# On soil textural classifications and soil texture-based estimations

Miguel Ángel Martín[1], Yakov. A. Pachepsky[2], Carlos García-Gutiérrez[1], and Miguel Reyes[1]

[1]Universidad Politécnica de Madrid, Madrid, Spain.
[2]USDA-ARS Environmental Microbial and Food Safety Laboratory.

*Correspondence to:* Miguel Ángel Martín (miguelangel.martin@upm.es)

**Abstract.** The soil texture representation with the standard textural fraction triplet 'sand-silt-clay' is commonly used to estimate soil properties. The objective of this work was to test the hypothesis that other fraction sizes in the triplets may provide better representation of soil texture for estimating some soil parameters. We estimated the cumulative particle size distribution and bulk density from entropy-based representation of the textural triplet with experimental data for 6300 soil samples. Results supported the hypothesis. For example, simulated distributions were not significantly different from the original ones in 25 and 85 % of cases when the 'sand-silt-sand' and 'very coarse+coarse + medium sand - fine +very fine sand - silt+clay', were used, respectively. When the same standard and modified triplets were used to estimate the average bulk density, the coefficients of determination were 0.001 and 0.967, respectively. Overall, the textural triplet selection appears to be application- and data-specific.

## 1 Introduction

The particle size distribution is one of the essential controls of soil structure and functioning. Soil processes, properties and specific features are usually related to these distributions, commonly named soil texture. To express these relationships, the continuous particle size distributions are commonly replaced by their discrete representation with several textural fractions. The fractions are defined as particles within a range of sizes, e.g. medium sand, or fine silt, etc. Then the percentages of textural fractions are used as attributes to classify soils and as predictors to estimate soil properties of parameters.

Different countries have employed different number of textural fractions and different ranges of sizes for each of the fraction. *Nemes et al.* (1999) reviewed definitions of textural fractions in 14 European countries and reported the number of ranges varying from three in Italy and France to 8 in the Netherlands and Germany, and 9 in Belgium. These authors also observed a large variability in size ranges, For example, while the minimum size of the second smallest fraction was 2 $\mu$m in most cases, the maximum size in such fraction varied from 6 $\mu$m in Greece to 60 $\mu$m in England and Wales. In 1967, the Committee of the Soil Science Society of America noted that the current system of particle size boundaries arose due to geographic accident (*Whiteside et al.*, 1967). The committee noted that there is "no narrowly definable natural particle size boundaries that would be equally significant in all soil materials" The authors noted that the boundary between clay and silt was originally set at 10 $\mu$m, then changed to 5 $\mu$m and finally established at 2 $\mu$m (*Whiteside et al.*, 1967).

There were indications that setting the boundaries between textural fractions might depend on the purpose of further textural data use as well as on the specifics of the dataset under consideration. *Twarakavi et al.* (2010) demonstrated that, soils are not

classified well from a hydraulic standpoint if the USDA textural fractions of sand, silt, and clay are used. They also noted that this conclusion is conditioned to the database used for the hydraulic classification evaluation. Reasons for selection of size boundaries varied. *Whiteside et al.* (1967) noted that for several reasons a scale based on 1 mm with subdivisions at 0.315 mm, 0.1 mm, etc. would seem to be the ideal scale for agricultural purposes, but advantages were not deemed to be sufficient to

5 outweigh the radical departure from the existing textural classification. Also, physics-based reasoning influenced the selection of size boundaries between textural fractions. For example, the 2 $\mu$m boundary for clay was chosen originally as 10 $\mu$m, and then moved to 5 $\mu$m. Around 1936 a switch of the clay limit from 0.005 to 0.002 mm was proposed based on the realization that at 0.002 mm a significant break in mineralogical properties of soil separates occurs (*Truog et al.*, 1936a, b) and that soil surveyors in the field tend to consider the 0.002 to 0.005 mm fraction as silt rather than clay (*Shaw and Alexander*, 1937).

