



1 **Micromorphological characteristics of sandy forest soils** 2 **recently impacted by wildfires in Russia**

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7 **Abstract.** Two fire affected soils have been studied using micromorphological methods. The objective of the paper
8 is to assess and compare fire effects on micropedological organization of soils in a forest steppe zone of the Central
9 Russia (Volga Basin, Togliatti city). Samples were collected in the green zone of Togliatti city. The results showed
10 that both soils were rich in quartz, feldspar. Mica was highly present in soils affected by surface fires, while
11 calcium carbonates were identified in the soils affected by crowns. The type of plasma is humus-clay, but the soil
12 assemblage is plasma-silt with the prevalence of the silt. Angular and subangular grains are the most dominant in
13 soil particulates. No evidence of intensive weathering was detected. There was a decreasing of the porosity in soils
14 affected by fires as consequence of soil pores filled with ash and charcoal.

15 **1 Introduction**

16 Fire has an important impact on soil properties as identified by previous works (Certini, 2005; Certini, 2013;
17 Guénon et al., 2013; Mataix-Solera et al., 2011; Jain Theresa et al., 2012; Bergeron et al., 2013; Dymov, Gabov,
18 2015; Zharikova, 2015; Maximova, Abakumov, 2015). Soil processes in postfire soils environments is quite
19 different from those in soils of natural landscapes or in thechnogenous ones. In general, changes in morphological
20 organization and soil mineralogy are well known in soils after fires produced at high temperatures. After the fire,
21 there is an accumulation of ash on topsoil (Pereira et al., 2014), leaching of some nutrients into deeper horizons
22 (Bodi et al., 2014), over-compaction of the surface and accumulation of crusts, transformation of soil structure
23 (Mataix et al., 2011).

24 Micromorphological methods are known as useful tool for investigation of soil transformation under natural and
25 human-impacted conditions (Stoops, 2009). Methodology of classical micropedology provides with required
26 information about soil development in micro level, such as changes in fine earth composition and soil plasma
27 evolution (Kubiena, 1938, 1967; Stoops and Eswaran, 1986). These methods are widely used for analysis of soil
28 paleo processes (Sedov et al., 2013), soil restoration on post-mining environments (Abakumov et al., 2005), soil
29 elementary process in different environments (Lebedeva et al., 2010; Abakumov et al., 2013) and specification
30 soil classification aspects (Kubiena, 1967). Micromorphological investigations related to fire effect of soil crusts
31 and fine earth (Greene et al., 1990) and aggregate dynamics in post-fire soil (Mataix-Solera et al., 2002) have been
32 undertaken. However, the micromorphological methods have never been applied to study post-fire soils
33 transformation in the Russian wild-fire environments. Moreover, this paper deals with comparison of different
34 post-fire scenarios (surface and crown fires) and provide 5-year monitoring of fire affected soils and ecosystems
35 in a whole. Few researches were carried out about this topic.



1 The objective of this work was to characterize the micromorphological indices of microstructure transformation
2 in soils affected by different types of wild-fire, as compared microstructure of mature, unaffected soil of pine
3 forests in Central Russia.

4 **2 Materials and methods**

5 **2.1 Study area**

6 The study was conducted in Samara region situated near the Volga River, in the central part of European Russia,
7 Samara region (N 58°39'44.55"; E 39°17'48.95", 179 m asl). The extremely hot weather in summer of 2010 in
8 Russia (especially the most difficult situation was on whole Russian European area and also Ukraine and Eastern
9 Europe) resulted in drought and eventually catastrophic forest fires on the vast territories of European and Siberian
10 Russia. The forest fire studied occurred in 2010 and affected more than 8000 hectares. Fire severity was very high.
11 Parent material is composed by old (Pleistocene) alluvial-dune landscape.

12 The affected ecosystem is characterized by a forest-steppe environment with higher pedodiversity. The vegetation
13 was composed mostly by pine forests *Pinus sylvestris* L. There are xerophyte species at dry locations like *Veronica*
14 *spicata* L., *Sedum acre* L., *Antennaria dioica* L., *Calamagrostis epigeios* (L.) Roth, *Centaurea marschalliana*
15 *Spreng* and hardwoods (*Quercus robur* L., *Betula pendula* Roth., *Populus tremula* L.) in more humid conditions.
16 The herbaceous vegetation consists of rhizomatous and loose-bunch gramin (*Bromus inermis* Leyss., *Elytrigia*
17 *repens* L., some species of *Poa*, and *Agrostis canina* L.) at post-fire plots.

