Responses to Prof. Federico Rossetti

Dear Prof. Federico Rossetti,

We thank your further reviews, it is very kindly of you to revise the English language. Based on your suggestions and comments, the manuscript have been revised.

The responses to your comments are as follows (comments in bold):

(1) It mostly concerns the English language and manuscript organisation of the revised sections.

We revise the English language and re-organise the revised sections. The titles of the 2.1 and 2.2 sections are deleted (Page 3, Line 19; Page 4, Line 32).

(2) Please check for quality English editing before the re-submission.

In order to check the quality English editing, **t**his revised manuscript has been edited again by a special English editing company (Stallard Scientific Editing) in New Zealand. They help to edit the English language and the format of the manuscript. After that, We check the whole manuscript again.

We look forward that the present version can be put in publication process.

Sincerely,

Yini Wang, Wenliang Xu, Feng Wang, Xiaobo Li

New insights on the early Mesozoic evolution of multiple tectonic regimes in the northeastern North China Craton from the detrital zircon provenance of sedimentary strata

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Abstract. To investigate the timing of deposition and provenance of early Mesozoic strata in the northeastern North China Craton (NCC), and to understand the early Mesozoic paleotectonic evolution~~tectono-paleogeography~~ of the region, we combine stratigraphy, U–Pb zircon geochronology, and Hf isotopic analyses~~LA-ICP-MS detrital zircon U-Pb dating, Hf isotopic data~~. Early Mesozoic strata include the Early Triassic Heisonggou, Late Triassic Changbai and Xiaoyingzi, and Early Jurassic Yihe formations. Detrital zircons in the Heisonggou Formation ~~comprise~~ yield ~58% Neoarchean to Paleoproterozoic ages and ~42% Phanerozoic ~~grains~~ ages, and ~~that~~ were sourced from areas to the south and north of the basins within the NCC, respectively. This indicates that Early Triassic deposition was controlled primarily by the southward subduction of the Paleo-Asian oceanic plate beneath the NCC~~,~~ and collision between the NCC and the Yangtze Craton (YC). Approximately 88% of the sediments within the Late Triassic Xiaoyingzi Formation were sourced from the NCC to the south, with the remaining ~12% from the Xing'an–Mongolia Orogenic Belt (XMOB) to the north. This implies that Late Triassic deposition was related to the final closure of the Paleo-Asian Ocean during the Middle Triassic and the rapid exhumation of the Su–Lu Orogenic Belt between the NCC and YC. In contrast, ~88% of sediments within the Early Jurassic Yihe Formation were sourced from the XMOB to the north, with the remaining ~12% from the NCC to the south. We therefore infer that rapid uplift of the XMOB and the onset of subduction of the Paleo-Pacific Plate beneath Eurasia occurred in the Early Jurassic.

1 Introduction

The Mesozoic tectonic evolution of the East Asian continental margin has been commonly interpreted as ~~the~~a response to multiple and overprinting tectonic, commonly referred to~~one of the research hotspots in earth scientific field owing to overprinting and evolution of multiple tectonic regimes such~~ as (i) the closure of the Paleo-Asian Ocean (Sengör and Natal'in, 1996; Li, 2006; Zhang et al., 2009, 2010, 2014; Zhou et al., 2010, 2015; Tang et al., 2013; Xu et al., 2013, 2014, 2015; Zhao et al., 2017), (ii) the subduction of the Paleo-Pacific Plate~~regimes~~ (Lin et al., 1998; Li et al.,

1999; Wu et al., 2000, 2004, 2007b; Jia et al., 2004; Zhang et al., 2004; Shen et al., 2006; Xu et al., 2009, 2013), and (iii) as well as subduction and collision between the North China Craton (NCC) and the Yangtze Craton (YC) (Yang et al., 2007; Pei et al., 2008). The northeastern NCC, a component of the East Asian continental margin, is located at the intersection among three tectonic regimes (Fig. 1). Therefore, the northeastern NCC can be considered as an ideal area to reveal the overprinting and evolution of multiple tectonic regimes during the Mesozoic. Within such tectonic scenarios, the northeastern NCC (as a component of the East Asian continental margin; Fig. 1) is a key area for deciphering the spatio-temporal Mesozoic tectonic evolution at a regional scale.

Two main aspects are still debated in the literature regarding the Mesozoic tectonic evolution of the northeastern NCCTwo problems that has been controversial is the key to understand the evolution of the northeastern NCC: One is(1) the timing of final closure of the Paleo-Asian Ocean, the and other is(2) the onset timing of onset of subduction of the Paleo-Pacific Plate beneath Eurasia. Up to now, various opinions exist regardingAt present, several opinions have been provided for the final closure timing of the Paleo-Asian Ocean, referred to the such as late Permian (JBGMR, 1988; Shi, 2006), early–middle Permian to Middle Triassic (Wang et al., 2015b), and Early–Middle Triassic (Sun et al., 2004; Li et al., 2007; Cao et al., 2013), respectively. Similarly, regarding the timing of For the onset timing subduction onset of the Paleo-Pacific Plate beneath Eurasia, multiple agesdates have been proposed so far, including the early Permian (Ernst et al., 2007; Sun et al., 2015; Yang et al., 2015b; Bi et al., 2016), Late Triassic (Wu et al., 2011; Wilde and Zhou, 2015), latest Triassic to Early Jurassic (Zhou and Li, 2017), and Early Jurassic (Xu et al., 2012, 2013; Guo, 2016; Tang et al., 2016; Wang et al., 2017). These contrasting reconstructions were mainly derived fromThe considerations above-mentioned were gained based mainly on the studies of the Mesozoic igneous rocks (especially granitoids) in the northeastern China (Li et al., 1999; Wu et al., 2002, 2007a, 2007b; Ge et al., 2007; Zhao and Zhang, 2011; Xu et al., 2012, 2013; Dong et al., 2014, 2016; Yang et al., 2015a, 2015b, 2016). However, the diversity in the compositions of these granitoids resulted in the different conclusions above-mentioned.