10 One application of the data on textural fraction content is the reconstruction of the particle size distribution from the data on small number of fractions. *Martín and Taguas* (1998) proposed to use the hypothesis of self-similarity and iterated function formalism to generate the particle size distribution from small number of textural fractions. In applications of this technique, they used sand, silt, and clay fraction contents with size boundaries defined by the USDA textural classification. Another application of data on textural fractions is to compute the information entropy as the metric of the particle size heterogeneity

15 and derive the relationship between the bulk density (BD) and information entropy (IE) (*Martín et al.*, 2017a). Seven textural fractions were used in the latter work where a strong linear correlation between the respective average values was shown. This fact together with computational results obtained in (*Martín et al.*, 2017b), seemed to reinforce the entropy self-similarity approach which is used in the PSD reconstruction. Self-similarity, commonly expressed by scaling laws, actually means that the content of information obtained at the coarse scale keeps its average value at smaller scales (*Pastor-Satorras and Wagensberg*,

20 1998),which agrees with the driving idea of the PSD representation used.

The objective of this work was to test hypotheses that (a) the reconstruction of the particle size distribution can be more accurate if the textural fraction size boundaries will be changed from the USDA sand-silt-clay sizes to other size ranges, (b) a satisfactory relationship between the information entropy and packing density can be achieved with three textural fractions with boundaries between fraction change sizes other than in the USDA 'sand-silt-clay' triplet.

## 2 Materials and Methods

### 2.1 The dataset

The data set used to evaluate the influence of soil texture heterogeneity in soil bulk density values was extracted from the USKSAT database. This data base is comprised from journal publications and technical reports containing coupled data on Saturated Hydraulic Conductivity (Ksat), soil texture, bulk density, and organic matter content obtained across the United

30 States. Detailed information can be found in (*Pachepsky and Park*, 2015). We selected the dataset from Florida ((*Carlisle et al.*, 1978) and (*Carlisle et al.*, 1981)). This dataset is the largest of dataset in UKSAT obtained in the same laboratory with the same methods. The dataset was filtered to exclude samples for which data on seven textural fractions, or on bulk density were not available. Samples with inconsistent textural data (the sum of mass texture fractions not agreeing with the total mass) were

rejected. The selection criteria used in *Martín et al.* (2017a) were followed. Under these selection criteria, total of 6240 soil samples were included in the study. By USDA textural classes, sands, loamy sands, sandy loams, loams, silt loams, silts, sandy clay loams, clay loams, silt clay loams, sandy clays, and clays were represented by 3956, 570, 698, 27, 27, 4, 667, 26, 3, 118 and 144 samples.

## 2.2 Reconstruction of the particle size distributions from data on textural fraction content

The reconstruction of the particle size distribution (PSD) is based on the assumption that the entropy as the measure of heterogeneity of these distributions is preserved across the support scales (*Martín and Taguas*, 1998). Assuming the texture interval divided into $k$ textural size ranges and that the respective textural fraction contents $p_1, p_2, \ldots, p_k$, $1 \le i \le k$, $\sum_{i=1}^{k} p_i = 1$ the Shannon Information Entropy (IE) (*Shannon*, 1948) is defined by

$$IE = -\sum_{i=1}^{k} p_i \log_2 p_i \tag{1}$$

where $p_i \log p_i = 0$ if $p_i = 0$. The IE is a widely accepted measure of the heterogeneity of distributions (*Khinchin*, 1957). The IE values for three textural size classes range from 1 when only one fraction is present to $-\log_2 \frac{1}{3} = 1.585$ when all three fractions are represented equally.

*Martín and Taguas* (1998) proposed a self-similarity model which allows generating the PSD from commonly available textural data which consist of mass percentages of small number of textural fractions. The driving idea of such proposal was that the heterogeneity that textural data show at the coarse scale - quantified in terms of Information Entropy - is also reproduced in a similar way inside any rescaled textural fractions at smaller scales (i.e. the heterogeneity of sieved fractions would resemble that observed at the coarse scale). From this single hypothesis, a mathematically precise representation of the PSD is then obtained by the iterated function formalism. A brief illustration of this technique for the case of three textural fractions looks as follows. Let us denote by $I_1$, $I_2$ and $I_3$ the subintervals of sizes corresponding to the three size classes and $p_1$, $p_2$ and $p_3$ the relative proportions of mass for the intervals $I_1$, $I_2$ and $I_3$, respectively. These proportions are treated as probabilities, $p_1 + p_2 + p_3 = 1$. Let $\varphi_1$, $\varphi_2$ and $\varphi_3$ be the linear functions (similarities) which transform the whole size interval $I$ into the subintervals $I_1$, $I_2$ and $I_3$, respectively. The set $\{\varphi_1, \varphi_2, \varphi_3, p_1, p_2, p_3\}$ is called an iterated function system (IFS) (*Barnsley and Demko*, 1985).The hypothesis of entropy self-similarity of the PSD states that the IE, now computed on the successive rescaled subintervals $\varphi_j(I_i) = I_{ij}$, $\varphi_k(I_{ij}) = I_{ijk}, \ldots$ and so on $(i, j, k = 1, 2, 3)$, is scale-invariant. Then at each scale, the mass proportion of soil in an interval may be exactly computed (see *Martín and Taguas* (1998) for further details).

