



## **What seismicity offshore Sicily suggests about lithosphere dynamics and microplate fragmentation models in the Central Mediterranean**

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9 **Abstract**

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11 We analyze an updated dataset of earthquakes of Southern Italy, focusing in particular on  
12 hypocenter locations and seismogenic stress distributions in the southern and eastern offshores of  
13 Sicily, the two sectors of the study region where seismic and geodetic information needed for  
14 geodynamic modeling is still poor because of poor geometry of monitoring networks. Using  
15 Bayesian non-linear methods for hypocentral locations and hypocenter error estimates we improve  
16 the earthquake locations performed by more traditional linearized techniques, and this helps us to  
17 make significant progress in the interpretation of seismicity and seismogenic stress distributions  
18 especially where seismometric network geometry is more critical. Epicenter maps and hypocenter  
19 vertical sections, together with (i) best quality focal mechanisms coming from seismic waveform  
20 inversion and (ii) orientations of stress principal axes estimated by inversion of focal mechanisms,  
21 help us to better recognize geodynamic engines and plate margin deformation in the study area.  
22 NW-trending convergence between Africa and Eurasia is recognized as the main source of tectonic  
23 stress in the study region, producing clearly detectable signatures in terms of  $\sigma_1$  orientations also in  
24 the offshore sectors of the western Ionian and the Sicily Channel. Seismicity and seismogenic stress  
25 tensor highlight nearly uniform compressional dynamics related to plate convergence in the Sicily  
26 Channel, in contrast to rifting and microplate divergence proposed in that sector by other



27 investigators. In the western Ionian, seismicity and stress inversion results reveal superposition of  
28 convergence-related compression and extensional dynamics. The latter, characterized by minimum  
29 compressive stress oriented SW-NE, can be related to a rifting process (opening SW-NE)  
30 hypothesized by previous investigators on the basis of marine geophysics analyses performed  
31 between the Alfeo-Etna and the Ionian Faults. The seismicity and seismogenic stress detected in the  
32 Western Ionian show that assumptions of microplate rigidity in this area made by previous workers  
33 when modeling poor geodetic data available can be inappropriate. Our findings indicate that more  
34 complex rheologic models should be adopted for reconstruction of tectonic deformation and  
35 microplate relative motions in the Central Mediterranean region.  
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## 37 **1 Introduction**

38

39 The progressive increase in the last few decades of available data concerning crustal motions and  
40 strains has allowed the researchers to intensify the investigations on the geometry, kinematics and  
41 dynamics of the plates, microplates and tectonic units in the Mediterranean region, in particular in  
42 the central part of it corresponding to south Italy and the Tyrrhenian and Ionian seas (Figs 1 to 3). In  
43 spite of the progresses made in the geophysical data acquisition and analysis, different views still  
44 exist, however, concerning the fragmentation of lithosphere and microplate architecture and  
45 kinematics in this portion of the Africa-Eurasia convergent margin (see e.g. Anderson and Jackson,  
46 1987; Oldow et al., 2002; Battaglia et al., 2004; Serpelloni et al., 2007; Nocquet, 2012; Sani et al.,  
47 2016, among others). A wide description of the debate and state of art of knowledge on these topics  
48 is given in Sect. 3.

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50 The relatively large number of models and hypotheses still resisting to checks by new data and  
51 analyses is mainly due to poor distribution of GPS stations in the offshore areas and to the often  
52 concomitant assumption of internally rigid crustal blocks adopted when modeling the  
53 plate/microplate kinematics. The puzzle is complicated by too weak evidence of potential  
54 microplate boundaries furnished by earthquake activity resulting scarce as regard to (i) number and  
55 energy of seismic events and (ii) geometrical continuity of seismolineaments. Other kinds of data  
56 and information from geophysics and geology have not, yet, solved the ambiguities (Oldow et al.,  
57 2002; D'Agostino et al., 2008; Nocquet, 2012).

58

59 In the present study we use an updated set of earthquake data with the purpose of contributing to the  
60 current debate on microplate architecture and dynamics in the Central Mediterranean, paying  
61 particular attention to a few sectors offshore Sicily, namely the Western Ionian and the Sicily  
62 Channel (WI and SC in Fig. 4a), resulting of crucial relevance for the solution of some still existing



63 uncertainties. The potentialities of the methods of seismological analysis we have planned to use, in  
64 part of recent conception and in all cases already proven to be effective in the study region (see e.g.  
65 Neri et al., 2005; Billi et al., 2010; D'Amico et al., 2010; Presti et al., 2013), make us confident that  
66 the above declared goal of the investigation may be reached.

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## 69 **2 Geodynamic frame of the study region**

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71 In the Mediterranean region (Fig. 1), after a long period between Late Paleogene and Neogene of  
72 Africa NW-ward subduction beneath Eurasia, subduction has almost ceased (Billi et al., 2011).  
73 With the progression of Africa-Eurasia convergence, a tectonic reorganization of the plate boundary  
74 has started to accommodate contraction, in particular contractional deformation in several segments  
75 of the boundary (such as Sicily) has shifted from the former subduction zone to the margins of the  
76 back-arc oceanic basins (Fig. 2). The tectonic reorganization of the boundary is still strongly  
77 controlled by the inherited tectonic fabric and rheological attributes, which are strongly  
78 heterogeneous along the boundary.

79

80 In Italy (Fig. 1) the Neogene convergence and associated subduction between Africa and Eurasia  
81 resulted in the NW-trending Apennine and the W-trending Maghrebic fold-thrust belts in  
82 peninsular Italy and Sicily, respectively (Malinverno and Ryan, 1986). The two belts are connected  
83 through the Calabrian Arc (Minelli and Faccenna, 2010), below which a narrow remnant of the  
84 former subducting slab seems to be still active, but close to cessation (Neri et al., 2009; 2012;  
85 Orecchio et al., 2015; Chiarabba and Palano, 2017).

86

87 Contraction in Sicily is, at present, mainly accommodated at the rear of the fold-thrust belt, in the  
88 southern Tyrrhenian sea (Fig. 2), where a series of contractional earthquakes located in an E-



89 trending belt between Ustica and Eolian islands have been recorded in the last decades (Pondrelli et  
90 al., 2004; Billi et al., 2007; Presti et al., 2013; Orecchio et al., 2017). Toward the east, the  
91 compressional seismic belt is delimited by the seismically active Tindari Fault System (TFS in Fig.  
92 2), to the east of which, both earthquakes and GPS data provide evidence for an ongoing  
93 extensional tectonics possibly connected with the residual subduction beneath the Calabrian Arc  
94 and related back-arc extension (Palano et al., 2015b). The age for the onset of the ongoing  
95 contraction in the south-Tyrrhenian margin is unknown, but the cessation of volcanism at Ustica  
96 (i.e., the already mentioned volcanic island located along the south-Tyrrhenian contractional belt)  
97 during Middle-Late Pleistocene may be connected with the onset of contractional tectonics in this  
98 area. This age corresponds approximately with the cessation of contractional displacements along  
99 the outermost Gela nappes in southern Sicily (Fig. 2; Ghisetti et al., 2009).

100

101 The plate margins strongly simplified in Fig. 1 have undergone significant changes over time as a  
102 consequence of the above processes and of the whole geodynamic activity occurring in the  
103 Mediterranean region. Recent investigations performed thanks to the increasing set of geodetic data  
104 available have brought the researchers to propose different scenarios of microplate fragmentation  
105 and reorganization along the boundary, in particular along the south Italy part of the boundary (see  
106 references quoted in the introduction). The next section presents a list of microplate architecture  
107 scenarios proposed in the most relevant literature and evidences the doubts and open questions still  
108 existing in the scientific community concerning the microplate geometry and kinematics in the  
109 Italian part of the Africa-Eurasia plate boundary.

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111

### 112 **3 The microplate puzzle of the Central Mediterranean**

113



114 Two of the former studies reporting on the possible fragmentation of lithosphere in the Italian  
115 portion of the Africa-Eurasia boundary were those by Anderson (1987) and Anderson and Jackson  
116 (1987). These authors explained the focal mechanisms and related slip vectors of the most  
117 significant earthquakes located around the Adriatic sea assuming that a rigid block (Adria) is  
118 independent from the two main plates and rotates anticlockwise around a pole located in northern  
119 Italy (Fig. 3b). They, however, remarked the main weakness of their model represented by the lack  
120 of significant seismicity at the presumed southern border of Adria microplate. Oldow et al. (2002)  
121 and Oldow and Ferranti (2006) proposed a quite different architecture of microplates and blocks in  
122 the Italian region, with a southeastern ‘Adria’ block separated from a northwestern ‘Adria’ one by a  
123 line crossing south Italy and the northern Adriatic sea (Fig. 3c).

124

125 Closer to Anderson and Jackson’s (1987) reconstruction, Battaglia et al. (2004) explained GPS data  
126 re-introducing an Adriatic domain strictly corresponding to the Adriatic sea, independent from  
127 Africa and Eurasia, but separated in two blocks, Northern and Southern Adria, confining in the  
128 Central Adriatic sea (Fig. 3d). Serpelloni et al. (2007) proposed an even more complex  
129 fragmentation of the whole plate boundary, suggesting in particular that a Sicilian domain is  
130 moving independently from Africa according to the presence of a right-lateral and extensional  
131 decoupling zone corresponding to the Tunisia-Libya and Sicily Channel deformation zone (Fig. 3e).  
132 According to the same authors, the tectonics and kinematics of the Italian region are further  
133 complicated by Ionian oceanic lithosphere subducted beneath Calabria and by Adria independent  
134 microplate rotating anticlockwise with respect to Africa and Eurasia main plates. Serpelloni and  
135 coworkers remark, however, that lack of GPS data and poor seismic network geometry in wide  
136 offshore sectors like the Ionian basin leave some uncertainties in the geodynamic modeling, in  
137 particular they suspect (but cannot prove) the existence of an independent Ionian microplate  
138 rotating counterclockwise between Africa and Adria plates (Fig. 3e). In this connection the same  
139 authors leave open the question “where strain locates between Hyblean and Apulia domains?”.



140 D'Agostino et al. (2008) interpreted their GPS and earthquake slip data (i) by limiting the  
141 anticlockwise rotating rigid Adria microplate to the northern Adriatic region and (ii) by introducing  
142 a larger, newly defined microplate to the south, including the Apulia promontory, the Ionian sea and  
143 the Hyblean region in southern Sicily (Fig. 3f). According to these authors, the hypothesis of  
144 Hyblean region belonging to such a hypothetical microplate would be supported by apparently low  
145 GPS-derived deformation in the western Ionian. In any case the authors admit that lack of data in  
146 the Ionian offshore of Sicily does not allow decisive checking of their assumption of two rigid  
147 microplates located between Africa and Eurasia, one rotating clockwise (Apulia-Ionian-Hyblean)  
148 and the other anticlockwise (Adria).

149

150 More recently, by analysis of a relatively long period of 18 years of GPS observations, Palano et al.  
151 (2012) supported the thesis of an Hyblean block independent with respect to Africa and Apulia (Fig.  
152 3g-h). They located the regional contraction existing between the Tyrrhenian and Hyblean blocks in  
153 two distinct belts identified in the northern Sicily offshore (Ustica-Eolie) and across the Sicily front  
154 (Sicilian basal thrust according to Lavecchia et al., 2007). The authors discussed also the role  
155 played by the Ionian domain and suggested two possible scenarios, one assuming that the Ionian is  
156 rigidly connected with the Hyblean block (Fig. 3g), the other assuming that the Ionian domain  
157 diverges from the Hyblean block and moves to northeast wrt to Eurasia (Fig. 3h). They concluded  
158 that the lack of islands (i.e. of data) in the Ionian offshore does not allow to make a choice among  
159 these scenarios. Based on the analysis of GPS velocities and earthquake focal mechanisms in the  
160 Central and Eastern Mediterranean, Pérouse et al. (2012) proposed that the area including the  
161 Hyblean Plateau, the Ionian basin, the Apulian peninsula, the south Adriatic sea and the Sirte plain  
162 may be considered as a single rigid block rotating clockwise wrt Africa, inducing an opening of a  
163 couple of mm/yr in the Sicily Channel - Pelagian rift (Fig. 3i). The same authors, similar to  
164 D'Agostino et al. (2008 and 2011), suggested that the 2-2.5 mm/yr slow trenchward motion of the  
165 Calabrian Arc wrt to the Apulian-Ionian-Hyblean-Sirte domain may be explained in terms of



166 ultraslow residual subduction or, alternatively, by pure gravitational collapse into an inactive  
167 subduction scenario. More recently, the Oldow et al.'s (2002) scheme of Fig. 3c was reposed  
168 with modifications by Sani et al. (2016) (Fig. 3j).

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170 In his review of papers and investigations regarding the crustal kinematics in the Mediterranean  
171 region, Nocquet (2012) drew the conclusion that it is quite difficult to state from the available data  
172 and analyses where stable Africa finishes and other eventual blocks like Apulia begin in the Central  
173 Mediterranean area at the longitude of Italy. In a very recent analysis performed by integration of  
174 multibeam, seismic reflection, magnetic and gravity data, Polonia et al. (2017) have concluded that:  
175 (i) tearing at the southwestern edge of the SEward-retreating Calabria subduction slab may be the  
176 deep source of shallow deformation detected in correspondence of the transtensional Ionian Fault in  
177 the Ionian basin (Fig. 2); (ii) the NW-SE trending belt comprised between the Ionian Fault and the  
178 Alfeo-Etna Fault (Fig. 2) hosts a rift zone opening SW-NE where serpentinite diapirs have been  
179 identified by the authors. On their hand, Gutscher et al. (2017), by analysis of multi-beam  
180 bathymetric data and seismic profiles, proposed the Alfeo-Etna right-lateral fault system as shallow  
181 tectonic expression of the retreating slab tear or STEP fault. In the view of these authors some  
182 sinistral lateral component appearing in the southeasternmost part of the Ionian Fault should  
183 exclude the latter as potential expression of the STEP fault. Another major fault system marking the  
184 transition between the Ionian basin and the Hyblean plateau, e.g. the Malta escarpment (Fig. 2),  
185 shows not to be currently active along most of its length and shows signs of recent faulting only in  
186 its northernmost segment (see, e.g., Argani, 2009; Gutscher et al., 2016). West of the Malta  
187 escarpment (Fig. 2), a detailed analysis of reflection profiles and stratigraphic and structural data  
188 allowed Cavallaro et al. (2016) to evidence clear time evolution from tensional to compressive  
189 regimes in the Sicily Channel WNW-trending main structural system, until its present-day  
190 behaviour as transcurrent fault system under compression due to NW-SE Africa-Eurasia  
191 convergence. Cavallaro et al. (2016) proposed, in particular, that local volcanism stopped in late



192 Miocene and Miocene-time normal faults were reactivated during Zanclean-Piacenzian age as right-  
193 lateral strike-slip faults.

194

195 The above description of scenarios and findings highlights uncertainties still existing concerning the  
196 plate boundary evolution and the present-day architecture and kinematics of microplates and  
197 lithospheric blocks in the Central Mediterranean. Poor distribution of GPS and seismic networks in  
198 the offshore sectors of Western Ionian and Sicily Channel cause most of these uncertainties (see,  
199 among others, the already quoted papers by Serpelloni et al., 2007; D'Agostino et al., 2008; Palano  
200 et al., 2012; Nocquet, 2012). In the present study, we start with a regional-scale analysis of recent  
201 seismicity updated to the end of 2016. Then, we focus on the most critical sectors of the Western  
202 Ionian and the Sicily Channel where new seismic data and improvements of knowledge concerning  
203 earthquake and seismogenic stress distribution can help answering the questions left open by the  
204 previous investigations.

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#### 207 **4 Data, methods of analysis and results**

208

209 We have taken from the Italian national seismic catalog ([http://istituto.ingv.it/l-ingv/archivi-e-](http://istituto.ingv.it/l-ingv/archivi-e-banche-dati/)  
210 [banche-dati/](http://istituto.ingv.it/l-ingv/archivi-e-banche-dati/)) and from the databases of the local seismic networks operating in Sicily and Calabria  
211 (Orecchio et al., 2011) the data of the earthquakes of magnitude over 2.5 that occurred between  
212 1981 and 2016 at depths less than 100 km in the area  $10^{\circ}$ - $20^{\circ}$ E  $35^{\circ}$ - $41^{\circ}$ N bounded by the dashed  
213 rectangle in Fig. 4a. We have selected only the events for which a minimum of 15 P+S arrival times  
214 were available. For these events we have performed hypocenter locations by the standard, linearized  
215 location method Simulps (Evans et al., 1994) and the 3D seismic velocity structure proposed for the  
216 study region by Orecchio et al. (2011). The epicenter maps of the earthquakes located with this  
217 procedure, corresponding to different hypocenter depth ranges, are shown in Fig. 4b-d. Then, we



218 have selected from the literature and the international catalogs all the fault-plane solutions estimated  
219 by waveform inversion for the earthquakes of magnitude over 2.5 occurring in the period 1977-  
220 2016 at depths less than 70 km in the same area of the above locations (dashed rectangle of Fig. 4).  
221 The map of these fault-plane solutions is shown in Fig. 5, the list of focal parameters is furnished in  
222 Table A1, Appendix A.

223

224 A new step of analysis was that of using the Bayesian non-linear location algorithm named Bayloc  
225 (Presti et al., 2004 and 2008) for relocation of the earthquakes located by Simulps in the two sectors  
226 indicated by WI (Western Ionian) and SC (Sicily Channel) in Fig. 4a. The same 3D local velocity  
227 structure by Orecchio et al. (2011) has been used also in this new phase of analysis. Starting from  
228 seismic phase arrival times at the recording stations, Bayloc computes for an individual earthquake  
229 a probability cloud marking the hypocenter location uncertainty. Then, Bayloc estimates the spatial  
230 distribution of probability relative to a set of earthquakes by summing the probability densities of  
231 the individual events. This method has been shown to help detection of the seismogenic structures  
232 through better hypocenter location and more accurate estimation of location errors compared to  
233 linearized methods (Presti et al., 2008) but computational reasons make its application easier when  
234 carried out in small areas (Presti et al., 2004). For this reason we have used Simulps for locating the  
235 whole dataset of events of Fig. 4 and Bayloc for locations in the two sectors of more crucial  
236 relevance in the present study, namely WI and SC. Details on the methodological aspects of Bayloc  
237 can be found in the above quoted papers. The epicenter maps and hypocenter vertical sections  
238 obtained by Bayloc are shown in the Figs 6 and 7. Epicenter and hypocentre errors of the order of 3  
239 km and 5 km have been estimated for the earthquakes of Sector WI, rising to values of 4 km and 9  
240 km in Sector SC.

