RAINFALL AND HUMAN ACTIVITY IMPACTS ON SOIL LOSSES AND RILL EROSION IN VINEYARDS (RUWER VALLEY, GERMANY)

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Rodrigo Comino, J.¹², Brings, C.¹, Lassu, T.¹, Iserloh, T.¹, Senciales, JM.², Martínez Murillo, JF.², Ruiz Sinoga, JD.², Seeger, M.¹, Ries, JB¹.

¹Department of Physical Geography, Trier University (Germany).

6 ²Department of Geography, Malaga University (Spain).

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8 Abstract

9 Vineyards are one of the most conditioned eco-geomorphological systems by human activity in Germany. The vineyards of the Ruwer Valley (Germany) are characterized by high soil erosion rates and rill problems on steep slopes (between 23-26°) caused by the increasingly frequent heavy rainfall events, what is sometimes deteriorate by incorrect land use managements. Soil tillage before and after vintage, application of vine training systems and anthropic rills generated by wheel tracks and footsteps are observed along these cultivated area.

The objective of this paper is to determine and to quantify the hydrological and erosive 16 phenomena in chosen vineyards in Germany, during the different seasons and under 17 18 different management conditions (before, during and after vintage). For this purpose, a 19 combined methodology was applied. Investigating climatic, pedological, geomorphological and botanic-marks variables was planned on the two experimental 20 plots in the village of Waldrach (Trier, region of Rhineland-Palatinate). 21

High infiltration rates (near 100%) and subsurface flow was detected by rainfall 22 simulations performed at different times of the year. To investigate the 23 24 geomorphological response of slope inclination, two 10 m and one 30 m long rills were measured using geometrical channel cross-section index, depth and width. The highest 25 variations (lateral and frontal movements) were noted before and during vintage, when 26 footsteps occurred concentrated during a short period of time. Finally, two maps of soil 27 loss were generated, indicated by botanic marks on the graft union of the vines. 62.5 t 28 ha⁻¹ yr⁻¹soil loss was registered on the experimental plots of the new vineyards (two 29 years), while $3.4 \text{ t ha}^{-1} \text{ yr}^{-1}$ on the old one (35 years). 30 31

32 **1. Introduction**

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Traditionally vineyards are among eco-geomorphological systems mostly conditioned by human activity. Cerdan et al. (2006, 2010) claimed, after studying 1350 experimental plots from several authors, that among cultivated areas, vineyards possess the highest erosion rates in Europe (12.2 t ha⁻¹ yr⁻¹). These problems appear at marginal anthropic environments with steep slopes, with bare soil cover and unsustainable land management activities (Martínez-Casasnovas et al., 2003; Paroissien et al., 2010).

On steep slopes in the European viticulture, terracing was the dominant correcting measure (Petit et al., 2012). However, the erosive processes affect with high intensity by several causes. Flow direction and rhythms of erosive process are manifested with several rills (with similar sizes), which divide the hillslopes in different transects

44 (Bryan, 2000; Prashun, 2011). This pattern of parallel rills (Ludwig et al., 1995) shows

45 the degradation processes on the vineyards, caused by water and anthropic erosion

46 (Sánchez-Moreno et al., 2012). Vandekerckhove et al. (1998) concluded that erosion
 47 rates are enhanced by incorrect land practices by vine-growers, and they are particularly

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higher after heavy and concentrated rainfall events. For example, Mediterraneanvineyards have the highest soil losses as a result of the increased surface flow rates by

50 the soil texture (Kosmas et al., 1997).

According to the methodology and the specific study areas, erosion rates are very variable: Martínez-Casasnovas and Poch (1998) and Martínez Casasnovas et al. (2002) respectively in north Spain observed 207 and 302-405 t ha⁻¹ yr⁻¹ respectively; in northwest Italy, Tropeano (1983) estimated between 40 and 70 t ha⁻¹ yr⁻¹; Wicherek (1991) and Wainwright (1996) validated 30 t ha⁻¹ yr⁻¹ in France.

56 In particular, Germany has a long tradition in viticulture and terraces on hillslopes along

57 the Mosel, Ahr and Rhine Valleys. However, Unwin (1996) and Auerswald et al. (2009)

- 58 reported several problems caused by erosion processes. The results of soil loss rates
- 59 from German vineyards reflected several differences, from 0.2 t ha⁻¹ yr⁻¹ (Richter, 1991)

60 to $151 \text{ t ha}^{-1} \text{ yr}^{-1}$ (Emde, 1992).

For the Mosel Valley, different studies with experimental plots to explain the
connection between precipitations (water and snow) and the soil loss behaviour by
surface flow mechanisms carried out by the researchers of this department (Richter and

Negendank, 1977; Richter, 1975, 1980a, 1980b, 1991). Soils were characterized by increased infiltration rates, gravel and fine mobilized elements, high organic matter proportions and intensive use of agricultural machinery (Hacisalihoglu, 2007).

