



1 **Spatial variability of some soil properties in west coastal area of**
2 **India having oil palm (*Elaeis guineensis* Jacq.) plantations**

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11 **Running title:** Soil property distribution in oil palm plantations of coastal India

12

13 **Abstract.** Mapping spatial variability of soil properties is the key to efficient soil resource
14 management for sustainable crop yield in coastal areas. Therefore, the present study was
15 conducted to assess the spatial variability of soil properties like – acidity (pH), salinity
16 (Electrical Conductivity (EC)), organic carbon, available K, available P, exchangeable Ca²⁺,
17 exchangeable Mg²⁺, available S and hot water soluble B in surface (0-20 cm) and subsurface (20-
18 40 cm) soil layers of oil palm plantations in south Goa and north Goa districts of Goa situated in
19 west coastal area of India. A total of 128 soil samples were collected from 64 oil palm
20 plantations of Goa located at an approximate interval of 5-7 km and analyzed. Soil was acidic to
21 neutral in reaction. Other soil properties varied widely in both the soil layers. Correlations
22 between soil pH and exchangeable Ca²⁺, between soil EC and available K, between available P



23 and available S and between exchangeable Ca^{2+} and exchangeable Mg^{2+} in both the soil layers
24 were found to be positive and significant ($P = 0.01$). Geostatistical analysis revealed different
25 spatial distribution pattern for the measured soil properties. Best fit models of measured soil
26 properties were spherical, linear, exponential, circular and Gaussian with weak to strong spatial
27 dependency. The results revealed that site-specific fertilizer management options needed to be
28 adopted in the oil palm plantations of the study area owing to variability in soil properties.

29 *Keywords:* Soil management, Spatial distribution, Precision agriculture, Soil fertility, Coastal
30 zone

31 **1 Introduction**

32 Soil is the key part of the earth system which controls hydrological, biological, and
33 geochemical cycles and it offers goods, resources and services to mankind (Keesstra et al., 2012;
34 Smith et al., 2015; Decock et al., 2015; Brevik et al., 2015; Berendse et al., 2015). Un-
35 sustainable soil management practices lead to soil degradation, which is a worldwide topic,
36 mainly because of loss of soil organic matter (SOM), soil erosion, changes in soil structure,
37 degradation of the biota in the soils and soil chemical degradation (Cerdeira et al., 2009; Mupenzi
38 et al., 2011; Novara et al., 2013; Mukherjee et al., 2014; Lieskovsky and Kenderessy, 2014;
39 Stanchi et al., 2015; Seutloali and Beckedahl, 2015; Novara et al., 2015;). Soil degradation along
40 with natural processes results in degradation of coastal areas, which covers more than 10% of the
41 earth surface area with 35, 6000 and 7517 km coast line in world and India respectively
42 (Misdorp, 1990; Sanil Kumar et al., 2006). Therefore, there is a need to describe and characterize
43 these areas for adoption of effective land use practices including application of agri-inputs
44 (Arakel et al., 1993; Guneroglu et al., 2015).



45 Geographical distribution maps of soil properties, obtained from soil surveys, help in
46 correct management of soil nutrients (Brevik et al., 2016). These maps are required to understand
47 the patterns and processes of soil spatial variability, which is the combined effect of soil
48 physical, chemical and biological processes operating at different spatio-temporal scales
49 combined with anthropogenic activities (Goovaerts, 1998). The distribution maps are prepared
50 by analysing spatial distribution pattern of soil properties. Geostatistical tools are useful in
51 preparation of the maps based on limited number of samples collected from agricultural
52 landscapes. These tools predict the values at un-sampled locations by spatial correlation and
53 reducing variance of estimation error and investigation costs (Saito et al., 2005; Pereira et al.,
54 2015). Spatial variability of soil properties is assessed effectively by geostatistical methods
55 (Mueller et al., 2003) for site-specific management of nutrients through variable rate fertilizer
56 application to avoid over and under application of nutrients. Li et al., 2011, Behera and Shukla,
57 2014 and Behera and Shukla, 2015 have reported different spatial variability pattern of soil
58 properties and soil nutrients in eroded areas of south China and some cultivated acid soils of
59 India. Information regarding variability of soil properties in soil profile is helpful to assess the
60 contribution of sub-surface soil layers to crop nutrition and potential capacity of the soil to
61 supply nutrients during crop growth. It also helps in understanding the effect of different
62 management practices, under a given cropping system, on the downward movement as well as
63 recycling of nutrients to the surface layers (Behera and Shukla, 2013, Parras-Alcantara et al.,
64 2015).

65 Oil palm (*Elaeis guineensis* Jacq.) is a high oil yielding crop (Lamade et al., 2015). On
66 average, it produces ten times more oil than any leading oil seed crop from a hectare of land and
67 some efficient farmers get as high as eight tonnes of oil yield per hectare. World-wide, oil palm



68 produces 32% oils and fats output from 5.5% land use for cultivation (Palm Oil Research, 2016).
69 Indonesia and Malaysia are the leading producer of oil palm. According to Rethinam et al.
70 (2012), oil palm can be cultivated as irrigated crop in 1.93 million ha area in 18 states of India.
71 At present, oil palm is being grown in an area of about 2, 68, 000 ha covering twelve states of the
72 country, having different soil types, with productivity levels reaching as high as 30-35 Mg fresh
73 fruit bunches (FFB) ha⁻¹ year⁻¹ (Kalidas et al., 2015).