Compared with studies on of igneous rocks, basin stratigraphy and sedimentary provenance studiessedimentary basins, especially the provenance from sedimentary strata could provide direct evidence for reconstructing the regional tectonic and paleogeographic evolution tectono-paleogeography during early Mesozoic. Some examples on reconstructing the paleotectonic setting by studying the evolution of sedimentary basins have been carried out in the northern and western–central NCC (Meng, 2003; Meng et al., 2011, 2014; Li and Huang, 2013; Li et al., 2015; Liu et al., 2015; Xu et al., 2016; Meng, 2017), but the formation timing, evolution, and tectonic setting of the early Mesozoic strata in the northeastern NCC remain poorly constrained.

Recent studies have indicated that detrital zircon geochronology has become a powerful tool for provenance analysis and can aid in constraining paleogeography, tectonic reconstructions, and crustal evolution (Ross and Bowring, 1990; Gehrels and Dickinson 1995; Gehrels et al., 1995, 2002; Cawood and Nemchin, 2001; Meng et al., 2010; Wang

et al., 2012a). Therefore, detrital zircon geochronology is critical in paleogeographic and tectonic reconstruction.

The northeastern NCC contains a series of early Mesozoic faulted basins (referred to as the Fusong–Changbai basin Basin group Group; JBGMR, 1976; Fig. 1), which that are filled by volcanics and coal-bearing clastic sediments. The sedimentary from this basins has the potential to decode the tectonic evolution of the northeastern NCC in the frame of the Mesozoic regional geodynamic scenario.

The establishment of the Mesozoic stratigraphic framework in of the northeastern NCC is based primarily on lithostratigraphic correlation (JBGMR, 1963, 1976, 1988, 1997). The lack of precise age data has resulted in some questions uncertainties regarding studies on the sedimentary strata in the northeastern NCC, including (1) the timing of deposition when did these strata form?; (2) what is thethe correct lithostratigraphic framework?; (3) where were thethe sediment source sediments offor the basins sourced from?; and (4) what is the relationship between the basin evolution of basins and the regional tectonic setting? .

In this contribution, we focus on the sedimentary strata from the basins in the northeastern NCC (Figs 1–2) and use U–Pb age and Hf isotopic data from detrital and magmatic zircons, combined with biostratigraphic data, to constrain the formation timing of deposition and the provenance of the early Mesozoic strata, and to establish a new stratrigraphic framework in for the northeastern NCC. Furthermore, provenance analysis provides new insights to intoreconstruct the paleotectonic evolutiontectono-paleogeography and reveals the early Mesozoic evolution of multiple tectonic regimes in the northeastern NCC.

2 Geological background

2.1 Geological background

The study area is located at in the northeastern NCC, at the intersection between three main tectonic domains (Fig. 1): the Paleo-Asian Tectonic Regime (the eastern Central Asian Orogenic Belt (CAOB)) to the north, and the Su–Lu Orogenic Belt to the south, and the CircumPaleo-Pacific Tectonic RegimeOrogen to the east. The eastern CAOB, which is also referred to as the Xing'an–Mongolia Orogenic Belt (XMOB), and consists of a collage of microcontinental massifs, which are (fFrom southeast to northwest), they are the Khanka, Jiamusi, Songnen–Zhangguangcai Range, Xing'an, and Erguna massifs (Sengör and Natal'in, 1996; Li et al., 1999; Jahn et al., 2004; Li, 2006; Fig. 1). The Su–Lu Orogenic Belt and its eastward extension, (the Jing–Ji Orogenic Belt (JJOB) in Korea,) formed during the early Mesozoic subduction and collision between the NCC and the YC.

The tectonic evolution of the northeastern NCC is characterized by early Paleozoic arc–continent collision in the early Paleozoic (Pei et al., 2014), late Paleozoic subduction of the Paleo-Asian oceanic plate beneath the NCC (Cao et al., 2012), and middle–late Permian to Middle Triassic closure of the Paleo-Asian Ocean (Sun et al., 2004; Li et al., 2009a; Wang et al.,

2015b). ~~Furthermore~~During the early Mesozoic, the northeastern NCC was influenced by subduction, collision, and subsequent rapid exhumation between the NCC and YC ~~in the early Mesozoic~~ (Zheng et al., 2003; Yang et al., 2007; Pei et al., 2008; Liu et al., 2009). Subduction of the Paleo-Pacific Plate beneath Eurasia controlled the Mesozoic tectonic evolution of the East Asian continental margin (Xu et al., 2013). The Dunhua–Mishan and Yitong–Yilan faults occur in the northwestern part of the study area.

The northeastern NCC is composed primarily of Archean and Paleoproterozoic metamorphic basement (including the An'shan, Ji'an, and Laoling groups), Neoproterozoic and Paleozoic sedimentary cover sequences, and several Mesozoic basins (referred to as the Fusong–Changbai B~~b~~asin ~~group~~Group) (Fig. 1). Neoproterozoic strata comprise sandstone and limestone with minor stromatolites. Cambrian–Middle Ordovician strata are mainly epicontinental carbonate sediments, whereas late Carboniferous to early Permian units are characterized by marine and coal-bearing terrestrial sequences~~, which~~ that are unconformably overlain by Mesozoic volcano-sedimentary formations. ~~Due to uplift of the craton in the middle Paleozoic, t~~The region lacks Silurian–Devonian and early Carboniferous strata due to uplift of the craton during the middle Paleozoic (SBGMR, 1989). Cenozoic basalts are common in the eastern part of the study area (JBGMR, 1988, 1997; Fig. 2).

The interior of the NCC contains widespread Neoarchean to Paleoproterozoic magmatic rocks (Wu et al., 2007b), but lacks evidence ~~for~~ of late Neoproterozoic or Paleozoic magmatism (except for early Paleozoic kimberlite) (JBGMR, 1988; LBGMR, 1989; IMBGMR, 1991). However, the northern margin of the NCC contains abundant Paleozoic igneous rocks (Zhang et al., 2004~~; Zhang et al.~~, 2010; Cao et al., 2013; Pei et al., 2016). In the northeastern NCC, Paleozoic igneous rocks are concentrated to the north of the study area and are characterized by negative (–20 to –2) and positive (0 to +17, ~~a few with~~with minor –2 to 0) $\varepsilon_{Hf}(t)$ values within the NCC and XMOB (Yang et al., 2006), respectively ~~(Yang et al., 2006)~~. The northern NCC contains Ordovician (467 Ma, dated by laser ablation–inductively coupled plasma–mass spectrometry, LA–ICP–MS) medium-K calc-alkaline pyroxene andesites and middle Permian (270 Ma, LA–ICP–MS) garnet-bearing monzogranites with zircon $\varepsilon_{Hf}(t)$ values of –5.96 to –2.43 (Pei et al., 2016) and –17.1 to –14.1 (Cao et al., 2013), respectively. In contrast, the XMOB contains late Cambrian (493 Ma, LA–ICP–MS) low-K tholeiitic meta-diabase and late Permian (260 Ma, LA–ICP–MS) biotite monzogranite that yield zircon $\varepsilon_{Hf}(t)$ values of +9.42 to +14.89 (Pei et al., 2016) and +8.31 to +9.80 (Wang et al., 2015b), respectively.