67 From an economic point of view, vineyards are a traditional form of land use, which constitute one of the main and substantial economic bases of this region (Ashenfelter 68 69 and Storchmann, 2010). The agricultural cultivation started in Roman times and 70 continued with the constructions of monasteries in the middle Ages along Central Europe (Urhausen et al., 2011). This dynamic was significantly increased by the 71 72 intensification of production and harmful tilling of the soil, which led to a reduction of 73 fertility (Boardman et al., 2003; Raclot et al., 2009). The process of expansion began in 74 the 1950s and continued until the 1990s with some substantial transformations in the 75 production methods by the introduction of new machinery (Martínez-Casasnovas et al., 76 2010). As a consequence, the presence of gullies and rills, soil compaction and alteration of the local biochemical cycle was increased (Van Oost et al., 2007; Quinton 77

78 et al., 2010).

The importance of land morphology (Fox and Bryan, 2000; Martínez-Casasnovas et al., 2010), soil surface components (Corbane et al., 2008; Ruiz-Sinoga and Martínez-Murillo, 2009) and the influence of hydrological properties (Arnáez et al., 2012) on cultivated and abandoned areas are noted by several authors. All this occurs as a trigger for the increased volume of soil loss and the heterogeneity of intra-plot situations (Brenot et al., 2008; Casalí et al., 2009). Erosive dynamics are revealed through

different forms, for example natural or anthropic rills and gullies (Poesen et al., 1998) or

86 modern technics, like rainfall simulation. Small portable rainfall simulators, designed by

Cerdà et al. (1997) or the innovations by Ries et al. (2009; 2013a; 2013b) and Iserloh
(2012), are essential tools to analyse the process dynamics of soil erosion and surface
runoff in situ and in the laboratory. It provides the possibility to quantify soil erosion

90 rates and to investigate the impact of several factors (slope, soil type, splash effect, 91 raindrops, aggregate stability, surface structure and vegetation cover) on soil erosion

with quick and reproducible measurements (Seeger, 2007; Iserloh et al. 2012, 2013a, 2013b).

94 In practice, almost all manifestations of erosion forms in the vineyards have the origins

95 especially in footsteps and wheel tracks, which can significantly modify the natural

96 dynamic of the hillslope (Van Dijck and van Asch, 2002; Materechera, 2009; Arnáez et

al., 2012). Rills, inter-rills (Bryan, 2000; Fox and Bryan, 2000), and ephemeral gullies

Comentario [RCJ2]: This sentence has been improved

Comentario [RCJ3]: The bibliography about small portable rainfall simulators has been changed.

(Nachtergaele, 2001) show a connection between the lateral or vertical expansion (from 98 0.15 to 0.35 m yr⁻¹) and headcut retreat (about 0.7 m yr⁻¹) (Martínez-Casasnovas, 2003). 99 The purpose of this study is to characterize the soil erosion process of vineyards in the 100 Ruwer Valley (Germany). The objectives are: i) to determinate the hydrological and 101 erosive response of soil; ii) to describe and quantify the spatial and temporal 102 development of rills during a particular period with natural rainfall events; iii) to 103 evaluate the impact of land use management before, during and after vintage in 104 connection with rill erosion process; iv) to compare the soil erosion rates between the 105 recent (2 years) and ancient (35 years) vine cultivation, with the results of other 106 locations with similar geomorphological characteristics. Finally, two spatial and 107 temporal scales of analyses and, consequently, of erosive processes are considered: i) 108 local scale with simulated rainfalls and ii) field scale with the monitoring of rills and 109 quantification of soil loss through the botanic-marks. 110

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113 2. Methods and data collection

114 115 *2.1 Study area*

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The study area (Fig. 1) is located in the traditional vineyard village of Waldrach in the
Ruwer Valley, an affluent of the Mosel River in West-Germany (Trier-Saarburg,
Rhineland-Palatinate). It is part of the Rhenish Slate Mountains. The Ruwer Valley
descends from a plateau formed at the Hünsrück Mountains, from about 500 m.a.s.l. in
the south to approximately 200 m.a.s.l. in the north (Richter, 1980b).

Two types of vineyards were studied: i) old (cultivated more than 35 years); ii) young (planted less than a year ago). Both have the same lithological characteristics. The parent material is composed by: i) primary basin of no calcareous lithology under undulating reliefs with devonic greywackes, slates and quartzites; ii) fines sediments near the Pleistocene rivers (Schröder, 1991).

The work area ranges from between 220 up to 250 m.a.s.l. The exposures of the
hillslopes are fundamentally south-southwest oriented, for maximizing the insolation
intensity and favouring the phenology of crops (Menzel, 2005).

In essence, the soil management techniques of vine-growers are composed of
(Eggenberger et al., 1990; Vogt and Schruft, 2000): i) soil tillage before and after
vintage (end of October and beginning of November); ii) the presence of grass cover
along the inter-rows and below grapevines (between 10-35 cm height); iii) the use of
vine training systems to find equilibrium between leaves and the graft, to maximize

photosynthesis and sugar creation, using all of the useful space possible along difficult steep slopes for tilling (between 23 and 36°).