The reconstruction of distributions was performed using three size fractions - coarse, intermediate, and fine. The dataset contained experimental data on seven fractions - very coarse sand, coarse sand, medium sand, fine sand, very fine sand, silt, and clay. We used all possible triplets formed from seven textural fractions that were available. The symbols for triplet showed how the fractions were grouped. For example for the triplet '3-2-3' had "coarse" included very coarse sand, coarse sand and medium sand, "intermediate" included fine sand and very fine sand , and "fine" included silt and clay; triplet '5-1-1' was the standard one where "coarse" included all five sand fractions, "intermediate" included silt, and "fine" included clay. Total of 15 triplets were available.

For all textural triplets we generated PSD and compared experimental particle size distributions (built from seven known fractions) with simulated ones. The Kolmogorov-Smirnov test has been applied to find the probability that ~~that~~ the samples are drawn from the same distribution.

### 2.3 Information Entropy/bulk density relation

5  Following Eq (1), the Information Entropy of soil texture is computed for all triplets in order to analyze how differences in the Information Entropy explain differences in the typical soil bulk density value of related soils. The range of information entropy values was divided into 10 subintervals of equal length. The mean bulk density value of soil samples binned into IE ranges was computed for each of the subintervals. The linear regression of the average information entropy vs. average bulk density value was computed.

10 ## 3 Results and Discussion

### 3.1 Results

Examples of ternary graphs ~~of~~ showing locations of samples in the 'coarse-intermediate-fine' textural fraction content coordinates are shown in Figures 1a and 1b. The standard triangle in Figure 1a shows the majority of points in the left bottom corner. This reflects the fact that soils in the database are mostly coarse textured in terms of the USDA textural classification.

15 The '3-2-2' textural triangle in the Figure 1b shows that the soil samples represent well both samples low in fines particles and samples low in coarse particles, whereas soils with low intermediate fraction contents are not represented well in the database. Table 1 shows total numbers of samples by the ranges of the IE values for standard and '3-2-2' triplets in Figure 1. The standard triplet assigns small values of the information entropy to the majority of samples and thus interprets the majority of samples as heterogeneous. On the contrary, applying the '3-2-2' triplet leads to the conclusion that the majority of samples has a moderate

20 level of textural heterogeneity.

The reconstruction of particle size distributions with the iterated function algorithm showed large difference between the applications different triplets. Data on the statistical difference between generated and measured distributions are shown in Table 2 for all samples and for textural classes where the number of available samples exceeded 100. Triplets where the group of fines includes fine sand, silt and clay, i.e. '1-3-3', '2-2-3', '3-1-3' provide the best results. Results for fine textured soils do

25 not depend on the triplet because the proportions of coarse particles are small and do not affect results. The differences between triplets become more pronounced as the textures become coarser. The worst results are obtained for triplets '1-5-1, '2-4-1', and '5-1-1' having clay as a separate file fraction, Using the standard triplet '5-1-1' leads to absolutely worst results. The simulated and the experimental cumulative particle- size distributions are not statistically different at the 95% probability level for 25% of soil samples when the standard triplet of fraction contents is used as input in the reconstruction of the PSD. Instead, when

30 using the ~~any~~ of the triplets '1-3-3', '2-2-3' or '3-1-3' the percentage of soils whose simulated particle size distribution is not statistically different ~~to~~ the original ~~one~~ is bigger than 97% of the total, for the same probability level.

Results of building linear regressions of Mean Information Entropy values versus mean bin bulk density values are shown in Table 3. Different triplets cause different efficiency in estimating BD from IE by textural classes. Overall best relationships were found for sands. Efficiency of estimation was worse in textural classes where there was no single dominant fraction. The sandy clay and sandy clay loam classes provide examples of the above. Noticeably triplets with clay, silt, and fine sand combined in the fine fraction do not result in good R2 for non-sandy soils (Table 3). This is opposite to the PSD reconstruction where fines consisting of clay, silt, and very fine sand provide the best results (Table 2).