Mesozoic magmatism was widespread along the East Asian continental margin (including the northeastern NCC and eastern XMOB). Within the northeastern NCC, Mesozoic magmatism occurred during the Late Triassic, Early Jurassic, Late Jurassic, Early Cretaceous, and Late Cretaceous (Yu et al., 2009; Xu et al., 2013; Zhang et al., 2014; Wang et al., 2017). Mesozoic igneous rocks in the NCC and XMOB yield negative (–20 to –2) and positive (0 to +17, ~~a few~~ with ~~minor~~few values of –2 to 0) zircon $\varepsilon_{Hf}(t)$ values (Yang et al., 2006; Pei et al., 2008; Wang et al., 2015b), respectively ~~(Yang et al., 2006; Pei et al., 2008; Wang et al., 2015b)~~.

~~2.2 Early Mesozoic basin-filling sedimentary sequence~~

Mesozoic strata are well preserved and exposed in the study area, and are characterized by coal-bearing volcano-

sedimentary formations containing abundant animal and plant fossils. Studies of Mesozoic strata in the northeastern NCC and the establishment of their lithostratigraphic sequence began during geological surveying in the 1960s–1970s. ~~In the 1990s, t~~These strata were then reclassified and a general lithostratigraphic sequence was established by JBGMR (1997). Early Mesozoic strata in the study area include, from bottom to top, the Heisonggou ~~Fm~~, Changbai ~~Fm~~, Xiaoyingzi ~~Fm~~, and Yihe ~~Fm~~ formations (Fig. 3). However, the ~~formation~~ timing of deposition and the stratigraphic sequence of these Triassic–Early Jurassic ~~units~~ formations are based mainly on lithostratigraphic correlations and remain ~~controversial~~debated. The Fusong–Changbai Basin Group represents a series of small- to medium-sized basins in the northeastern NCC. In this study, we focus on three small basins ~~known as,~~ (from north to south);~~,~~ the Fusong, Yihe, and Yantonggou basins (Figs~~.~~ 2~~,~~–3). The Fusong Basin is located ~~within~~ around Songshu Village~~,~~ and is filled by the Late Triassic Xiaoyingzi Formation (~~herein,~~ "Formation" is hereafter abbreviated to "Fm"), which is overlain by the Early Cretaceous Guosong Fm. The Yihe Basin, located ~~in~~ around Yihe Village, is filled by the Late Triassic Changbai ~~Fm~~ and ~~the~~ Early Jurassic Yihe Formations. The Yantonggou Basin is located ~~in~~ around Heisonggou Village and is filled by the Early Triassic Heisonggou Fm. These strata represent the early Mesozoic sedimentary sequence of the Fusong–Changbai Basin ~~group~~Group, and record deposition ~~that was~~ controlled by the tectonics of the northeastern NCC. The lithostratigraphic units are described in detail below (Fig. 3).

The Heisonggou Formation comprises conglomerate, sandstone, siltstone, and shale, and contains plant fossils. The stratotype is exposed in Heisonggou Village, strikes E–W, and extends westward into Korea (Fig. 2, section 3). The stratotype displays faulted contacts with the under- and overlying Mesoproterozoic strata (JBGMR, 1963, 1976). The basal conglomerate contains clasts of quartzite, marble, phyllite, and schist derived from the underlying Mesoproterozoic units. Sandstone, siltstone, and shale are concentrated within the middle and upper strata. The upper part of the section is intruded by andesites (Fig. 4).

The Changbai Formation comprises a lower member of andesites and andesitic volcaniclastics, and an upper member of rhyolite and rhyolitic volcaniclastics. The stratotype is located in Naozhi Village. The formation overlies Paleoproterozoic strata with a faulted contact, and is unconformably overlain by the Early Jurassic Yihe Fm (JBGMR, 1997). The formation strikes E–W and is also observed in Naozhi and Yihe villages (Fig. 2, section 2).

The Xiaoyingzi Formation comprises a ~30-m-thick lower member of conglomerate and sandstone (representing a complete depositional cycle) and an upper member of sandstone, siltstone, shale, mudstone, and coal. The conglomerate contains clasts of stromatolite-bearing dolomite that ~~was~~ were sourced primarily from Neoproterozoic units. Diabase porphyrite ~~is~~ commonly ~~observed to~~ intrudes along bedding planes. A thin layer of tuffaceous siltstone occurs in the middle of the section, and abundant plant and animal fossils are preserved within the middle and upper parts of the formation (Fig. 5). Volcanics and organic-rich beds have been used to constrain the ~~timing of formation~~ age of the Xiaoyingzi Fm. The stratotype is exposed in Xiaoyingzi Village and the strata strike NW–SE (Fig. 2, section 1). The base of the section is not exposed and the formation is overlain by Early Cretaceous volcanics of the ~~(~~Guosong Fm~~)~~ (JBGMR, 1963, 1976; Fig. 5).

The Yihe Formation comprises dominantly conglomerate, sandstone, siltstone, shale, coal, and minor tuffaceous siltstone (JBGMR, 1976, 1997, 1998; Fig. 6). ~~The Yihe Fm~~ It unconformably overlies andesites of the Late Triassic

Changbai Fm, which is present as gravels within the basal conglomerate of the Yihe Fm, thereby illustrating a conformable relationship between the two formations. The stratotype is exposed in Yihe Village and the strata strike E–W (Fig. 2, section 2).