74 Rationale use of fertilizer results in environmentally sustainable and economically viable
75 oil palm yield (Goh et al., 2003). Oil palm uses about 162, 30, 217, 38 and 36 kg of N, P, K, Mg
76 and Ca ha⁻¹ year⁻¹ respectively, to produce 2.5 Mg of oil ha⁻¹ year⁻¹ (Mengel and Kirkby, 1987).
77 Considering oil to bunch ratio of 1:4, 2.5 Mg oil ha⁻¹ is equivalent to 10 Mg FFB ha⁻¹ year⁻¹, but
78 average FFB yield in well-managed plantations is much higher (Narsimha Rao et al., 2014).
79 Nutrient content in 1 Mg of FFB obtained from Dura palms is 2.94, 0.44, 3.71, 0.77, 0.81 kg of
80 N, P, K, Mg and Ca, respectively, whereas, Mn, Fe, B, Cu and Zn content per 1 Mg of FFB is
81 1.51, 2.47, 2.15, 4.76 and 4.93 g, respectively of Mn, Fe, B, Cu and Zn (Ng and Thamboo,
82 1967). According to Narsimha Rao et al. (2014), nutritional problems like N/K imbalance, K
83 deficiency, Mg deficiency and B deficiency affect oil palm production in oil palm plantations of
84 India. Calibrated soil and leaf analysis helps in effective fertilizer recommendations in most of
85 the crops (Smith and Loneragan, 1997; McLaughlin et al., 1999). In oil palm, leaf nutrient
86 analysis is commonly used for estimating fertilizer requirement (Fairhurst and Mutert 1999;
87 Corley and Tinker, 2003). The relationship between leaf analysis and palm productivity is
88 generally evident, and an assessment of fertilizer needs can be based on such an analysis.
89 However for a cost-effective approach, leaf analysis has to be integrated with soil analysis (Goh



90 et al., 2009). It is therefore pertinent to assess soil and leaf nutrient status for effective and
91 sustainable fertilizer management programme in oil palm.

92 The nutrient management recommendations in oil palm plantations in India in general
93 and oil palm plantations in the area under study are generic ones. Prasad et al. (2013) reported
94 wide range in quantity of fertilizer applied indicating that oil palms were either under-fertilized
95 or over-fertilized. Also, low cost and easy availability of some fertilizers have encouraged
96 farmers to make excessive applications with the belief that high yields would be ensured.
97 However, this management adversely affects soil fertility, productivity, fruit quality and ground
98 water quality. It is therefore pertinent for the farmers to economize on fertilizer adopting a
99 strategy for site-specific and/or area-specific management based on spatial variability of soil
100 properties to make oil palm production environmentally sustainable and economically viable.
101 Spatial variability of soil properties in oil palm plantations have to be carefully evaluated to
102 carryout sustainable soil management practices. Thus, the present study was carried out in soils
103 of oil palm plantations of Goa state of India with the following objectives, (i) to estimate the
104 spatial variability of some soil properties through semivariogram analysis, (ii) to develop spatial
105 maps for soil properties using the parameters of the best fitted semivariogram model and
106 interpolation by ordinary kriging technique and (iii) to assess the relationship among the
107 estimated soil properties.

108 **2 Material and methods**

109 **2.1 Study site**

110 A survey was carried out in Goa state of India during 2012-13 to find out soil and plant
111 nutritional status in randomly selected 64 tenera oil palm plantations (with 5 to 21 years of age)



112 (Figure 1). Oil palm is cultivated in an area of approximately 1000 ha which is 1% of agricultural
113 land in the state. The sampling area lies between $15^{\circ} 6.8 96$ N to $15^{\circ} 41.7 26$ N latitudes and
114 $74^{\circ} 76 60$ to $73^{\circ} 56 78$ E longitudes with altitude ranging from 4 to 90 meter above sea level.
115 The climate of the area is tropical monsoon type. Hot and humid climate prevails for most of the
116 year. Annual mean rainfall (average of 30 years) is 2926 mm, concentrated from early June to
117 late September. On average, May is the warmest month, with temperature peaks over 35°C
118 (during 24 h) and relative humidity of 70%. Goa experiences short winter seasons between mid-
119 December and February and these months are marked by mean night temperature of
120 approximately 21°C and mean day temperature of around 28°C with relative humidity of
121 65%. According to Bhattacharyya et al., (2013), the main soils in the study area are Inceptisols
122 (26, 000 ha), Ultisols (4, 000 ha), Entisols (3,000 ha) and Alfisols (3, 000 ha) (classified as in
123 Soil Survey Staff, 2014), sandy loam to silty loam texture, developed from granite, granite-
124 gneiss, quartzite/schistose and basalt.

125 **2.2 Soil sampling, processing and analysis**

126 A total of 128 soil samples i.e. 64 from 0-20 cm (surface) and 64 from 20-40 cm (sub-
127 surface) depths were collected at random points inside 3-m radius from the palm during the
128 survey to assess soil properties of oil palm plantations at an approximate interval of 5 to 7 km.
129 All the samples were collected with a hand auger. The latitude, longitude, and elevation at each
130 sampling point were recorded using a hand held global positioning system (GPS). The soil
131 samples were dried at room temperature ($25 \pm 3^{\circ}\text{C}$). Roots and debris were removed from the
132 samples by hand. Samples were processed following standard procedures. The processed soil
133 samples were tested for acidity (pH), salinity (EC), organic carbon (OC) content, available K
134 ($\text{NH}_4\text{OAc-K}$), available P (Bray's P-1) (Bray's-P), exchangeable Ca^{2+} , exchangeable Mg^{2+} ,



135 available S ($\text{CaCl}_2\text{-S}$) and hot water extractable B (HWB). Determination of soil pH and EC
136 (1:2 soil water ratio (w/v) suspension) were carried out using pH-meter and conductivity meter
137 (Jackson, 1973). Walkley-Black method (Walkley and Black, 1934) was followed for assessing
138 soil OC content. $\text{NH}_4\text{OAc-K}$ was estimated after extracting soil samples with neutral 1 N
139 ammonium acetate solution (Hanway and Heidel, 1952) followed by flame photometry
140 estimation. Available P was extracted using Bray's P-1 reagent (Bray and Kurtz, 1945) and
141 estimated through spectrophotometry. Ca^{2+} and Mg^{2+} were extracted using neutral normal
142 ammonium acetate solution (Jones, 1998) and estimated through atomic absorption spectrometry.
143 Available S was estimated by the turbidity method (Williams and Steinbergs, 1969). HWB
144 content was estimated through Azomethine-H reagent (Gupta, 1967) using spectrophotometry.