18 A variety of Luvisols and Chernozems prevails in the watershed sections (Nosin, 1949; Vasil'eva, Baranova, 2007;
19 Abakumov, Gagarina, 2008; Abakumov et al., 2009; Urusevskaja et al., 2000). Whereas, Calcaric Chernozems
20 (southern type) dominate in the south of the region (steppe zone of the Samara region), accompanied with some
21 polypedons of Kastanozems (Nosin, 1949). Soils of investigated area are sandy and sandy loam textured. In the
22 studied area, soils were classified as sandy loam soils on Late Pleistocene alluvial Volga sands – Protoargic
23 Arenosols according to World reference base (2015), and they have weak features of illuvial phenomena without
24 formation of separate horizons. Sand content in these soils is 70.5-86.4%; clay content is 0.3-2.6 %.

25 Three soil pits were sampled in two different fire affected areas: one in a site affected by a surface forest fire and
26 another affected by a crown fire. A sample was collected in an unburned area to serve as control. Soil types and
27 vegetation were the same. Soils were sampled as fast as possible after the removal of a state of emergency from
28 the territory in summer of 2010, and also during the period 2011-2015. Three soil pits were sampled at each studied
29 area made at each plot. Undisturbed soil samples were collected using Kubiena-type boxes of 5x3.5x1.5 cm sizes
30 at the depth of 0-10 cm and taken to the laboratory.

31 **2.2 Micromorphological analysis**

32 Soil samples were air-dried and being passed through a 1-mm sieve. Fine sections of soil material were prepared
33 form micro monoliths of soils, sampled in field. Samples were dried and saturated with resin.

34 Thin sections were investigated with use of polarization microscope Leico DFC 320 in transmitted light and
35 crossed nicols. The following soil micromorphological indexes have been investigated: soil microfabric, spatial
36 arrangements of fabric units, soil particles distribution, elements of microstructure and character of organic matter.



1 Terminology, used in this paper are published by Stoops (2003), also by Gagarina (2004) manuals and Gerasimova
 2 (2011) review, where details of micro organization of soil were described in details.

3 Results and discussion

4 3.1 Soil profile analysis and physico-chemical properties

5 The soil profile organization can be described as: A – AC – C. The A horizon has a black layer due the deposition
 6 of an ash and charcoal layer on soil surface, contrary to the observed in the control plot (fig. 1,2).



7
 8 **Figure 1: Fire affected soil.**



9 **Figure 2: Control soil.**

10 Data on the general characteristics of the soils are given in Table 1.

11 Organic matter is lost from the surface horizons of the soil, which is related to the destruction of the organic
 12 horizons, the mineralization of root residues, and the almost complete absence of fresh plant waste, which could
 13 be a material for humification. In 2010, the content of organic matter in the ash on the soil surface after the surface
 14 fire ($2.31 \pm 0.27\%$) was lower than that after the crown fire ($3.19 \pm 0.19\%$). A similar tendency was observed in the
 15 pyrogenically transformed humus-enriched horizons: the content of organic carbon was $1.21 \pm 0.50\%$ after the
 16 surface fire and $1.42 \pm 0.31\%$ after the crown fire. Thus, a surface fire, which leads to the complete burning out of
 the litter and the upper horizons, results in larger losses of organic matter.

17 The acidity in the upper horizons of the burnt soils decreases significantly, and the burnt litters have an alkaline
 18 reaction (pH 7.9-8.0), while the lower horizons have a weakly acid reaction close to that in the corresponding
 19 horizon of the undisturbed forest soil (pH 5.7-5.9). The increase in the pH values of the soils after fires is related
 20 to the fact that the water-soluble ash components penetrate into the soil and saturate the soil exchange complex
 21 with alkaline-earth elements, which shifts the reaction toward the neutral value. No differences in the pH changes
 22 between the effects of the crown and surface fires were observed in the first year of the study. The reaction of the
 23 parent rock is similar in all three plots and is characterized as weakly acid. The partial or sometimes complete
 24 mineralization of organic residues because of pyrolysis resulted in the synchronous input of ash elements onto the
 25 soil surface and into the litter, which neutralized the organic acids arriving in the soil solution during the
 26 decomposition of the litters. It is therefore obvious that the higher the ash yield (i.e., the more intensive the fire),
 27 the more complete and active the neutralization of the litters.