3 Materials and Methods

The Fusong–Changbai Bbasin ~~group~~ Group studied here is chosen from a series of ~~the~~ early Mesozoic faulted basins in the northeastern NCC. ~~The t~~Three stratotypes ~~from~~ of the early Mesozoic sedimentary strata ~~are~~ were ~~explored~~studied, ~~and~~ which are~~they~~ exposed in Xiaoyingzi ~~Village~~ (Xiaoyingzi Fm: E42°01'54.26", N127°10'03.89"), Naozhi ~~Village~~ (Yihe Fm: E41°56'41.17", N127°04'05.67"), and Heisonggou ~~Village~~ (Heisonggou Fm: E41°47'34.79", N126°58'33.10") villages ~~respectively~~ (~~details of the sections in Geological background~~see section 2 for further details). ~~Detrital and magmatic zircons from 10 samples are collected from these sections.~~

Stratigraphic logs~~Both of the stratigraphy~~ and zircon U–Pb geochronology and Hf isotopic constraints~~isotope geochronology~~ are used to study the age and provenance of the sedimentary strata. The research rationale is based on the following steps~~The process is as follows~~: (1) the stratotypes ~~are~~ were measured in the field ~~in order~~ to understand the lithostratigraphic succession and to guide the sampling~~sequence accurately and obtain more exact sampling bed~~; (2) detrital and magmatic zircon geochronological data, ~~together~~combined with biostratigraphic data from previous studies, and stratigraphic correlation are used to establish a new stratigraphic framework for the northeastern NCC; (3) detrital zircon age distributions~~in this precise chronostratigraphic framework, the relative probability of detrital zircons~~ and Hf isotopic ~~analyses~~ compositions are used to identify the source of sedimentary strata and (4) based on provenance ~~analysis~~data, we reconstruct~~ion of~~ and discuss the paleography and ~~–~~evolution of multiple tectonic regimes in the northeastern NCC~~ are discussed~~.

3.1 Sample descriptions

For zircon U–Pb dating, we collected detrital and magmatic zircons from 10 early Mesozoic sandstone and igneous samples (Fig. 3). Three samples were collected from the Heisonggou Fm, one from the Changbai Fm, three from the Xiaoyingzi Fm, two from the Yihe Fm, and one from the Guosong Fm, which overlies the Xiaoyingzi Fm. Details of their stratigraphy and petrography are presented in Figures 4–7 and ~~are~~ described below.

Sample 16LJ6-1 is a medium-grained feldspathic quartz sandstone from the lower Heisonggou Fm (Fig. 4). The sample is gray–white in color, displays clastic texture, and ~~is~~ ~~has~~shows bedded structure. Grains ~~range~~ are ~~in size from~~ 0.3 ~~to~~ –0.6 mm in size and comprise plagioclase and alkali-feldspar (~14 vol.%), quartz (~78 vol.%), lithic fragments (~2 vol.%), and matrix (~5 vol.%) (Fig. 7a).

Sample 15LJ4-11 is a fine-grained feldspathic quartz sandstone from the upper Heisonggou Fm. The sample is gray–white in color, exhibits clastic texture, and ~~has~~shows bedded structure. Grains are angular–subangular and ~~range in size from~~ 0.1 ~~to~~ –0.2 mm in size, comprising plagioclase and alkali-feldspar (~13 vol.%), quartz (~78 vol.%), lithic fragments (~2

vol.%), and calcareous cement (~6 vol.%) (Fig. 7b).

Sample 15LJ4-6 is an andesite that intrudes the Heisonggou Fm (Fig. 4). It is light gray–green in color, displays pilotaxitic texture, and ~~hasis~~ shows massive structure (Fig. 7c).

Sample 15JFS1-1 is a medium-grained feldspathic quartz sandstone from the lower Xiaoyingzi Fm (Fig. 5). ~~The sample~~It is yellow–white in color, displays clastic texture, and ~~is~~ shows bedded structure. Grains ~~range in size from~~are 0.4 ~~to~~ –0.8 mm in size and comprise quartz (~80 vol.%), plagioclase and alkali-feldspar (~12 vol.%), lithic fragments (~3 vol.%, volcanic fragments), and matrix (~4 vol.%) (Fig. 7d).

Sample 15JFS2-1 was collected from a diabase porphyrite intruding ~~into~~ the Xiaoyingzi Fm (Fig. 5). It is gray–green in color, displays porphyritic texture, and ~~hasis~~shows massive structure. The phenocrysts are dominantly plagioclase (~5 vol.%) and the ~~matrix~~ groundmass displays pilotaxitic texture (Fig. 7e).

Sample 16LJ8-1 is a tuffaceous siltstone from the middle Xiaoyingzi Fm (Fig. 5). The sample is white in color, displays clastic texture, and ~~is~~ shows bedded structure.

Sample 15JFS10-1 is a pyroxene andesite that unconformably overlies the Xiaoyingzi Fm. (Fig. 5). The andesite belongs to the Guosong Fm and is gray–green in color, displays porphyritic texture, and shows massive structure (Fig. 7f).

Sample 16LJ1-1 is an andesite from the Changbai Fm, which underlies the Yihe Fm. The andesite is gray–green in color, displays porphyritic texture, and ~~hasis~~shows massive structure. The phenocrysts are mainly plagioclase (~10 vol~~.%).~~ .%) and t~~T~~he ~~matrix~~ groundmass exhibits pilotaxitic texture (Fig. 7g).

Sample 15LJ1-2 is a fine-grained feldspathic quartz sandstone from the middle Yihe Fm (Fig. 6). It is gray–white in color, displays clastic texture, and ~~is~~ shows bedded structure. ~~Grains comprise~~ The rock comprises plagioclase and alkali-feldspar (~13 vol.%), quartz (~78 vol.%), lithic fragments (~3 vol.%), and matrix (~5 vol.%).

Sample16LJ3-1 is a tuffaceous siltstone from the upper Yihe Fm. It is white in color, displays clastic texture, and ~~is~~ shows bedded structure (Fig. 7h).