Along the embankments and inter-rows, anthropic rills by wheel tracks and footsteps
are noted. For example, the monitored rills of this investigation (R1, R2 and R3) are
emplaced on the stony embankment (Fig. 1) and were generated by these causes.

140 Due to the lack of a complete climatic station in the study area, values of rainfall and

temperature (all with more than 30 years of data) must be extrapolated. Peripheral
climatical stations in Mertesdorf (211 m; 49.7722, 6.7297), Hermeskeil (480 m;

- 143 49.6556, 6.9336), Trier-Zewen (131.5 m; 49.7325, 6.6133), Trier-Petrisberg (265 m;
- 144 49.7492, 6.6592), Trier-Irsch (228 m; 49.7259, 6.6957), Deuselbach (480.5 m; 49.7631,
- 145 7.0556), Konz (180 m; 49.6883, 6.5731), Bernkastel-Kues (120 m; 49.9186, 7.0664)
- and Weiskirchen (380 m; 49.5550, 6.8125) were applied. Data were obtained from the

147 German Meteorological Service (Deutscher Wetterdienst -DWD-), which allowed to

contextualize this territory with a Cfb climate (Köppen and Geiger, 1954). So, the
obtained annual rainfall depth was 765 mm and was concentrated in the summer months
(65-72 mm per month). The lowest monthly precipitation is observed between
February-April (50-60 mm per month). Annual average temperature is 9.3°C, with
average maximum values in June, July and August (16.2-17.6°C), and minimum values
along January and December (1.5-2.3°C).

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155 2.2. Soil analysis

The soil samples were collected from four different positions on a space with no more 157 than 0.5 m^2 and with different depths: 0-5 and >5 cm (maximum to 15-25 cm). Along 158 two inter-rows of old and young vineyards and two from the embankments of old 159 grapevines with rills (top and bottom). Each sample was analysed with two replicates 160 161 and they were taken in order to determine the soil properties, as grain size (<2 mm and 162 >2 mm), pH, total organic carbon (TOC) and inorganic carbon (TIC) content by ignition 163 (550° C and 1050° C respectively in muffle furnace). Saturation and absorption capacity was measured with a simplification of the "Counting the Number of Drop-Impact 164 method" and by Emerson (1967), Imeson and Vis (1984) and Herrick et al. (2001). 165 Finally, bulk density using steel cylinder were calculated. 166

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168 2.3. Description of rainfall and agricultural events during the monitoring

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170 Climatic and agricultural actions (during the monitoring) were monitored to describe the important events in the study area. In order to obtain the rainfall data, an extrapolation 171 172 of the gradients data at surface level was made, by using the data from the peripheral agro-climatic stations of the German Meteorological Service (Deutscher Wetterdienst -173 174 DWD-) and the Dienstleistungszentrum Ländlicher Raum/Rheinland-Pfalz (DLR-RLP). 175 Calculations were linear estimations and intersections with the axis, using rainfall and elevation data (Rodrigo Comino, 2013; Senciales and Ruiz Sinoga, 2013). Rainfall 176 177 events were frequent during all the research period. The daily intensity in this period (September to December) was 2.2 mm d^{-1} and the days with rainfall were 4.6 days in 178 each interval of this monitoring period (between 6 to 7 days). 179

- 180
- 181 2.4. Statistical and spatial analysis

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183 A continuation, K (with the soil analysis) and R factor of RUSLE (Wischmeier and 184 Smith, 1978; Dabney et al., 2014) were added to complete the soil analysis and rainfall erosivity respectively (Martínez Casasnovas et al., 2002). For this purpose, R factor 185 186 (54.31) was calculated with the index for Germany with better results than adjusted equation for Rhineland-Palatinate region (Sauernborn, 1994; Casper et al., 2013). After 187 188 that, following the example of Arnáez et al. (2007), recurrence periods with Poisson method were included to justify the intensity of rainfall simulation and classify rainfall 189 events on the study area (Mays, 2011). Results are presented as percentage of days per 190 191 year using a co-Kriging extrapolation with GIS from the peripheral agro-climatic stations. 192

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- 194 <u>2.4.1. Rainfall simulations</u>

195 In alternate varying months, eight rainfall simulations were carried out under different 196 soil moisture conditions. During the first four simulations in August (2012) the soil **Comentario** [**RCJ4**]: All information about soil analysis (methods, table 1, bibliography...) has been improved and changed.

moisture were between 50-70%, while October and December (2013) between 20-40%. 197 198 The objectives were to quantify the soil losses, the degree of infiltration, runoff coefficients, the suspension and concentration of sediments. All simulations were 199 carried out on the inter-rows of old-vineyards with the same rainfall intensity (40 mm h 200 ¹) for two reasons. Firstly, when the return period is calculated, 40 mm h⁻¹ of intensity is 201 the least usual. Therefore, the rainfall greater than 40 mm h⁻¹ would have a little 202 probability to happen. So the different reactions of the soil to extreme rainfall could be 203 204 recorded. Secondly, the simulator was exactly calibrated to control splash effects 205 following Iserloh et al., (2012, 2013). The defined area of experiments coincided with a metal ring of 0.28 m². To measure the quantity of water a flow control (Type KSK-206 207 1200HIG100, 0-125 L h⁻¹, Kobold Company) and a manometer (with a calibrated pressure of 0.2 bar) were applied. In each simulation (30 minutes), we were using 208 intervals of 5 minutes to collect runoff. 209