The best results by considering both sand and non sandy samples are obtained with triplets 2-4-1 and 3-3-1, i.e. triplets where fines are represented only by clays and there is a certain balance between the coarse and the intermediate fractions. Where this balance is not present (1-5-1 and 5-1-1) separation of clay in the fine fraction does not help. The standard triangle seems to work only for non sandy soils. Also, this triplet's IE relates well to the BD of sandy clays, sandy loams and sandy clay loams, but it gives unsatisfactory results for sands.

### 3.2 Discussion

The triplets having the fine fraction consisting of very fine sand, silt, and clay appeared to be superior in serving as the input for PSD reconstruction. One possible explanation is that mass size scaling is not scale-invariant across all particle sizes. Rather it has ranges particle sizes within which the power law scaling dependencies are applied and the boundaries between these ranges are between reflected by the modified textural triplet rather than by the original '5-1-1' sand-silt-clay triplet. Breaks in particle size distribution scaling were first highlighted by *Tyler and Wheatcraft* (1992) who noted that the strict fractal or self-similar behavior in soil PSDs is restricted to a narrow spectrum of soils found in nature. For the soils tested, the power law scaling was observed in only limited portions of their PSDs. Data on soils B to F from their work are shown in Figure 2. The diameter of the break in scaling varied between 100 $\mu$m and 400 $\mu$m and on average in this group of soils occurred at diameters of 220 $\mu$m which is close to the boundary 250 mm between fine and medium sand. Later the break in scaling was demonstrated by other authors, e.g. *Kravchenko and Zhang* (1998) who noted that "The critical particle size [radii - M. A.] at which the fractal dimension values are changing, is about 100 to 200 $\mu$m for most of the soils. The result is consistent with that reported in the literature (*Wu et al.*, 1993)"

Another reason for the better simulations of particle size distributions can be the better representation of the soil texture, i.e. distribution of samples by the ranges of IE where the majority of soil are found (Table 1). When the IE is computed with the standard triplet a great amount of soils have low IE value (have unbalanced contents respect to those texture fractions). This may be an obstacle for reconstruction of the PSD under the entropy self-similarity. In particular, because of the meaning of self-similarity itself, if the input contents are very unbalanced, it cause a multiplicative effect of more unbalanced distribution in the "sub-fractions" at lower scales and probably a more unreliable simulation. On the contrary, in the case of the modified triplet a great amount of soils have medium/high IE, which means that have more balanced contents respect to the respective new fractions: a greater power of discriminating texture, texture based properties and of obtaining better PSD simulations are expected. This can be an interesting avenue to explore.

The large difference between the 'IE - Bulk density' relationships developed for different textural classes indicates that the IE computed for different triplets has the potential to reflect the effect of soil texture on particle packing in soils. The theoretical analysis of *Assouline and Rouault* (1997) and *Martín et al.* (2017b) shows that the pore space arrangement can be related to the type of distribution of particles sizes. The IE parameter is related to packing but cannot reflect aggregation that is characteristic

5   for soils where fine particles are present in substantial amounts. We note that when IE was computed using the seven texture fraction contents with the same database, the determination coefficient of the regression 'average IE vs. average bulk density' was equal to 0.99 (*Martín et al.*, 2017b). Thus, results shows that the modified triplet provides almost the same information respect to the bulk density values that the provided by the seven texture fractions altogether.

The best triplets were different for reconstruction of the particle size distributions and for establishing relationships between
10  information entropy and bulk density after binning samples. Different triplets may be most informative to characterize results of fragmentation and sedimentation that manifest themselves in particle size distributions, and results of packing that manifest themselves in 'IE - BD' relationships. Finally, some process affecting the particle size distributions and 'IE - BD' relationships may not be elucidated by textural data only, aggregation and weathering being examples.