241

242 The information furnished in the bibliographic sources of the focal mechanisms of Fig. 5 indicates  
243 that these focal mechanisms should be characterized by fault parameter errors of the order 10-15



244 degrees, then generally smaller than errors of focal mechanisms computed by inversion of P-onset  
245 polarities in areas of critical network geometry like ours (D'Amico et al., 2011; Presti et al., 2013;  
246 Musumeci et al., 2014). This level of uncertainty makes the dataset of Fig. 5 suitable for application  
247 of the method by Gephart and Forsyth (1984) and Gephart (1990) for calculating the seismogenic  
248 stress tensor directions in the study region. This method searches for the stress tensor showing the  
249 best agreement with the available focal mechanisms (FMs). Four stress parameters are calculated:  
250 three of them define the orientations of the main stress axes; the other is a measure of relative stress  
251 magnitudes,  $R = (\sigma_2 - \sigma_1) / (\sigma_3 - \sigma_1)$ , where  $\sigma_1$ ,  $\sigma_2$  and  $\sigma_3$  are the values of the maximum, intermediate  
252 and minimum compressive stresses, respectively. In order to define discrepancies between the stress  
253 tensor and observations (FMs), a misfit variable is introduced: for a given stress model, the misfit of  
254 a single focal mechanism is defined as the minimum rotation about any arbitrary axis that brings  
255 one of the nodal planes, and its slip direction and sense of slip, into an orientation that is consistent  
256 with the stress model. Searching through all orientations in space by a grid technique operating in  
257 the whole space of stress parameters, the minimum sum of the misfits of all FMs available is found.  
258 The confidence limits of the solution are computed by a statistical procedure described in the papers  
259 by Parker and Mc Nutt (1980) and Gephart and Forsyth (1984). The size of the average misfit  
260 corresponding to the best stress model provides a guide as to how well the assumption of stress  
261 homogeneity is fulfilled (Michael 1987). In the light of results from a series of tests carried out by  
262 Wyss et al. (1992) and Gillard et al. (1996) to identify the relationship between FM uncertainties  
263 and average misfit in the case of uniform stress, we will assume that the condition of homogeneous  
264 stress distribution is fulfilled if the misfit,  $F$ , is smaller than  $6^\circ$ , and that it is not fulfilled if  $F > 9^\circ$ . In  
265 the range  $6^\circ < F < 9^\circ$ , the solution is considered as acceptable, but may reflect some heterogeneity.  
266 For the application of Gephart and Forsyth's (1984) method in the present study, we have focused  
267 on the area contoured by the thin line in Fig. 5 including southern Sicily, the Sicily Channel and the  
268 Western Ionian, e.g. the sectors where the knowledge of seismogenic stress distributions is poorer



269 than elsewhere in the region (see, among others, Totaro et al., 2016). The stress inversion results  
270 obtained in the present study are reported in Table 1 and Fig. 8.

271

272

## 273 **5 Discussion**

274

275 The epicenter and focal mechanism maps of Figs 4 and 5 evidence two main features of the regional  
276 seismicity already known from previous investigations. The first of these features is represented by  
277 clear spatial grouping of shallow earthquakes in correspondence of the Apennine-Maghrebian chain  
278 from the southern Apennines to Sicily (Fig. 4c) marking response to perpendicular-to-chain  
279 extensional stress (Fig. 5). The second feature is the nearly east-trending compressional seismic belt  
280 appearing from Figs 4c and 5 in the southern Tyrrhenian sea offshore Sicily, ending to east in the  
281 Eolian Islands area where the orientation of the belt and the faulting style clearly change to NW-SE  
282 and dextral strike-slip, respectively. Moderate activity can be detected from the same figures in the  
283 Sicily Channel (compressional regime) and in the Western Ionian sea (compression prevailing but  
284 not exclusive). Also, a clear drop of activity can be noted in the Ionian offshore of Southern  
285 Calabria (Figs 4c and 5). These features confirm the picture of regional seismicity given by the  
286 previous investigators who distinguished a compressional domain of the southern Tyrrhenian and  
287 Ionian seas due to Africa-Eurasia NW-trending convergence and an extensional one along the  
288 Apennine-Maghrebian chain due to different factors in the different segments of the chain (see e.g.,  
289 Presti et al., 2013; Totaro et al., 2016; Orecchio et al., 2017).

290

291 The previous sections have described the efforts made by many investigators to identify opening  
292 zones between diverging microplates in the Sicily Channel and the Ionian sea with the purpose of  
293 explaining the space variations of crustal motions measured by GPS networks in the Central  
294 Mediterranean region. Lack of data in wide offshore sectors is the main reason why the debate on



295 microplate geometry and kinematics in the region is still open. The results of the present study may,  
296 in our opinion, furnish a useful contribution to this debate. Stress tensor inversion of earthquake  
297 fault-plane solutions in the area including Southern Sicily, the Sicily Channel and the Western  
298 Ionian sea (Fig. 8a, set ALL) reveals moderate stress heterogeneity (F-value of  $8.3^\circ$ ) around a best  
299 model of stress characterized by a sub-horizontal, NW-trending  $\sigma_1$  clearly reductible to Africa-  
300 Eurasia convergence (Fig. 8a and Table 1). Starting from this result, we have sub-divided the  
301 dataset of focal mechanisms of Fig. 8a in tens of subsets according to the epicenter distribution,  
302 focal depth and magnitude of the earthquakes, in order to search for subsets satisfying the condition  
303 of stress homogeneity. As explained in the previous Section, this condition can be considered  
304 reasonably satisfied in the present study when the F-value of inversion is lower than  $6^\circ$ . In order to  
305 guarantee the significance of stress computations in the different subsets we have decided to fix a  
306 minimum number of 20 earthquakes (= focal mechanisms) for the creation of the individual subset.  
307 The stress inversion results obtained after the first step of partitioning of the dataset into different  
308 subsets indicated new strategies of partitioning or data grouping. For sake of conciseness, we do not  
309 report the stress inversion results obtained for all the subsets investigated, we only report in Fig. 8  
310 and Table 1 the results that we consider more meaningful, that is, able to better outline the stress  
311 patterns and tectonic features in the study area.

312

313 Fig. 8b and Table 1 (lines W and E) report the stress inversion results obtained by subdividing the  
314 study area in two sectors W and E located, respectively, west and east of a NW-trending separation  
315 line indicated as AB in the same Fig. 8b. A F-value of  $5.9^\circ$  shows that stress is homogeneous or  
316 close to homogeneity in the W sector (Table 1). The best model of stress in this western sector is  
317 similar to that obtained by inversion of the mechanisms of the whole study area (Fig. 8a and Table  
318 1). The 95% confidence limits of stress orientations in W (Fig. 8b) show an acceptable level of  
319 constraint of  $\sigma_1$  orientation. On the other hand, the confidence area of  $\sigma_3$  orientation in the same  
320 sector is relatively large and extends from vertical to SW-NE horizontal direction, i.e.  $\sigma_3$  and  $\sigma_2$  are



321 unconstrained on the SW-NE vertical plane, plausibly in relation to co-existence of reverse and  
322 strike-slip seismic faulting in the study volume under SE-NW compression due to plate  
323 convergence. These results are in good agreement with the geostructural and geodynamic  
324 reconstruction of this area proposed by Cavallaro et al. (2016) who evidenced, in particular, long  
325 term evolution from tensional to compressive regimes in the Sicily Channel until the present-day  
326 state of compression led by NW-SE Africa-Eurasia convergence.

327

328 The inversion of the fault-plane solutions available in sector E of Fig. 8b leads to a F-value of  $8.3^\circ$   
329 indicating a certain degree of stress heterogeneity as already inferred from the analysis of the whole  
330 dataset of Fig. 8a. The best model of stress in sector E (Fig. 8b and Table 1) is quite similar to those  
331 of sets ALL and W (Table 1), and this indicates that NW-SE compression from plate convergence is  
332 again detectable in this eastern part of the study area. However, the 95% confidence limits of stress  
333 orientations in E reveal that  $\sigma_1$  orientation is practically unconstrained from horizontal NW-SE to  
334 vertical, and this leads us to suppose that some extensional process opening SW-NE acts together  
335 with NW-trending plate convergence in this sector.

336

337 By comparing the location of the separation line between the E and W domains of Fig. 8b with the  
338 structural map of Fig. 2 we can note that the separation line approximately corresponds with the  
339 location of the Alfeo-Etna Fault. Again, by comparing the Bayloc's epicenter distribution of the  
340 earthquakes shallower than 30 km in Sector WI (Fig. 6c) with the structural map of Fig. 2 we may  
341 note (i) a certain degree of activity in correspondence with the northern segment of the Malta  
342 escarpment, the only considered active in this structural system by several authors (see, e.g.,  
343 Argnani, 2009; Gutscher et al., 2016), (ii) a clear belt of activity trending NW-SE between the  
344 Alfeo-Etna and the Ionian faults and (iii) a drop of seismicity going to NE across the Ionian Fault.  
345 In the 30-70 km depth range (Fig. 6d) there is no seismicity in correspondence with the Malta  
346 escarpment and most of seismic activity is dispersed around the Ionian Fault. The SW-NE vertical



347 section of Fig. 6b highlights that seismicity is as shallow as 20 km west of the Alfeo-Etna Fault but  
348 deepens to depths of the order of 40 km between the Alfeo-Etna and the Ionian faults, and remains  
349 stable at this depth level for several tens of km around the Ionian Fault. In the same vertical section  
350 the presence of seismic activity can also be noted in the upper 20 km halfway between the Alfeo-  
351 Etna and the Ionian faults (AEF and IF).

352

353 Previous investigators (see, e.g., Polonia et al., 2011; 2016; Palano et al., 2015b) have proposed that  
354 the Ionian fault zone corresponds with the southwestern edge of the Ionian subducting slab dipping  
355 to NW and presumably still affected by slow SE-ward residual rollback. The Ionian fault zone  
356 would represent the shallow expression of a STEP fault (Polonia et al., 2016). The above  
357 commented epicenter and hypocenter distributions of Fig. 6 support this interpretation, evidencing  
358 however that the dislocation process between the subducting slab (northeast) and the adjacent  
359 lithosphere (southwest) is distributed over a relatively wide zone probably because the subduction  
360 and slab rollback kinematics are very slow and do not mimic fast STEP dynamics (Gallais et al.,  
361 2013; Orecchio et al., 2014). The earthquake activity detected in the upper 20 km halfway between  
362 the Alfeo-Etna and Ionian faults (Fig. 6b) occurs in a NW-trending highly heterogeneous and  
363 fractured zone identified by Polonia et al. (2017) as the site of a rifting process with SW-NE  
364 opening direction where serpentinite diapirs rise from deeper depths. This extensional process can  
365 be a cause of stress heterogeneity evidenced by stress inversion in the E domain of Fig. 8b where  
366 SW-NE extension appears to be superimposed to NW-SE compression related to Africa-Eurasia  
367 convergence. The rifting process identified in this part of the Ionian sea by Polonia et al. (2017)  
368 may, therefore, contribute to explain our results from analysis of seismicity and stress inversion of  
369 earthquake fault-plane solutions. This interpretation appears also to match well with the space  
370 variation of GPS vectors in Sicily and Calabria (Fig. 9; vector data from Palano et al., 2012)  
371 showing clear rotation of crustal motions when crossing the onshore prolongation of the NW-  
372 trending zone between the Alfeo-Etna and Ionian faults. Having more than 20 focal mechanisms in



373 the NW-trending zone where Polonia et al. (2017) detected the rifting process, we have performed a  
374 stress inversion run in this zone (RZ in Fig. 8c). The results (shown in the same Fig. 8c and in Table  
375 1) are similar to those obtained in the E domain of Fig. 8b and confirm our hypothesis of  
376 combination of NW-SE compression due to plate convergence and SW-NE rifting. We have also  
377 attempted to detect possible depth variations of the stress field in this rifting zone RZ and found that  
378 the convergence-related compression seems to prevail at depths deeper than 30 km, while the  
379 combination of the two stress factors appears clearer at shallower depth. However, this latter result  
380 must be considered as uncertain or highly preliminary because the number of data available in the  
381 0-30 km and 30-70 km depth ranges is 15 and 12, respectively, then lower than the threshold of 20  
382 adopted in this work for stress inversion. We do not report these latter results in graphical and  
383 numerical form, we wait for future analyses with more data before drawing conclusions on the  
384 depth variation of stress in the rifting zone of Polonia et al. (2017). We conclude our discussion on  
385 stress inversion results by remarking that partitioning of the dataset according to earthquake  
386 magnitude has not evidenced any relevant change of stress or variation of heterogeneity level  
387 among subsets. An implication of this result is that a quite frequent situation documented since  
388 decades in the literature is not observed in the case of our dataset: we specifically refer to the  
389 situation of stronger earthquakes marking a uniform regional stress locally disturbed, however, by  
390 smaller scale stress heterogeneities marked by weaker events (see, e.g., Rebaï et al., 1992;  
391 Christova, 2015).

392

393 The Bayloc's maps of epicenters in Sector SC (Fig. 7) show that seismicity is mainly located in  
394 correspondence with a ca. NNE-SSW trending fault system crossing the WNW-ESE Pelagian rift  
395 area. The depths of these events cover all the range of investigation of our analysis, i.e. 0-70 km.  
396 The seismic activity of this NNE-SSW trending fault system covering this relatively large depth  
397 range has already been detected by previous investigators (Calò and Parisi, 2014). Our results show  
398 also that some activity occurs in the Hyblean corner of Sicily and south of it, but in this case the



399 seismicity is shallow and never exceeds the depth of 30 km. Minor activity is located in  
400 correspondence with the fault system of the Pelagian rift. These results, together with the nearly  
401 homogeneous stress tensor marking SE-NW plate convergence in sector W of Fig. 8b, do not  
402 support active rifting in the Sicily Channel. In other words our results do not support the hypothesis  
403 of microplate separation and divergence in the Sicily Channel advanced in previous works based on  
404 analysis of poorly distributed geodetic data and microplate rigidity assumptions (see, e.g.,  
405 D'Agostino et al., 2008; Perouse et al., 2012). On the other hand, our results match well with the  
406 results recently obtained in the same area by detailed analysis of reflection profiles and stratigraphic  
407 and structural data indicating clear time evolution from tensional to compressive regimes in the  
408 Sicily Channel WNW-trending main structural system, until its present-day behaviour as  
409 transcurrent fault system under compression due to NW-SE Africa-Eurasia convergence (Cavallaro  
410 et al., 2016). According to these authors, Miocene-time extensional structures of the Sicily Channel  
411 WNW-trending structural system were reactivated during Zanclean-Piacenzian age as right-lateral  
412 strike-slip faults under convergence-related compression.

413

414 If we re-enlarge the frame of analysis to include the southern Tyrrhenian region (Fig. 9) we find  
415 additional signature of the SE-NW Africa-Eurasia convergence in the southern Tyrrhenian east-  
416 trending compressional belt located offshore northern Sicily (see also Presti et al., 2013 and Totaro  
417 et al., 2016 for stress orientations in the southern Tyrrhenian seismic belt). On this scale of analysis  
418 our results match well with the very recent Nijholt et al's (2018) conclusion according to which in  
419 this south-central part of the Mediterranean region "the Calabrian arc is now further transitioning  
420 towards a setting dominated by Africa-Eurasia plate convergence, whereas during the past 30 Myrs  
421 slab retreat continually was the dominant factor". Also, our investigation leads us to evidence that:  
422 (i) perpendicular-to-convergence opening in the Ionian sea (between the Alfeo-Etna and Ionian  
423 faults) introduces some seismogenic stress heterogeneity in the eastern compartment of our stress  
424 inversion area (Figs 8b, 8c and 9); (ii) residual, very slow subduction and SE-ward rollback of the



425 Ionian lithosphere northeast of the Ionian Fault reduces convergence-related compression in the  
426 shallow Ionian offshore of southern Calabria and causes there a detectable drop of seismicity  
427 indicated by the arrow in Fig. 4c. This process of reduced compression can also influence stress  
428 parameters in this sector with particular reference to the Ionian offshore of Calabria, but with the  
429 data available we cannot analyze this aspect in a greater detail. We look at a future availability of  
430 additional data in the eastern sector of Fig. 8b to explore in a greater detail the local space variations  
431 of stress and the contributions by the different tectonic factors.

432

433 Our results may contribute to answer some questions put by previous investigators such as, for  
434 example, Serpelloni et al. (2007) (where strain locates between Hyblean and Apulia domains?) or  
435 Palano et al. (2012) (is Ionian rigidly connected with the Hyblean block or diverging from it and  
436 moving to northeast wrt to Europe?). Our results allow us to state that (i) the Western Ionian is a  
437 main site of strain release between Hyblean and Apulia domains and (ii) the Ionian is not “rigidly  
438 connected with the Hyblean block”. Following the reasoning of Palano et al. (2012) who concluded  
439 with the proposal of two alternative scenarios reported in our Fig. 3g-h, we tend to favour the  
440 hypothesis of a Ionian block diverging from the Sicilian-Hyblean-Malta block (3h), with the  
441 boundary between the respective blocks located, however, as in Fig. 3g or slightly eastward. Our  
442 results lead us to believe that the rifting process suggested by Polonia et al. (2017) between Alfeo-  
443 Etna and Ionian faults in the Ionian basin should be taken in consideration in the future modeling of  
444 microplate geometry and kinematics. Concerning the doubts of D’Agostino et al. (2008) whether  
445 their assumption of a rigid Apulian-Ionian-Hyblean microplate is correct or not, our results in the  
446 Western Ionian showing evidence of a rifting process as proposed by Polonia et al. (2017) suggest  
447 that it is not.

448

449 We strongly feel that the usual assumption of rigid blocks or microplates in the physical modeling  
450 of geodynamic processes should be overcome and different crustal rheologies should be tested



451 when modeling tectonic deformation and microplate relative motions in the Central Mediterranean  
452 region. For this purpose, we have recently started a research collaboration with the geophysical  
453 team of University of Milan for Finite Element Modeling of tectonic stress and strain distributions  
454 through high-resolution 3D thermo-rheological representation of lithosphere in the region of our  
455 interest. This line of research is grounded on methodological developments described in the papers  
456 by Splendore and Marotta (2013) and Marotta et al. (2015), among others.