210 Since October until November the same results were obtained: total infiltration. 211 Therefore, the last one was carried out in December (2013) and to understand the reason 212 of the 100% infiltration, the stoney A horizon was removed inside the metal ring. The 213 main purposes were to: i) confirm the increased infiltration; ii) investigate the 214 relationship between the process and the soil surface components. A hydrophilic nylon fabric was used to protect the soil from the splash effect. A vertical soil profile was 215 216 caved underneath the simulator (50 cm depth and 150 cm width) in order to observe the 217 infiltration dynamic. In this manner, subsurface flow was observable (Fig.2) by the 218 profile and the metal collector, however it was impossible to quantify it.

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220 <u>2.4.2. Geometrical rills monitoring</u>

221 Three rills with different geomorphological origins were chosen for the monitoring (R1, 222 R2 and R3). The rills were divided into one meter sections. Between September and 223 December, the width, depth and slope angle of the sections along rills were measured. 224 The first rill (R1) was caused by the wheel tracks and it was nearly 30 meters long (30 sections), starting from the bottom of the embankment. The average inclination of the 225 rill was 28° and had approximately a contributing catchment area of 600 m². The second 226 (R2) and third (R3) rill were located on the embankments with steeper slopes (34° and 227 31.7°) and had smaller contributing catchment areas (19 m² and 25 m²). R2 (near a wall 228 229 and drainage channel) was 7 meter length (7 sections), and R3 around 10 meters (10 230 sections). Both were caused by the footsteps of vine workers. The methods of Govers and Poesen (1987), Takken et al. (1999), Vandekerckhove et al. (2003) and Wirtz et al. 231 232 (2012) were followed to measure their changes in geometry. In order to calculate 233 weekly the geometrical variation of transects, the geometrical channel cross-section index was calculated (Dingman, 2008; Quiquerez et al. 2008):

234 index was calculated (Dingman, 2008; Quiquerez et al. 2008

$$235 \quad TSI = \frac{W}{Y}$$

Where W represents the width and Y the depth (both in centimetres). Note, while the quotient is more elevated, the widening process of rills is faster than the deepening process. Furthermore, the standard deviation was added to distinguish when averages were obtained with equal or unequal values. Consequently, two types of analyses with the geometrical channel cross-section index (Dingman, 2008; Quiquerez et al. 2008) were elaborated. Inclination was measured with a clinometer.

242 Firstly, the total average values per section were used to detect the most vulnerable

transects, which were mostly modified by geomorphological changes both temporally

and spatially. The second calculation aimed to show the geometrical variation of each

rill between the monitoring phases with the standard deviation (before, during and aftervintage).

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248 <u>2.4.3. Frontal botanic marks on the graft union</u>

The distance between frontal marks on the graft union and the visible actual rootstock of 249 250 grape-vines were measured (Fig. 3) on a total area of 0.065 ha (with old grapevines) and on 0.043 ha (with young grapevines). Graft union can be defined as unearthing or 251 buried signal, which could show the theoretical ancient topsoil (Brenot et al., 2008). 252 253 This analysis aims at confirm the theory about the "botanic marks" as indicators of soil loss (Brenot et al. 2008; Casalí et al., 2009; Paroissien et al. 2010). Vitis vinifera after 254 255 the Phylloxera crisis was grafted with the American scion of controlled species as the Vitis rupestris, Vitis riparia and Vitis berlandieri (Unwin, 1996). Several authors 256 (Brenot et al. 2008; Casalí et al., 2009; Paroissien et al. 2010) demonstrated that these 257 258 signals were correct indicators of soil movements in the vineyards (erosion, transport 259 and sedimentation). The conditions described in Brenot et al. (2008), were previously confirmed with the vine-growers and those were: i) there is no vertical growth of the 260 261 graft after the vineyard plantation; ii) the recommendations concerning the graft union 262 elevation at the vineyard are followed so that this elevation can be considered to be constant over the studied region; iii) the measurement errors are negligible compared to 263 264 the observed unearthing or burying of vine-rootstock.

265 Furthermore all graft unions near 2 cm from the topsoil were planted during the first 266 year. In total 1200 graft unions were measured with a subtraction of 2 cm, from which 267 720 were cultivated 35 years ago on the study area (coinciding with the monitored rills). The other 480 were planted in 2012. The average inclination of the hillslope is almost 268 269 constant from 22 to 24°. It is important to note that a little contention wall with a 270 drainage vertical collector (adjacent to R2) divides the study area in two parts. This 271 infrastructure was planned to reduce accumulation of the eroded materials along the road and to drain the possible surface flow. Below two isoline maps are presenting the 272 soil erosion level, according to the geomorphological conditions of the plots. The co-273 274 kriging method (Dirks, 1998; Goovaerts, 1999; Wang et al., 2013) was applied with 0.1 275 precision intervals (quartiles) and two variables: botanic marks and digital elevation 276 model with a resolution of 1x1 meter.

