderma*, **124,** 3-22, 2005.

Cambardella, C.A., Moorman, T.B., Novak, J.M., Parkin, T.B., Karlen, D.L., Turco, R.F. and Konopka, A.E.: Field-scale variability of soil properties in Central Iowa soils. *Soil Sci. Soc. Am. J.*, **58**, 1501-1511, 1994.

Celik, I.: Land-use effects on organic matter and physical properties of soil in a southern Mediterranean highland of Turkey. *Soil Till. Res.*, **83,** 270–277, 2005.

Cerdà, A. and Doerr, S.H.: Soil wettability, runoff and erodibility of major dry-Mediterranean land use types on calcareous soils. *Hydrological Processes*, 21, 2325-2336, 2007.

Cerdà, A., Giménez-Morera, A. and Bodí, M.B.: Soil and water losses from new citrus orchards growing on sloped soils in the western Mediterranean basin. *Earth Surface Processes Landforms*, **34**, 1822-1830, 2009.

Chiles, J.P. and Delfiner, P.: Geostatistics: Modelling Spatial Uncertainty. John Wiley & Sons, New York, 1999.

FAO: *World reference base for soil resources*. A framework for international classification, correlation and communication. World Soil Resources Reports 103, Rome, 2006.

Gabriel, K.R.: The biplot graphic display of matrices with application to principal component analysis. *Biometrika*, **58**, 453-467, 1971.

Gallego-Álvarez, I., Rodríguez-Domínguez, L. and García-Rubio, R.: Analysis of the environmental issues worldwide: a study from the biplot perspective. *J. Cleaner Production*, **42,** 19-30, 2013.

Garcia-Talegon, J., Vicente, M.A., Molina-Ballesteros, E. and Vicente-Tavera, S.: Determination of the origin and evolution of building stones as a function of their chemical composition using the inertia criterion based on an HJ-biplot. *Chemical Geology*, **153**, 37-51, 1999.

González-Cabrera, J., Hidalgo-Martínez, M., Martín-Mateos, E. and Vicente-Tavera, S.: Study of the evolution of air pollution in Salamanca (Spain) along a five-year period (1994–1998) using HJ-Biplot simultaneous representation analysis. *Environ. Model. Soft.,* **21**, 61–68, 2006.

Goovaerts P.: Geostatistics for Natural Resources Evaluation. Oxford University Press, New York, 1997.

Grewal, H. and Williams, R.: *Lucerne varieties differ in their response to liming on an acid soil.* In: Proceedings of the Australian Agronomy Conference. Gosford, Australia: Australian Society of Agronomy, pp 92–95, 2001.

Haregeweyn, N., Poesen, J., Verstraeten, G., Govers, G., de Vente, J., Nyssen, J., Deckers, J. and Moeyersons, J.: Assessing the performance of a spatially distributed soil erosion and sediment delivery model (WATEM/SEDEM in Northern Ethiopia. *Land Degrad. Dev.,* **24,** 188- 204, 2013.

Lal, R.: Soil degradation by erosion. *Land Degrad. Dev.,* **12,** 519-539, 2001.

Leh, M., Bajwa, S. and Chaubey, I.: Impact of land use change on erosion risk: and integrated remote sensing geopraphic information system and modeling methodology. *Land Degrad. Dev.,* **24,** 409- 421, 2013.

Lopez-Bermudez, F., Romero-Diaz, A., and Martinez-Fernandez, J.: Vegetation and soil erosion under a semi-arid Mediterranean climate: a case study from Murcia (Spain), *Geomorphology*, **24**, 51-58, 1998.

Mandal, D. and Sharda, V.N.: Appraisal of soil erosion risk in the Eastern Himalayan region of India for soil conservation planning, *Land Degrad. Dev.*, **24**, 430-437, 2013.

Martín-Rodríguez, J., Galindo-Villardón, P., Vicente-Villardón, J.L.: Comparison and integration of subspaces from a biplot perspective, *J. Stat. Plan. Inference* **102**, 411-423, 2002.

Morgan, R.P.C.: *Soil Erosion and Conservation*. Blackwell Publishing. Oxford, 2005.