457

458

## 459 **6 Conclusion**

460

461 Seismicity spatial patterns, earthquake focal mechanisms and seismogenic stress tensor orientations  
462 in and around Sicily, analyzed by use of seismometric data recorded in the last few decades, mark  
463 the dominant action of Africa-Eurasia NW-oriented plate convergence in this part of the  
464 Mediterranean region. Evidences of other tectonic factors acting together (or superimposed to) plate  
465 convergence are found in the Western Ionian sea where (i) residual slow subduction and SE-ward  
466 trench retreat reduce plate coupling in the Southern Calabria offshore and (ii) a rifting process with  
467 opening direction perpendicular to convergence at the southwestern edge of the subducting slab  
468 adds extensional stress to convergence-related compression in the offshore of Eastern Sicily. No  
469 seismic evidence of active rifting or microplate separation and divergence is detected in the  
470 Pelagian area of the Sicily Channel, where clear signatures of plate convergence are found, in  
471 agreement with findings of recent analyses of reflection profiles and structural data performed in the  
472 specific area (Cavallaro et al., 2016). In other words, our results do not support the hypothesis of  
473 Sicily Channel rifting dynamics advanced by previous investigators using geodetic data under  
474 microplate rigidity assumption (e.g. D'Agostino et al., 2008; Perouse et al., 2012). Also, our results  
475 answer several open questions on tectonic strain in the Western Ionian left by Serpelloni et al.  
476 (2007) (where strain locates between Hyblean and Apulia domains?), or by Palano et al. (2012) (is



477 Ionian rigidly connected with the Hyblean block or diverging from it and moving to northeast wrt to  
478 Europe?) or, again, by D’Agostino et al. (2008) who admit that lack of GPS data in the Ionian  
479 offshore does not allow decisive checking of their model assuming a rigid Apulia-Ionian-Hyblean  
480 microplate. Even taking into account the intrinsic limitations of our seismic datasets relative to  
481 (possibly short) time intervals of a few decades, but also considering their significance where GPS  
482 data are lacking or poor, the seismic data in our possess indicate that (i) strain “between Hyblean  
483 and Apulia domains” mainly locates around the southwestern edge of the Ionian subduction slab in  
484 the westernmost Ionian, (ii) the Ionian is not “rigidly connected with the Hyblean block” and (iii)  
485 the assumption of “a rigid Apulia-Ionian-Hyblean microplate” needs to be revised. Our results show  
486 that the assumption of rigid blocks or microplates made by previous investigators in their kinematic  
487 reconstructions should be overcome and different crustal rheologies should be tested when  
488 modeling tectonic deformation and microplate relative motions in this region. High-resolution 3D  
489 thermo-rheological representations of lithosphere in the frame of Finite Element Modeling of  
490 tectonic stress and strain distributions can, in our opinion, be an appropriate road towards  
491 geodynamic modeling of Southern Italy and the Central Mediterranean region. We are starting to  
492 work in this direction.  
493

494 **Appendix A**

495 **Table A1:** Database of earthquake focal mechanisms of southern Italy and surroundings reported in  
 496 Fig. 5. ID is the order number. O.T., Lon, Lat and Depth are the GMT origin time, the longitude E  
 497 (°), the latitude N (°) and the focal depth (km) of the earthquake, respectively. Strike, dip, and rake  
 498 are the fault parameters in degrees of the focal solution. M is the earthquake magnitude. Source is  
 499 the bibliographic source of the solution (Italian CMT= <http://rcmt2.bo.ingv.it/Italydataset.html>,  
 500 CMT= <http://www.globalcmt.org/>, RCMT= <http://rcmt2.bo.ingv.it/>, TDMT=  
 501 <http://cnt.rm.ingv.it/tdmt>; the other sources are reported in the reference list at the end of the  
 502 article).

503

<b>Id</b>	<b>Data</b>	<b>O.T.</b>	<b>Lon</b>	<b>Lat</b>	<b>Depth</b>	<b>Strike</b>	<b>Dip</b>	<b>Rake</b>	<b>M</b>	<b>Source</b>
1	19770605	13:59:23	14.46	37.84	11.3	61	26	-139	4.6	Italian CMT
2	19770815	21:10:40	16.98	38.85	40.0	307	38	120	5.2	CMT
3	19780311	19:20:49	16.03	38.10	15.0	270	41	-72	5.2	Italian CMT
4	19780415	23:33:47	14.63	37.77	34.0	135	60	-176	6.0	Italian CMT
5	19790120	13:49:59	12.86	38.67	9.0	72	29	53	5.2	Italian CMT
6	19791208	00:06:33	11.49	37.95	15.0	235	45	67	5.3	CMT
7	19800220	02:34:03	16.21	39.30	12.0	14	43	-78	4.8	Italian CMT
8	19800309	12:03:40	16.12	39.94	19.0	157	35	-80	4.6	Italian CMT
9	19800514	01:41:04	15.85	40.46	24.0	119	38	-112	4.5	Italian CMT
10	19800528	19:51:19	14.25	38.48	19.0	83	43	99	5.7	Italian CMT
11	19800601	02:32:52	14.33	38.39	10.0	65	39	91	4.9	Italian CMT
12	19801123	18:34:54	15.39	40.82	14.0	135	41	-80	6.9	Italian CMT
13	19801124	00:24:00	15.26	40.89	10.0	131	29	-110	4.9	Italian CMT
14	19801124	03:03:54	15.33	40.90	10.0	115	44	-125	5.1	Italian CMT
15	19801125	17:06:44	15.47	40.70	10.0	122	30	-119	5.1	Italian CMT
16	19801125	18:28:21	15.36	40.15	15.0	129	26	-65	5.4	Italian CMT
17	19801203	23:54:24	15.48	40.74	10.0	148	36	-76	4.9	Italian CMT
18	19810116	00:37:47	15.23	40.13	15.0	115	30	-93	5.2	Italian CMT
19	19810607	13:00:57	12.47	37.67	18.0	48	29	48	4.9	Italian CMT
20	19810622	09:36:18	14.09	38.49	13.0	71	47	116	4.8	Italian CMT
21	19811129	05:06:47	15.64	40.74	33.0	104	41	-138	4.9	Italian CMT
22	19820321	09:44:00	15.64	39.70	18.9	15	39	-127	5.0	Italian CMT
23	19820815	15:09:50	15.36	40.81	10.0	158	48	-45	4.8	Italian CMT
24	19821116	23:41:29	19.35	40.12	10.0	297	35	54	5.6	CMT
25	19870128	05:33:22	15.47	40.95	10.0	160	45	-79	4.6	Italian CMT
26	19870813	07:22:10	15.06	37.90	35.9	352	42	-10	4.8	Italian CMT
27	19880108	13:05:46	16.01	40.08	10.0	148	30	-86	4.8	Italian CMT
28	19880109	01:02:48	19.49	40.37	15.0	321	12	62	5.9	CMT



29	19880518	05:17:40	19.88	37.70	23.0	163	38	95	5.3	CMT
30	19890103	16:52:24	11.80	35.79	15.0	247	90	180	5.1	CMT
31	19890824	02:13:12	19.72	37.05	15.0	356	38	131	5.2	CMT
32	19900505	07:21:19	15.58	40.24	26.0	184	73	13	5.8	Italian CMT
33	19900505	07:38:12	15.81	40.75	15.0	282	83	173	5.0	Italian CMT
34	19901029	08:16:14	14.67	36.23	23.0	198	72	-13	4.5	Italian CMT
35	19901213	00:24:24	14.90	37.25	15.0	274	64	174	5.6	Italian CMT
36	19910526	12:26:00	15.77	40.73	8.0	183	71	-9	5.2	Italian CMT
37	19920123	04:24:20	19.97	38.22	15.0	351	42	97	5.6	CMT
38	19920406	13:08:34	14.61	37.83	21.0	100	37	-97	4.7	Italian CMT
39	19930626	17:47:54	14.21	37.92	10.0	170	53	6	4.4	Italian CMT
40	19940416	23:09:39	19.99	36.99	15.0	340	18	134	5.5	CMT
41	19950529	06:52:27	12.07	37.90	11.0	82	70	-180	4.8	Italian CMT
42	19960201	17:58:02	19.57	37.84	15.0	156	48	49	5.3	CMT
43	19960403	13:04:34	15.49	40.76	10.0	123	30	-110	4.9	Italian CMT
44	19961214	00:18:45	13.84	37.81	40.0	123	23	-43	4.7	Italian CMT
45	19970112	12:10:51	19.67	40.96	10.0	4	70	-167	4.9	RCMT
46	19970119	19:42:38	19.67	40.82	10.0	2	65	-177	4.7	RCMT
47	19970325	00:46:14	16.03	36.93	33.0	104	78	179	4.5	Italian CMT
48	19971202	19:22:47	19.40	36.68	15.0	206	49	-20	5.2	CMT
49	19980117	12:32:51	12.90	38.40	10.0	58	29	71	4.8	Italian CMT
50	19980620	02:25:47	13.08	38.46	10.0	69	22	76	5.2	Italian CMT
51	19980621	08:59:47	13.10	38.50	10.0	69	36	77	4.6	Italian CMT
52	19980621	12:59:04	12.67	38.43	10.0	88	38	102	4.6	Italian CMT
53	19980909	11:27:59	16.07	39.67	15.0	139	29	-83	5.6	Italian CMT
54	19980914	05:24:47	13.60	38.46	10.0	72	30	80	5.0	Italian CMT
55	19990214	11:45:54	15.06	38.17	33.0	18	39	-108	4.7	Italian CMT
56	19991030	07:09:09	19.98	37.76	5.0	80	8	-173	4.8	RCMT
57	19991124	21:10:49	19.76	40.27	10.0	141	39	101	4.3	RCMT
58	20000426	13:37:48	10.10	40.98	10.0	179	39	83	4.8	RCMT
59	20000428	17:41:02	19.70	37.83	33.0	305	64	-4	4.5	RCMT
60	20000627	04:07:56	10.03	40.95	10.0	184	27	82	4.3	RCMT
61	20010411	00:10:28	16.12	40.43	16.6	130	20	-90	3.0	Orecchio&al.(2014)
62	20010417	00:42:36	19.47	40.64	10.0	343	43	97	4.2	RCMT
63	20010422	13:56:36	15.10	37.72	10.0	316	56	27	4.2	Italian CMT
64	20010526	06:02:20	16.34	37.46	33.0	71	54	134	4.5	Italian CMT
65	20010930	23:44:58	16.04	40.22	12.6	30	70	-80	2.7	Orecchio&al.(2014)
66	20011018	11:02:44	16.61	39.10	10.0	332	44	-88	4.3	Italian CMT
67	20011125	19:34:20	13.96	37.91	20.0	137	31	-57	4.7	Italian CMT
68	20020405	04:52:24	14.74	38.48	10.0	90	41	108	4.4	Italian CMT
69	20020417	06:42:54	16.67	39.37	15.0	141	28	13	5.3	Totaro&al.(2016)
70	20020418	20:56:48	15.58	40.69	10.0	340	49	-52	4.4	Italian CMT
71	20020624	01:20:43	10.29	36.03	15.0	28	48	128	5.2	CMT
72	20020906	01:21:29	13.57	38.42	15.0	26	50	40	5.9	Italian CMT
73	20020906	01:45:30	13.73	38.44	4.0	252	48	126	4.7	Italian CMT
74	20020909	06:06:37	19.98	37.63	33.0	12	61	-165	4.4	RCMT



75	20020910	02:32:51	13.70	38.47	5.0	71	29	126	4.4	Italian CMT
76	20020920	23:06:04	13.74	38.46	5.0	46	33	77	4.7	Italian CMT
77	20020927	06:10:45	13.66	38.41	15.0	35	24	65	5.2	Totaro&al.(2016)
78	20020928	02:46:46	13.71	38.47	5.0	79	39	103	4.6	Italian CMT
79	20021002	22:57:26	13.72	38.46	5.0	33	41	59	4.9	Italian CMT
80	20021027	02:50:26	15.16	37.79	10.0	320	60	171	4.9	Italian CMT
81	20021027	07:32:09	15.18	37.92	10.0	67	54	19	4.5	Italian CMT
82	20021029	10:02:22	15.27	37.67	10.0	316	61	-173	4.7	Italian CMT
83	20021029	16:39:48	15.56	37.69	10.0	207	54	-28	4.2	Italian CMT
84	20021209	09:35:06	19.97	37.87	15.0	255	28	-20	5.2	CMT
85	20030616	08:27:51	19.73	37.69	10.0	321	31	43	4.7	RCMT
86	20030707	15:08:12	14.90	36.01	10.0	350	62	4	4.3	Italian CMT
87	20041011	07:31:41	15.48	37.88	6.6	89	90	-45	3.6	D'Amico&al.(2010)
88	20041022	21:10:13	15.32	38.08	10.7	78	61	-37	3.4	D'Amico&al.(2010)
89	20041218	09:12:48	10.15	40.89	10.0	144	49	51	4.6	RCMT
90	20050131	01:05:35	19.97	37.37	12.0	346	16	120	5.7	CMT
91	20050131	10:44:50	16.86	39.66	30.0	23	79	-41	4.1	D'Amico&al.(2011)
92	20050207	20:05:37	10.91	36.10	10.0	67	44	128	4.8	RCMT
93	20050207	20:46:26	10.87	36.22	10.0	51	41	124	5.1	RCMT
94	20050212	12:13:45	19.70	39.92	0.0	187	42	53	4.2	RCMT
95	20050419	22:36:23	15.66	38.14	7.1	220	42	-10	3.1	D'Amico&al.(2010)
96	20050423	19:10:48	15.82	38.43	13.6	120	50	-64	2.8	D'Amico&al.(2010)
97	20050423	19:11:43	16.71	39.47	23.0	128	58	14	4.1	D'Amico&al.(2011)
98	20050602	03:05:51	15.31	39.59	7.0	278	73	-172	3.7	Li&al.(2007)
99	20050721	15:41:43	14.85	39.40	7.0	184	68	41	3.8	Li&al.(2007)
100	20050818	22:02:27	15.12	37.80	6.7	82	50	-18	3.1	D'Amico&al.(2010)
101	20050907	12:40:33	16.32	38.71	16.0	80	90	-42	3.6	D'Amico&al.(2011)
102	20050927	22:33:09	17.10	38.62	29.0	38	79	141	3.9	Li&al.(2007)
103	20051030	19:09:47	15.93	38.53	22.0	241	66	-84	3.4	Li&al.(2007)
104	20051118	18:35:25	17.07	39.17	23.0	120	34	3	3.6	D'Amico&al.(2011)
105	20051121	10:57:41	14.16	37.61	61.0	102	79	179	4.6	RCMT
106	20051203	08:33:02	17.00	39.20	15.0	290	64	-18	3.8	Presti&al.(2013)
107	20060107	22:08:44	17.21	39.27	16.0	174	69	61	3.2	Orecchio&al.(2014)
108	20060117	03:33:58	17.13	39.20	34.0	146	62	-21	3.7	Orecchio&al.(2014)
109	20060227	04:34:01	15.20	38.15	9.0	62	50	-71	4.1	D'Amico&al.(2010)
110	20060227	09:11:59	15.18	38.14	10.5	39	48	-90	3.1	D'Amico&al.(2010)
111	20060227	14:16:06	15.18	38.14	9.1	76	48	-58	3.1	D'Amico&al.(2010)
112	20060329	20:20:00	13.89	37.73	10.0	338	80	-42	3.9	TDMT
113	20060417	02:44:06	17.14	39.57	10.0	114	74	-3	4.4	D'Amico&al.(2011)
114	20060423	14:42:38	15.02	37.04	24.0	100	88	147	3.9	TDMT
115	20060510	00:44:33	19.31	40.23	5.0	292	72	10	4.5	RCMT
116	20060510	21:33:06	19.78	40.06	18.0	289	59	7	4.3	RCMT
117	20060520	07:05:56	14.95	37.65	12.0	280	75	47	3.7	TDMT
118	20060525	23:14:41	19.91	36.55	23.0	346	23	129	5.2	CMT
119	20060530	11:30:40	16.52	37.63	46.0	347	85	0	4.5	RCMT
120	20060613	14:15:38	19.96	40.27	10.0	283	67	25	4.7	RCMT



121	20060619	20:55:35	14.89	37.83	16.0	350	28	11	3.8	Neri&al.(2014)
122	20060619	21:20:13	14.88	37.83	18.0	354	32	2	3.1	Totaro&al.(2016)
123	20060619	21:27:12	14.87	37.82	20.0	0	37	9	3.1	Totaro&al.(2016)
124	20060620	13:16:36	14.86	37.83	18.0	0	26	22	3.3	Totaro&al.(2016)
125	20060621	07:17:50	14.83	37.83	18.0	354	23	11	3.3	Totaro&al.(2016)
126	20060622	19:34:58	16.60	39.73	15.0	110	33	-33	4.4	D'Amico&al.(2011)
127	20060702	17:52:00	15.10	38.13	10.0	70	59	-49	2.6	D'Amico&al.(2010)
128	20060718	07:42:40	15.17	38.12	9.1	90	41	-48	3.1	D'Amico&al.(2010)
129	20060730	09:53:36	16.31	37.99	6.0	292	64	-7	2.7	Orecchio&al.(2014)
130	20060805	20:47:19	14.73	38.55	10.0	40	30	39	3.3	Totaro&al.(2016)
131	20060806	07:49:49	19.53	40.05	16.0	20	44	146	4.8	CMT
132	20060808	21:20:12	19.60	40.04	20.0	13	38	138	4.8	CMT
133	20060809	02:10:01	19.59	40.11	19.0	312	69	16	4.2	RCMT
134	20060819	16:29:11	14.43	38.58	26.0	229	78	28	3.3	Totaro&al.(2016)
135	20060830	22:45:03	15.72	37.32	30.0	190	64	-23	3.1	Orecchio&al.(2014)
136	20060901	21:50:41	19.87	40.05	50.0	129	64	63	4.4	present work
137	20060907	15:31:43	16.19	40.57	34.0	178	55	35	4.1	Italian CMT
138	20060909	15:45:23	14.23	38.70	8.0	222	68	13	3.3	Totaro&al.(2016)
139	20061006	21:16:23	15.57	38.10	9.6	18	52	-90	3.2	D'Amico&al.(2010)
140	20061022	05:13:10	16.70	39.05	10.0	312	38	-30	3.0	Orecchio&al.(2014)
141	20061104	05:59:22	15.01	38.03	10.6	59	49	-36	3.0	D'Amico&al.(2010)
142	20061107	11:13:36	18.26	35.91	35.0	201	30	93	4.3	RCMT
143	20061118	00:01:56	17.23	39.06	24.0	113	46	-51	3.0	Orecchio&al.(2014)
144	20061123	13:31:56	12.94	35.97	10.0	357	70	-2	4.8	RCMT
145	20061124	04:37:40	15.76	36.26	11.0	188	82	0	4.4	Italian CMT
146	20061219	14:58:06	14.91	37.78	23.0	18	16	-40	4.2	Italian CMT
147	20061220	11:38:08	14.26	38.54	6.0	201	64	13	3.6	Neri&al.(2014)
148	20061226	00:49:00	16.17	39.22	2.0	223	38	-12	3.1	Orecchio&al.(2014)
149	20070130	22:18:07	16.14	39.91	8.0	84	90	19	3.6	Orecchio&al.(2014)
150	20070202	06:51:01	16.35	39.55	12.0	31	48	-61	3.2	Totaro&al.(2013)
151	20070326	13:55:26	16.96	39.28	20.0	301	61	8	3.7	Totaro&al.(2016)
152	20070410	19:17:23	12.91	36.93	22.0	100	75	164	4.1	TDMT
153	20070421	19:41:27	13.43	38.57	12.0	250	52	10	3.9	Neri&al.(2014)
154	20070426	00:49:36	16.37	39.54	16.0	290	8	20	3.8	Orecchio&al.(2014)
155	20070503	18:43:56	17.63	39.02	18.0	92	25	-31	3.5	Orecchio&al.(2014)
156	20070517	05:48:13	14.69	38.57	8.0	22	50	8	3.5	Presti&al.(2013)
157	20070525	09:39:45	16.83	39.66	12.0	91	29	-48	4.2	D'Amico&al.(2011)
158	20070609	05:56:38	16.62	39.18	18.0	71	49	-59	3.1	Orecchio&al.(2014)
159	20070615	22:56:01	15.29	36.97	18.0	12	87	20	3.6	TDMT
160	20070617	12:11:58	15.79	38.37	10.0	262	38	-43	2.9	D'Amico&al.(2010)
161	20070706	23:28:43	17.25	39.18	28.0	118	38	-35	3.5	Orecchio&al.(2014)
162	20070714	18:13:03	14.75	38.63	4.0	30	31	38	3.1	Presti&al.(2013)
163	20070731	06:53:16	14.74	37.47	32.0	142	78	-21	3.4	Totaro&al.(2016)
164	20070801	00:07:54	17.18	39.00	18.0	80	67	-45	4.1	D'Amico&al.(2011)
165	20070818	14:04:07	15.17	38.22	12.0	44	50	-23	3.9	D'Amico&al.(2010)
166	20070818	14:21:11	15.12	38.19	10.0	26	69	18	3.4	D'Amico&al.(2010)