277 The total soil loss was calculated from the volume of an imaginary polygon and then it

was extrapolated to m^3 ha⁻¹ and t ha yr⁻¹ with an estimation. The sides of the polygon were the distance between each vine-stock (0.9x1 m), while the height was the distance between the botanic marks on the graft union and the visible actual rootstock. Total soil

- loss (t ha^{-1}) was estimated with the erosion-deposition (ER) equation (Paroissien et al.
- 282 2010):

$$ER = \frac{Vol \ x \ Ds}{St \ x \ Av}$$

The volume (*Vol*), the total area field (*St*), the age of the vines (*Av*) and the bulk density data (*Ds*) were applied. For the young vineyards 1.14 g cm³ and for the old one 1.4 g cm³ were used, both the average of the two soil samples in different depth (0-5 cm and >5cm). At this level, this method also requires the assumption that the study area is absolutely even. However, due to the rills, footsteps and wheel tracks it is rough.

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201	2	Doculto
291	э.	Results

292 293 *3.1. Soil analysis* **Comentario** [RCJ5]: Information of calculation by Paroissien et al. (2010) was included.

Comentario [RCJ6]: All conclusions and results about soil analysis (methods, table 1, bibliography...) has been improved and changed.

294 295 Laboratory analysis data (Table 1) show chemical and physical properties of the soils, which are relevant to introduce the context of the geomorphological processes. The old 296 (>68%) and young (>70%) vineyards have the highest concentration of grain size larger 297 298 than 2 mm, which could be classified as stony soils. The highest concentration of organic matter (10-13%) was noted on the surface horizon (0-5 cm) along the old 299 vineyards and the upper embankment, according with the most elevate rates of bulk 300 density (1.4 g cm³). The young vineyards and the below embankment have lower 301 302 organic matter (< 6%) and more fine sediment concentration (<31%). The most elevated 303 point of saturation and water absorption capacity of the soil samples were calculated along the subsurface horizons (>5 cm). The results of K factor indicating the erodibility 304 of soil following Wischmeier and Smith (1978) and Dabney et al., (2014) showed 0.22 305 306 and 0.37 for old and young vineyards respectively. Finally, Cambisol leptic-humic was 307 classified using the methodology of FAO (2006, 2007, 2014).

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309 *3.2. Rainfall events and land management during the study period*310

Soil surface characteristics, during and after the agricultural activity, and the
extrapolated rainfalls in 2013 (total and intensity) from the nearby climate stations were
described to add more information (Table 2). The probability of the return period (Table
3) is added to include the recurrence of different rainfall depth and intensities per day.

During and one week after vintage a powerful anthropic action was observed. This situation coincided with the elevated soil moisture rates. The increased footsteps of the workers disturbed the soil (sub and superficially) and therefore rills appeared. This dynamic was observed at areas without vegetation cover or without cultivation (e.g.

319 embankments).

320 After vintage the number of footsteps was decreased, coinciding with the decreasing of 321 rainfall depth and intensity (mm d^{-1}). Accordingly, less soil movement was observed 322 and the rills began to widen. However, currently every morning the soil was frozen and 323 along the day a thaw was occurred.

The precipitation between 20 and 5, and 5-0.1 mm d^{-1} have the highest probability (36.1-36.3% and 22.6-23% respectively). The more intense rainfall events (>40 mm d^{-1}) have the lowest possibility. The probability of rainfall events at this season could be

327 classified between a 22.7-23%.

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329 *3.3. Rainfall simulations*

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In total, eight rainfall simulations were carried out during August, October, November and December (Table 4), but only the summer simulations gave quantifiable result

about runoff and soil loss (Figure 4). During the other simulations 100% infiltration ratewas observed.

For the four simulations of August, runoff and sediment suspension data appeared. The maximum runoff coefficient and suspended sediment load were 15.2±7.8% and 25.81 g

 m^{-2} respectively. These values were lower compared to the infiltration averages (near

100%). In each experiment, only one increase interval of soil loss and at the same timemore surface runoff was noted. Consequently, the sediment concentration decreased.

340 Principally, this situation happened in the central minutes of the rainfall simulation

341 (between 10 and 20 minutes), when the soil became saturated and expelled water as

342 surface flow. After this saturation point, the A horizon was being eliminated and it

seemed that the water could be moving as subsurface flow by gravity.

344 This supposition was confirmed in the next three simulations (October-November), 345 because the rainfalls were completely infiltrated and fine sediments were not eroded. Finally, for the last simulation in December, a soil profile of 0.5 meters below the 346 347 simulator was excavated (Fig. 2) in order to observe the intensity and direction of a possible subsurface-flow. From the beginning of the simulation, this hydrodynamic 348 behaviour of total infiltration was noted across the profile. However, we could not 349 350 calculate the intensity and observe the direction in situ, because the water flowed across 351 an area larger than the rainfall simulator collector.