3.2 Analytical methods

3.2.1 Zircon U–Pb dating

Zircons were separated from samples using ~~the~~ conventional heavy liquid and magnetic techniques, and purified by handpicking under a binocular microscope at the Langfang Yantuo Geological Survey, Langfang, Hebei Province, China. The handpicked zircons were examined under transmitted and reflected-light ~~with~~ using an optical microscope, and ~~in order to reveal their internal structures,~~ cathodoluminescence (CL) images were obtained to reveal their internal structures, using a JEOL scanning electron microscope housed at the State Key Laboratory of Continental Dynamics, Northwest University, Xi'an China. Based on CL images, ~~Distinct~~ distinct domains within the zircons were selected for analysis~~, based on their CL images~~. An Agilent 7500a ICP–MS equipped with a 193 nm laser, housed at the State Key Laboratory of Geological Processes and Mineral Resources, China University of Geosciences, ~~(~~Wuhan, China~~)~~, was used to measure the U–Pb ages of zircons. Zircon 91500 was used as external standard for age calibration and the NIST SRM 610 silicate glass was ~~applied~~

used for the instrument optimization. The An crater analysisThe spot diameter was ofwas 32 μm was used during for the analyses. The For details of the instrumental analytical parameters and detail procedures followed, were seedescribed by Yuan et al. (2004). The ICPMSDataCal 7.0 (Liu et al., 2010) and Isoplot 3.0 (Ludwig, 2003) programs were used for data analysis and plottingreduction. Correction for common Pb was made following Andersen (2002). Errors on individual analyses by LA-–ICP-–MS analyses are quatedquoted at the 1σ level, while errors on pooled ages are quoted at the 95% (2σ) confidence level. The dating results are presented in Table S1.

Due to the paucity of zircons recovered from samples of 15JFS2-1 (diabase porphyrite) and 15JFS10-1 (pyroxene andesite), provide on 52 and 32 zircons grains, respectively. Considering the few zircon grains, measurement for theanalysis of these two samples were was conducted using a Cameca 1280 secondary ion mass spectrometer (SIMS) at the Institute of Geology and Geophysics, Chinese Academy of Sciences, in Beijing, China, using operating and data processing procedures similar to those described by Li et al. (2009b).

3.2.2 Hf isotopic analyses

In situ zircon Hf isotope analyses were conducted using a Neptune Plus multi-collector (MC)–ICP–MS (Thermo Fisher Scientific, Germany) equipped with a 193 nm excimer ArF laser ablation system (Lambda Physik, Göttingen, Germany) that was hosted at the State Key Laboratory of Geological Processes and Mineral Resources, China University of Geosciences. The Laser ablation analyses used an energy density of laser ablation that was used in this study was 5.3 J cm^{-2}. Helium was used as the carrier gas within the ablation cell and was merged with argon (makeup gas) after the ablation cell. A simple Y-Y-junction downstream from the sample was used to add small amounts of nitrogen (4 ml min^{-1}) to the argon makeup gas flow (Hu et al., 2008a, 2008b). Compared with the standard arrangement, the addition of nitrogen, in combination with the use of a newly designed X skimmer and Jet sample cones in the Neptune Plus instrument, improved the signal intensities of Hf, Yb, and Lu by factors of 5.3, 4.0, and 2.4, respectively. All data were acquired using ain single single-spot ablation mode with a 44 μm spot size. Each measurement consisted of 20 s of acquisition of the background signal followed by 50 s of ablation signal acquisition. Details For further details of the operating conditions for the laser ablation system and the MC–ICP–MS instrument, and analytical method procedures followed, are given insee Hu et al. (2012). The dating results are presented in Table S2.

4 Analytical results

4.1 Zircon U–Pb dating

4.1.1 Heisonggou Formation

Zircons from sample 16LJ6-1 are euhedral–subhedral, and display fine oscillatory growth zoning in cathodoluminescence (CL) images (Fig. 8a).), Majority of zZircons and show yield Th/U values of 0.07-–1.31(>0.1),

indicating a magmatic origin (Corfu et al., 2003). A total of 68 spots provide ages ranging from 2485 to 248 Ma (ages of >1000 Ma use $^{207}Pb/^{206}Pb$ ages, whereas ages of <1000 Ma use $^{206}Pb/^{238}U$ ages), yielding four major age populations ~~that give~~ with ~~weighted~~ weighted-mean ages of 252 ± 1 Ma (MSWD = 1, n = 21), 293 ± 2 Ma (MSWD = 1.2, n = 10), 323 ± 2 Ma (MSWD = 0.39, n = 15), and 2402 ± 12 Ma (MSWD = 8.5, n = 14) (Fig. 9a). Other grains yield ages of 339 (2 grains), 594, 1757, 1791, 2006, 2075, and 2171 Ma (Table S1). The youngest zircon age population ~~age~~ of 252 ± 1 Ma constrains the maximum depositional age of this sample (i.e., the medium-grained feldspathic quartz sandstone was deposited after ~252 Ma).

Zircon grains from sample 15LJ4-11 are euhedral–subhedral, ~~and~~ display fine oscillatory growth zoning in CL images (Fig. 8b), ~~. Z~~Majority ~~of zircons~~ and ~~show~~yield ~~high~~ Th/U ~~ratios, indicating a magmatic origin~~values~~ratios~~ of 0.01–4.31. Some grains are round in shape and exhibit a core–rim texture. A total of 72 spots provide ages ranging from 2832 to 248 Ma and ~~yield~~ nine age populations with ~~weighted~~ weighted-mean ages of 253 ± 3 Ma (MSWD = 0.4, n = 7), 270 ± 3 Ma (MSWD = 0.24, n = 7), 304 ± 3 Ma (MSWD = 0.38, n = 5), 323 ± 3 Ma (MSWD = 0.12, n = 6), 360 ± 5 Ma (MSWD = 0.35, n = 3), 382 ± 7 Ma (MSWD = 0.03, n = 3), 1845 ± 9 Ma (MSWD = 1.7, n = 14), 2337 ± 23 Ma (MSWD = 4.8, n = 4), and 2504 ± 8 Ma (MSWD = 4.0, n = 20) (Fig. 9b). Other grains yield ages of 426, 2152, and 2838 Ma (Table S1). The youngest ~~population~~ zircon age population of 253 ± 3 Ma represents the maximum depositional age of the fine-grained feldspathic quartz sandstone.