Nunes, J.P., Seixas, J., Keizer, J.J. and Ferreira, A.J.D.: Sensitivity of runoff and soil erosion to climate change in two Mediterranean watersheds: Part I. Model parameterization and evaluation. *Hydrological Processes* **23,** 1202-1211, 2009.

Pagliai, M., Vignozzi, N. and Pellegrini, S.: Soil structure and the effect of management practices, *Soil & Till. Res.*, **79,** 131–143, 2004.

Panagopoulos, T., Jesus, J., Antunes, M.D.C., Beltrao, J.: Analysis of spatial interpolation for optimising management of a salinized field cultivated with lettuce. *Eur. J. Agronomy*, 24, 1-10, 2006.

Panagopoulos, T. and Antunes, M.D.C.: Integrating geostatistics and GIS for assessment of erosion risk on low density Quercus suber woodlands of South Portugal. *Arid Land Res. Manag.*, 22, 159-177, 2008.

Panagopoulos, T., Jesus, J., Blumberg, D. and Ben-Asher, J.: Spatial variability of wheat yield as related to soil parameters in an organic field. *Comm. Soil Sci. Plant Anal.*, 45, 2018-2031, 2014.

Pandey, A., Chowdary, V.M. and Mal. B.C.: Identification of critical erosion prone in the small agricultural watershed using USLE, GIS and remote sensing. *Water Resources Manag.*, **21,** 729-746, 2007.

Paz-González, A.V.S., Vieira, S.R. and Castro, M.T.: The effect of cultivation on spatial variability of selected properties of an umbric horizon, *Geoderma,* **97**, 273-292, 2000.

Pérez-Rodríguez, R., Marques, M. and Bienes, R.: Spatial variability of the soil erodibility parameters and their reaction with the soil map at subgroup level. *Sci. Total Envir.,* **378,** 166-173, 2007.

Renard, K.G., Foster, G.R., Weesies, G.A., McCool, D.K. and Yoder, D.C.: *A Guide to conservation planning with the Revised Universal Soil Loss Equation (RUSLE).* Agricultural Handbook No 703, United States Department of Agriculture: Agricultural Regional Service. Washington, DC, 1997.

Tesfahunegn, G.B., Tamene, L. and Vlek, P.L.G.: Catchment-scale spatial variability of soil properties and implications on site-specific soil management in northern Ethiopia. *Soil Till. Res.*, **117**, 124–139, 2011.

Varennes, A.: *Produtividade dos solos e Ambiente*. Escolar Editora, Lisbon, Portugal, 2003.

Vicente-Villardón, J.: *MULTBIPLOT: A package for Multivariate Analysis using Biplot*, Salamanca: Departamento de Estadística. Universidad de Salamanca, 2014.

Wackernagel, H.: *Multivariate geostatistics: An introduction with applications*. Springer, Berlin, 1995.

Walkley, A. and Black, I.A.: An examination of Degtjareff method for determining soil organic matter and a proposed modification of the chromic acid titration method, *Soil Sci.,* **37,** 29-37, 1934.

Wang, Y.Q. and Shao, M.A.: Spatial variability of soil physical properties in a region of the loess plateau of PR China subject to wind and water erosion. *Land Degrad. Dev.,* **24,** 296- 304, 2013.

Watson, M.E. and Brown, J.R.: *pH and Lime Requirement.* In: J.R. Brown (ed.) Recommended Chemical Soil Test Procedures for the North Central Region, North Central Regional Research

Publication Number 22, Agricultural Experiment Station SB 1001, Columbia, Missouri, USA, pp 13-16, 1998.

Yang, D., Kanae, S., Oki, T., Koikel, T., Musiake, T.: Global potential soil erosion with reference to land use and climate change, *Hydrological Processes,* **17,** 2913–2928, 2003.

Yimer, F., Ledin, S. and Abdelkadir, A.: Concentrations of exchangeable bases and cation exchange capacity in soils of cropland, grazing and forest in the Bale Mountains, Ethiopia, *Forest Ecol. Manag.,* **256,** 1298–1302, 2008.