145 **2.3 Statistical and geostatistical analysis**

146 The descriptive statistics like minimum, maximum, mean, standard deviation (SD),
147 coefficient of variation (CV), and skewness for soil properties were computed using the SAS 9.2
148 software pack (SAS, 2011). Relationship among the estimated soil properties were established
149 using Pearson's correlation coefficient analysis at $p = 0.05$ and $p = 0.01$.

150 ArcMap 10.1 (ESRI, 2012) was used to analyze the spatial structure of soil properties.
151 Before using geostatistics, normality of data distribution were checked by Shapiro-Wilk test at 5%
152 (Shapiro and Wilk, 1965). Soil properties like pH and OC content in both the soil layers and $\text{CaCl}_2\text{-}$
153 S content in subsurface soil layers exhibited normal distribution (Table 1). While, data
154 transformation to normal distribution was carried out for rest of the soil properties. Trend of the
155 data set was checked and removed. The semivariogram models of soil properties were derived as
156 described by Goovaerts (1997) and Tesfahunegn et al. (2011).



$$\gamma(h) = \frac{1}{2m(h)} \sum_{i=1}^{m(h)} [Z(X_i + h) - Z(X_i)]^2 \quad (1)$$

157
158 Where $\gamma(h)$ is the experimental semivariogram, $m(h)$ is number of sample value pairs,
159 $Z(X_i)$, $Z(X_i+h)$ are sample values at two points. Best fitted semivariogram model for each soil
160 property was selected by using the cross validation technique.

161 Semivariogram parameters like nugget/sill ratio and range were obtained for soil
162 properties. The nugget/sill ratio expressed in percentage was used to classify the spatial
163 dependence of variables (Oliver and Webster, 2014). Ratio values less than or equal to 25%,
164 between 25 and 75%, more than 75% were considered strongly, moderately and weakly spatially
165 dependent, respectively (Behera et al., 2011). Best-fit semivariograms models were selected by
166 cross-validation technique. Mean square error (MSE) was estimated to predict the accuracy of
167 models (Utset et al., 2000).

$$\text{MSE} = \frac{\sum_{i=1}^n [Z(x_i, y_i) - Z^*(x_i, y_i)]^2}{n} \quad (2)$$

169 Accuracies of interpolated maps were checked by the goodness-of-prediction criterium G
170 (Agterberg, 1984). According to Parfitt et al. (2009), positive and negative and close to zero
171 values of G indicate that the map obtained by interpolating data from the samples is more
172 accurate than average value of the area and the average value predicts the values at un-sampled
173 locations as accurately as or even better than the sampling estimates, respectively. Ordinary
174 kriging interpolation was carried out to develop spatial distribution maps for soil properties.

175 3 Results and discussion

176 3.1 Descriptive statistics of soil properties



177 The descriptive statistics revealed considerable variability of soil properties in both
178 surface and sub-surface soil layers of oil palm plantations (Table 1). The values of CV for soil
179 acidity in both the soil layers revealed their low variability (CV < 10%) (Nielsen and Bouma,
180 1985). The rest of the soil properties exhibited moderate (CV 10 to 100%) variability except
181 salinity in surface soil layers and Bray's-P in both the soil layers, which had high (CV > 100%)
182 variability. Low CV values for soil acidity was due to transformed measurement of hydrogen ion
183 concentration. Skewness coefficient values of 0.18 to 3.89 for different soil properties revealed
184 that some soil properties were not normally distributed. This variation and non-normal
185 distribution of soil properties in the studied areas may be due to adoption of different soil
186 management practices including variation in fertilizer application and other crop management
187 practices (Tesfahunegn et al., 2011; Srinivasarao et al., 2014; Ferreira et al., 2015).

188 The mean values of soil pH were acidic in both surface (5.35) and subsurface (5.28) soil
189 layers (Table 1). The acidic nature of soil in the studied area may be due to acidic parent material
190 and prevailing rainfall pattern. The values of soil EC indicate the non-saline nature of soils. Soil
191 OC contents varied widely in both surface and subsurface soil layers. Principal reason for
192 variation in soil OC content may be due to adoption of different cultural practices including
193 addition of crop biomass to the soils. Surface soil layers had slightly higher OC content (mean
194 value 19.8 g kg⁻¹) than OC content in subsurface soil layers (mean value 13.2 g kg⁻¹). Surface
195 soil layers had higher NH₄OAc-K, Bray's-P, CaCl₂-S and HWB content compared to that in
196 subsurface soil layers (Table 1). The content of these nutrients varied greatly among the soils
197 because of heterogeneity in fertilizer application in the area. The mean values of exchangeable
198 Ca²⁺ were 914 and 795 mg kg⁻¹ for surface and subsurface soil layers, respectively. Whereas,
199 surface soil layers were having 203 and 225 mg kg⁻¹ of mean exchangeable Mg²⁺ content,



200 respectively. Other studies reported similar results highlighting different distribution pattern of
201 soil properties, primary, secondary and micronutrients under different soil-crop management
202 situations (Franzlubbers and Hons, 1996; Sharma et al., 2005; Behera and Shukla, 2013).