Zircons from sample 15LJ4-6 are euhedral–subhedral and display fine oscillatory growth zoning in CL images (Fig. 8c), with ~~high~~ Th/U ~~ratio~~ values of 0.07–1.03~~s, indicating a magmatic origin~~. A total of 20 spots provide ages ranging from 2435 to 243 ~~to 2435~~ Ma, ~~yielding~~ with two age populations with ~~weighted~~ weighted-mean ages of 246 ± 2 Ma (MSWD = 1.6, n = 8) and 297 ± 3 Ma (MSWD = 2.2, n = 6) (Fig. 9c). Other grains yield~~ed~~ ages of 268, 313, 1788, 1819, 2411, and 2435 Ma (Table S1). The mean age of 246 ± 2 Ma is interpreted as the crystallization age ~~for~~ of the andesite, whereas the other ages are interpreted ~~to~~ as captured zircons.

4.1.2 Changbai Formation

Zircon grains from the andesite (sample 16LJ1-1) are euhedral–subhedral and display typical oscillatory growth zoning in CL images (Fig. 8d), with ~~their~~ Th/U values ~~ratios~~ of 0.46–1.06. Twenty-five analyses yield~~ed~~ a ~~weighted~~ weighted-mean $^{206}Pb/^{238}U$ age of 227 ± 1 Ma (MSWD = 0.74, n = 25) (Fig. 9d; Table S1), which is interpreted as the crystallization age of the andesite.

4.1.3 Xiaoyingzi Formation

Zircons from sample 15JFS1-1 are euhedral–subhedral and show fine oscillatory growth zoning in CL images, with ~~high~~ Th/U ~~ratios~~values of 0.05–1.58~~(>0.1), indicating a magmatic origin~~. Some rare grains are rounded in shape and display fine oscillatory zoning in CL images (Fig. 8e). A total of 79 spots provide ages ranging from 3285 to 220~~ to 3285~~ Ma, yielding five age populations with ~~weighted~~ weighted-mean ages of 224 ± 3 Ma (MSWD = 0.39, n = 7), 232 ± 5 Ma

(MSWD = 0.27, n = 3), 257 ± 5 Ma (MSWD = 2.3, n = 3), 1880 ± 3 Ma (MSWD = 2.9, n = 42), and 1982 ± 21 Ma (MSWD = 1.2, n = 4) (Fig. 9e). Other grains yielded ages of 275, 277, 294, 296, 315, 364, 369, 392, 397, 433, 453, 496, 569, 691, 708, 1768, 2242, 2434, 2500, and 3285 Ma (Table S1). The youngest zircon age population of 224 ± 3 Ma constrains the maximum depositional age of the medium-grained feldspathic quartz sandstone.

Zircons from the diabase porphyrite (sample 15JFS2-1) are euhedral–subhedral and dominantly display mainly oscillatory growth zoning, although some grains exhibit core–rim structures (Fig. 8f). A total of 17 spots provide ages ranging from 2516 to 111 to 2516 Ma, with a peak zircon age population at 113 ± 2 Ma (MSWD = 3.4, n = 5) (Fig. 9f). Other grains yield ages of 173, 218, 270, 332, 744, 1862, 1884, 1887, 1891, 2059, 2412, and 2516 Ma (Table S1). The youngest weighted-mean $^{206}Pb/^{238}U$ age of 113 ± 2 Ma is considered to represent the age of intrusion of the diabase porphyrite.

Zircon grains from the tuffaceous siltstone (sample 16LJ8-1) are typically euhedral–subhedral and display oscillatory growth zoning in CL images, although some grains are rounded and/or subrounded (Fig. 8g). Zircons show yield high Th/U ratiovalues of 0.04–1.2(>0.1). A total of 65 spots provide ages ranging from 2602 to 223 to 2602 Ma, yielding defining six zircon age populations with weighted weighted-mean ages of 227 ± 2 Ma (MSWD = 4.1, n = 3), 259 ± 2 Ma (MSWD = 4.5, n = 5), 1770 ± 9 Ma (MSWD = 1.9, n = 18), 1831 ± 9 Ma (MSWD = 2.1, n = 16), 2212 ± 18 Ma (MSWD = 7.7, n = 5), and 2486 ± 15 Ma (MSWD = 8.3, n = 7) (Fig. 9g). Other grains yielded ages of 240 (two grains), 316, 330, 505, 914, 1469, 1950, 2023, 2061, and 2602 Ma (Table S1). The youngest $^{206}Pb/^{238}U$ age of 223 ± 2 Ma is interpreted as the maximum depositional age of the tuffaceous siltstone.

4.1.4 Yihe Formation

Zircon grains from sample 15LJ1-2 are primarily euhedral–subhedral and display fine oscillatory zoning in CL images, whereas other grains are rounded and display oscillatory zoning (Fig. 8h), which, together with their Th/U ratiosvalues of 0.11–1.72, indicates a magmatic origin (Table S1). A total of 77 spots provide ages ranging from 2472 to –177 to 2472 Ma, yielding six zircon age populations with weighted weighted-mean ages of 184 ± 2 Ma (MSWD = 1.7, n = 7), 233 ± 3 Ma (MSWD = 0.82, n = 7), 245 ± 1 Ma (MSWD = 1.17, n = 17), 254 ± 2 Ma (MSWD = 0.72, n = 20), 266 ± 2 Ma (MSWD = 1, n = 11), and 1831 ± 17 Ma (MSWD = 9.9, n = 4) (Fig. 9h). Other grains yielded ages of 191, 200, 211, 286, 293, 323, 363, 460, 1468, 2471, and 2472 Ma (Table S1). These results indicate that the fine-grained feldspathic quartz sandstone was deposited after 184 Ma.

Zircon grains from sample 16LJ3-1 are typically euhedral–subhedral and display fine oscillatory growth zoning in CL images (Fig. 8i), ; thatwhich, together with their Th/U ratiosvalues of 0.26–1.16, are indicative of, indicating a magmatic origin. A total of 65 spots provide ages ranging from 2477 to 178 to 2477 Ma, yielding three zircon age populations with weighted weighted-mean ages of 182 ± 1 Ma (MSWD = 3.3, n = 20), 253 ± 1 Ma (MSWD = 0.74, n = 33), and 1831 ± 20 Ma (MSWD = 1.8, n = 4) (Fig. 9i). Other grains yielded ages of 212, 222, 237, 263, 276, 340, 2457, and 2477 Ma (Table S1). The youngest age population of 182 ± 1 Ma represents the maximum depositional age of the tuffaceous siltstone.