The utility of textural fractions different from traditional sand-silt-clay triplet appears to have an application in the devel-
15  opment of pedotransfer functions. Boundary of new fraction sizes can be parameters of pedotransfer functions along with the regression coefficients. *Nemes and Rawls* (2006) experimented with the boundary between silt and sand ~~in the~~ in the range from 20 $\mu$m and 63 $\mu$m and developed pedotransfer functions for water retention at -33 kPa and -1500 kPa matric potential values. They could not point out the boundary size between silt and sand that would clearly provide better results in estimating the selected soil hydraulic properties. Our work indicates that the boundary may be moved to the range of much larger particle
20  diameters.

Usability of triplets other than standard ~~one indicates~~ the opportunity of a more efficient use of t~~he~~ existing results of textural analysis. Although these results traditionally consist of seven fractions including five fractions of sand, in the majority of applications all sand fractions have been lumped together. For example, the overwhelming majority of pedotransfer functions in soil hydrology use the elements of the standard triplet 'sand-silt-clay' (*Pachepsky and Rawls*, 2004). The use of different
25  'coarse-intermediate-fine' triplets in pedotransfer studies allows the use of ~~the~~ available detailed data on fractions of sand and revisiting existing databases appears be interesting avenue to explore.

When analyzing the utility of traditional sand-silt-clay triplet for classifying soils by their hydraulic properties, *Twarakavi et al.* (2010) concluded that from a philosophical perspective, the research further stresses the need to revisit and reevaluate the results from the past in order to successfully move ahead into the future of soil physics. Consideration of textural fraction boundaries
30  as parameters that can be task and dataset specific can provide additional insights on the role of texture in soil functioning and ecological services.

## 4    Conclusions

Having three textural size ranges, i.e. coarse, intermediate, and fine particle sizes, undoubtedly appears to be convenient for data presentation and textural class definition. Currently the coarse, intermediate, and fine fractions are identified as sand, silt, and clay, respectively. However, it is not warranted that current sand, silt, and clay size ranges can provide the best representation

5    of soil texture when these three size ranges are used for estimating soil properties. We hypothesized the cumulative particle size distribution and soil bulk density can be more accurately estimated from the triplet 'coarse-intermediate-fine', if the boundaries of the coarse, intermediate, and fine size ranges will be different from those in the 'sand-silt-clay' triplet. The entropy-based representation of particle size distributions was used to convert the triplet particle size representations into particle size distributions and to define ranges of soil textural heterogeneity. Experimental data on seven size fraction contents

10    and bulk density for 6300 predominantly coarse-textured soil samples were extracted from the USKSAT database

It appeared that the redefining the triplet 'coarse-intermediate-fine' may lead to a very substantial improvement of soil property estimates from soil textural data. Overall, the drastic improvement in predictions of both cumulative particle size distribution and mean bulk density for heterogeneity ranges occurred when the standard 'sand - silt - clay' triplet was replaced with the modified textural triplet that was defined as very coarse, coarse, medium sand (coarse fraction), fine and very fine sand

15    (intermediate fraction) and clay and silt (fine fraction). The modified triplet apparently provided more information about the particle size heterogeneity and particle packing. Different modified triplets provided the best inputs for different soil textural classes.

Results of this work indicate that the detailed information about soil particle size distributions has the potential to enhance estimation of soil properties with soil texture as a predictor. Analyses of both existing and developing soil databases as well as

20    the pedotransfer methodologies may benefit from exploring modifications of textural triangles. Compression of information on textural heterogeneity in textural triangles into a single entropy-based parameter may provide additional advantages.

*Acknowledgements.*    This research work was funded by Spain's Plan Nacional de Investigación Científica, Desarrollo e Innovación Tecnológica (I+D+I), under ref. AGL2015-69697-P.

# References

Assouline, S., and Y. Rouault (1997), Modeling the relationships between particle and pore size distributions in multicomponent sphere packs: Application to the water retention curve, *Colloids and Surfaces A: Physicochemical and Engineering Aspects*, 127(1-3), 201-210, doi:10.1016/S0927-7757(97)00144-1.

5   Barnsley, M. F., and S. Demko (1985), Iterated functions systems and the global construction of fractals, *Proc. R. Soc. Lon.*, A 399, 243-275, doi:10.1098/rspa.1985.0057.

Carlisle, V. W., R. E. Caldwell, F. Sodek, III, L. C. Hammond, F. G. Calhoun, M. A. Granger, and H. L. Breland (1978), Characterization Data for Selected Florida Soils, Univ. of Florida, Gainesville, Fla.