203 **3.2 Relationship among soil properties**

204 The exchangeable Ca^{2+} content increased with pH (Table 2). Behera and Shukla (2015)
205 also recorded positive and significant relationship of soil pH and soil OC with K, exchangeable
206 Ca^{2+} and exchangeable Mg^{2+} content in some cropped acid soils of India. Soil OC content in
207 surface layers was positively and significantly correlated with exchangeable Ca^{2+} and HWB (P
208 0.05). Most of the soil properties which influence nutrient storage and availability to plants are
209 influenced by soil organic matter (SOM) type and content (Foth and Turk, 1972). Increased soil
210 EC content led to higher $\text{NH}_4\text{OAc-K}$ in both soil layers ($P = 0.01$), and higher $\text{CaCl}_2\text{-S}$ in surface
211 layer and Bray's-P in subsurface layer ($P = 0.05$). Soil EC does not directly affect plant growth
212 but has been used as an indirect indicator of the amount of nutrients available for plant uptake
213 and salinity levels (Corwin and Lesch, 2005). EC has been used as a surrogate measure of salt
214 concentration, organic matter, cation-exchange capacity, soil texture, soil thickness, nutrients,
215 water-holding capacity, and drainage conditions. In site-specific management and high-intensity
216 soil surveys, EC is used to partition units of management, differentiate soil types, and predict soil
217 fertility and crop yields.

218 **3.3 Spatial structure and distribution of soil properties**

219 The best-fit semivariogram models and parameters of studied soil properties are given in
220 Table 3. The best fit models for soil properties of studied areas were spherical, linear,



221 exponential, circular, and Gaussian depending on soil layer and parameter. Our findings are in
222 line with the observations made by Tesfahunegn et al. (2011).

223 Cross-validation technique was used to select semivariograms models for soil properties
224 with the lowest MSE values (Table 3). Lowest MSE values indicate that kriging predictions of
225 soil properties are closer to measured values. The accuracy of kriged interpolation maps of soil
226 properties was also measured by the G values (Table 3) which varied from 26 (for exchangeable
227 Ca^{2+} in subsurface layer) to 76% (for HWB in subsurface layer). Positive G values for all the soil
228 properties revealed the developed maps are more accurate than the maps generated using the
229 average value of the area.

230 Soil pH in both the soil layers of oil palm plantations was having moderate spatial
231 dependency class. Soil EC had strong and moderate spatial dependency for surface and sub-
232 surface soil layers respectively. Soil OC content in oil palm plantations had weak spatial
233 dependency for surface soil layers and moderate spatial dependency for sub-surface soil layers.
234 Spatial dependency classes were weak for $\text{NH}_4\text{OAc-K}$ and strong for exchangeable Ca^{2+} for both
235 the soil layers of oil palm plantations. Bray's-P and $\text{CaCl}_2\text{-S}$ had weak spatial dependency for
236 surface layers and moderate spatial dependency for sub-surface layers. Whereas, exchangeable
237 Mg^{2+} and HWB had moderate spatial dependency in surface soil layers and weak spatial
238 dependency for sub-surface soil layers in oil palm plantations. Weak spatial dependency of soil
239 properties like $\text{NH}_4\text{OAc-K}$, Bray's-P and OC (in surface layer) in oil palm plantations is
240 ascribed to the anthropogenic activities like adoption of cultural practices including application
241 of fertilizers. In these oil palm plantations, activities like application of irrigation water, weeding,
242 basin cleaning, mulching and application of N, P, K and Mg fertilizer are carried out at regular



243 intervals. Whereas, moderate and strong spatial dependency of soil pH, EC and exchangeable
244 Ca^{2+} is due to soil type and parent material.

245 According to Webster and Oliver (1990), range value is a measure of the spatial
246 extension within which autocorrelation exists. Spatially related samples were separated by
247 distances closer than range values. The range values of soil properties in studied area varied
248 widely (Table 3). The range values for surface soil properties were 554 to 4530 m and for
249 subsurface soil properties were 581 to 4530 m. Among the soil properties, higher range values
250 were recorded for $\text{NH}_4\text{OAc-K}$ and $\text{CaCl}_2\text{-S}$ for both the soil layers. The possible causes for
251 spatial variability of soil properties in studied areas are adoption of different soil management
252 practices (Bodi et al., 2013; Pereira et al., 2013; Ochoa-Cueva et al., 2015). The difference in
253 annual average temperature in the state of Goa was more than 12 °C, indicating temperature
254 could be important factor influencing soil nutrient mineralization and accumulation. Moreover,
255 this area is having rising slope from the coast line towards *ghats* i.e. from western side to eastern
256 side, which could also affect distribution of nutrients probably wash by surface runoff or
257 subsurface water movements.

258 Interpolation maps (Figure 2) of different soil properties revealed that oil palm
259 plantations of the area could be divided into homogenous small zones depending upon the
260 different nutrient ranges. Overlying of the spatial distribution maps on map of Goa revealed that
261 the spatial distribution map of pH in surface soil layers revealed almost all the area having pH of
262 5.00 to 6.00. Low pH values occurred in north-western and south-eastern parts. In sub-surface
263 soil layers, low pH of 4.75 to 5.00 occurred in south-eastern part whereas relatively higher pH
264 prevailed in north-western part. Areas having low pH values compared to other areas may be due
265 acidic parent material from which the soil developed and different soil management practices.