4.1.5 Guosong Formation

Zircon grains from the pyroxene andesite (15JFS10-1) are euhedral–subhedral and display oscillatory growth zoning and striped absorption in CL images (Fig. 8j), with ~~their~~the ~~Th/U ratios~~values of 0.19–2.05~~, implying a magmatic origin (Fig. 8j)~~. A total of 14 spots provide ages ranging from 1647 to ~~112 to 1647 Ma, yielding~~ defining two zircon age populations with ~~weighted~~ weighted-mean ages of 113 ± 3 Ma (MSWD = 1.19, n = 2) and 227 ± 3 Ma (MSWD = 0.74, n = 7) (Fig. 9j). Other grains yield~~ed~~ ages of 156, 425, 438, 946, and 1647 Ma (Table S1). The youngest age of 113 ± 3 Ma is interpreted ~~to represent~~as the crystallization age of the pyroxene andesite.

4.2 Zircon Hf isotopes

We performed *in situ* Hf isotopic analysis on the same spots as ~~as~~ used for U–Pb dating on samples from the Heisonggou ~~Fm~~ (15LJ4-11), Xiaoyingzi ~~Fm~~ (15JFS1-1), and Yihe ~~Fm~~ (15LJ1-2) formations. The results are listed in Table S2 and shown in Fig. 10.

4.2.1 Heisonggou Formation

We determined the Hf isotopic compositions of 22 detrital zircons from the Heisonggou Fm. Three zircon grains (~13.6%) with ages of 2513–2466 Ma yielded $\varepsilon_{Hf}(t)$ values of –4.2 to +4.9, four Paleoproterozoic (1862–1812 Ma) grains (~18.2%) yielded $\varepsilon_{Hf}(t)$ values of –1.9 to +0.2, and fifteen Phanerozoic (383–250 Ma) grains (~68.2%) yielded $\varepsilon_{Hf}(t)$ values of –16.3 to –7.0 (Table S2; Fig. 10a) and two-stage model (T_{DM2}) ages of 2.3–1.7 Ga, similar to ~~those of~~ zircons from Phanerozoic igneous rocks in the NCC (Yang et al., 2006).

4.2.2 Xiaoyingzi Formation

We determined the Hf isotopic compositions of 25 detrital zircons with ages of 1928–221 Ma from the Xiaoyingzi Fm. Most zircon grains (~88%) yielded negative $\varepsilon_{Hf}(t)$ values ~~ranging from~~of –21.9 to –0.2 and T_{DM2} ages of 2.7–1.6 Ga, whereas three zircon grains (~12%) with ages of 569, 496, and 453 Ma yielded positive $\varepsilon_{Hf}(t)$ values ~~ranging from~~of +2.9 to +8.5 and T_{DM2} ages of 1.2–0.98 Ga (Table S2; Fig. 10b). Two grains within the latter, with ages of 708 and 691 Ma, yielded $\varepsilon_{Hf}(t)$ values of –10.9 and –6.1 and T_{DM2} ages of 2.3 and 2.0 Ga, respectively.

4.2.3 Yihe Formation

We determined the Hf isotopic compositions of 17 detrital zircons with ages of 255–177 Ma from the Yihe Fm. Most grains (~88%) yielded positive $\varepsilon_{Hf}(t)$ values of +2.4 to +12.6 and T_{DM2} ages of 1.1–0.4 Ga, whereas two zircon grains (~12%) with ages of 248 Ma yielded negative $\varepsilon_{Hf}(t)$ values of –5.8 and –1.9 and T_{DM2} ages of 1.6 and 1.4 Ga (Table S2; Fig. 10c).

5 Discussion

5.1 Age of early Mesozoic strata in the northeastern NCC

~~We now~~Here we combine the ages of youngest concordant detrital zircons ~~ages~~, ~~the ages of~~ interbedded volcanic and intrusive rocks, ~~biostratigraphic ages,~~ and ~~the ages of~~ overlying strata, as well as biostratigraphy, to constrain the ages of ~~the~~ Mesozoic strata in the northeastern NCC.

5.1.1 Heisonggou Formation

The Heisougou Fm was first ~~established~~ defined in Heisonggou Village and was assigned to the Early Jurassic (JBGMR, 1963). It was subsequently reclassified as the Shiren Fm and assigned to the Early Cretaceous (JBGMR, 1997). Therefore, the age of the Heisonggou Fm remains uncertain. In this study, samples from the lower (16LJ6-1) and upper (15LJ4-11) Heisonggou Fm yield youngest zircon age populations of 252 ± 1 Ma and 253 ± 3 Ma, respectively, indicating that deposition of the Heisonggou Fm occurred after 252 ± 1 Ma. Furthermore, andesite intruding this unit (Fig. 4, sample 15LJ4-6) yield~~sed~~ a ~~weighted~~ weighted-mean age of 246 ± 2 Ma, thereby constraining the deposition of the Heisonggou Fm to between 252 and 246 Ma. This Early Triassic age contrasts with the previously proposed Early Jurassic (JBGMR, 1963) and Early Cretaceous ~~ages~~ (JBGMR, 1997) ages.

5.1.2 Changbai Formation

The Changbai Fm was first ~~established~~ defined in ~~the~~ Caiyuanzi and Ergulazi villages and was assigned to the Late Triassic. The formation comprises intermediate–acidic volcanics and has been ~~subdivided~~ reassigned into the Ergulazi ~~Fm~~ and Naozhigou ~~Fm~~ formations (JBGMR, 1976, 1997). Zircon U–Pb dating of andesite (16LJ1-1) from the Changbai Fm yielded a ~~weighted~~ weighted-mean age of 227 ± 1 Ma. Combined with a previously reported zircon $^{206}Pb/^{238}U$ age of 222 ± 1 Ma (Yu et al., 2009) from the Naozhigou Fm in the northern area of Caiyuanzi Village, we confirm that the Changbai Fm was deposited in the Late Triassic.