Zhao, G., Mu, X., Wen, Z., Wang, F. and Gao, P.: Soil erosion, conservation, and eco-environment changes in the Loess Plateau of China. *Land Degrad. Dev.,* **24,** 499-510, 2013.

Ziadat, F.M. and Taimeh, A.Y.: Effect of rainfall intensity, slope and land use and antecedent soil moisture on soil erosion in an arid environment. *Land Degrad. Dev.,* **24,** 582- 590, 2013.

Table 1- Descriptive statistics of soil properties and parameters of the fitted variogram models and the cross validation results.

	Classic Statistics					Geostatistics					
	Mean	**CV (%)**	**Min**	**Max**	**Skewness**	**Variogram**	**Nugget**	**Sill**	**Nugget /Sill**	**ME**	**RMSSE**
	Montado grassland (n=25)					**Montado grassland (n=25)**					
Clay (%)	17.29	37.7	5.68	29.62	0.07	Exponential	0	38.30	0.00	0.0055	1.01
Silt (%)	29.55	17.2	12.99	39.72	-0.99	Exponential	0	36.00	0.00	0.0238	1.04
Sand (%)	53.16	13.5	39.68	70.34	0.33	Pentaspherical	0	57.60	0.00	0.0223	0.99
VFS (%)	11.13	25.6	4.49	19.04	0.16	Stable	0	12.00	0.00	-0.0188	0.99
OM (%)	5.22	32.1	2.25	10.35	1.19	Exponential*	0.031	0.07	0.44	-0.0003	1.04
N (%)	0.19	43.2	0.07	0.42	1.13	Exponential*	0.056	0.17	0.32	0.0001	1.04
EC (dS/m)	0.100	38.1	55.5	217.5	1.28	Exponential*	0.012	0.13	0.09	0.5640	0.95
pH	5.90	4.2	5.38	6.30	0.01	Exponential	0	0.06	0.00	0.0022	0.99
HC_{sat} (cm/h)	4.56	42.9	1.20	7.20	-0.57	-	-	-	-	-	-
K (t ha h ha^{-1} MJ^{-1} mm^{-1})	0.021	31.4	0.006	0.039	0.43	Stable	0	0.001	0.00	0.0001	1.00
	Lucerne cultivation (n=27)					**Lucerne cultivation (n=27)**					
Clay (%)	13.29	28.8	5.65	22.28	0.32	Stable	0	15.30	0.00	0.0017	1.02
Silt (%)	33.79	26.6	8.35	47.29	-1.48	Stable*	0	44.20	0.00	0.0073	0.97
Sand (%)	52.93	17.7	39.32	79.99	1.00	Exponencial	0	92.00	0.00	0.0297	0.98
VFS (%)	15.28	37.0	2.59	25.17	-0.39	Exponencial	15.60	25.0	0.62	0.0347	1.04
OM (%)	2.08	52.8	0.45	5.44	1.21	Exponencial*	15.90	119	0.13	0.0036	0.94
N (%)	0.11	70.2	0.02	0.35	1.43	Circular*	0.10	0.52	0.20	0.0017	1.01
EC (dS/m)	0.107	45.9	40.5	205.0	0.64	Exponential	1.15	1.79	0.64	0.2240	0.96
pH	7.14	4.3	6.53	7.85	0.02	Exponencial	0.04	0.07	0.57	0.0052	1.07
HC_{sat} (cm/h)	5.95	26.7	0.65	1.30	-0.29	-	-	-	-	-	-
K (t ha h ha^{-1} MJ^{-1} mm^{-1})	0.039	21.9	0.013	0.052	-0.88	Stable	0	0.01	0.00	0.0001	1.03
	Olive tree orchard (n=52)					**Olive tree orchard (n=52)**					
Clay (%)	9.83	28.8	5.40	16.66	0.52	Stable	0	8.04	0.00	0.0001	0.99
Silt (%)	24.37	46.8	3.82	43.36	-0.41	Pentaspherical	50.00	89.80	0.55	0.0001	0.90
Sand (%)	65.81	18.2	40.6	89.66	0.21	Exponential	0	16.10	0.00	0.0002	0.91
VFS (%)	18.14	32.5	4.49	19.04	0.16	Exponencial	0.01	33.70	0.00	0.0037	1.05
OM (%)	2.10	52.8	0.62	8.35	3.54	Exponential*	0.07	0.16	0.44	-0.0006	1.02
N (%)	0.10	45.3	0.04	0.29	2.02	Exponential*	0.02	0.15	0.12	0.0028	1.10
EC (dS/m)	0.182	61.3	53.50	583.50	1.80	Exponential	0	1.4	0.00	0.6820	1.02
pH	5.48	7.6	4.30	6.21	-0.43	Exponential	0	0.21	0.00	-0.0002	0.95
HC_{sat} (cm/h)	2.60	64.9	0.00	0.67	-0.45	-	-	-	-	-	-
K (t ha h ha^{-1} MJ^{-1} mm^{-1})	0.038	33.6	0.012	0.061	-0.36	Exponencial	0.00	0.001	0.51	-0.0001	0.92