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353 3.4. Geometrical monitoring of rills with anthropic origin

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Size variations of the rills are presented in graphics with data from the monitoring period (Figs. 5 and 6).

357 The highest variations were observed before and during vintage. When no footsteps occurred in a concentrated short time, soil reacted without lateral and frontal 358 movements. Of this situation, the behavior of the soil could be deduced with the 359 deepening and widening process of the rills. In general, using the geometrical channel 360 361 cross-section index, four intervals were detected with relevant weekly changes for all rills: i) between 0-1 and 1-2 meters (below the hillslope) irregularities were noted in 362 363 little alluvial fans on the border of the embankment and the road; ii) from 3-4 to 4-5 meters fracture appeared in the slope as a micro-terraces (between 32° to 36° of slope), 364 365 in which small slide scars by the soil movements were noticed below the A horizon; iii) along 7th and 9th meter at the top of the embankment, where the vines grapes were 366 cultivated (the slopes were 30° to 23°); iv) only for R1 (originated by wheel tracks), it 367 was noted an increase of the values of geometrical cross-section index from 26-27 368 metres and a maintaining of the gradient (27-28°). In this section, in contrast to 369 370 deepening process weeding was favoured, especially during the vintage. Moreover, 371 average values (Fig. 7 and 8) in each rill with this index was noted.

For R2, higher value (5.3±2.9 cm) was obtained than for R3 (4.9±2.5 cm). In this regard, the most inconstant rill (R2) was located near a little contention wall with a drainage channel and it was significantly modified by several footsteps.

At R1 (5.3 ± 2.2 cm) between 1 and 10 meters elevated data were observed (5.5 ± 2.9 cm), but from here the values were descending (5.3 ± 1.8 cm). Finally, the highest parameters were measured from 27 meters (10.7 ± 4.6 cm), during the weeding processes (confluence of two or more rills).

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380 *3.5. Soil loss level maps*

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Fig. 9 and 10 present the soil losses and the trend of movements. Annual average soil loss per row, on each side of the contention wall and on the total study area was added to the final table 5.

At each side of the channel and the contention wall at both vineyards, diverse dynamics was noted. The highest erosion rates (dark colors) were located on the top at the left side

of the hillslope. This situation was increased near the channel in contact with the

embankment on the left side (for the young grapevines 134.1 t ha^{-1} and 124.3 t ha^{-1} in

the old vineyard). The behavior is more in accordance with the natural conditions on the

right side, because the soil loss was lower (light colors) and below the accumulation

391 was predominant (during two years 116 t ha^{-1} and in 35 years 113 t ha^{-1}).

In two years of cultivation very high total soil loss was calculated (125 t ha⁻¹ and 62.5 t ha⁻¹ yr⁻¹). However, for the old vineyards (35 years), 118.7 t ha⁻¹ erosion rates with an

Comentario [RCJ7]: An explantion about figures 7 and 8 was included.

Comentario [RCJ8]: This paragraph was rewritten.

annual rate of 3.4 t ha^{-1} yr⁻¹ was calculated. Again, on the left side losses were higher than on the right side (3.6 and 3.2 t ha^{-1} yr⁻¹).

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398 4. Conclusions and discussion

Due to the stony soil surface (between 58.3 and 70.7% larger than 2 mm) and the active
cultivation work (wheel tracks and footsteps along the inter-rows), high infiltration rates
(near 100%) were observed. During this study, applying different tools and experiments
on the soil, was observed: i) subsurface processes, such as micro-piping or creeping ii)
concrete pedological conditions, like variations between agricultural seasons or
geomorphological instability after soil tilling.

Firstly, the highest organic matter content and bulk density were noted in the old
vineyards, which could explain a lower sediment transport or soil movements across the
hillslope than at the young vineyards and the embankments. On the other hand, the most
saturation and absorption capacity rates were located from 5 cm depth. Subsurface flow
dynamics could be analyzed, maybe according a high porosity rates due to these stony
soils.

Secondly, spatial and temporal geometrical evolutions of rills were monitored with the 412 413 geometrical channel cross-section index before, during and after the agricultural 414 activities (vintage) in the study area. Accordingly soils had three different responses in 415 the three different situations. The highest variability (in width and depth) of the rills was 416 observed on the embankment close to the contention wall and drainage channel. Due to the soil tilling (land removal), no plants with their roots holding the soil and the 417 418 uncorrected located wall, the development of the rills was increased. The footsteps and 419 wheel tracks before and during vintage increased the dynamic of these processes. This 420 was coinciding with the frequent and intensive rainfall events.