266 Accordingly, different management options may be adopted in different parts of the area with
267 different levels of pH. Soil EC had irregular distribution pattern in surface soil layers whereas
268 low values of EC were recorded in north-western part. This may be due to sandy loam soil
269 texture and presence of low OC in north-western part. Higher EC values in other parts of
270 surveyed area probably due to silt loam soil texture with high water table. Higher amount of soil
271 OC was found to be distributed in the south-eastern parts in surface as well as sub-surface soil
272 layers. This may be ascribed to prevalence of higher slope and low rate of SOM mineralization
273 in south-eastern parts compared to other areas. Higher amount of NH_4OAc -K and CaCl_2 -S was
274 found to be distributed in almost all parts in both the soil layers. Higher amount of Bray's-P was
275 found to be distributed in almost all parts in surface soil layers whereas low amount of Bray's-P
276 occurred in north-central and south-western parts in sub-surface soil layers. Build up of P in
277 surface layers may be due to continuous P addition and their fixation in soil which is acidic in
278 nature. Exchangeable Ca^{2+} exhibited irregular distribution pattern in both the soil layers. In
279 surface as well as sub-surface soil layers, lower amount of Exch. Mg was found to be distributed
280 in southern parts as compared to that in northern parts. Similar distribution of exchangeable Ca^{2+}
281 and exchangeable Mg^{2+} was recorded in these soils which corroborate our finding of significant
282 and positive correlation between exchangeable Ca^{2+} and exchangeable Mg^{2+} in both the soil
283 layers. Higher amount of HWB was found to be distributed in north-eastern part in surface soil
284 layers and in central and south-western parts in sub-surface soil layers. The different distribution
285 variability of the soil properties in oil palm plantations of this area is predominantly due to
286 climate and landscape along with farm practices including application of different quantities of
287 nutrients through fertilizers. The kriged distribution maps for different soil properties providing
288 quantitative information about soil properties in both the soil layers is of great use for plantation



289 staff, farm managers, extension officers and farmers. This will help in visualizing soil fertility
290 status for planning appropriate strategies for efficient site specific soil nutrient management and
291 variable-rate fertilizer application technology. It leads for obtaining optimum output and oil palm
292 yield which can provide environmentally sustainable maximum return to famers with optimum
293 input utilization combined with best management practices (Fu et al., 2010; Behera et al., 2012).
294 The areas with low and medium nutrient status require more amount of fertilizer application as
295 compared to areas having high nutrient status. For example, exchangeable Mg^{2+} status is low in
296 southern part of the area compared to northern part. Therefore, the requirement of Mg fertilizer
297 application is more in southern part compared to northern part.

298 **4 Conclusions**

299 Geostatistical analysis is the key for studying the spatial variability of soil properties for
300 sustainable soil resource management. The present study divulged that the measured soil
301 properties had large variability in spatial distribution pattern in both surface and subsurface soil
302 layers of oil palm plantations of the studied area. Positive and significant correlations were
303 recorded between soil pH and exchangeable Ca^{2+} , soil EC and NH_4OAc-K , Bray's-P and $CaCl_2-$
304 S and exchangeable Ca^{2+} and exchangeable Mg^{2+} in both the soil layers. The prediction maps
305 generated by geostatistical analysis are useful for site-specific soil nutrient management in oil
306 palm plantations of the area by delineating management zones and adoption of variable fertilizer
307 application strategies.

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Table 1. Soil properties of surface (0-20 cm) and sub-surface (20-40 cm) layers (n = 64 at each case).

Variable	Soil layer	Mean ± SD	CV (%)	Minimum	Maximum	Skewness coefficient	Distribution
pH	Surface	5.35± 0.45	8.64	4.25	6.77	0.18	Normal
	Subsurface	5.28± 0.46	8.63	4.53	6.52	0.65	Normal
EC	Surface	0.13±0.17	125	0.05	1.06	4.06	Transformed
	Subsurface	0.08±0.06	75.3	0.03	0.41	3.02	Transformed
OC	Surface	19.8±8.77	44.4	5.07	48.4	0.83	Normal
	Subsurface	13.2±7.33	55.5	1.95	31.2	0.75	Normal
NH ₄ OAc-K	Surface	270±29.9	88.7	58.1	1167	1.80	Transformed
	Subsurface	199±165	82.8	16.1	856	2.16	Transformed
Bray's-P	Surface	24.7±3.39	127	0.86	141	2.14	Transformed
	Subsurface	9.78±13.2	135	0.90	42.3	2.52	Transformed
Ca ²⁺	Surface	914±588	64.3	200	2997	1.56	Transformed
	Subsurface	795±724	91.1	194	5177	3.89	Transformed
Mg ²⁺	Surface	203±141	69.3	36.0	744	1.75	Transformed
	Subsurface	225±156	69.4	24.0	720	1.27	Transformed
CaCl ₂ -S	Surface	23.2±16.4	70.7	3.00	87.7	1.60	Transformed
	Subsurface	16.3±10.1	62.0	1.50	43.5	0.93	Normal
HWB	Surface	0.70±0.38	54.7	0.09	2.10	1.43	Transformed
	Subsurface	0.64±0.44	68.6	0.04	2.56	1.70	Transformed

SD-standard deviation; CV-coefficient of variation; EC-electrical conductivity, dS m⁻¹; OC-organic carbon, g kg⁻¹; K, mg kg⁻¹; P, mg kg⁻¹; exchangeable Ca²⁺, mg kg⁻¹; exchangeable Mg²⁺, mg kg⁻¹; S, mg kg⁻¹; HWB, hot water soluble B, mg kg⁻¹.



Table 2. Pearson's correlation coefficients between soil properties at the surface (0-20 cm) and subsurface (20-40 cm) layers. Only significant coefficients are shown (*, $p < 0.05$; **, $p < 0.01$) ($n=64$).