*Transformation for normal distribution.

CV – Coefficient variation; Min – minimum; Max – maximum; VFS – Very fine sand; N – Nitrogen; OM – Organic matter; EC – Electrical conductivity; HC_{sat}- Saturated hydraulic conductivity; K – Soil erodibility; ME- Mean error; RMSSE - Root-mean-square standardized error

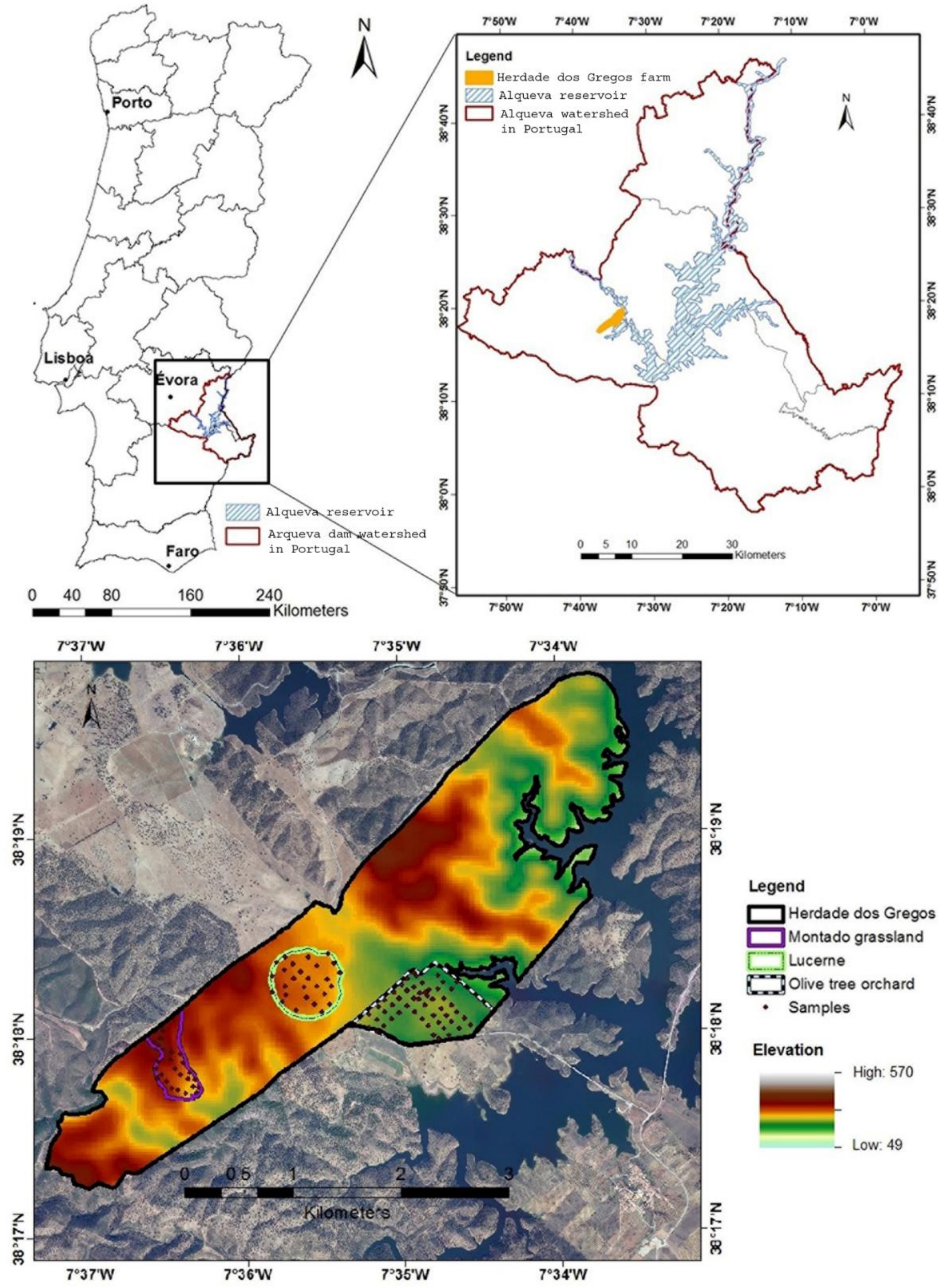

Figure 1 – Location of the study area at the Alqueva dam watershed in Portugal.

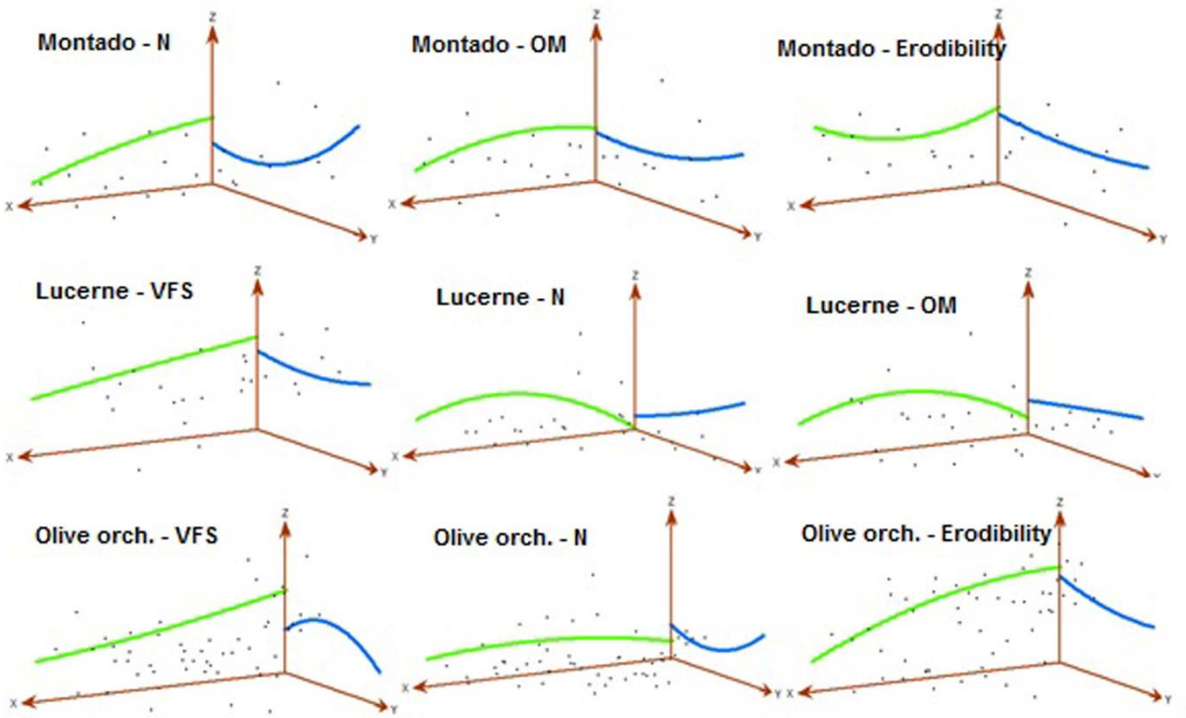

Figure 2 –Three-dimensional perspective of the trends in the input datasets.