28 A year after the fire, the pH of the burnt litters decreased from 7.9-8.0 (in 2010) to 6.4-6.6 (in 2011), and its
 29 absolute values approached the control level. This is easily explicable: rain and snowmelt waters almost completely



1 removed the soluble ash components for a year; i.e., the alkali elements were removed from the ash at the fire
2 sites.

3 One-Way Anova test has shown that significant differences were revealed for carbon content between surface fire
4 and crown fire ($p < 0.02$) and also between control and surface fire plot ($p < 0.01$). As for silt-and-clay fraction there
5 were differences only between control and surface fire ($p < 0.04$). The same situation was characteristic for sand
6 fraction ($p < 0.02$).

7 The WRB system (World reference base..., 2015) does not have any horizons or diagnostic parameters of
8 pyrogenic soils. But WRB system has Pretic horizon as dark, high content of organic matter and phosphorous, low
9 biological activity, high contents of exchangeable calcium and magnesium, with remnants of charcoal and/or
10 artefacts. A pyrogenic horizon with abundance of charcoal is formed after wildfires. It can resist to degradation
11 when vegetation has not started to recover yet at burned places and charcoal has not started to redistribute while
12 erosion and infiltration processes. However, black carbon decomposition is controversial, and there are different
13 views about this issue. Some studies argue that black carbon decomposes very slowly (Liu et al., 2008) or
14 practically non-degradable (González-Pérez et al., 2004), while others show that it successfully affected by
15 chemical (Cheng C.H. et al., 2006, 2008) and microbial (Knicker et al., 2013; Marschner et al., 2008) oxidation.
16 The assumption of black carbon complete stability in soils is doubtful because its content varies considerably in
17 different soils that is explained not only by a difference of pyrogenic activity in different natural zones, but also
18 by a difference of humidity (Nguyen, Lehmann, 2009) and temperature (Cheng C.H. et al., 2008; Nguyen et al.,
19 2010), by various physical and chemical soil characteristics, different biological activity and land use practices
20 (Czimczik, Masiello, 2007).

21 **3.2 Micromorphological characteristics**

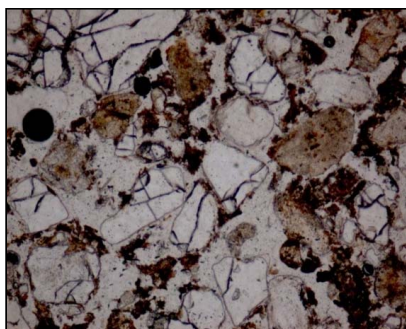
22 Differences of postpyrogenic and unburned soils are well-shown in morphological, chemical, physical and
23 biological properties of horizon A. Morphological organization of solum in burned soils differ from unburned in
24 a number of parameters: wide distribution of charcoal pieces, absence of litter and its transformation into ash that
25 is a mix of mineral soil components, burned-down plant residues, and small pieces of charcoal (Bodi et al., 2014;
26 Pereira et al., 2014, 2015); an also reduction of the humic horizon's depth. There is a soil erosion because of
27 rainfall characterized by decrease of black surface horizon thickness after several years of investigation as observed
28 in early studies (Francos et al., 2016).

29 Data on soil micro morphological features are presented in table 2, fig. 3. The results showed that elementary
30 assemblage of crown fire soil is plasma-silt (with small content of plasma and prevalence of silt particles). In case
31 of the soil affected by surface fire, the type of assemblage is plasma-silt, which is a result of the accumulation of
32 humus-type organic matter. In both cases the type of the plasma classified as humus-clay. The structure of all 3
33 soils investigated is crumb, inherited from the previous stages of soil formation. There are not evident features of
34 mass aggregation in the fire affected soils. The type of the microstructure of all the soils is angular blocky or
35 subangular blocky, which is caused by low intensity of current weathering in soil mass. The particle shape is
36 subidiomorphic or idiomorphic in cases investigated with weak or medium degree of corrodness.