Layer		pH	EC	OC	P	Ca ²⁺
Surface	K		0.45**			
	P					
	Ca ²⁺	0.67**		0.26*		
	Mg ²⁺					0.37**
	S		0.31*		0.44**	
Sub-surface	HWB			0.30*		
	K		0.48**			
	P		0.32*			
	Ca ²⁺	0.42**				
	Mg ²⁺					0.33**
	S				0.36**	

EC-electrical conductivity, dS m^{-1} ; OC-organic carbon, g kg^{-1} ; K, mg kg^{-1} ; P, mg kg^{-1} ; exchangeable Ca²⁺, mg kg^{-1} ; exchangeable Mg²⁺, mg kg^{-1} ; S, mg kg^{-1} ; HWB, hot water soluble B, mg kg^{-1} .

**Table 3.** Semivariogram parameters of soil properties of studied areas.

Variable	Soil layer	Model	Nugget	Sill	Nugget: Sill ratio	Spatial class	Range (m)	MSE	G (%)
pH	Surface	Spherical	0.098	0.130	0.715	Moderate	1416	0.754	62
	Subsurface	Spherical	0.110	0.160	0.687	Moderate	1468	0.681	58
EC	Surface	Spherical*	0.001	0.004	0.025	Strong	554	0.0003	55
	Subsurface	Linear*	0.003	0.004	0.750	Moderate	2186	0.0002	51
OC	Surface	Exponential	54.10	67.70	0.797	Weak	1131	2.31	48
	Subsurface	Circular	20.80	51.10	0.407	Moderate	581	3.12	56
NH ₄ OAc-K	Surface	Spherical*	36371	40122	0.906	Weak	4530	28.31	65
	Subsurface	Linear*	21523	22506	0.956	Weak	4530	30.01	60
Bray's-P	Surface	Gaussian*	875.0	940.0	0.930	Weak	1996	40.02	53
	Subsurface	Gaussian*	97.60	149.9	0.651	Moderate	770	39.58	50
Ca ²⁺	Surface	Linear*	0.000	263780	0.000	Strong	1585	221.01	33
	Subsurface	Exponential*	0.000	330416	0.000	Strong	581	198.65	26
Mg ²⁺	Surface	Gaussian*	11244	21059	0.533	Moderate	885	89.56	50
	Subsurface	Exponential*	19839	20685	0.959	Weak	1114	70.04	53
CaCl ₂ -S	Surface	Linear*	234.0	245.0	0.955	Weak	4530	0.067	45
	Subsurface	Gaussian	62.10	93.20	0.666	Moderate	4530	0.071	42
HWB	Surface	Gaussian*	0.046	0.073	0.630	Moderate	1424	0.023	71
	Subsurface	Linear*	0.111	0.147	0.755	Weak	1148	0.018	76

*Transformation for normal distribution.

EC-electrical conductivity, dS m⁻¹; OC-organic carbon, g kg⁻¹; K, mg kg⁻¹; P, mg kg⁻¹; exchangeable Ca²⁺, mg kg⁻¹; exchangeable Mg²⁺, mg kg⁻¹; S, mg kg⁻¹; HWB, hot water soluble B, mg kg⁻¹; MSE-mean square error; G-goodness-of-prediction criterium.



Figure 1. Spatial distribution of sampling points in Goa state (western India)

Figure 2. Kriged interpolation maps of soil properties in surface (0-20 cm) and subsurface (20-40 cm) soil layers