Carlisle, V. W., C. T. Hallmark, F. Sodek, III, R. E. Caldwell, L. C. Hammond, and V. E. Berkeiser (1981), Characterization Data For Selected
10   Florida Soils, Soil Conservation Service, University of Florida, Institute of Food and Agricultural Sciences, Soil Sciences Department, Soil Characterization Laboratory.

Khinchin, A. I., (1957), Mathematical Foundation of Information Theory. Dover Publications, New York.

Kravchenko, A., and R. Zhang (1998), Estimating the soil water retention from particle-size distributions: a fractal approach, *Soil Science*, 163(3), 171-179.

15   Martín, M. A., and F. J. Taguas (1998), Fractal modeling, characterization and simulation of particle-size distributions in soil, *Proc. R. Soc. Lond.*, A 454, 1457-1468, doi:10.1098/rspa.1998.0216.

Martín, M. A., M. Reyes, and F. J. Taguas (2017a), Estimating soil bulk density with information metrics of soil texture, *Geoderma*, 287, 66-70, doi:10.1016/j.geoderma.2016.09.008.

Martín, M. A., M. Reyes, and F. J. Taguas (2017b), An entropy-like parameter of particle size distributions as packing density index in
20   complex granular media, *Granular Matter*, 19:9, doi:10.1007/s10035-016-0692-3.

Nemes, A., J. H. M. Wösten, A. Lilly, and J. O. Voshaar (1999), Evaluation of different procedures to interpolate particle-size distributions to achieve compatibility within soil databases, *Geoderma*, 90(3), 187-202, doi:10.1016/S0016-7061(99)00014-2.

Nemes, A., and W. J. Rawls (2006), Evaluation of different representations of the particle-size distribution to predict soil water retention, *Geoderma*, 132(1), 47-58, doi:10.1016/j.geoderma.2005.04.018.

25   Pachepsky, Y. A., and W. J. Rawls (eds.) (2004), Development of Pedotransfer Functions in Soil Hydrology, Elsevier, Amsterdam.

Pachepsky, Y. A., and Y. Park, (2015), Saturated hydraulic conductivity of US soils grouped according to textural class and bulk density, *Soil Sci. Soc. Am. J.*, 79, 1094-1100, doi:10.2136/sssaj2015.02.0067.

Pastor-Satorras, R., and Wagensberg, J., (1998), The maximum entropy principle and the nature of fractals, *Physica A*, 251, 291-302, doi:10.1016/S0378-4371(97)00571-2.

30   Shannon, C. E., (1948), A mathematical theory of communication, *I. Bell Syst. Technol. J.*, 27, 379-423.

Shaw, T. M., and L. T. Alexander (1937), A note on mechanical analysis and soil texture, *Soil Sci. Soc. Amer. Proc.*, 1, 303-304, doi:10.2136/sssaj1937.03615995000100000053x.

Truog, E., J. R. Taylor, R. W. Pearson, M. E. Weeks, and R. W. Simonson (1936a), Procedure for special type of mechanical and mineralogical soil analysis, *Soil Sci. Soc. Amer. Proc.*, 1, 101-112, doi:10.2136/sssaj1937.03615995000100000013x.

35   Truog, E., J. R. Taylor, R. W. Simonson, and M. E. Weeks (1936b), Mechanical and mineralogical subdivisions of the clay separate of soils, *Soil Sci. Soc. Amer. Proc.*, 1, 175-179, doi:10.2136/sssaj1937.03615995000100000025x

Twarakavi, N. K. C., J. Šimůnek, and M. G. Schaap (2010), Can texture-based classification optimally classify soils with respect to soil hydraulics?, *Water Resources Research*, 46(1), doi:10.1029/2009WR007939

Tyler, S. W., and S. W. Wheatcraft (1992), Fractal scaling of soil particle-size distributions: analysis and limitations, *Soil Sci. Soc. Amer. J.*, 56(2), 362-369, doi:10.2136/sssaj1992.03615995005600020005x.

5  Whiteside, E. P., K. W. Flach, V. C. Jamison, W. D. Kemper, E. G. Knox, and A. C. Orvedal (1967), Considerations relative to a common particle size scale of earthy materials, *Soil Sci. Soc. Amer. Proc.*, 31(4), 579-584, doi:10.2136/sssaj1967.03615995003100040046x.