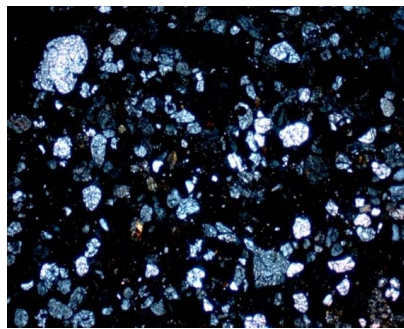
37 Mineral composition of the soils can be described as follows: uncoloured minerals, quartz, orthoclase and
38 carbonates (crown fire); quartz and orthoclase with many crystals of muscovite with absence of carbonates (surface
39 fire) and predominance of uncoloured minerals with quartz and muscovite in case of the control plot. The



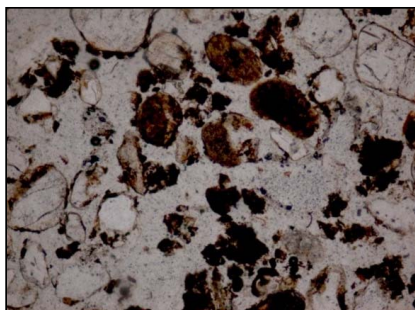
1 possibility of new minerals formation under the strong heating effect has been reported in previous works (Nobles,
2 2010; Leon et al., 2014), however, carbonate accumulation is not resulted from heating process in scenario of the
3 crown fire. In some arrangements, soil microstructure can be classified as skeletal, which is caused by high content
4 of not weathered soil particles. No evidences of current weathering (alteration) have been fixed, but some pores
5 infilling were recognized as a result of ash and charred material accumulation in fire affected soils (fig. 4). Not
6 decomposed organic tissues and residues infills the porous media in these soils. This is result of increment of raw
7 forms of organic material in burned soils. Organic matter under the effect of fire was polymorphic (fig. 4, a, b)
8 (Stoops, 1986), while it is monomorphic in the soils of the control plot. The porous media infilling after the fires
9 was described previously by Nobles (2010), however in this case was an accumulation of Mn-Fe enriched
10 materials. Balfour and Woods (2007) observed similar results in fire affected soils. So, taking our data into account
11 it is possible to conclude that infilling of porous media by material of different composition is typical in burned
12 soils. Decreasing of the porous media area was also described as result of accumulation of ash (Balfour, Woods,
13 2007), and this explanation of porous media infilling is more appropriate to our case. Investigated soils
14 characterizes by developed system of porous media, this is important in sense of heating penetration into soil and
15 soil sustainability to heating. Soils investigated are not so sustainable to heating than clay textures ones, where the
16 porous media is not so developed. Porous, cracks and other forms of space give the possibility to the combustion
17 products penetrate into the soil and affect the PAHs and other products accumulations.