167	20070905	21:24:13	14.84	38.56	6.0	10	54	-2	3.3	Totaro&al.(2016)
168	20070913	15:19:52	15.16	38.25	8.0	246	82	-60	2.9	Totaro&al.(2016)
169	20070923	07:12:46	14.79	38.59	8.0	27	60	28	3.6	Presti&al.(2013)
170	20070930	15:41:20	14.80	38.59	6.0	70	73	16	3.1	Presti&al.(2013)
171	20071213	23:38:24	16.61	38.93	30.0	98	9	-17	3.2	Orecchio&al.(2014)
172	20071214	00:42:55	16.62	38.93	26.0	331	70	-60	3.0	Orecchio&al.(2014)
173	20071217	09:44:39	16.36	39.39	40.0	162	90	-71	3.5	Orecchio&al.(2014)
174	20071220	03:25:32	16.19	39.36	2.0	210	66	-71	3.5	Orecchio&al.(2014)
175	20080115	02:38:31	16.33	39.81	18.0	327	80	39	3.2	Orecchio&al.(2014)
176	20080118	13:01:00	16.53	39.14	12.0	57	78	-67	3.9	TDMT
177	20080209	07:46:36	15.56	37.84	6.9	40	90	-10	3.0	D'Amico&al.(2010)
178	20080221	05:00:09	17.97	37.82	30.0	333	27	134	4.7	Italian CMT
179	20080305	04:08:21	19.67	40.22	0.0	173	47	152	4.1	RCMT
180	20080310	10:33:27	16.85	39.66	20.0	121	39	-7	3.5	Presti&al.(2013)
181	20080408	17:20:01	16.66	39.16	10.0	235	49	-35	4.4	Italian CMT
182	20080413	10:10:02	16.52	39.16	14.0	205	69	-90	3.6	D'Amico&al.(2010)
183	20080413	13:06:57	15.70	38.25	14.3	6	47	-36	2.8	Orecchio&al.(2014)
184	20080414	18:44:34	16.52	39.15	12.0	48	41	-62	3.1	Orecchio&al.(2014)
185	20080419	21:41:11	17.47	39.13	16.0	107	42	-39	3.6	Presti&al.(2013)
186	20080426	22:23:06	16.53	39.14	18.0	256	60	-31	3.2	Orecchio&al.(2014)
187	20080501	21:05:49	15.07	37.80	2.0	97	76	-2	2.8	D'Amico&al.(2010)
188	20080513	21:28:30	15.06	37.80	12.0	76	46	-20	3.5	D'Amico&al.(2010)
189	20080702	17:43:33	16.23	38.97	30.0	266	69	-30	3.2	Orecchio&al.(2014)
190	20080703	20:56:52	13.71	38.45	24.2	182	68	27	3.3	TDMT
191	20080705	17:04:36	15.87	38.20	2.0	311	59	2	2.6	D'Amico&al.(2010)
192	20080709	23:08:27	16.23	38.97	24.0	268	76	-32	3.3	Orecchio&al.(2014)
193	20080710	01:45:45	16.24	38.97	12.0	76	72	-40	3.0	Orecchio&al.(2014)
194	20080710	12:50:20	16.24	38.96	14.0	71	75	-39	3.3	Orecchio&al.(2014)
195	20080711	07:15:00	16.24	38.96	16.0	50	52	-29	3.0	Orecchio&al.(2014)
196	20080711	07:20:21	16.25	38.96	12.0	78	80	-58	2.9	Orecchio&al.(2014)
197	20080813	13:39:30	16.42	37.48	34.0	181	71	11	3.2	Orecchio&al.(2014)
198	20080901	14:45:40	15.06	37.97	8.1	70	31	-80	3.1	D'Amico&al.(2010)
199	20080902	09:16:45	15.06	37.99	10.3	279	64	-44	3.3	D'Amico&al.(2010)
200	20080902	21:57:20	15.69	38.25	34.0	351	72	-65	3.1	Orecchio&al.(2014)
201	20080910	11:30:48	17.13	39.18	34.0	81	69	-90	4.6	Orecchio&al.(2014)
202	20080912	20:12:11	17.37	39.17	34.0	332	85	-2	3.6	Orecchio&al.(2014)
203	20080927	08:28:27	17.21	39.18	30.0	123	71	-8	4.0	Totaro & al. (2016)
204	20081024	16:55:37	16.44	38.61	30.0	0	60	70	3.3	Orecchio&al.(2014)
205	20081024	18:47:54	16.47	38.59	28.0	323	38	-24	3.4	Orecchio&al.(2014)
206	20081027	10:55:55	15.13	38.11	2.0	50	28	-71	3.5	D'Amico&al.(2010)
207	20081102	06:46:44	16.49	37.64	40.0	141	67	-79	3.6	Orecchio&al.(2014)
208	20081107	15:00:59	16.46	39.15	14.0	139	72	-65	3.3	Orecchio&al.(2014)
209	20081120	14:09:21	17.49	39.14	15.0	166	82	-2	4.5	Italian CMT
210	20081128	08:04:47	17.02	39.89	34.0	97	50	21	3.6	Orecchio&al.(2014)
211	20081128	23:39:21	13.69	37.54	35.0	337	74	9	4.4	Italian CMT
212	20081209	12:55:27	17.20	39.04	24.0	104	21	-31	3.6	Orecchio&al.(2014)



213	20081225	18:55:58	15.96	40.34	6.0	100	59	-23	2.6	Orecchio&al.(2014)
214	20090205	14:50:14	16.03	37.39	28.0	167	78	18	3.3	Orecchio&al.(2014)
215	20090316	00:28:06	15.96	37.67	28.0	34	60	-24	3.0	Orecchio&al.(2014)
216	20090319	08:27:54	12.72	36.52	28.0	255	48	-180	4.4	Italian CMT
217	20090325	12:23:25	19.63	40.34	10.0	3	35	90	4.3	RCMT
218	20090407	20:24:54	16.81	39.19	14.0	161	70	-33	3.2	Orecchio&al.(2014)
219	20090413	11:39:58	16.39	39.53	8.0	260	40	-7	3.3	Totaro&al.(2013)
220	20090427	09:42:16	15.08	38.07	30.0	69	78	-19	3.6	Presti&al.(2013)
221	20090701	17:58:54	15.01	38.34	2.0	40	90	19	3.1	Presti&al.(2013)
222	20090727	22:15:14	15.69	37.12	30.0	353	48	-13	3.2	Orecchio&al.(2014)
223	20090804	16:17:16	15.71	37.12	18.0	22	73	-13	3.6	Orecchio&al.(2014)
224	20090829	06:55:17	15.47	37.92	8.0	56	80	-47	2.9	Orecchio&al.(2014)
225	20090907	21:26:31	13.98	38.73	18.0	63	39	65	4.8	Totaro&al.(2016)
226	20090917	22:53:00	19.96	39.95	10.0	307	55	44	4.5	RCMT
227	20091012	20:07:49	15.96	37.23	30.0	204	82	12	3.4	Orecchio&al.(2014)
228	20091108	06:51:16	14.55	37.83	15.0	310	21	-54	4.5	RCMT
229	20091125	06:20:07	16.45	38.05	16.0	341	62	-43	3.2	Orecchio&al.(2014)
230	20091215	11:49:07	15.57	38.96	24.0	195	61	-54	3.7	Orecchio&al.(2014)
231	20091219	09:01:19	15.09	37.76	40.0	112	44	176	4.4	Italian CMT
232	20100101	22:01:13	16.29	39.20	36.0	267	58	-60	3.8	Orecchio&al.(2014)
233	20100208	07:23:58	16.77	39.50	34.0	94	73	-31	3.6	Presti&al.(2013)
234	20100317	11:01:11	14.73	38.57	8.0	78	56	81	3.3	Presti&al.(2013)
235	20100325	17:30:18	15.86	40.03	2.0	0	51	-67	3.2	Orecchio&al.(2014)
236	20100402	20:04:47	15.11	37.76	2.0	274	55	10	4.2	Italian CMT
237	20100404	15:40:28	16.82	39.35	22.0	314	76	-18	3.3	Presti&al.(2013)
238	20100413	12:12:14	17.15	39.35	18.0	128	29	-26	3.5	Presti&al.(2013)
239	20100415	20:05:47	17.22	39.35	18.0	137	39	-15	3.6	Presti&al.(2013)
240	20100511	10:28:47	16.22	39.75	6.0	152	56	-90	2.8	Orecchio&al.(2014)
241	20100511	18:09:43	17.47	39.31	22.0	126	40	-28	3.8	Orecchio&al.(2014)
242	20100606	16:49:53	15.11	38.27	10.0	237	82	-34	3.5	Presti&al.(2013)
243	20100616	22:39:41	16.14	38.83	15.0	109	50	-38	4.1	Italian CMT
244	20100801	21:31:53	14.46	38.61	4.0	37	60	78	3.1	Presti&al.(2013)
245	20100816	12:54:46	14.92	38.42	10.0	218	66	42	4.5	Presti&al.(2013)
246	20100822	10:23:05	19.95	37.27	17.0	325	16	99	5.5	CMT
247	20100910	19:19:48	16.21	38.54	26.0	198	53	-60	3.3	Orecchio&al.(2014)
248	20100910	21:39:20	15.82	38.20	28.0	204	69	-70	3.2	Orecchio&al.(2014)
249	20101008	17:26:58	16.33	36.91	38.0	190	79	17	3.6	Orecchio&al.(2014)
250	20101014	14:18:28	16.69	38.84	28.0	60	60	31	3.2	Orecchio&al.(2014)
251	20101015	05:21:20	16.66	38.87	15.0	287	62	173	4.4	Italian CMT
252	20101109	08:43:20	15.93	40.05	10.0	329	61	-57	3.5	Orecchio&al.(2014)
253	20101127	08:45:49	15.64	38.08	38.0	332	22	-12	3.7	Orecchio&al.(2014)
254	20110325	16:18:12	16.94	38.87	6.0	87	70	-53	3.3	Orecchio&al.(2014)
255	20110325	18:31:31	16.96	38.87	6.0	281	53	-19	3.6	Orecchio&al.(2014)
256	20110424	13:02:12	14.88	35.82	20.3	28	34	-76	4.2	RCMT
257	20110426	21:02:30	15.16	38.15	2.0	33	40	-90	3.2	Totaro&al.(2016)
258	20110503	22:24:52	16.68	37.78	36.0	323	49	-41	3.6	Orecchio&al.(2014)



259	20110506	15:12:35	14.96	37.78	22.2	13	57	15	4.0	Italian CMT
260	20110609	16:16:36	19.75	40.69	10.0	171	78	-170	4.4	RCMT
261	20110623	22:02:47	14.78	38.06	10.0	315	90	-1	4.7	Neri&al.(2014)
262	20110624	09:00:08	16.61	39.61	22.0	330	21	-20	3.1	Totaro&al.(2013)
263	20110627	05:23:41	14.74	38.02	8.0	313	63	-2	3.2	Totaro&al.(2016)
264	20110627	22:13:45	14.74	38.04	10.0	299	44	-31	3.4	Totaro&al.(2016)
265	20110629	09:04:17	14.73	38.05	8.0	305	90	2	3.2	Totaro&al.(2016)
266	20110629	19:15:15	14.74	38.06	10.0	139	81	-10	3.5	Totaro&al.(2016)
267	20110706	09:08:39	14.78	38.05	10.0	124	42	-41	3.7	Neri&al.(2014)
268	20110707	01:01:15	14.79	38.04	10.0	315	90	10	3.3	Totaro&al.(2016)
269	20110727	04:03:14	14.77	38.06	8.0	113	46	-61	3.2	Totaro&al.(2016)
270	20110817	01:20:32	14.47	38.55	6.0	209	70	38	3.0	Totaro&al.(2016)
271	20110831	16:33:20	14.70	37.11	2.0	7	80	1	3.1	Totaro&al.(2016)
272	20111103	14:37:10	14.58	38.43	10.0	8	90	22	3.5	Totaro&al.(2016)
273	20111109	17:00:48	16.02	39.91	9.0	10	50	-51	2.7	Totaro&al.(2015)
274	20111115	04:59:00	14.67	38.27	10.0	18	74	9	4.1	Neri&al.(2014)
275	20111119	10:19:16	14.35	37.81	14.0	121	70	-25	3.4	Totaro&al.(2016)
276	20111123	14:12:34	16.01	39.92	9.0	7	40	-48	3.5	Totaro&al.(2015)
277	20111201	14:01:20	16.01	39.93	9.8	7	48	-52	3.3	Totaro&al.(2015)
278	20111202	21:25:38	16.01	39.92	9.8	156	49	-90	3.2	Totaro&al.(2015)
279	20111214	17:59:49	16.20	39.38	6.0	178	39	-43	3.1	Orecchio&al.(2014)
280	20111217	23:20:15	16.18	39.37	26.0	345	62	-79	3.6	Orecchio&al.(2014)
281	20111218	15:01:06	12.68	36.10	12.0	183	85	2	4.7	CMT
282	20111224	20:17:50	16.02	39.92	9.1	348	44	-90	3.2	Totaro&al.(2015)
283	20111227	01:07:45	16.92	39.58	20.0	121	28	-26	3.6	Totaro&al.(2016)
284	20120103	23:47:41	16.84	39.64	16.0	156	29	-29	3.1	Orecchio&al.(2014)
285	20120129	11:14:50	14.27	37.88	10.0	37	61	-23	3.0	Totaro&al.(2016)
286	20120201	14:28:38	14.28	37.89	30.0	119	90	14	3.6	Neri&al.(2014)
287	20120208	16:15:56	14.30	37.89	12.0	200	50	-51	3.1	Totaro&al.(2016)
288	20120225	20:34:35	13.55	38.54	12.0	20	27	19	4.3	Neri&al.(2014)
289	20120226	16:17:23	16.01	37.31	36.0	338	70	-40	3.7	Polonia&al.(2016)
290	20120324	20:34:59	15.88	37.59	32.0	158	84	-9	3.1	Polonia&al.(2016)
291	20120405	03:01:06	16.42	39.56	18.0	275	42	-53	3.0	Totaro&al.(2013)
292	20120409	23:29:02	15.46	40.41	14.0	165	58	-77	3.1	Totaro&al.(2013)
293	20120412	13:20:28	15.62	37.89	10.0	319	90	81	3.1	Polonia&al.(2016)
294	20120413	06:21:33	13.30	38.35	12.0	319	29	-71	4.4	Neri&al.(2014)
295	20120528	01:06:27	16.12	39.89	7.0	146	49	-90	4.3	Totaro&al.(2015)
296	20120528	01:32:10	16.10	39.90	9.3	236	51	-27	3.1	Totaro&al.(2015)
297	20120531	03:16:22	15.57	39.89	8.0	211	45	2	3.0	Orecchio&al.(2014)
298	20120531	20:18:23	16.93	38.96	4.0	258	82	-74	2.9	Totaro&al.(2016)
299	20120615	06:27:25	16.29	37.45	38.0	190	80	7	3.8	Polonia&al.(2016)
300	20120618	03:17:12	16.11	39.90	9.5	162	29	-90	2.7	Totaro&al.(2015)
301	20120625	10:52:51	15.05	37.01	18.0	164	77	-16	3.2	Totaro&al.(2016)
302	20120627	01:14:20	15.03	37.00	4.0	182	90	3	3.5	Totaro&al.(2016)
303	20120627	01:20:59	15.03	36.99	4.0	10	81	-3	2.9	Totaro&al.(2016)
304	20120627	02:48:02	15.03	37.00	4.0	172	90	-6	2.9	Totaro&al.(2016)