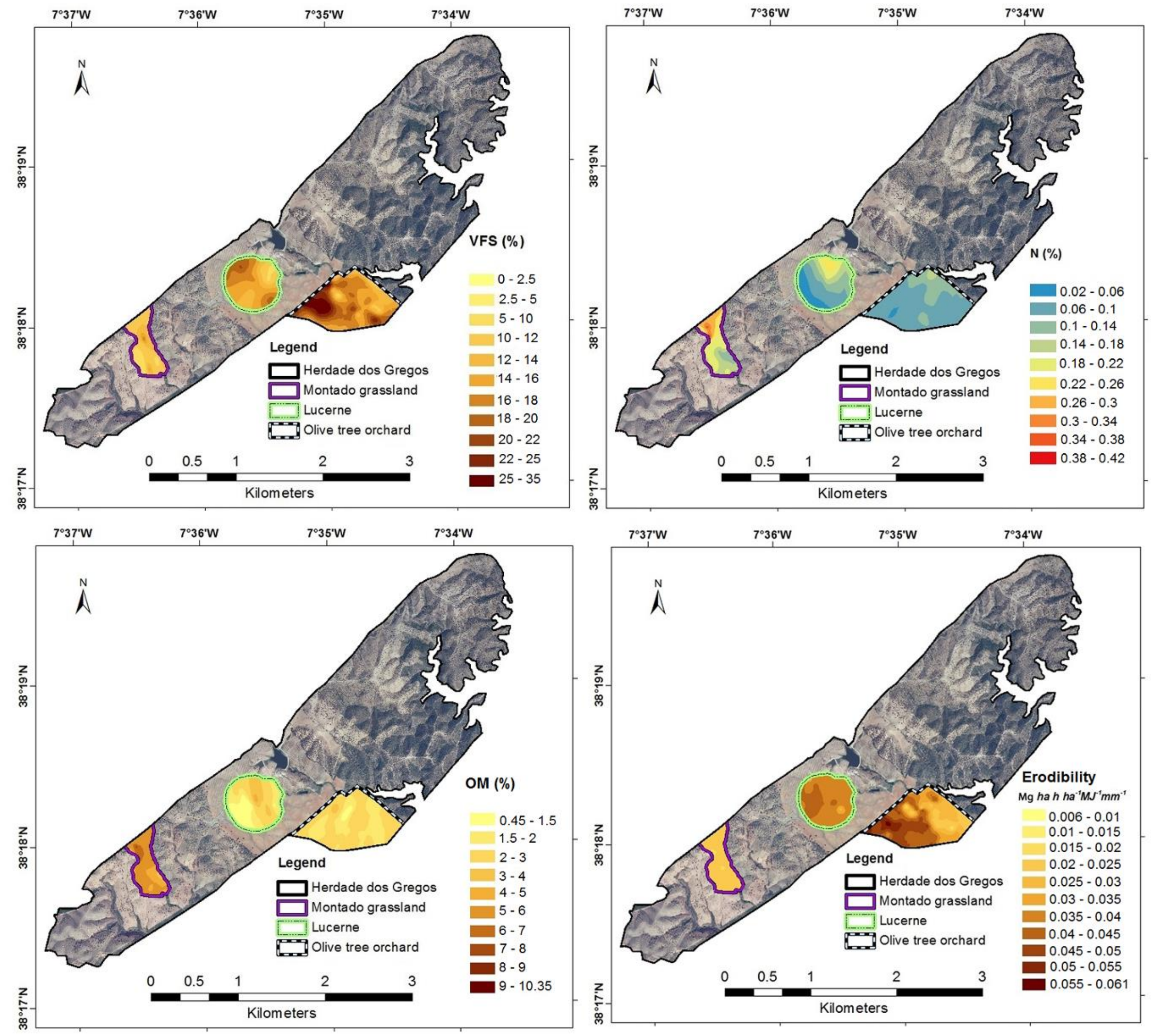

Figure 3 - Prediction map of very fine sand (VFS), total nitrogen (N), organic matter (OM) and soil erodibility (K factor).