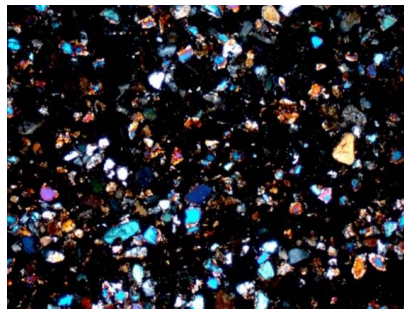
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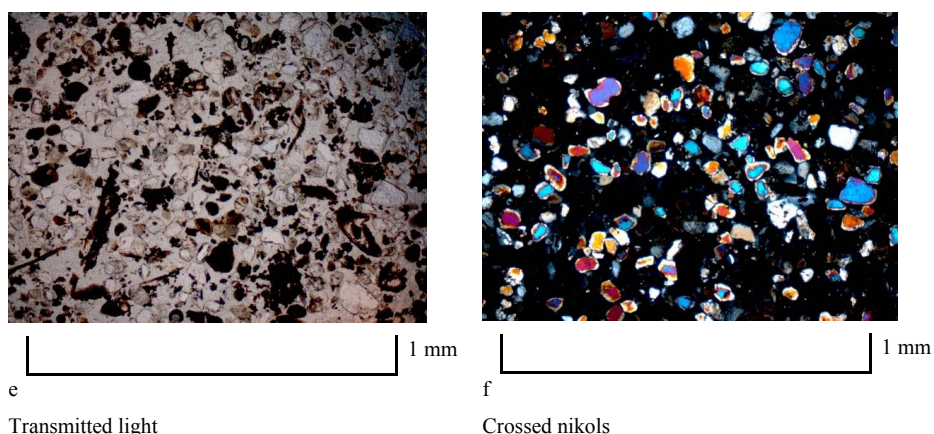
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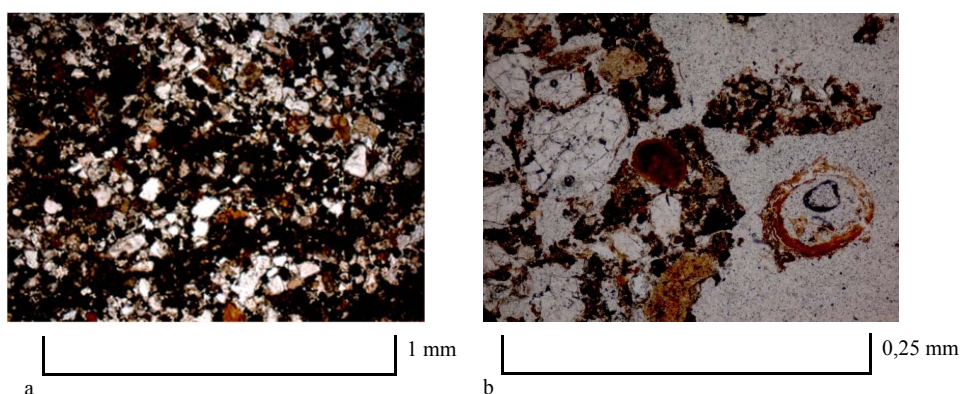
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1 **Figure 3: Thin sections. a, b – crown fire, c, d- surface fire, e, f- nature plot. Left column if transmitted light, right**
2 **column – crossed nikols.**



3 **Figure 4: a – humons (organic particles with not humified organic matter), transmitted light; b – organic tissues of not**
4 **decomposed remnants.**

5 The type of fire affected the soil organic micromorphology. The quantity of not humified organic matter in the
6 burned soils, especially in the samples affected by surface fire was identified as increased in burned soils (fig. 4,
7 a). It is also evident that the size of organic residues is higher in fire affected soils than in mature ones. In mature
8 soils transformation and humification of organic matter is gradual process, while new, relatively fresh not
9 decomposed organic matter has come into upper soil horizons in fire affected areas (Gagarina, 2004).

10 **4 Conclusions**

11 The 2010 catastrophic natural fires in the urban forests of Tol'yatti resulted in the formation of pyrogenically
12 transformed soils, the morphological parameters and the main chemical and physical properties of which
13 significantly differ from those of the undisturbed soils.

14 The burnt soils differ from the control soil on the macromorphological level only in the upper part of the profile,
15 where the litter is transformed to ash identified as a dim-gray organomineral mixture. Processes of soil erosion are
16 clearly manifested a year after the fire under the effect of precipitation and the illuviation of organic matter to the
17 medium part of the profile and will probably continue for several years.



1 The fires significantly affect the physicochemical and chemical properties of the soils. However, the effect of fires
2 on the properties of the studied soils usually does not spread deeper than 10 cm.
3 The results of this work showed that mineral composition of all the soils studied is presented by quartz, feldspar
4 (orthoclase); in case of surface forest fire there was more mica (muscovite), and calcium carbonates appeared in
5 soils affected by crown forest fire, the reasons of this compound accumulation are still not well understood. The
6 type of plasma is humus-clay, but the soil assemblage is plasma-silt with the prevalence of the silt. Angular and
7 subangular grains form the main soil carcass and no evidence of intensive weathering alteration has been revealed.
8 At the same time, decreasing of the porous media was recognized as a main soil development process after the
9 fires. This is result of soil porous infilling by ash and charred organic material of different nature: some organic
10 remnants (tissues) come into the porous media after fire and some transformed, coaled and dark coloured part also
11 appears in post fire horizons. Partially decomposed fire affected particles of soil organic matter accumulate in
12 postfire soils, especially in soil porous media, which is a result of the soil organic matter accumulation and
13 transformation in postfire environments.