305	20120704	11:12:12	16.87	37.69	40.0	186	74	3	4.6	RCMT
306	20120715	11:51:55	16.91	39.65	18.0	118	27	-17	3.0	Orecchio&al.(2014)
307	20120726	14:20:03	16.34	37.90	16.0	134	83	-19	3.1	Polonia&al.(2016)
308	20120813	07:30:51	13.73	38.52	26.5	19	24	63	4.2	RCMT
309	20120819	17:45:08	16.02	39.89	8.8	154	46	-90	3.5	Totaro&al.(2015)
310	20120819	21:28:29	16.03	39.89	8.3	341	37	-62	2.7	Totaro&al.(2015)
311	20120828	23:12:15	15.71	38.25	45.4	130	10	-18	4.6	present work
312	20120901	14:02:45	16.03	39.89	8.4	178	60	-69	3.4	Totaro&al.(2015)
313	20120904	03:48:03	16.03	39.90	8.3	161	55	-81	2.8	Totaro&al.(2015)
314	20120907	12:40:51	16.03	39.89	7.9	177	52	-70	3.3	Totaro&al.(2015)
315	20120907	15:10:07	16.02	39.89	6.0	176	61	-62	2.8	Totaro&al.(2016)
316	20120914	03:50:11	16.03	39.90	7.9	156	57	-90	3.6	Totaro&al.(2015)
317	20120922	01:45:02	16.02	39.91	8.0	139	58	-84	2.7	Totaro&al.(2015)
318	20120922	05:10:35	16.61	39.78	14.0	128	90	71	3.5	Totaro&al.(2016)
319	20120923	06:13:56	16.02	39.91	8.4	331	32	-90	2.7	Totaro&al.(2015)
320	20120924	20:48:36	16.03	39.92	6.4	231	59	-42	2.7	Totaro&al.(2015)
321	20120928	05:56:46	16.10	39.91	7.3	22	41	-80	2.8	Totaro&al.(2015)
322	20121001	20:28:28	16.03	39.91	7.7	343	39	-82	3.5	Totaro&al.(2015)
323	20121001	21:27:51	16.02	39.91	7.8	13	40	-43	3.3	Totaro&al.(2015)
324	20121002	00:08:57	16.03	39.91	8.6	331	40	-80	3.3	Totaro&al.(2015)
325	20121002	04:35:18	16.03	39.91	8.2	140	58	-78	2.8	Totaro&al.(2015)
326	20121004	09:32:33	16.02	39.90	8.0	159	52	-84	2.9	Totaro&al.(2015)
327	20121005	11:12:28	16.03	39.90	7.6	0	40	-73	3.0	Totaro&al.(2015)
328	20121014	14:49:24	16.02	39.91	8.7	20	42	-40	2.7	Totaro&al.(2015)
329	20121018	02:51:57	16.03	39.90	7.8	350	34	-90	3.3	Totaro&al.(2015)
330	20121023	10:40:24	16.03	39.91	8.4	324	30	-82	3.1	Totaro&al.(2015)
331	20121025	23:05:25	16.03	39.89	8.8	166	50	-77	5.0	Totaro&al.(2015)
332	20121026	00:31:53	16.00	39.89	10.0	349	29	-49	3.0	Totaro&al.(2016)
333	20121026	02:25:09	16.03	39.92	6.6	352	40	-81	2.9	Totaro&al.(2015)
334	20121026	02:40:08	16.02	39.88	8.1	73	50	-50	2.8	Totaro&al.(2015)
335	20121026	16:08:58	16.03	39.89	8.9	11	56	-23	2.7	Totaro&al.(2015)
336	20121028	13:52:18	16.02	39.93	8.5	12	75	27	3.1	Totaro&al.(2015)
337	20121102	01:59:34	16.47	38.78	14.0	69	22	-50	3.1	Orecchio&al.(2014)
338	20121102	17:50:44	16.03	39.91	7.9	244	58	-66	3.0	Totaro&al.(2015)
339	20121102	17:58:47	16.03	39.92	7.8	39	52	-56	2.7	Totaro&al.(2015)
340	20121105	12:06:32	16.01	39.94	8.8	21	71	-11	3.4	Totaro&al.(2015)
341	20121108	11:11:57	16.10	39.91	8.4	158	16	-79	3.1	Totaro&al.(2015)
342	20121112	03:03:53	16.01	39.92	8.7	229	42	20	3.0	Totaro&al.(2015)
343	20121121	06:43:25	16.02	39.92	8.2	63	60	-39	2.9	Totaro&al.(2015)
344	20121122	01:59:52	16.02	39.92	9.0	0	41	-78	3.2	Totaro&al.(2015)
345	20121122	09:10:41	14.96	37.80	10.0	258	65	154	4.1	RCMT
346	20121122	11:25:52	14.99	37.77	20.0	6	57	23	4.2	RCMT
347	20121124	22:24:26	16.02	39.92	7.9	0	47	-79	2.8	Totaro&al.(2015)
348	20121125	08:28:39	16.02	39.92	9.4	360	42	-72	3.5	Totaro&al.(2015)
349	20121125	08:42:25	16.03	39.93	5.5	0	41	-74	2.9	Totaro&al.(2015)
350	20121125	08:53:33	16.01	39.89	7.6	168	45	-90	3.0	Totaro&al.(2015)



351	20121125	17:48:02	16.01	39.92	9.5	7	43	-69	3.1	Totaro&al.(2015)
352	20121128	02:43:46	16.01	39.92	9.1	45	86	2	2.9	Totaro&al.(2015)
353	20121211	14:28:43	16.01	39.88	9.4	339	29	-70	3.3	Totaro&al.(2015)
354	20121213	04:44:03	16.03	39.88	8.8	20	70	-31	3.2	Totaro&al.(2015)
355	20121218	11:03:18	16.17	39.84	2.0	130	49	-90	3.3	Totaro&al.(2016)
356	20121218	11:05:43	16.17	39.84	4.0	158	62	-30	3.0	Totaro&al.(2016)
357	20121226	08:22:48	16.31	39.50	8.0	41	29	-41	3.2	Totaro&al.(2016)
358	20130104	07:50:06	14.72	37.88	12.0	318	43	41	4.4	Neri&al.(2014)
359	20130104	10:50:21	14.70	37.88	31.0	301	81	-14	3.2	Totaro&al.(2016)
360	20130106	07:50:19	14.72	37.87	18.0	118	63	-22	3.2	Totaro&al.(2016)
361	20130109	16:10:34	14.72	37.88	6.0	312	78	-11	3.0	Totaro&al.(2016)
362	20130205	22:08:04	15.86	40.07	18.0	176	71	-66	3.1	Totaro&al.(2016)
363	20130223	19:14:18	14.98	38.31	4.0	220	79	-7	3.4	Totaro&al.(2016)
364	20130303	23:39:13	15.83	38.13	8.0	237	57	-82	3.3	Totaro&al.(2016)
365	20130307	22:36:59	14.52	37.97	2.0	0	50	-31	3.6	Neri&al.(2014)
366	20130317	14:22:15	15.89	39.62	32.0	31	79	48	3.3	Totaro&al.(2016)
367	20130319	07:50:06	14.51	37.98	4.0	164	78	-43	3.4	Totaro&al.(2016)
368	20130319	08:37:04	14.51	37.98	4.0	159	69	-39	3.3	Totaro&al.(2016)
369	20130319	08:38:45	13.56	37.84	2.0	183	69	-8	2.8	Totaro&al.(2016)
370	20130324	15:47:22	16.50	37.76	30.0	257	87	178	4.6	RCMT
371	20130401	03:07:13	16.67	39.68	28.0	127	68	90	3.4	Totaro&al.(2016)
372	20130402	01:10:52	15.59	37.79	12.0	219	85	-10	2.9	Totaro&al.(2016)
373	20130406	04:14:11	15.21	40.41	2.0	249	47	22	2.9	Totaro&al.(2016)
374	20130412	17:50:00	14.92	38.16	16.0	0	67	15	3.1	Totaro&al.(2016)
375	20130509	20:41:22	16.05	39.18	31.0	344	82	-110	3.8	TDMT
376	20130622	08:41:00	19.58	40.21	10.0	324	21	55	4.4	RCMT
377	20130704	13:56:06	15.61	40.49	4.0	211	41	-42	2.9	Totaro&al.(2016)
378	20130717	04:26:36	15.83	40.02	10.0	149	90	-72	2.9	Totaro&al.(2016)
379	20130804	02:47:47	12.85	38.69	4.0	12	72	-11	3.3	Totaro&al.(2016)
380	20130815	23:04:58	14.91	38.14	12.0	78	82	57	4.5	Neri&al.(2014)
381	20130815	23:06:51	14.92	38.15	12.0	77	82	55	4.6	Neri&al.(2014)
382	20130819	05:48:23	14.26	37.70	20.0	23	59	-14	3.0	Totaro&al.(2016)
383	20130828	09:07:00	14.31	38.85	6.0	210	72	12	3.6	Totaro&al.(2016)
384	20130917	22:56:38	15.80	40.79	8.0	339	69	1	3.3	Totaro&al.(2016)
385	20130917	23:38:49	15.81	40.79	10.0	133	72	-83	3.4	Totaro&al.(2016)
386	20130921	13:18:02	15.78	40.79	16.0	342	57	-34	3.2	Totaro&al.(2016)
387	20130926	12:19:59	17.23	39.13	22.0	188	12	52	3.1	Totaro&al.(2016)
388	20131003	18:22:25	13.49	38.47	14.0	20	51	0	3.4	Totaro&al.(2016)
389	20131007	04:44:05	15.08	38.13	6.0	84	63	-29	2.8	Totaro&al.(2016)
390	20131008	10:33:20	15.82	40.02	2.0	55	60	-10	2.7	Totaro&al.(2016)
391	20131009	08:14:49	15.09	37.61	4.0	33	81	8	2.9	Totaro&al.(2016)
392	20131009	08:33:22	15.07	37.61	10.0	258	90	1	3.2	Totaro&al.(2016)
393	20131018	08:03:44	14.90	38.11	14.0	258	79	-61	3.0	Totaro&al.(2016)
394	20131018	11:05:21	14.97	36.79	2.0	173	82	3	2.9	Totaro&al.(2016)
395	20131018	15:08:31	10.83	35.70	12.0	99	51	-179	4.8	CMT
396	20131018	20:50:52	15.22	40.80	14.0	319	41	-76	3.1	Totaro&al.(2016)



397	20131019	11:08:03	15.18	40.59	14.0	112	30	58	3.0	Totaro&al.(2016)
398	20131021	19:37:00	10.93	35.60	10.0	89	71	165	4.5	RCMT
399	20131103	14:33:41	15.18	40.60	2.0	87	58	-80	2.7	Totaro&al.(2016)
400	20131105	05:06:39	14.91	37.69	22.0	84	57	-11	3.1	Totaro&al.(2016)
401	20131105	05:29:58	14.92	37.70	18.0	99	72	-14	3.0	Totaro&al.(2016)
402	20131105	17:25:23	16.01	39.88	8.0	80	51	-45	3.0	Totaro&al.(2016)
403	20131105	17:26:45	16.01	39.88	10.0	328	32	-79	3.4	Totaro&al.(2016)
404	20131210	20:39:39	15.45	39.72	22.0	172	19	69	3.1	Totaro&al.(2016)
405	20131214	21:49:05	14.74	37.76	30.0	53	67	-27	3.4	Totaro&al.(2016)
406	20131215	03:57:33	14.94	36.67	15.0	83	47	143	4.1	RCMT
407	20131223	04:20:39	15.57	38.22	2.0	31	61	-60	3.5	Neri&al.(2014)
408	20131223	16:17:11	15.04	38.18	16.0	113	63	-33	3.1	Totaro&al.(2016)
409	20140102	06:13:18	15.04	38.18	12.0	103	38	8	3.0	Totaro&al.(2016)
410	20140114	03:43:42	14.92	38.37	12.0	315	78	171	4.1	RCMT
411	20140114	04:35:00	14.94	38.36	11.0	308	47	159	4.2	RCMT
412	20140122	19:35:01	15.12	40.44	8.0	61	32	-61	3.6	Totaro&al.(2016)
413	20140127	21:39:32	15.93	40.12	10.0	38	58	-19	3.0	Totaro&al.(2016)
414	20140218	21:44:19	13.82	37.63	4.0	250	79	33	3.3	Totaro&al.(2016)
415	20140219	06:58:05	15.11	38.17	8.0	34	62	17	3.0	Totaro&al.(2016)
416	20140301	01:48:50	15.22	40.24	4.0	19	27	40	2.9	Totaro&al.(2016)
417	20140301	11:51:25	15.29	40.25	6.0	305	68	8	2.7	Totaro&al.(2016)
418	20140303	06:26:04	15.01	40.40	8.0	287	47	-64	3.0	Totaro&al.(2016)
419	20140308	14:19:05	15.81	39.77	16.0	280	79	46	3.0	Totaro&al.(2016)
420	20140308	20:52:51	14.89	37.96	32.0	13	47	-40	3.8	Neri&al.(2014)
421	20140314	03:32:24	14.81	38.22	2.0	19	90	-1	3.0	Totaro&al.(2016)
422	20140314	03:37:16	14.80	38.20	6.0	331	62	4	2.8	Totaro&al.(2016)
423	20140323	18:31:52	16.48	37.47	38.0	177	61	21	3.6	Totaro&al.(2016)
424	20140325	07:56:52	16.54	39.29	28.0	276	59	-50	3.6	Totaro&al.(2016)
425	20140405	10:24:46	17.18	38.82	64.0	92	59	-168	4.8	RCMT
426	20140412	06:53:05	13.88	37.82	2.0	360	81	-1	3.1	Totaro&al.(2016)
427	20140417	21:52:26	15.21	38.23	2.0	12	90	2	2.8	Totaro&al.(2016)
428	20140503	20:04:06	16.11	39.90	4.0	232	31	-70	3.0	Totaro&al.(2016)
429	20140506	08:24:42	16.07	39.91	4.0	241	51	-90	3.0	Totaro&al.(2016)
430	20140507	06:20:00	19.78	37.75	10.0	181	68	16	4.3	RCMT
431	20140507	17:08:35	16.02	39.89	8.0	160	49	-58	2.9	Totaro&al.(2016)
432	20140517	22:38:44	15.61	37.41	30.0	191	78	-38	3.1	Totaro&al.(2016)
433	20140519	00:59:23	19.94	40.91	20.0	59	49	11	5.1	CMT
434	20140520	04:43:26	19.86	40.92	2.0	63	74	8	4.4	RCMT
435	20140604	21:20:41	16.00	39.88	8.0	159	58	-90	3.6	Totaro&al.(2016)
436	20140606	13:41:38	16.09	39.90	4.0	155	31	-90	3.9	Totaro&al.(2016)
437	20140607	15:00:49	15.10	38.09	10.0	285	73	7	3.7	Neri&al.(2014)
438	20140607	15:13:20	15.10	38.08	10.0	238	39	-49	2.6	Totaro&al.(2016)
439	20140617	12:25:02	16.98	38.96	2.0	239	40	-90	2.7	Totaro&al.(2016)
440	20140627	02:56:49	14.64	37.83	14.0	360	37	-26	3.3	Totaro&al.(2016)
441	20140629	04:24:28	15.60	39.92	14.0	248	22	1	3.8	Totaro&al.(2016)
442	20140629	05:32:22	15.60	39.91	6.0	47	65	-47	2.6	Totaro&al.(2016)



443	20140629	05:42:56	15.59	39.91	14.0	240	21	-21	3.3	Totaro&al.(2016)
444	20140702	18:49:46	15.17	40.59	2.0	71	49	-70	3.1	Totaro&al.(2016)
445	20140707	17:44:15	15.50	40.72	12.0	32	64	30	3.3	Totaro&al.(2016)
446	20140708	05:02:43	16.12	39.90	2.0	347	51	-83	2.7	Totaro&al.(2016)
447	20140711	21:57:58	15.39	40.49	10.0	344	82	-49	3.5	Totaro&al.(2016)
448	20140724	01:12:51	15.02	38.09	30.0	100	90	-16	3.2	Totaro&al.(2016)
449	20140731	03:29:29	16.11	39.91	4.0	349	58	-82	3.4	Totaro&al.(2016)
450	20140806	08:16:21	15.83	40.58	8.0	323	83	-9	3.0	Totaro&al.(2016)
451	20140812	20:15:34	16.42	40.45	16.0	330	44	-54	3.9	RCMT
452	20140813	10:08:09	16.43	40.44	4.0	160	63	-25	3.0	Totaro&al.(2016)
453	20140816	05:00:12	13.42	38.53	10.0	161	86	59	3.4	TDMT
454	20140826	01:19:46	14.33	37.95	14.0	298	20	-90	3.5	Totaro&al.(2016)
455	20140905	10:10:19	18.76	39.19	10.0	334	39	-56	4.1	RCMT
456	20140907	09:56:25	19.88	37.61	10.0	329	33	104	4.6	RCMT
457	20140919	05:32:38	14.81	38.49	14.0	66	78	25	3.2	Totaro&al.(2016)
458	20140924	15:39:09	15.85	39.74	24.0	51	80	-39	3.2	Totaro&al.(2016)
459	20140926	23:38:11	16.50	36.78	40.0	267	75	170	4.2	RCMT
460	20141009	22:58:28	14.85	38.51	10.0	76	40	80	4.1	RCMT
461	20141010	16:16:18	15.14	38.09	34.0	60	90	-76	3.1	Totaro&al.(2016)
462	20141010	16:27:13	15.14	38.09	30.0	349	72	-31	2.8	Totaro&al.(2016)
463	20141013	03:34:47	16.46	39.37	10.0	18	77	-73	2.8	Totaro&al.(2016)
464	20141014	00:50:55	16.63	38.95	14.0	150	74	-1	2.9	Totaro&al.(2016)
465	20141025	20:09:48	15.95	38.17	15.0	275	68	-79	3.3	TDMT
466	20141116	12:38:42	15.08	38.24	6.0	276	84	1	2.8	Totaro&al.(2016)
467	20141127	09:09:44	16.49	38.86	20.0	28	39	-81	3.2	Totaro&al.(2016)
468	20141228	21:43:38	16.36	39.29	10.0	179	49	-69	4.2	Totaro&al.(2016)
469	20150110	03:40:38	13.79	38.07	4.0	327	81	1	2.9	Totaro&al.(2016)
470	20150120	07:17:22	15.57	38.11	18.0	199	38	32	3.3	Totaro&al.(2016)
471	20150208	19:39:22	15.22	37.35	22.0	87	52	-50	3.1	Totaro&al.(2016)
472	20150211	01:42:09	14.75	38.05	4.0	127	37	-57	3.0	Totaro&al.(2016)
473	20150211	03:57:01	14.74	38.05	4.0	272	70	-37	3.1	Totaro&al.(2016)
474	20150312	15:29:04	16.23	38.43	2.0	230	59	-78	3.3	Totaro&al.(2016)
475	20150328	22:07:51	16.21	38.09	20.0	233	79	79	3.1	Totaro&al.(2016)
476	20150329	10:48:46	16.21	38.09	12.0	52	76	-83	3.5	Totaro&al.(2016)
477	20150420	01:07:43	15.12	37.80	3.0	178	83	-164	3.5	TDMT
478	20150430	05:35:21	15.39	37.86	26.0	160	61	44	2.9	Totaro&al.(2016)
479	20150511	08:26:32	16.80	37.33	40.0	184	62	20	4.5	RCMT
480	20150511	08:26:30	16.79	37.18	44.0	187	69	-9	4.2	present work
481	20150521	15:31:18	19.85	37.68	15.0	207	72	177	4.4	RCMT
482	20150521	22:13:22	14.72	38.45	2.0	281	53	-10	2.9	Totaro&al.(2016)
483	20150522	06:31:16	19.84	37.67	10.0	206	80	178	4.2	RCMT
484	20150524	06:00:00	16.03	37.96	62.0	23	62	-3	4.1	present work
485	20150611	04:31:45	14.01	37.88	8.0	148	60	-54	2.7	Totaro&al.(2016)
486	20150617	09:44:07	14.15	37.66	24.0	332	82	8	3.4	Totaro&al.(2016)
487	20150703	01:07:24	16.52	39.92	16.0	188	60	62	3.2	Totaro&al.(2016)
488	20150715	04:19:11	14.43	37.22	18.0	164	79	11	3.1	Totaro&al.(2016)