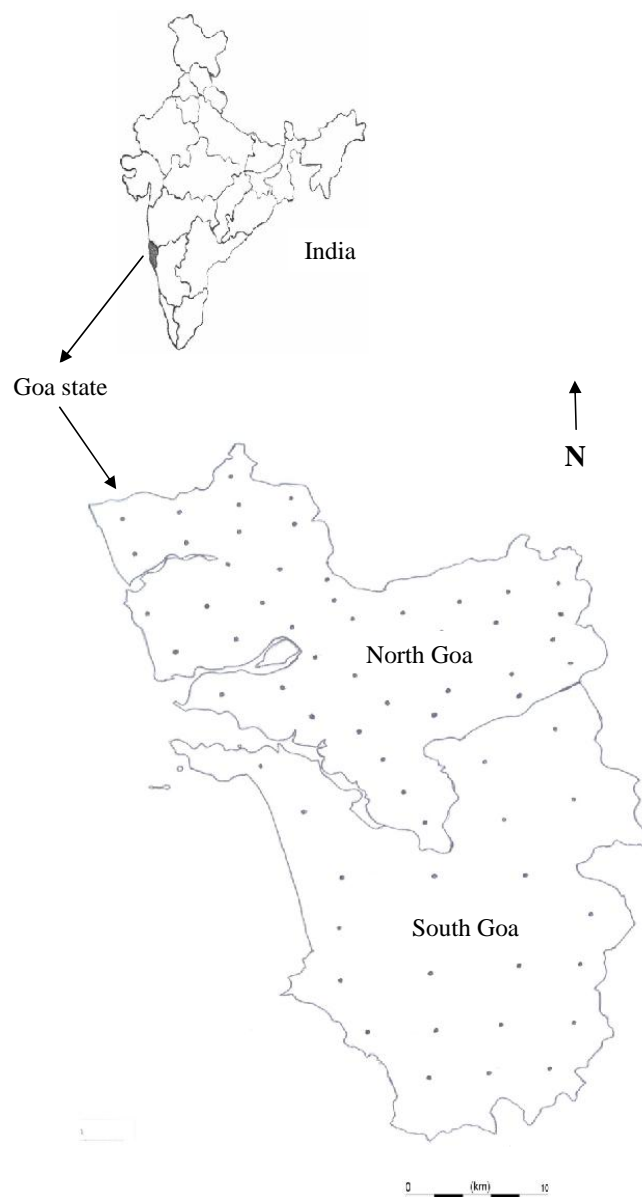
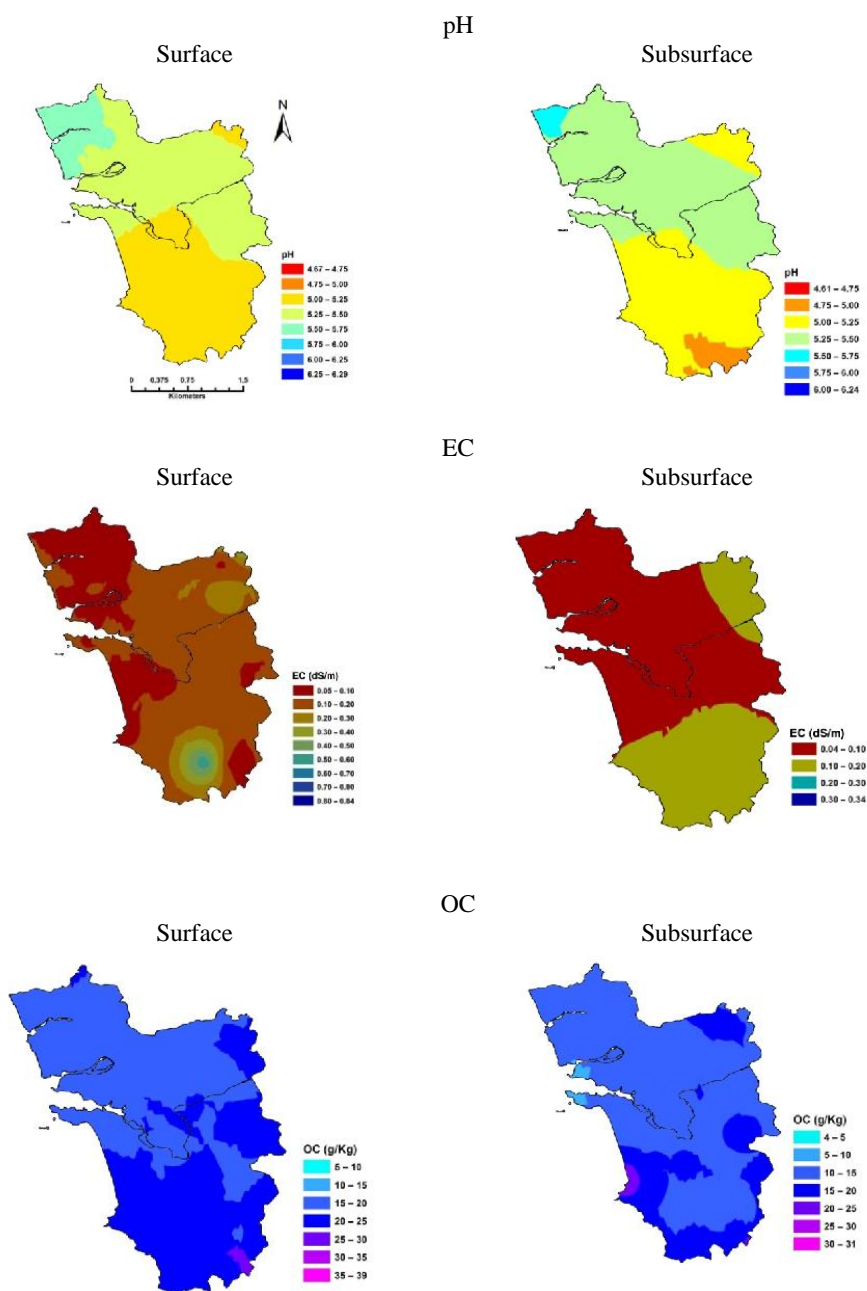
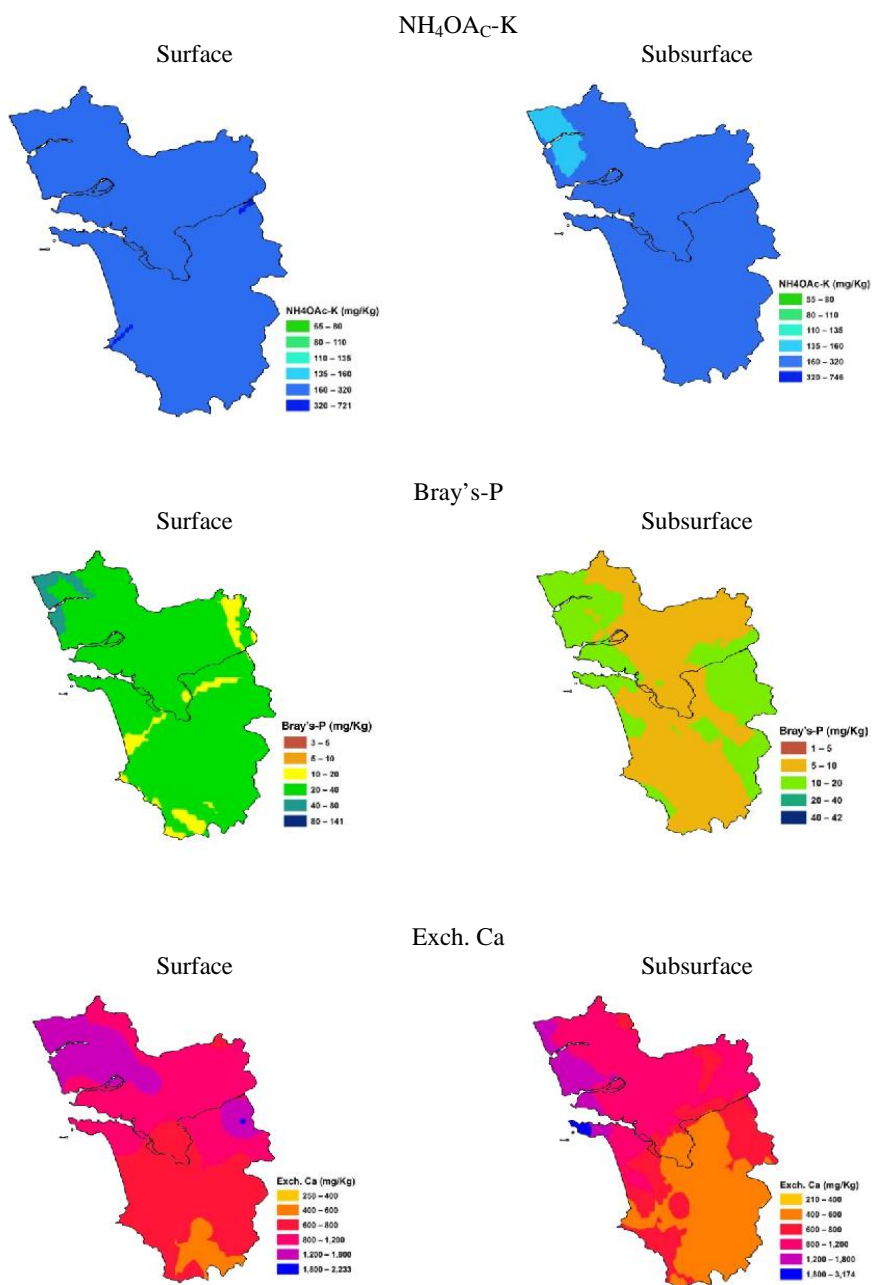