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 1 **Table 1. Morphological features and general properties of soils, ± after the mean value means SD**

Horizon	Depth, cm	Colour	Soil humidity, %	pH	Total organic carbon, %	Clay content, g·kg ⁻¹	Sand content, g·kg ⁻¹
Surface fire							
Burned litter	0-5	10 YR 3/2	2.85±0.79	8.0±0.06	2.31±0.27	20.0	705.0
AY	5-14	10 YR 4/2	1.45±0.19	6.2±0.32	1.21±0.50	25.0	807.0
AC	14-27	7.5 YR 5/4	1.38±0.36	6.0±0.21	0.75±0.48	26.0	835.0
AC	27-36	10 YR 6/4	1.02±0.27	5.8±0.21	0.31±0.13	19.0	852.0
AC	36-53	10 YR 6/4	0.98±0.42	5.3±0.31	0.22±0.05	24.0	866.0
C	53-73	7.5 YR 5/4	0.69±0.05	5.7±0.21	0.24±0.10	13.0	864.0
Crown fire							
Burned litter	0-3	10 YR 3/2	2.37±0.36	7.9±0.12	3.19±0.19	17.0	720.0
AY	3-10	10 YR 4/2	1.43±0.35	5.9±0.38	1.42±0.31	20.0	788.0
AC	10-15	7.5 YR 5/2	0.86±0.20	5.9±0.25	0.78±0.07	17.0	852.0
AC	15-24	10 YR 7/4	1.11±0.63	5.9±0.36	0.26±0.07	6.0	862.0
C	24-44	10 YR 6/3	0.52±0.05	5.7±0.12	0.14±0.05	4.0	867.0
C	44-64	10 YR 6/3	0.49±0.03	5.9±0.25	0.12±0.05	9.0	868.0
Control							
Litter	0-7	-	5.92±2.27	6.5±0.10	nd	nd	nd
AY	7-10	10 YR 4/2	1.60±0.47	6.3±0.06	1.94±1.35	21.0	787.0
AY	10-14	10 YR 6/4	0.78±0.15	6.2±0.23	0.78±0.33	18.0	837.0
AC	14-23	7.5 YR 3/2	0.78±0.42	6.1±0.23	0.33±0.10	13.0	867.0
AC	23-33	2.5 YR 8/6	0.42±0.01	5.9±0.20	0.15±0.02	4.0	888.0
C	33-50	2.5 YR 8/6	0.33±0.02	5.7±0.12	0.21±0.03	7.0	891.0
C	50-70	2.5 YR 8/6	0.32±0.03	5.8±0.31	0.15±0.05	3.0	895.0

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1 Table 2. Micromorphological characteristics of the soils under investigation

Horizon	Sampling depth (cm)	Skeleton grains							Type of plasma	Elementary assemblage	Plant residues		Structural separations	Soil microstructure	Coprofites			
		%	Character of distribution	Size, mm		Size sorting	Particle shape	Rounding			Corrodness	Mineral composition			%	Decomposition degree	Voids	Size, mm
Apir	0-15	40-50	Uniform	0.07/0.01	0.04	good-medium	subidiomorphic, xeromorphic, regenerate (some minerals in one)	rounded	weak	colorless minerals are prevail: quartz, feldspar (orthoclase), carbonates (calcispar)	humus-clay	plasma-silt (little plasma, most of all - silt particles)	10	weak-medium	aggregates	pore structure with separated aggregates		
Apir	0-14	40-50	Uniform	0.07/0.01	0.01	medium-weak	idiomorphic, subidiomorphic, orthorphic, xeromorphic	rounded-not rounded	medium	a lot of specular stone (muscovite), quartz, feldspar (orthoclase)	humus-clay	plasma-silt	5	weak	aggregates	pore structure with separated aggregates	0,4	brown, reddish, harsh with minerals (calcispar)
O	0-20	40-50	Uniform	0.09/0.01	0.02	good-medium	idiomorphic, subidiomorphic, orthorphic, xeromorphic	rounded-not rounded	weak	colorless minerals are prevail: quartz, feldspar (orthoclase), specular stone (muscovite)	humus-clay	plasma-silt (little plasma, most of all - silt particles)	30	weak-medium-strong is prevail	aggregates	pore structure with separated aggregates		

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