489	20150715	16:29:49	15.05	37.62	10.0	214	90	-11	3.0	Totaro&al.(2016)
490	20150726	13:39:38	14.76	38.50	6.0	204	84	48	2.9	Totaro&al.(2016)
491	20150801	02:46:51	15.86	37.64	42.0	195	90	9	3.4	Totaro&al.(2016)
492	20150803	07:27:49	16.49	39.15	16.0	242	19	-33	3.9	Totaro&al.(2016)
493	20150803	13:52:37	16.14	37.39	38.0	205	59	1	3.6	Totaro&al.(2016)
494	20150804	23:36:31	14.15	37.64	10.0	168	84	14	3.2	Totaro&al.(2016)
495	20150806	01:59:43	15.19	38.24	8.0	214	76	-61	3.1	Totaro&al.(2016)
496	20150808	22:46:24	14.27	38.55	8.0	173	90	14	3.8	Totaro&al.(2016)
497	20150818	05:59:15	15.45	40.64	12.0	326	70	-45	3.0	Totaro&al.(2016)
498	20150826	04:28:36	16.91	38.79	24.0	10	90	-33	3.3	Totaro&al.(2016)
499	20150829	20:25:13	12.12	38.54	14.0	332	81	8	4.0	Totaro&al.(2016)
500	20150920	22:27:58	15.61	37.16	30.0	226	58	-2	3.8	Totaro&al.(2016)
501	20151220	09:46:00	13.58	38.35	10.0	24	20	59	4.4	RCMT
502	20151229	14:15:53	19.89	37.42	10.0	107	58	-14	4.1	RCMT
503	20160102	12:36:24	12.03	36.46	10.0	268	79	-178	4.2	RCMT
504	20160113	17:01:30	14.67	36.14	10.0	262	79	175	4.2	RCMT
505	20160208	15:35:43	14.90	36.99	10.0	280	42	178	4.5	RCMT
506	20160306	08:12:36	16.74	38.20	39.0	248	75	158	4.0	TDMT
507	20160329	01:05:33	19.96	37.25	17.0	338	19	121	5.4	CMT

504

505

506



507 *Data availability.* Data used in the present study were collected from the databases of Istituto  
508 Nazionale di Geofisica e Vulcanologia ([www.ingv.it](http://www.ingv.it)) and from catalogs and bibliographic sources  
509 indicated in detail in the article.

510

511 *Author contributions.* All authors contributed to the scientific content, interpretations, message of  
512 the paper, and the discussions. CT primarily worked on stress inversion of earthquake fault –plane  
513 solutions, BO and DP on hypocenter locations in 3D velocity structures with the different  
514 algorithms, GN linked the different contributions and produced the final paper with the support of  
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516

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518

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522

## 523 **References**

524

525 Anderson, H.: Is the Adriatic an African promontory?, *Geology*, 15(3), 212-215, 1987.

526

527 Anderson, H., and Jackson, J.: Active tectonics of the Adriatic region, *Geophys. J. Int.*, 91(3), 937-  
528 983, 1987.

529

530 Argnani, A.: Evolution of the southern Tyrrhenian slab tear and active tectonics along the western  
531 edge of the Tyrrhenian subducted slab, Geological Society, London, Special Publications, 311(1),  
532 193-212, 2009.



533

534 Battaglia, M., Murray, M. H., Serpelloni, E., and Bürgmann, R.: The Adriatic region: An  
535 independent microplate within the Africa-Eurasia collision zone, *Geophys. Res. Lett.*, 31(9).  
536 doi:10.1029/2004GL019723, 2004.

537

538 Billi, A., Presti, D., Faccenna, C., Neri, G., and Orecchio, B.: Seismotectonics of the Nubia plate  
539 compressive margin in the south Tyrrhenian region, Italy: Clues for subduction inception, *J.*  
540 *Geophys. Res.*, 112(B8), 2007.

541

542 Billi, A., Presti, D., Orecchio, B., Faccenna, C., and Neri, G.: Incipient extension along the active  
543 convergent margin of Nubia in Sicily, Italy: Cefalù-Etna seismic zone, *Tectonics*, 29(4), 2010.

544

545 Billi, A., Faccenna, C., Bellier, O., Minelli, L., Neri, G., Piromallo, C., Presti, D., Scrocca, D., and  
546 Serpelloni, E.: Recent tectonic reorganization of the Nubia-Eurasia convergent boundary heading  
547 for the closure of the western Mediterranean, *Bull. Soc. Geol. Fr.*, 182(4), 279–303, 2011.

548

549 Cadet, J. P., and Funicello, R.: Carte Géodynamique de la Méditerranée: Geodynamique Map of  
550 the Mediterranean, Commission for the Geological Map of the World, 2004.

551

552 Calò, M., and Parisi, L.: Evidences of a lithospheric fault zone in the Sicily Channel continental rift  
553 (southern Italy) from instrumental seismicity data, *Geophys. J. Int.*, 199(1), 219-225, 2014.

554

555 Cavallaro, D., Monaco, C., Polonia, A., Sulli, A., and Di Stefano, A.: Evidence of positive tectonic  
556 inversion in the north-central sector of the Sicily Channel (Central Mediterranean), *Nat. Hazards*,  
557 86(2), 233-251, 2017.

558



- 559 Chiarabba, C., and Palano, M.: Progressive migration of slab break-off along the southern  
560 Tyrrhenian plate boundary: Constraints for the present day kinematics, *J Geodyn.*, 105, 51-61,  
561 2017.
- 562
- 563 Christova, C. V.: Spatial distribution of the contemporary stress field in the Kurile Wadati-Benioff  
564 zone by inversion of earthquake focal mechanisms, *J Geodyn.*, 83, 1-17, 2015.
- 565
- 566 D'Agostino, N., Avallone, A., Cheloni, D., D'Anastasio, E., Mantenuto, S., and Selvaggi, G.:  
567 Active tectonics of the Adriatic region from GPS and earthquake slip vectors, *J. Geophys. Res.*,  
568 113, B12413, doi:10.1029/2008JB005860, 2008.
- 569
- 570 D'Agostino, N., D'Anastasio, E., Gervasi, A., Guerra, I., Nedimovic, M.R., Seeber, L., and  
571 Steckler, M. Forearc extension and slow rollback of the Calabria Arc from GPS measurements,  
572 *Geophys. Res. Lett.*, 38, L17304, doi:10.1029/2011GL048270, 2011.
- 573
- 574 D'Amico, S., Orecchio, B., Presti, D., Zhu, L., Herrmann, R. B., and Neri, G.: Broadband  
575 waveform inversion of moderate earthquakes in the Messina Straits, southern Italy, *Phys. Earth  
576 Planet. Inter.*, 179(3-4), 97-106, 2010.
- 577
- 578 D'Amico, S., Orecchio, B., Presti, D., Gervasi, A., Zhu, L., Guerra, I., Neri, G., and Herrmann, R.  
579 B.: Testing the stability of moment tensor solutions for small earthquakes in the Calabro-Peloritan  
580 Arc region (southern Italy), *B. Geofis. Teor. Appl.*, doi:10.4430/bgta0009, 2011.
- 581
- 582 Evans, J.R., Eberhart-Phillips, D., and Thurber, C.H.: User's Manual for Simulps12for Imagingvp  
583 andVp/Vs: A Derivative of the "Thurber" Tomographic InversionSimul3 for Local Earthquakes and  
584 Explosions. Open-file Report. US GeologicalSurvey, Menlo Park, 94-431, 1994.



585

586 Gallais, F., Graindorge, D., Gutscher, M. A., and Klaeschen, D.: Propagation of a lithospheric tear  
587 fault (STEP) through the western boundary of the Calabrian accretionary wedge offshore eastern  
588 Sicily (Southern Italy), *Tectonophysics*, 602, 141-152, 2013.

589

590 Gephart, J. W.: FMSI: A FORTRAN program for inverting fault/slickenside and earthquake focal  
591 mechanism data to obtain the regional stress tensor, *Computers & Geosciences*, 16(7), 953-989,  
592 1990.

593

594 Gephart, J. W., and Forsyth, D. W.: An improved method for determining the regional stress tensor  
595 using earthquake focal mechanism data: application to the San Fernando earthquake sequence, *J.*  
596 *Geophys. Res.*, 89(B11), 9305-9320, 1984.

597

598 Ghisetti, F. C., Gorman, A. R., Grasso, M., and Vezzani, L.: Imprint of foreland structure on the  
599 deformation of a thrust sheet: The Plio-Pleistocene Gela Nappe (southern Sicily, Italy), *Tectonics*,  
600 28(4), 2009.

601

602 Gillard, D., Wyss, M., and Okubo, P.: Type of faulting and orientation of stress and strain as a  
603 function of space and time in Kilauea's south flank. *Hawaii, J. Geophys. Res.*, 101, 16025–16042,  
604 1996.

605

606 Gutscher, M. A., Dominguez, S., Lepinay, B. M., Pinheiro, L., Gallais, F., Babonneau, N.,  
607 Cattaneo, A., Le Faou, Y., Barreca G., Micallef, A., and Rovere, M.: Tectonic expression of an  
608 active slab tear from high-resolution seismic and bathymetric data offshore Sicily (Ionian Sea).  
609 *Tectonics*, 35(1), 39-54, 2016.

610



611 Gutscher, M. A., Kopp, H., Krastel, S., Bohrmann, G., Garlan, T., Zaragosi, S., Klaucke, I.,  
612 Wintersteller, P., Loubrieu, B., Le Faou, Y., San Pedro, L., Dominguez, S., Roverei, M., Mercier de  
613 Lepinay, B., Ranero, C., and Sallares, V. Active tectonics of the Calabrian subduction revealed by  
614 new multi-beam bathymetric data and high-resolution seismic profiles in the Ionian Sea (Central  
615 Mediterranean), *Earth Planet. Sci. Lett.*, 461, 61-72, 2017.

616

617 Lavecchia, G., Ferrarini, F., de Nardis, R., Visini, F., and Barbano, M. S.: Active thrusting as a  
618 possible seismogenic source in Sicily (Southern Italy): Some insights from integrated structural–  
619 kinematic and seismological data, *Tectonophysics*, 445(3-4), 145-167, 2007.

620

621 Li, H., Michelini, A., Zhu, L., Bernardi, F., and Spada, M. Crustal velocity structure in Italy from  
622 analysis of regional seismic waveforms, *Bull. Seism. Soc. Am.*, 97(6), 2024-2039, 2007.

623

624 Malinverno, A., and Ryan, W. B.: Extension in the Tyrrhenian Sea and shortening in the Apennines  
625 as result of arc migration driven by sinking of the lithosphere, *Tectonics*, 5(2), 227-245, 1986.

626

627 Marotta, A. M., Splendore, R., and Barzaghi, R.: An application of model uncertainty statistical  
628 assessment: A case study of tectonic deformation in the Mediterranean, *J Geodyn*, 85, 24-31, 2015.

629

630 Michael, A.J.: Use of focal mechanisms to determine stress: a control study, *J. Geophys. Res.* 92,  
631 357–368, 1987.

632

633 Minelli, L., and Faccenna, C.: Evolution of the Calabrian accretionary wedge (Central  
634 Mediterranean), *Tectonics*, 29, TC4004, doi: 10.1029/2009TC002562, 2010.

635



636 Musumeci, C., Scarfi, L., Palano, M., and Patanè, D.: Foreland segmentation along an active  
637 convergent margin: New constraints in southeastern Sicily (Italy) from seismic and geodetic  
638 observations, *Tectonophysics*, 630, 137-149, 2014.

639

640 Neri, G., Barberi, G., Oliva, G., and Orecchio, B.: Spatial variations of seismogenic stress  
641 orientations in Sicily, south Italy, *Phys. Earth Planet. Inter.*, 148(2-4), 175-191, 2005.

642

643 Neri, G., Orecchio, B., Totaro, C., Falcone, G., and Presti, D.: Subduction beneath southern Italy  
644 close the ending: results from seismic tomography, *Seismol. Res. Lett.*, 80, 63–70, 2009.

645

646 Neri, G., Marotta, A. M., Orecchio, B., Presti, D., Totaro, C., Barzaghi, R., and Borghi, A.: How  
647 lithospheric subduction changes along the Calabrian Arc in southern Italy: geophysical evidences.  
648 *Int. J. Earth Sci.*, 101(7), 1949-1969, 2012.

649

650 Nijholt, N., Govers, R., and Wortel, R.: On the forces that drive and resist deformation of the south-  
651 central Mediterranean: a mechanical model study, *Geophys. J. Int.*,  
652 <https://doi.org/10.1093/gji/ggy144>, 2018.

653

654 Nocquet, J.: Present-day kinematics of the Mediterranean: a comprehensive overview of GPS  
655 results, *Tectonophysics*, 579, 220–242, 2012.

656

657 Oldow, J.S., Ferranti, L., Lewis, D. S., Campbell, J. K., d'Argenio, B., Catalano, R., Pappone, G.,  
658 Carmignani, L., Conti, P., and Aiken, C. L. V.: Active fragmentation of Adria, the north African  
659 promontory, central Mediterranean orogen, *Geology*, 30, 779–782, 2002.

660



661 Oldow, J. S., and Ferranti, L.: Fragmentation of Adria and active decollement tectonics within the  
662 southern peri-Tyrrhenian orogen, Italy. In *The Adria Microplate: GPS Geodesy, Tectonics and*  
663 *Hazards*, 269-286, Springer, Dordrecht, 2006.

664

665 Orecchio, B., Presti, D., Totaro, C., Guerra, I., and Neri, G.: Imaging the velocity structure of the  
666 Calabrian Arc region (South Italy) through the integration of different seismological data, *Boll.*  
667 *Geofis. Teor. Appl.*, 52, 625–638, 2011.

668

669 Orecchio, B., Presti, D., Totaro, C., and Neri, G.: What earthquakes say concerning residual  
670 subduction and STEP dynamics in the Calabrian Arc region, south Italy. *Geophys. J. Int.*, 199(3),  
671 1929-1942, 2014.

672

673 Orecchio, B., Presti, D., Totaro, C., D'Amico, S., and Neri, G.: Investigating slab edge kinematics  
674 through seismological data: The northern boundary of the Ionian subduction system (south Italy), *J.*  
675 *Geodyn.*, 88, 23-35, 2015.

676

677 Orecchio, B., Aloisi, M., Cannavò, F., Palano, M., Presti, D., Pulvirenti, F., Totaro, C., Siligato, G.,  
678 and Neri, G.: Present-day kinematics and deformation processes in the southern Tyrrhenian region:  
679 new insights on the northern Sicily extensional belt, *It. Jour. Geosci.*, 136(3), 418-433, 2017.

680

681 Palano, M., Ferranti, L., Monaco, C., Mattia, M., Aloisi, M., Bruno, V., Cannavò, F., and Siligato,  
682 G.: GPS velocity and strain fields in Sicily and southern Calabria, Italy: updated geodetic  
683 constraints on tectonic block interaction in the central Mediterranean, *J. Geophys. Res.*, 117(B7),  
684 2012.

685



686 Palano, M., González, P. J., and Fernández, J.: The diffuse plate boundary of Nubia and Iberia in  
687 the Western Mediterranean: crustal deformation evidence for viscous coupling and fragmented  
688 lithosphere, *Earth Planet. Sci. Lett.*, 430, 439-447, 2015a.

689

690 Palano M., Schiavone D., Loddo M., Neri M., Presti D., Quarto R., Totaro C., and Neri G. Active  
691 upper crust deformation pattern along the southern edge of the Tyrrhenian subduction zone (NE  
692 Sicily): Insights from a multidisciplinary approach, *Tectonophysics*, 657, 205-218,  
693 doi:10.1016/j.tecto.2015.07.005, 2015b.

694

695 Parker, R. L., and McNutt, M. K. Statistics for the one-norm misfit measure, *J. Geophys. Res.*,  
696 85(B8), 4429-4430, 1980.

697

698 Pérouse, E., Chamot-Rooke, N., Rabaute, A., Briole, P., Jouanne, F., Georgiev, I., and Dimitrov,  
699 D.: Bridging onshore and offshore presentday kinematics of central and eastern Mediterranean:  
700 Implications for crustal dynamics and mantle flow, *Geochem. Geophys. Geosyst*, 13, 9, 2012.

701

702 Polonia, A., Torelli, L., Mussoni, P., Gasperini, L., Artoni, A., and Klaeschen, D.: The Calabrian  
703 Arc subduction complex in the Ionian Sea: Regional architecture, active deformation, and seismic  
704 hazard, *Tectonics*, 30(5), 2011.

705

706 Polonia, A., Torelli, L., Artoni, A., Carlini, M., Faccenna, C., Ferranti, L., Gasperini, L., Govers, R.,  
707 Klaeschen, D., Monaco, C., Neri, G., Nijholt, N., Orecchio, B., and Wortel, R.: The Ionian and  
708 Alfeo–Etna fault zones: New segments of an evolving plate boundary in the central Mediterranean  
709 Sea?. *Tectonophysics*, 675, 69-90, 2016.

710



711 Polonia, A., Torelli, L., Gasperini, L., Cocchi, L., Muccini, F., Bonatti, E., Hensen, C., Schmidt, M.,  
712 Romano, S., Artoni, A., and Carlini, M. Lower plate serpentinite diapirism in the Calabrian Arc  
713 subduction complex, *Nat. Commun.*, 8(1), 2172, 2017.