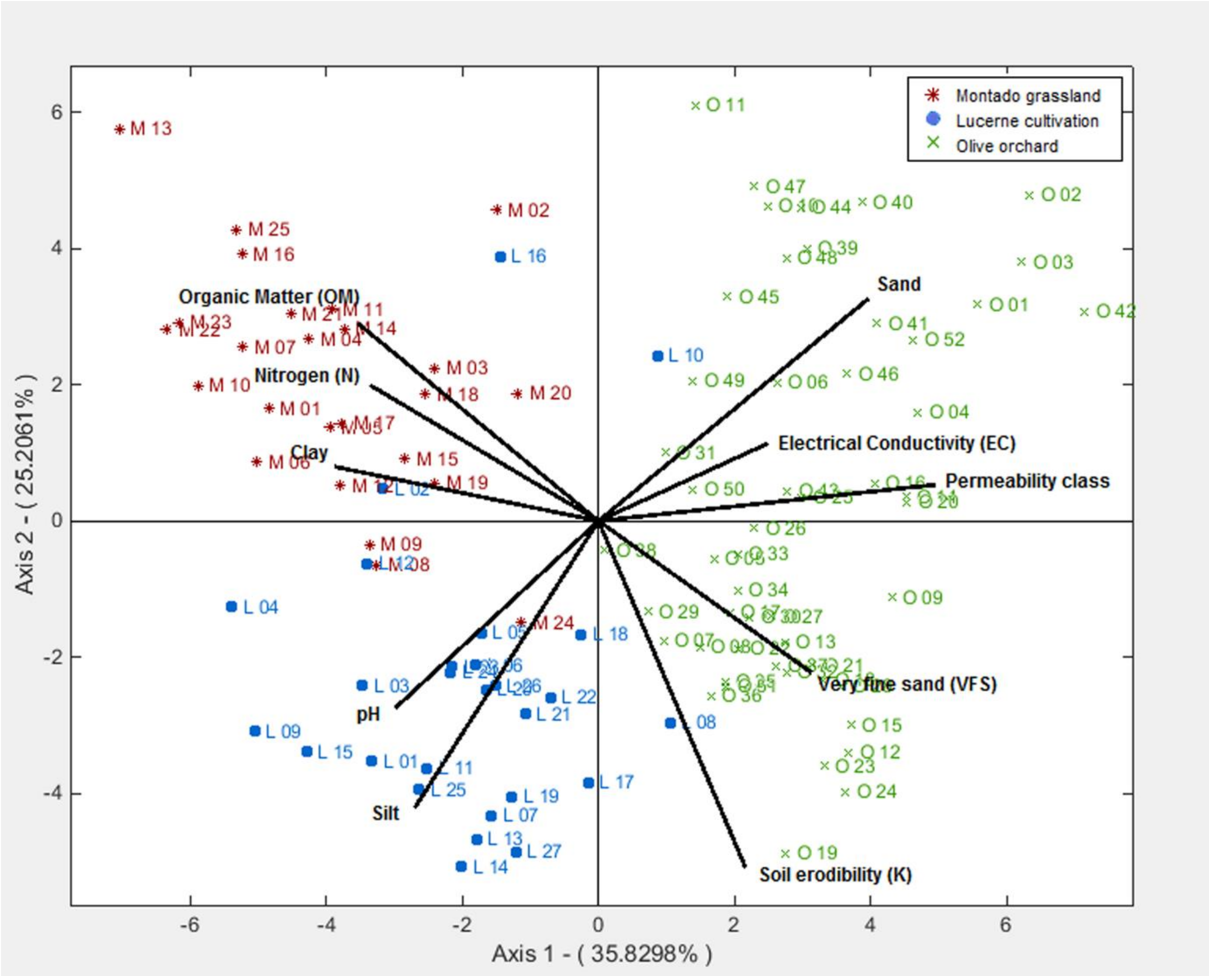

Figure 4 - The HJ-biplot representation matrix of soil samples and studied variables.