Wu, Q., M. Borkovec, and H. Sticher (1993), On particle-size distributions in soils, *Soil Sci. Soc. Amer. J.*, 57(4), 883-890, doi:10.2136/sssaj1993.03615995005700040001x.
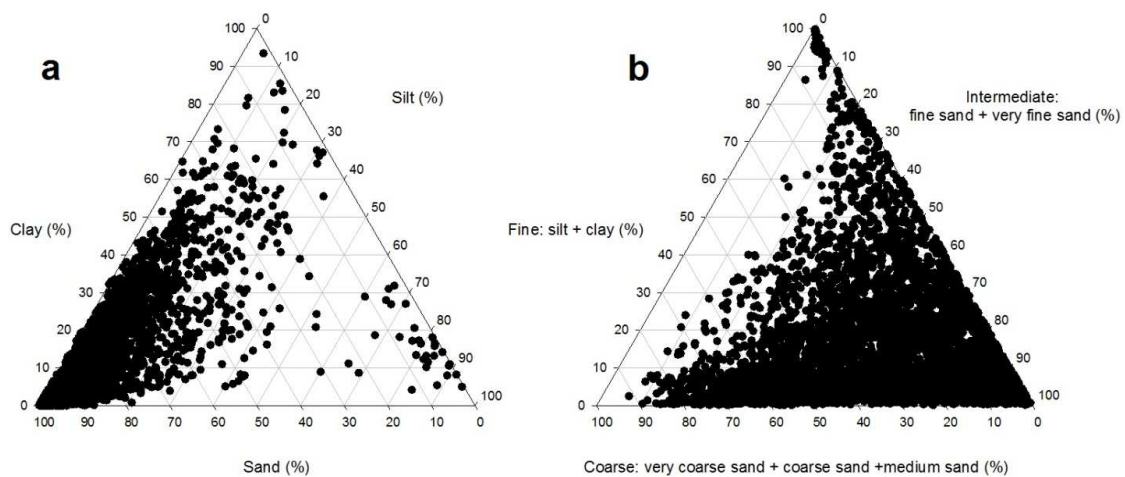
**Figure 1.** Texture of soil samples in the database shown in the standard USA (a) and modified textural triangles (b).
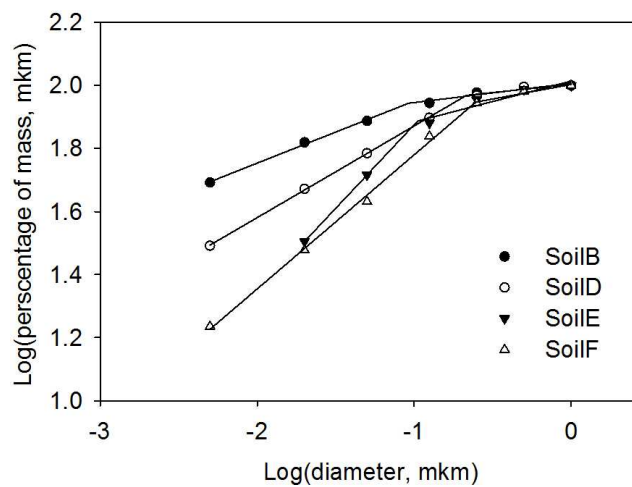
**Figure 2.** Scaling in cumulative particle mass of four soils studied by *Tyler and Wheatcraft* (1992)

**Table 1.** Total numbers of samples by ranges of the information entropy for two textural fraction triplets. The '3-2-2' triplet includes very coarse, coarse and medium sand (fraction 1), fine and very fine sand (fraction 2) and clay and silt (fraction 3); the standard triplet '5-1-1' includes sand (fraction 1), silt (fraction 2) and clay (fraction 3).

| | Number of samples by information entropy ranges | |
|---|---|---|
| Range of the Information Entropy | '3-1-1' triplet | '5-1-1' triplet |
| 0.00-0.16 | 23 | 698 |
| 0.16-0.32 | 101 | 1346 |
| 0.32-0.48 | 269 | 1169 |
| 0.48-0.64 | 407 | 754 |
| 0.64-0.80 | 539 | 485 |
| 0.80-0.96 | 772 | 499 |
| 0.96-1.12 | 1172 | 465 |
| 1.12-1.28 | 1250 | 416 |
| 1.28-1.44 | 926 | 272 |
| 1.44-1.60 | 781 | 136 |