714

715 Pondrelli, S., Piromallo, C., and Serpelloni, E.: Convergence vs. retreat in Southern Tyrrhenian Sea:  
716 insights from kinematics. *Geophys. Res. Lett.*, 31(6), 2004.

717

718 Presti, D., Troise, C., and De Natale, G.: Probabilistic location of seismic sequences in  
719 heterogeneous media, *Bull. Seism. Soc. Am.*, 94, 2239–2253, 2004.

720

721 Presti, D., Orecchio, B., Falcone, G., and Neri, G.: Linear versus nonlinear earthquake location and  
722 seismogenic fault detection in the southern Tyrrhenian Sea, Italy, *Geophys. J. Int.*, 172, 607–618,  
723 2008.

724

725 Presti, D., Billi, A., Orecchio, B., Totaro, C., Faccenna, C., and Neri, G.: Earthquake focal  
726 mechanisms, seismogenic stress, and seismotectonics of the Calabrian Arc, Italy, *Tectonophysics*,  
727 602, 153-175, 2013.

728

729 Rebaï, S., Philip, H., and Taboada, A.: Modern tectonic stress field in the Mediterranean region:  
730 evidence for variation in stress directions at different scales. *Geophys J Int*, 110(1), 106-140, 1992.

731

732 Sani, F., Vannucci, G., Boccaletti, M., Bonini, M., Corti, G., and Serpelloni, E. Insights into the  
733 fragmentation of the Adria Plate, *J Geodyn*, 102, 121-138, 2016.

734



735 Serpelloni, E., Vannucci, G., Pondrelli, S., Argnani, A., Casula, G., Anzidei, M., Baldi, P., and  
736 Gasperini, P.: Kinematics of the Western Africa-Eurasia plate boundary from focal mechanisms and  
737 GPS data, *Geophys J Int*, 169(3), 1180-1200, 2007.

738

739 Splendore, R., and Marotta, A. M. Crust-mantle mechanical structure in the Central Mediterranean  
740 region, *Tectonophysics*, 603, 89-103, 2013.

741

742 Totaro, C., Presti, D., Billi, A., Gervasi, A., Orecchio, B., Guerra, I., and Neri, G.: The ongoing  
743 seismic sequence at the Pollino Mountains, Italy, *Seismol. Res. Lett.*, 84(6), 955-962, 2013.

744

745 Totaro, C., Seeber, L., Waldhauser, F., Steckler, M., Gervasi, A., Guerra, I., Orecchio, B., and  
746 Presti, D.: An Intense Earthquake Swarm in the Southernmost Apennines: Fault Architecture from  
747 High-Resolution Hypocenters and Focal MechanismsAn Intense Earthquake Swarm in the  
748 Southernmost Apennines. *Bull. Seism. Soc. Am.*, 105(6), 3121-3128, 2015.

749

750 Totaro, C., Orecchio, B., Presti, D., Scolaro, S., and Neri, G.: Seismogenic stress field estimation in  
751 the Calabrian Arc region (south Italy) from a Bayesian approach. *Geophys. Res. Lett.*, 43(17),  
752 8960-8969, 2016.

753

754 Wortel, R., and Spakman, W.: Subduction and slab detachment in the Mediterranean-Carpathian  
755 region, *Science*, 290, 1910–1917, 2000.

756

757 Wyss, M., Liang, B., Tanigawa, W.R., and Xiaoping, W.: Comparison of orientations of stress and  
758 strain tensor based on fault plane solutions in Kaoiki, Hawaii, *J. Geophys. Res.* 97, 4769–4790,  
759 1992.

760



761 Zoback, M.L.: First-and second-order patterns of stress in the lithosphere: the world stress map

762 project, *J. Geophys. Res.*, 97(B8), 11 703–11 728, 1992.

763

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767 **Figures caption**

768

769 **Fig. 1.** Simplified sketch map of the Africa-Eurasia plate boundary. Black arrows indicate the  
770 present-day sense of motion of Africa with respect to Eurasia according to Palano et al. (2015a).  
771 Abbreviations: AP=Apennines, CA=Calabrian Arc, MA=Magrhebides. The dashed rectangle  
772 including southern Italy and relative off-shore sectors indicates the main area of interest in the  
773 present study (see Fig. 2). The top-left inset shows the western Mediterranean plate boundary  
774 evolution in the last 30 Myrs (redrawn from Wortel and Spakman, 2000, with modifications  
775 according to Neri et al., 2009). The solid curves with the sawtooth pattern indicate the location of  
776 the boundary at different times as a consequence of ESE-ward rollback of the WNW-ward  
777 subducting lithosphere. The sawteeth point in the direction of subduction or underthrusting. Black  
778 sawteeth indicate in-depth continuous subducting slab in contrast to white sawteeth marking plate  
779 boundary segments where slab detachment has occurred. The white arrow along the Apennines  
780 shows the inferred direction of lateral migration of slab detachment.

781

782 **Fig. 2.** Tectonic map of southern Italy and relative off-shore areas (redrawn from Palano et al.,  
783 2012, with integration of data in the Ionian Sea according to Polonia et al., 2011). Black arrows  
784 indicate the present-day sense of motion of Africa with respect to Eurasia according to Palano et al.  
785 (2015a). TFS stands for Tindari Fault System.

786

787 **Fig. 3.** Hypotheses of lithosphere fragmentation and microplate architecture proposed in the  
788 literature for the central Mediterranean region. (a) Simplified tectonic setting of the central  
789 Mediterranean region (redrawn from Cadet and Funicciello, 2004). (b) Main plate boundaries and  
790 Adria microplate contours according to Anderson and Jackson (1987). RP indicates the pole of  
791 rotation of Adria wrt Eurasia. (c) Macroplate and Adria representation in the view of Oldow et al.  
792 (2002). (d) Adria separation in two independent blocks located between the main plates as reported



793 by Battaglia et al. (2004). (e) Sketch of the main tectonic and kinematic features in the central  
794 Mediterranean (redrawn from Serpelloni et al., 2007). (f) Adria and Apulia-Ionian-Hyblean blocks  
795 according to D'Agostino et al. (2008). (g-h) The two alternative scenarios proposed by Palano et al.  
796 (2012): (g) the Ionian domain is rigidly connected with the Sicilian-Hyblean-Malta block; (h) the  
797 Ionian domain diverges from the Sicilian-Hyblean-Malta block and moves together with the  
798 Calabrian block. (i) According to Pérouse et al. (2012) a rigid single block including the Hyblean  
799 Plateau, the Ionian basin, the Sirte plain, the Apulian peninsula and the southern Adriatic sea can be  
800 hypothesized, rotating clockwise wrt Africa (the star indicates the rotation pole). (j) New  
801 reconstruction of Adria domain between the main plates according to Sani et al. (2016). Shaded  
802 areas indicate uncertain limits among blocks.

803

804 **Fig. 4.** Section (a) displays the map of seismic stations used for hypocenter locations in the present  
805 study. The dashed rectangle indicates the study area for the earthquake locations reported in  
806 sections (b) to (d) of this figure. The continuous rectangles indicate the sectors Western Ionian (WI)  
807 and Sicily Channel (SC) where hypocenter locations were also performed by the Bayloc method  
808 (the results are shown in Figs 6 and 7). The shadowed curved belt shows the approximate location  
809 of the Apennine-Maghrebian chain in south Italy and Sicily. Section (b) displays the epicentral map  
810 of earthquakes of magnitude over 2.5 that occurred between 1981 and 2016 at depths less than 100  
811 km in the area 10°-20°E 35°-41°N (circles are proportional to the earthquake magnitude, see  
812 legend). For these locations we used the linearized location method known as Simulps (Evans et al.  
813 1994) and the 3D seismic velocity structure proposed for the study region by Orecchio et al. (2011).  
814 Sections (c) and (d) display the earthquakes of Section (b) after separation according to hypocenter  
815 depth (0-30 km and 30-100 km, respectively).

816

817 **Fig. 5.** Fault-plane solutions of earthquakes of magnitude over 2.5 occurring in the period 1977-  
818 2016 at depths less than 70 km in the area countoured by the dashed rectangle in Fig. 4. Only



819 solutions estimated by waveform inversion are reported. The main parameters and the bibliographic  
820 source of each solution are given in Table A1, Appendix A. The different colors in the figure  
821 identify different types of mechanisms following Zoback's (1992) classification based on values of  
822 plunges of P and T axes: red = normal faulting (NF) or normal faulting with a minor strike-slip  
823 component (NS); green = strike-slip faulting (SS); blue = thrust faulting (TF) or thrust faulting with  
824 a minor strike-slip component (TS); black = unknown stress regime (U). "U" includes all focal  
825 mechanisms that do not fall in the other five categories (Zoback, 1992). The beach ball size is  
826 proportional to the earthquake magnitude (see legend). The curved line contours the study area for  
827 stress inversion (results in Fig. 8)

828

829 **Fig. 6.** (a-c-d) Epicentral maps obtained by the Bayloc probabilistic method for the earthquakes  
830 occurring during 1981–2016 in the sector WI, depth ranges 0-70 km (plot a), 0-30 km (plot c) and  
831 30-70 km (plot d). For comparison, the structural information of Fig. 2 is reported. (b) Bayloc's  
832 hypocentral vertical section along the AA' profile indicated in plot a, +/- 70 km around the profile.  
833 AEF and IF stand for Alfeo-Etna Fault and Ionian Fault, respectively.

834

835 **Fig. 7.** (a-c-d) Epicentral maps obtained by the Bayloc probabilistic method for the earthquakes  
836 occurring during 1981–2016 in the sector SC, depth ranges 0-70 km (plot a), 0-30 km (plot c) and  
837 30-70 km (plot d). For comparison, the structural information of Fig. 2 is reported. (b) Bayloc's  
838 hypocentral vertical section of earthquakes of plot (a) along a west-east profile.

839

840 **Fig. 8.** (a) Orientations of the principal stress axes (lower hemisphere stereographic projection)  
841 obtained by inversion of the earthquake focal mechanisms shown in map. Red, green, and blue dots  
842 indicate the orientations of the maximum ( $\sigma_1$ ), intermediate ( $\sigma_2$ ), and minimum ( $\sigma_3$ ) compressive  
843 stresses, respectively. Crosses and squares indicate the 95% confidence areas for the  $\sigma_1$  and  $\sigma_3$   
844 axes, respectively. F is the average of the individual earthquake misfits wrt the best model of stress



845 found by inversion (see text). (b) Stress inversion results obtained after subdivision of the study area  
846 in two sub-areas W and E, west and east of the black line AB, respectively. (c) Stress inversion  
847 results in the shadowed sector RZ where Polonia et al. (2017) have identified a rifting process with  
848 opening in the SW-NE direction, approximately. See Table 1 for numerical values of stress  
849 inversion results.

850

851 **Fig. 9.** This figure indicates the overall compressional domain caused by Africa-Eurasia  
852 convergence in southern Italy disturbed by (i) extensional processes in the Calabrian Arc, (ii) rifting  
853 in the westernmost Ionian offshore Sicily and (iii) subduction-related reduced compression in the  
854 trench retreat zone offshore eastern Calabria. Black-to-grey transition of GPS crustal motion vectors  
855 marks their clear orientation change from NW-ward to NE-ward and corresponds with the onshore  
856 prolongation of the Ionian NW-trending rifting zone (diverging arrows are taken from Polonia et al.,  
857 2017, GPS vectors are from Palano et al., 2012). Seismic data do not give significant information  
858 on stress regimes in the white field of the southern Tyrrhenian sea.

859

860 **Table 1.** Stress tensor inversion of earthquake focal mechanisms performed for the earthquake sets  
861 ALL, W, E and RZ described in the text and relative to the sectors indicated in Fig. 8. N is the  
862 number of earthquakes (= focal mechanisms) belonging to the inversion set. F is the average of the  
863 misfits of the individual earthquakes with respect to the best model of stress found by inversion. R  
864 is the amplitude ratio  $(\sigma_2 - \sigma_1) / (\sigma_3 - \sigma_1)$  where  $\sigma_1$ ,  $\sigma_2$ , and  $\sigma_3$  represent the amplitudes of the  
865 maximum, intermediate and minimum compressive stress, respectively. Pl and Az are the plunge  
866 and azimuth, respectively, of the three main stress axes.

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870  
 871 **Table 1.** Stress tensor inversion of earthquake focal mechanisms performed for the earthquake sets  
 872 ALL, W, E and RZ described in the text and relative to the sectors indicated in Fig. 8. N is the  
 873 number of earthquakes (= focal mechanisms) belonging to the inversion set. F is the average of the  
 874 misfits of the individual earthquakes with respect to the best model of stress found by inversion. R  
 875 is the amplitude ratio  $(\sigma_2 - \sigma_1) / (\sigma_3 - \sigma_1)$  where  $\sigma_1$ ,  $\sigma_2$ , and  $\sigma_3$  represent the amplitudes of the  
 876 maximum, intermediate and minimum compressive stress, respectively. Pl and Az are the plunge  
 877 and azimuth, respectively, of the three main stress axes.

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Set	N	F (°)	R	$\sigma_1$ Pl (°)	$\sigma_1$ Az (°)	$\sigma_2$ Pl (°)	$\sigma_2$ Az (°)	$\sigma_3$ Pl (°)	$\sigma_3$ Az (°)
All	72	8.3	0.5	3	320	76	217	14	51
W	32	5.9	0.5	3	150	84	275	5	60
E	40	8.3	0.5	3	319	75	216	15	50
RZ	27	6.5	0.4	64	162	25	345	3	79

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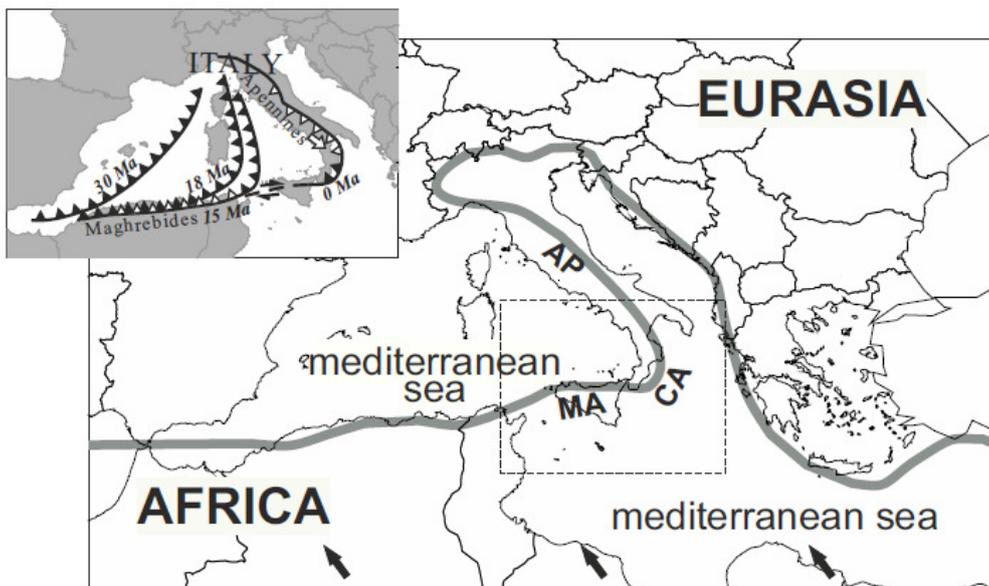


Figure 1

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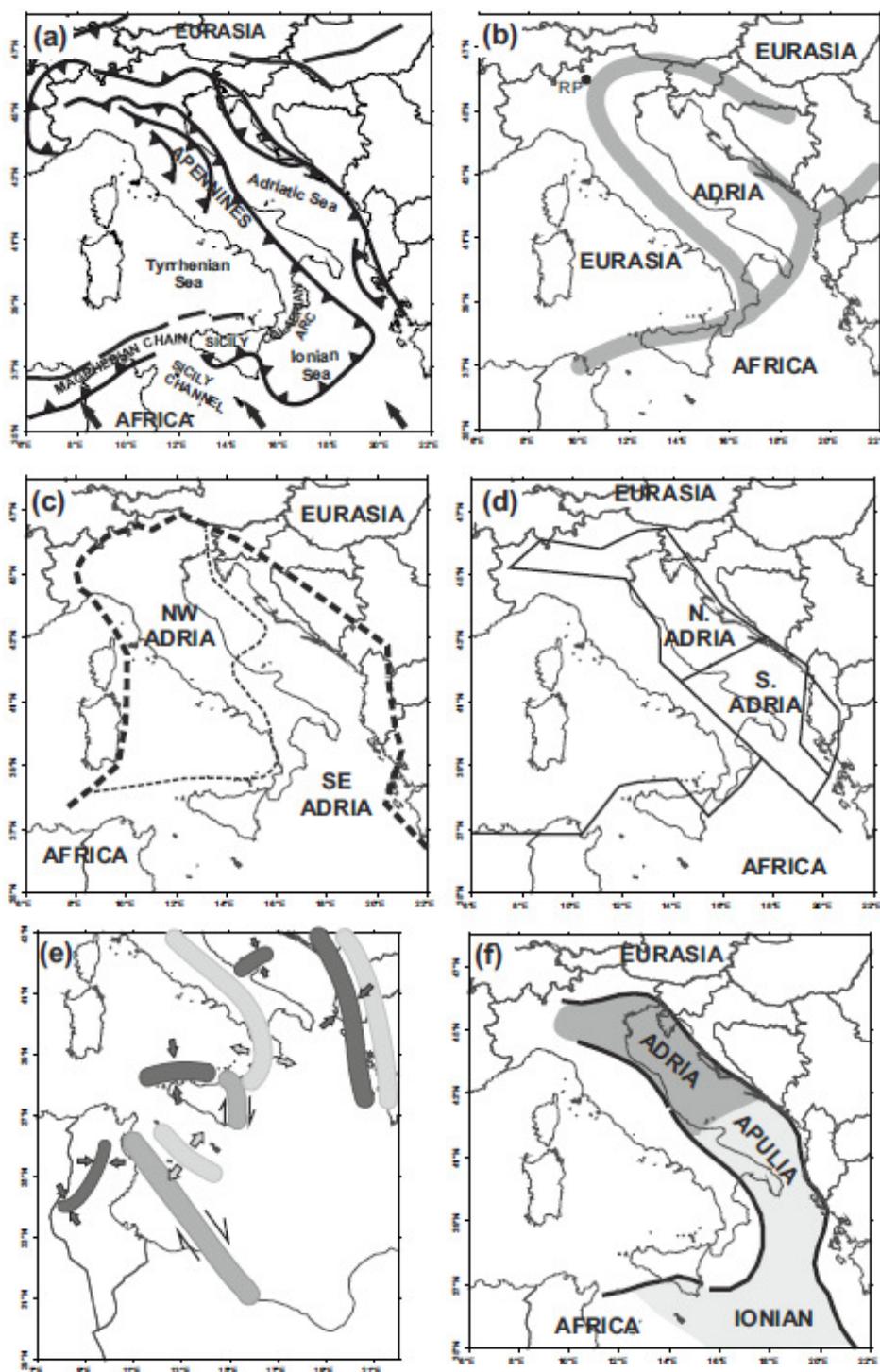