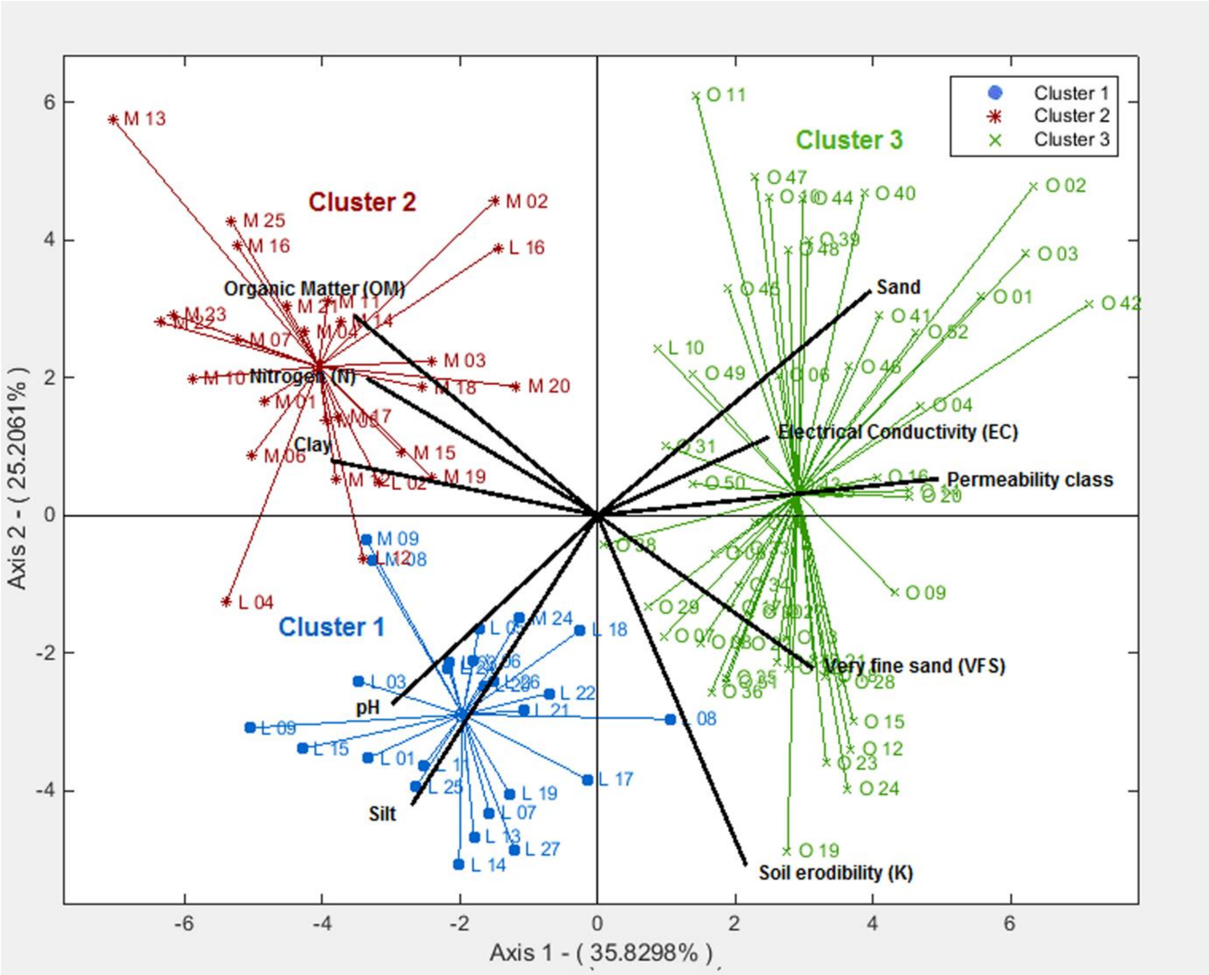

Figure 5 - Hierarchical clusters representation of soil samples and studied variables.