**Table 2.** Percentage of samples with simulated and measured particle size distributions being not different at the 0.05 significance level.

| Textural triplet | Clay | Sandy clay | Sandy clay loam | Sandy loam | Loamy sand | Sand | All samples |
|---|---|---|---|---|---|---|---|
| 115 | 100 | 100 | 82 | 45 | 37 | 8 | 27 |
| 124 | 100 | 100 | 97 | 83 | 76 | 51 | 64 |
| 133 | 100 | 100 | 100 | 100 | 100 | 97 | 98 |
| 142 | 100 | 100 | 98 | 84 | 57 | 15 | 40 |
| 151 | 100 | 100 | 84 | 32 | 5 | 1 | 19 |
| 214 | 100 | 100 | 97 | 84 | 78 | 53 | 66 |
| 223 | 100 | 100 | 100 | 100 | 100 | 100 | 99 |
| 232 | 100 | 100 | 100 | 95 | 79 | 28 | 52 |
| 241 | 100 | 100 | 93 | 59 | 38 | 8 | 31 |
| 313 | 100 | 100 | 100 | 100 | 100 | 100 | 100 |
| 322 | 100 | 100 | 100 | 99 | 94 | 77 | 85 |
| 331 | 100 | 100 | 99 | 90 | 79 | 55 | 68 |
| 412 | 100 | 100 | 100 | 100 | 100 | 82 | 88 |
| 421 | 100 | 100 | 100 | 100 | 100 | 83 | 89 |
| 511 | 97 | 100 | 99 | 78 | 0 | 0 | 25 |
| N | 144 | 115 | 660 | 685 | 565 | 3966 | 6224 |

**Table 3.** Determination coefficients of regressions of the average information entropy versus average bulk density for ten average entropy bins.

| Triplet | Clay | Sandy clay | Sandy clay loam | Sandy loam | Loamy sand | Sand | All samples | All sands | Not sandy |
|---|---|---|---|---|---|---|---|---|---|
| 115 | 0.633 | 0.076 | 0.426 | 0.725 | 0.520 | 0.509 | 0.571 | 0.567 | 0.035 |
| 124 | 0.579 | 0.085 | 0.817 | 0.807 | 0.797 | 0.908 | 0.915 | 0.899 | 0.167 |
| 133 | 0.695 | 0.143 | 0.460 | 0.463 | 0.111 | 0.448 | 0.932 | 0.854 | 0.445 |
| 142 | 0.432 | 0.037 | 0.027 | 0.024 | 0.318 | 0.799 | 0.412 | 0.772 | 0.590 |
| 151 | 0.762 | 0.345 | 0.219 | 0.630 | 0.789 | 0.433 | 0.125 | 0.663 | 0.742 |
| 214 | 0.682 | 0.177 | 0.834 | 0.894 | 0.545 | 0.919 | 0.824 | 0.868 | 0.239 |
| 223 | 0.713 | 0.141 | 0.337 | 0.517 | 0.682 | 0.942 | 0.786 | 0.978 | 0.474 |
| 232 | 0.604 | 0.275 | 0.004 | 0.053 | 0.803 | 0.931 | 0.960 | 0.959 | 0.748 |
| 241 | 0.837 | 0.256 | 0.305 | 0.689 | 0.743 | 0.756 | 0.891 | 0.933 | 0.897 |
| 313 | 0.594 | 0.482 | 0.053 | 0.767 | 0.496 | 0.963 | 0.823 | 0.964 | 0.508 |
| 322 | 0.388 | 0.055 | 0.000 | 0.216 | 0.777 | 0.968 | 0.909 | 0.968 | 0.704 |
| 331 | 0.878 | 0.164 | 0.312 | 0.872 | 0.822 | 0.966 | 0.967 | 0.971 | 0.790 |
| 412 | 0.361 | 0.077 | 0.080 | 0.603 | 0.546 | 0.307 | 0.943 | 0.873 | 0.731 |
| 421 | 0.701 | 0.161 | 0.019 | 0.438 | 0.061 | 0.308 | 0.833 | 0.865 | 0.774 |
| 511 | 0.703 | 0.812 | 0.814 | 0.873 | 0.542 | 0.372 | 0.000 | 0.218 | 0.745 |