Figure 3

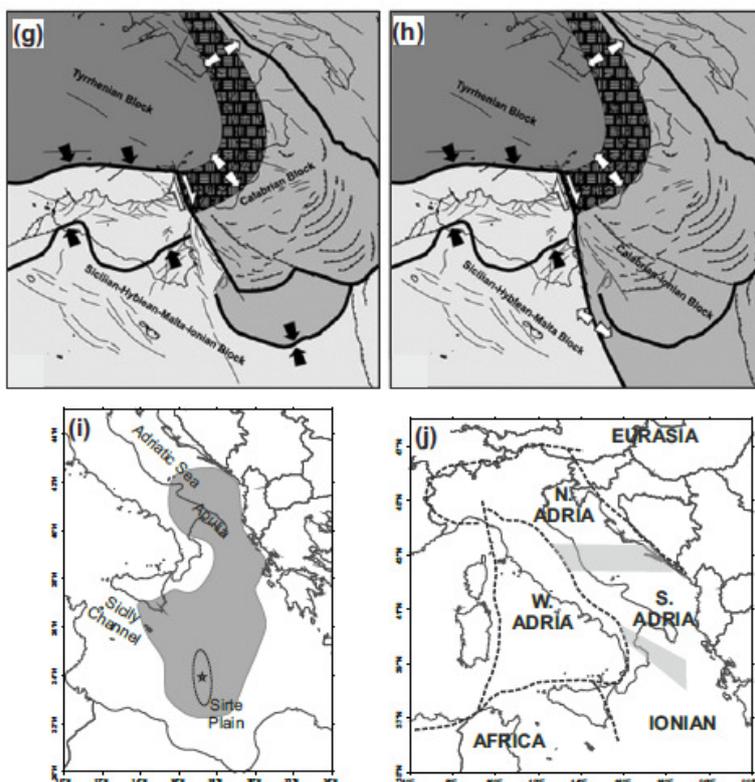


Figure 3 (continued)

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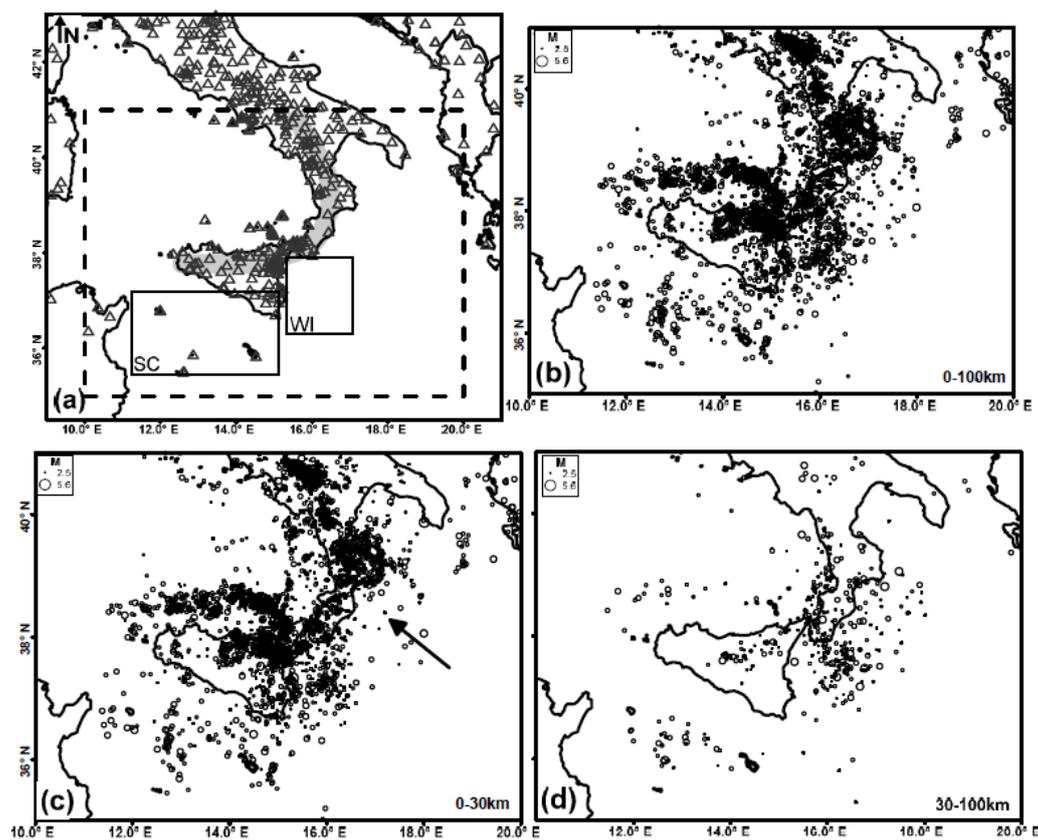


Figure 4

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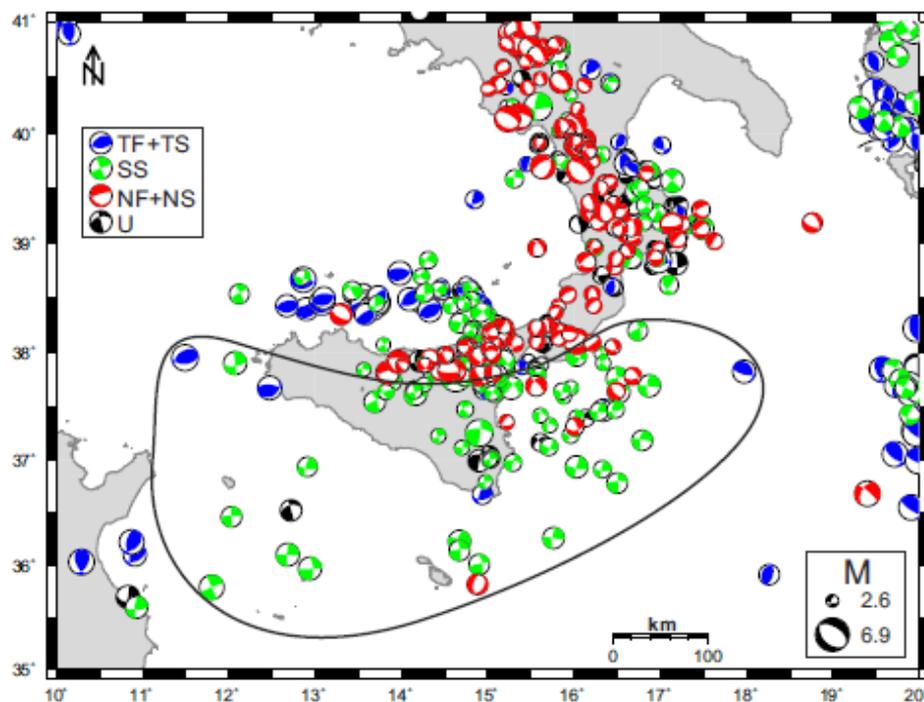


Figure 5

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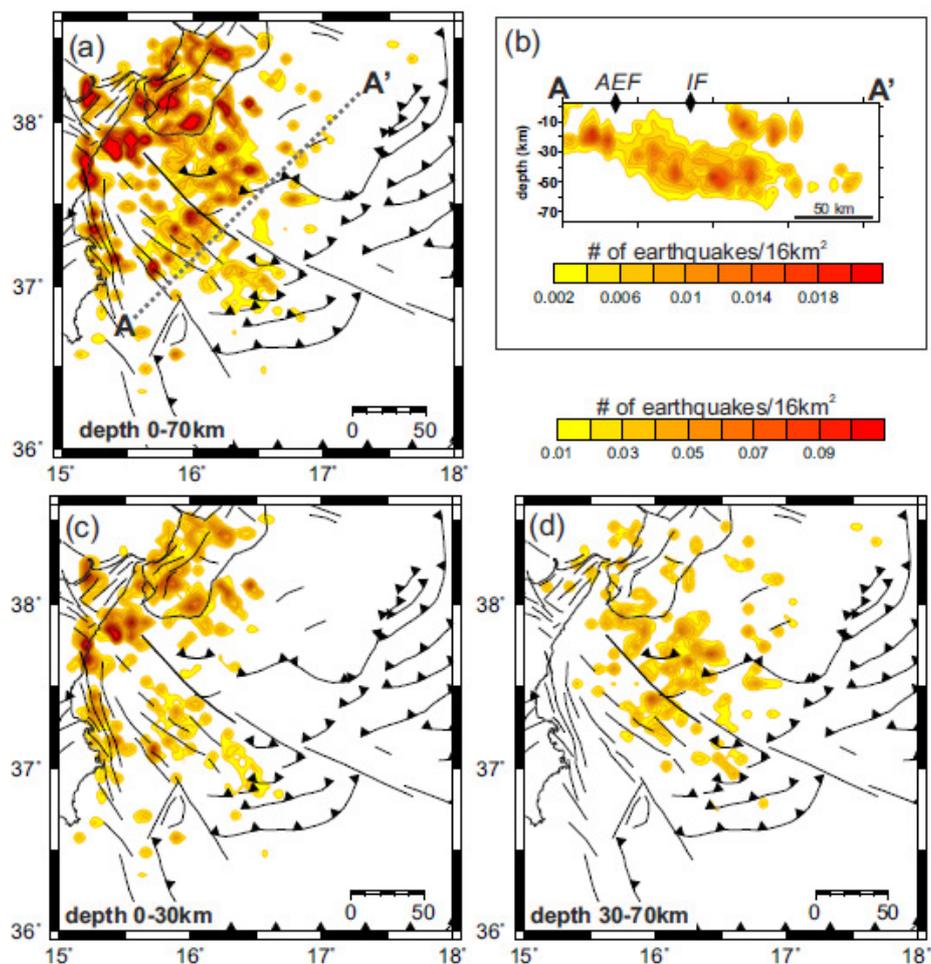


Figure 6

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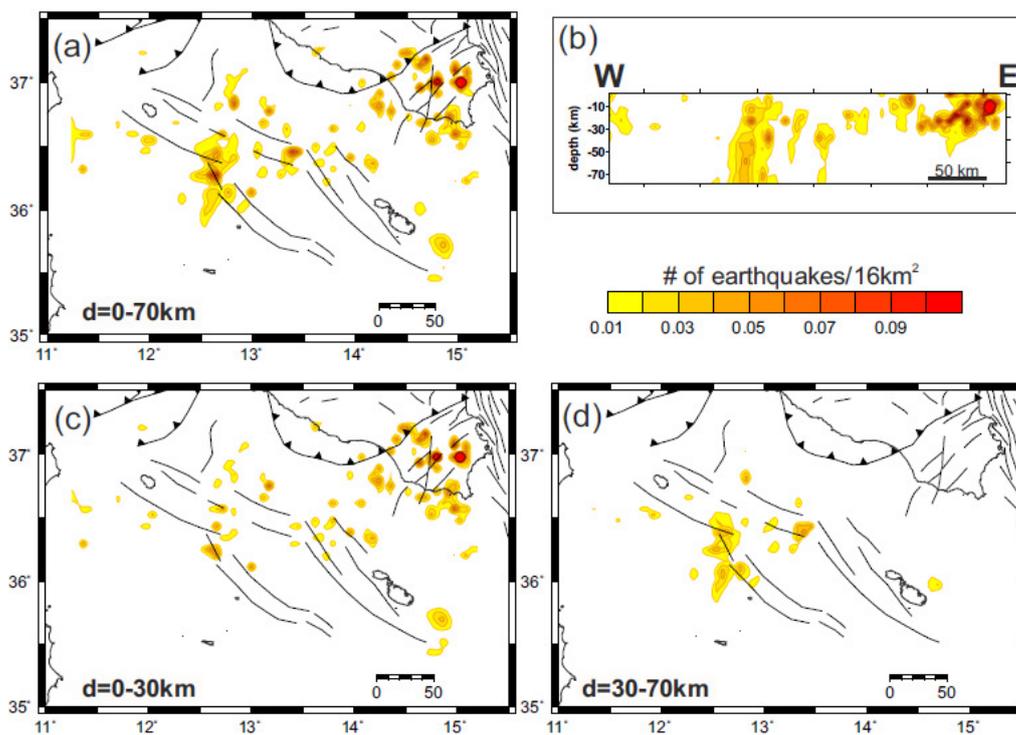


Figure 7

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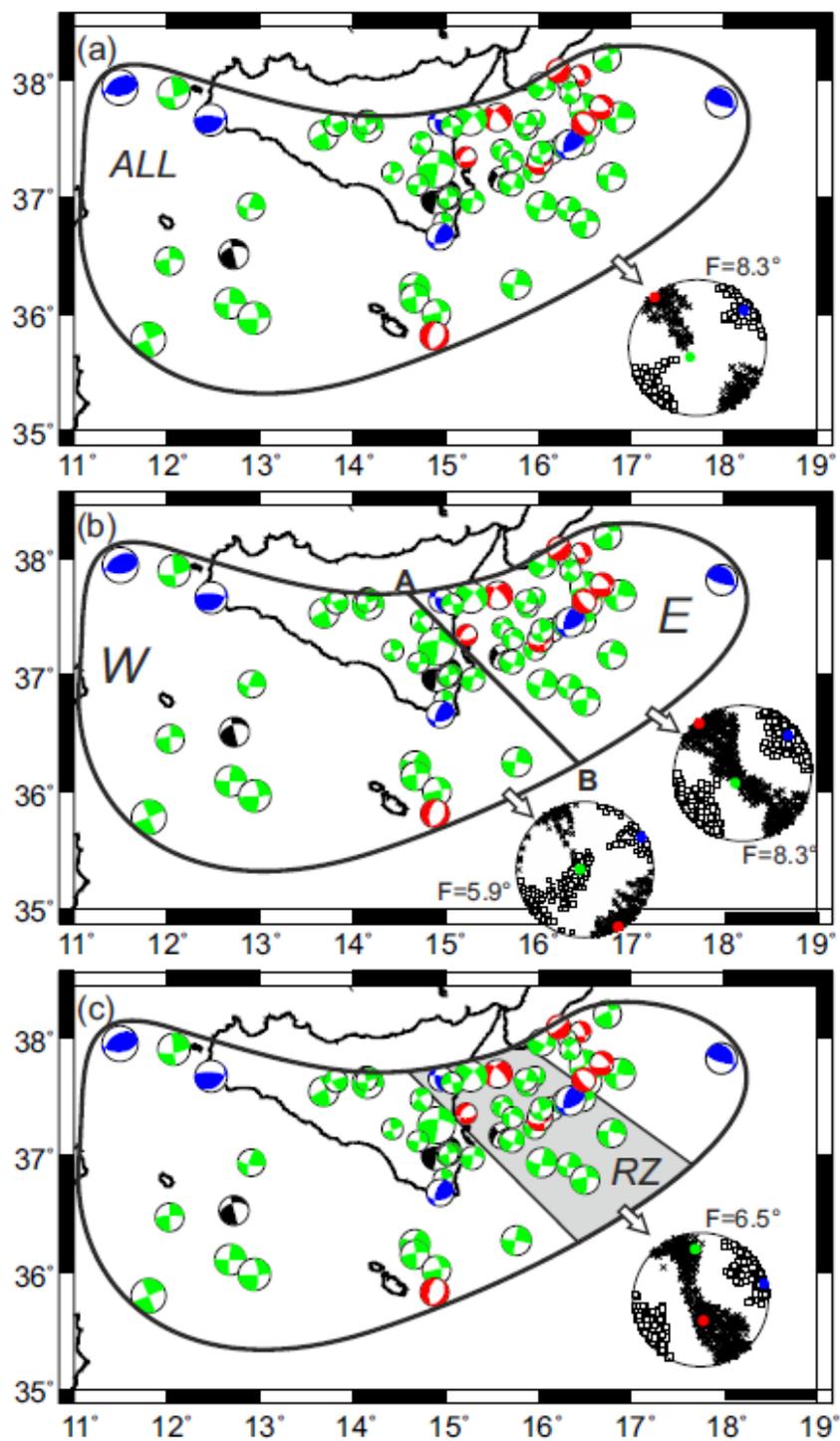


Figure 8

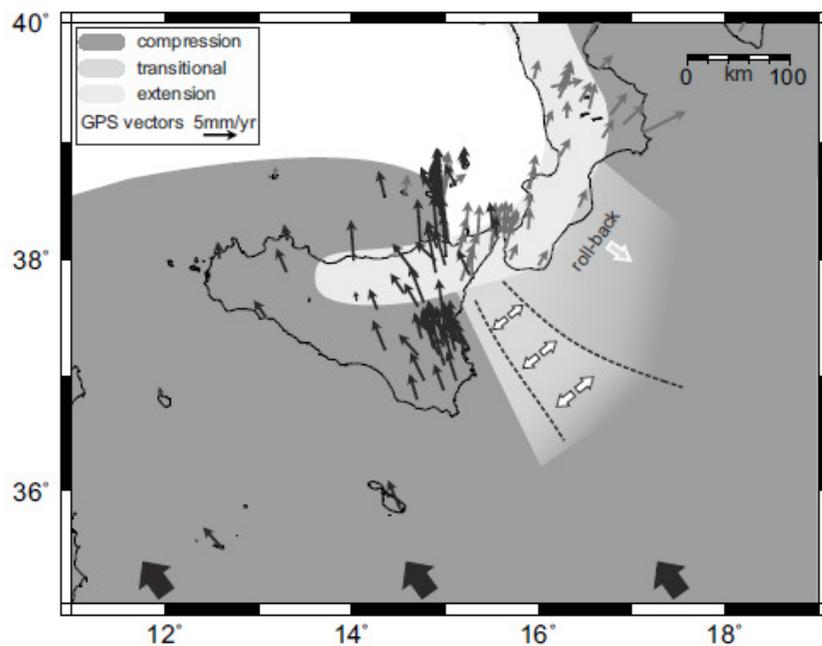


Figure 9

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