Figure 1. Spatial distribution of sampling points in Goa state (western India)

pH





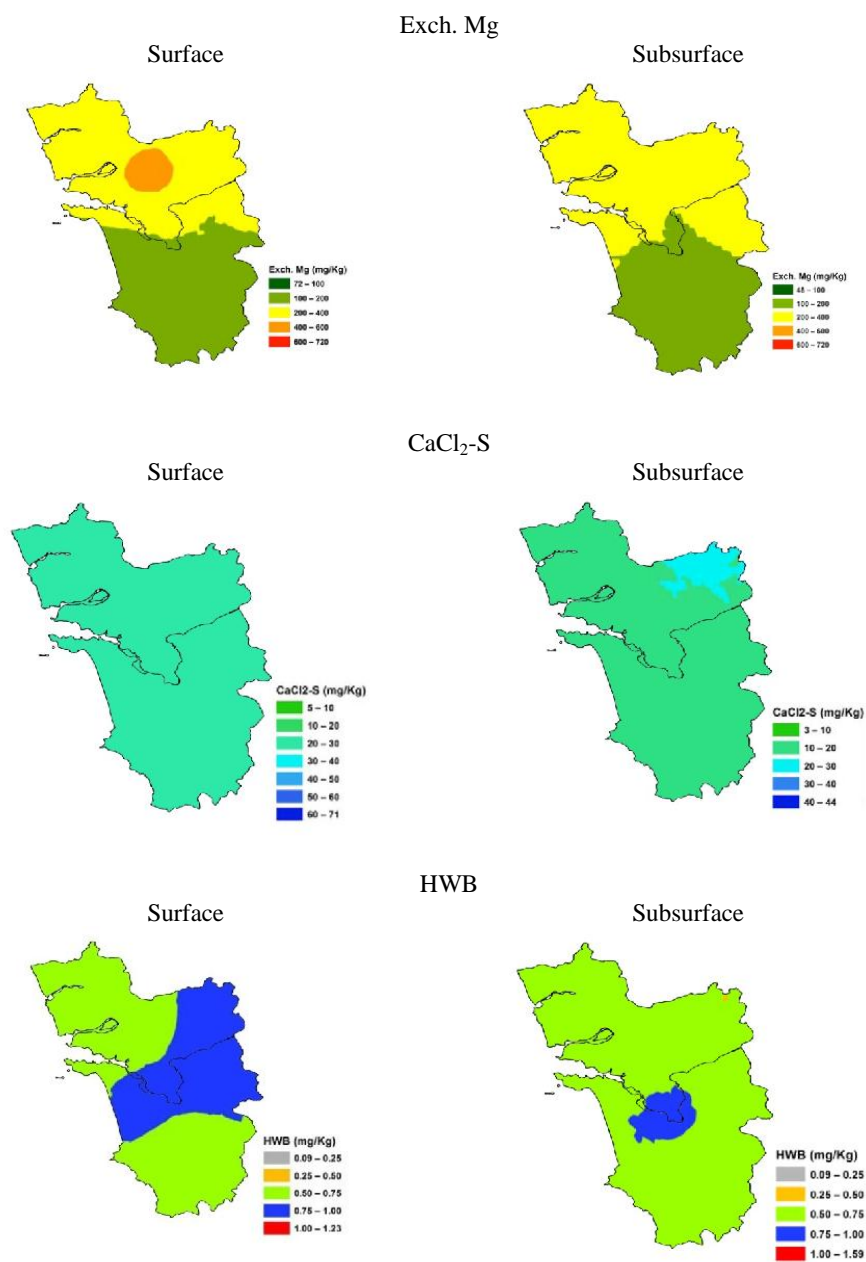


Figure 2. Kriged interpolation maps of soil properties in surface (0-20 cm) and subsurface (20-40 cm) soil layers