421 Moreover, the impact of land management was evaluated with the total soil losses rates, using the botanic marks of the grapes and the erosion-deposition (ER) equation 422 (Paroissien et al. 2010). The instructions of Brenot et al. (2008), Casalí et al. (2009) and 423 Paroissien et al. (2010) was followed to measure the difference in the graft union of 424 1200 grapevines. However, an elevate component of subjectivity is adverted by several 425 426 authors, because the method depends of arbitrary criteria. With this method 118.7 t ha⁻¹ soil loss was calculated on the old vineyard, which means $3.4 \text{ t} \text{ ha}^{-1} \text{ yr}^{-1}$. Respectively 427 428 on the young vineyard 125 t ha⁻¹, which supposes 62.5 t ha⁻¹ yr⁻¹, were measured. During the first years of plantation very high rates of soil losses were observed. 429 430 However, the next years the sediment transport descends considerably, possibly due to: i) the soil tillage against the erosion is increased; ii) the structural stability of soils is 431 432 improving continuously since the plantation (organic matter, bulk density, absorption 433 capacity...). Although it might be asked how much money could be saved by the vinegrowers, applying before directly correct land management measures on the hillslopes. 434 435 For a correct land management, the location, the quantification and the proposition of measures for the prevention of the destabilizations and modifications on hillslopes are 436 437 considered to be essential. Territories with intensive and mountain farming should be considered as vulnerable points by erosion problems. Policies must aim to protect 438 439 hillslope morphologies for terracing and to prohibit indiscriminate heavy machinery 440 use. Alterations can implicate changes with unappreciable consequences in short-term, 441 but irreversible in long-term (Piccarreta et al., 2006).

Comentario [RCJ9]: More ideas about conclusions and discusions were included (soil analysis, organic matter, erosiondeposition equation by Paroissien et al -2010-, soil tilling...). Finally, the erosion rates could be compared with other studies about vineyards in theMosel Valley, Germany and Europe by different authors (Table 6). The problem isnowadays relevant.

As this study, Richter (1975, 1991) and Hacisalihoglu (2007) worked also in the Ruwer Valley vineyards context, but with different methodologies. In these experiments they were using sediment boxes and empiric equations, but the measured soil loss rates were similar to the erosion rates of the old vineyards of this paper ($3.4 \text{ t} \text{ ha}^{-1} \text{ yr}^{-1}$) 0.2-6.6 and 6.47 t ha⁻¹ yr⁻¹, respectively. For other scales (Germany and Europe), Auerswald et al. (2009) and Cerdan et al. (2006, 2010) calculated similar soil erosion rates as well (5.2 and 12.2 t ha⁻¹ yr⁻¹), with extrapolations from different works.

452 Only Emde (1992) with USLE inferred a rate over 150 t ha^{-1} yr⁻¹, which is 453 approximated to the soil erosion of the young grapevines (62.5 t ha^{-1} yr⁻¹) in this paper.

The results of this paper contribute the validity of the available data, although the comparability with other studies is difficult, due to the different methodological approaches and the diverse climatic situations. Furthermore, all studies coincided in the same assumption: the vineyards soil erosion rates were the highest compared to other land uses (forest, grassland, shrubs or regeneration).

459 460

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462

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Table 1. Soil analysis of the study area								
Soil samples	>2 mm (%Total)	<2 mm (%Total)	pН	TOC (%) ^a	TIC (%) ^b	Saturation (%) ^c	Absorption capacity (%) ^d	Bulk density (gr/cm3)
Old grapevines (0-5 cm)	68.18	31.82	6.6	10.7	1.5	11.3	12.7	1.4
Old grapevines (5-15 cm)	66.19	33.81	6.6	6.5	1.5	9.8	10.9	1.4
Upper embankment (0-5 cm)	61.45	38.55	6.7	13.7	1.5	10.1	11.2	1.4
Upper embankment (5-15 cm)	58.25	41.75	6.3	6.6	2.2	11.2	12.6	1.4
Below embankment (0-5 cm)	61.82	38.18	6.4	5.7	1.3	9.1	10	1.3
Below embankment (5-15 cm)	68.85	31.15	6.7	5.7	1.4	11.8	13.4	1.4
Young grapevines (0-5 cm)	70.17	29.83	6.6	4.1	2.3	11	12.5	1.1
Young grapevines (5-15 cm)	70.68	29.32	6.1	5.5	1.5	12.7	14.5	1.2

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Comentario [RCJ10]: More information abou the methods and

(2.2., 3.1., 4).

conclusions was included along the text

a) TOC = Total Organic Carbon; b) TIC = Total Inorganic Carbon; c) Saturation (%) = (Water added to saturation/final weighted) x 100; d) Absorption capacity (%) = (Weighted of saturated aggregate – Initial weighted) / Initial weighted) x 100.

Monitoring phase	Date	Rainfall (mm)*	Days with rain	Intensity $(mm d^{-1})$	Activities
	24.09.2013	22.98	6.6	3.3	Leaves of the grapevines were cut
Before	1.10.2013	10.34	4.3	1.5	to improve the absorption of the
vintage	8.10.2013	1.25	2.5	0.2	sunlight and appearance of footsteps.
	15.10.2013	22.78	3.3	3.3	Several footsteps marks were
Vintage	22.10.2013	26.63	4.1	3.8	situated from the sections 0-1 to 8-
v intage	29.10.2013	8.78	4.9	1.3	9 meters. A lot of grapes and leaves stayed on the surface.
	6.11.2013	51.40	5.9	7.3	Several footsteps modified R2. Increasing of lateral enlargement (no deepening).
	12.11.2013	33.96	4.5	4.9	Many grape-leaves and branches on the surface. Footsteps began to dissolve on monitored rills (1, 2 and 3).
After vintage	19.11.2013	10.34	6.3	1.5	The soil was cleaned of leaves and branches. Footsteps developed to new rills by the rainfall.
	26.11.2013	1.95	4.0	0.3	Each morning soil freeze appeared.
	03.12.2013	7.96	4.8	1.1	After midday it was almost dry, but not the subsurface horizons.
	10.12.2013	3.24	4.3	0.5	Footsteps marks were visible only from the sections 0-1 to 1-2. Rills stayed without remarkable changes.

Table 2. Rainfall events and descriptions of agricultural activities during the monitoring.

756 * Rainfall (mm) means total mm after each measure, currently, each 6 or 7 days.

Table 3. Return period of rainfall events per year.

Rainfall depth (mm)	% probability of return period (d ⁻¹ yr ⁻¹)
>40	0.44-0.46
40-20	5.65-7.23
20-5	36.12-36.36
5-0.1	22.67-22.95
0	9.02-11.16

Table 4. Rainfall simulation parameters

ID	Pp (mm h ⁻¹)	Runoff (L/5min)	Runoff Coef./5min (%)	Infiltration/5min (%)	Concentration/5min (g/L)	Total erosion (g m ⁻² h ⁻¹)
1. August 2012	9.72	0.03±0.01	3.9±1.1	96.1±1.1	3.34±1.95	23.2
2. August 2012	10.32	0.004 ± 0.002	0.52±0.2	99.5±0.2	5.03±2.91	30.9
3. August 2012	13.2	0.17±0.09	15.2±7.8	84.8±7.8	7.77±3.07	51.5
4. August 2012	10.44	0.06 ± 0.04	6.7±4.8	93.3±4.8	7.01±8.03	30.5
5. October 2013	10.8	0	0	100	0	0
6. October 2013	10.68	0	0	100	0	0
7. November 2013	11.16	0	0	100	0	0
8. December 2013*	9.48	0	0	100	0	0

762 * = Rainfall simulation without A horizon.

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 Table
 5. Volume estimations of soil loss in young and old vineyards

	Parameters	m ³ ha ⁻¹	t ha ⁻¹ *	t ha ⁻¹ yr ⁻¹
Young vineyards (2 years)	Total soil loss	4.7	125	62.5
	Total on the left side	5.1	134.1	67.1
	Total on the right side	4.4	116	58
	Total soil loss	5.5	118.7	3.4
(35 years)	Total on the left side	5.8	124.3	3.6
	Total on the right side	5.2	113	3.2

764 $t = t = t^{-1}$: The soil loss is equivalent to the total erosion since the first moment of plantation.

Comentario [RCJ11]: The table was joined in only one.

Table 6. Comparison of soil losses rates between different uses, territories and methodologies. 767

Rates Authors Study area Method Types of land uses $(t ha^{\cdot 1} yr^{\cdot 1})$ Richter Mertesdorf Sediment boxes 0.2-6.6 Vineyards (1975, 1991) (Mosel Valley) Rheingau Emde (1992) USLE Vineyards 151 (Rhin Valley) 0.71 Regeneration "Algemeine Boden 0.67 Forest Hacisalihoglu Mertesdorf Abtrags 0.87 Shrubs Gleichung" (Mosel Valley) (2007)1.2 Grassland (ABAG) Vineyards 6.47 Extrapolations and 5.7 Annual arable land Auerswald et al. R factor of USLE 0.5 Grassland Germany (Universal Soil 0.2 Forest (2009) Loss Equation) 5.2 Vineyards Cerdan et al. Extrapolations 12.2 Vineyards Europe (2006, 2010) from other works Waldrach This study Botanic marks 3.4-62.5* Vineyards (Mosel Valley)

768 * = 3.4 t ha⁻¹ yr⁻¹ on the old vineyards (average in 35 years) and 62.5 t ha⁻¹ yr⁻¹ for the other area 769 with young grapevines (since 2012).



Figure 1. Study area in Waldrach (Ruwer Valley, Germany).



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- Figure 2. The rainfall simulation in December. a) A horizon eliminated (between 5- 7 cm). b) Before simulation. c) Profile to 0.5 m below (1.5 m x 0.5m) with the sediment collector. d) Situation of simulator ring. e) Concurrently rainfall simulation f) Subsurface flow during the experiment.
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776 777 778 779 Fig. 3. Monitoring of botanic marks and rills. a) Example of measured distance between the botanic mark and the actual topsoil (with 2 cm of the initial planting); b) Weekly geometrical rill monitoring: width and depth; c) Vintage: vine workers use rills to ascend or descend the vineyards; d) Imaginary polygon to calculate the soil loss with botanic marks. 780

781 Figure 4. Relationships between variables: surface flow, suspension and sediment concentration



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Figure 7. Diagram of the embankment with the rills on the old vineyards





Figure 8. Geometrical channel cross-section averages of the rills during the total monitoring period (R1, R2 and R3) on the old vineyards