5.1.3 Xiaoyingzi Formation

The Xiaoyingzi Fm was first ~~established~~ defined in ~~the~~ Hengdaohezi and Xiaoyingzi villages and assigned to the Early Jurassic (JBGMR, 1971). It was later correlated with the Xiaohekou Fm and assigned to the Late Triassic (JBGMR, 1988, 1997). In contrast, JBGMR (2007) classified ~~it~~ placed within the Yihe Fm and assigned ~~it to the~~an Early Jurassic age. Thus, the age of the Xiaoyingzi Fm remains controversial. Detrital zircon grains from the medium-grained feldspathic quartz sandstone (15JFS1-1) and tuffaceous siltstone (16LJ8-1) yielded youngest concordant ages of 224 ± 2 Ma and 223 ± 2 Ma, respectively, ~~suggesting~~ indicating that the Xiaoyingzi Fm was deposited after ~223 Ma. Freshwater bivalve fossils (e.g., *Ferganoconcha* sp. and *Sibireconcha* sp.) in this formation belong to the *Unio–Shaanxiconcha* assemblage (Zhu, 1991; JBGMR, 1997), which is similar to assemblages found in Late Triassic strata in China, South Australia, North American,

and South Africa (Zhu, 1991). Plant fossil assemblages (e.g., *Glossophyllum-Neocalamites*) in this formation are generally limited to the Late Triassic–Early Jurassic (JBGMR, 1997). We therefore conclude that the Xiaoyingzi Fm was deposited in the Late Triassic. The previously described "interbedded volcanic rocks" within the Xiaoyingzi Fm (JBGMR, 1971) are here reinterpreted as diabase porphyrite (15JFS2-1), which was ~~was~~ intruded at 113 ± 2 Ma. Zircon grains from the pyroxene andesite (15JFS10-1) of the Guosong Fm also yield an ~~formation~~ age of 113 ± 3 Ma. These results indicate that the Guosong Fm, which overlies the Xiaoyingzi Fm, and the diabase porphyrite have similar ages and were produced during coeval magmatism.

5.1.4 Yihe Formation

The Yihe Formation was first ~~established~~ defined in ~~the~~ Naozhigou and Yihe villages and was assigned to the Early Jurassic (JBGMR, 1976). Previous studies referred to this unit as the Yantonggou ~~Fm~~ and Shiren ~~Fm~~ formations~~(EGJS, 1975)~~, ~~and~~ which were assigned ~~it~~ to the Late Jurassic and Early Cretaceous (EGJS, 1975), respectively. Detrital zircon grains from the fine-grained feldspathic quartz sandstone (15LJ1-2) and tuffaceous siltstone (16LJ3-1) yielded youngest age populations of 184 ± 2 Ma and 182 ± 1 Ma, respectively. In addition, Early Jurassic plant fossils such as *Cladophlebis ukienesis*, *Marattia hoerensis*, and *Pterophyllum propinquum* are observed within the Yihe Fm (Si and Zhou, 1962). Based on ~~Thus~~these observations, combined with the absence of Middle–Late Jurassic strata in ~~the~~ southern Jilin Province, we conclude that the Yihe Fm was deposited in the Early Jurassic.

~~We use o~~Our new geochronological data, field relationships, and biostratigraphic data are used to establish a new early Mesozoic stratigraphic framework for the northeastern NCC (Fig. 3), which is summarized as follows. The Heisonggou Fm is assigned to the Early Triassic (252–246 Ma) based on the U–Pb ages of detrital and magmatic zircons. The Xiaoyingzi Fm is assigned to the Late Triassic based on zircon geochronological data and fossil assemblages. Deposition of this unit ~~was later than~~post-dated the Changbai Fm. The Yihe Fm was deposited in the Early Jurassic, consistent with biostratigraphic data (Si and Zhou, 1962).

5.2 Stratigraphic correlation with early Mesozoic strata in the northern NCC

Triassic–Middle Jurassic strata are well preserved and exposed in the northern NCC (Meng et al., 2013; Meng, 2017), and were deposited within early Mesozoic basins in the Yinshan–Yanshan O~~o~~rogenic ~~belt~~ Belt (e.g., the Beipiao, Xiabancheng, Jingxi, and Shiguaizi basins) (Fig. 11; Meng et al., 2018). The northern NCC shared a similar stratigraphic and sedimentological evolution to the interior of the craton ~~in~~ during the Mesoproterozoic to Paleozoic, but has undergone a ~~a~~ distinct evolution since the Mesozoic, involving alternating periods of contraction and extension (Davis et al., 2001; Cui et al., 2002).

Triassic–Middle Jurassic deposits in the northern NCC are dominated by volcano-sedimentary formations containing coal and abundant animal and plant fossils, and their stratigraphy, geochronology, and depositional processes are well constrained (Meng, 2003; Meng et al., 2011, 2014; Li and Huang, 2013; Li et al., 2015; Liu et al., 2015; Xu et al., 2016;

Meng, 2017). Here, we determine the relationship between early Mesozoic strata in the northern and northeastern NCC.

The late Permian Tiechang Fm in the northeastern NCC is characterized by gray–purple coarse-grained sandstone, pebbly sandstone, siltstone, and shale, which is a similar lithological association to to those ofthe late Permian strata of the northern NCC (e.g., the lower Tiechang Fm correlates with the Shihezi Fm, Naobaogou Fm, and Laowuopu Fmformations; Fig. 11). In contrast, Early Triassic strata (e.g., the Heisonggou Fm) are only observed in the northeastern NCC. In the northern NCC, Middle Triassic strata are rare and only observed within the central area (e.g., the Xiabancheng Basin), and are characterized by sandstone, siltstone, and shale. Such strata are not observed in the northeastern NCC. Late Triassic strata (e.g., the Changbai Fm and Xiaoyingzi Fmformations) are commonly observed in the northeastern NCC, and comprise intermediate–acidic volcanics and coal-bearing clastic sediments. In the northern NCC, coeval and comparable strata are represented by the Xiaolanwuo Fm, Wuchang Fm, and Shanggu Fm formations in the Xiabancheng Basin, as well as by the Yangcaogou Fm in the Beipiao Basin. For example, the Late Triassic Xiaoyingzi Fm in the northeastern NCC contains a similar lithological association and fossil assemblages to the Yangcaogou Fm and Shanggu Fm formations in the northern NCC. Furthermore, ages of volcanics in the Changbai Fm (227 and 222 Ma; Yu et al., 2009) are similar to those of volcanics in the Xiaolanwuo Fm (225 ± 1 Ma) and Wuchang Fm (227.6 ± 2 Ma) (Meng et al., 2018) formations in the Xiabancheng Basin. The Early Jurassic Yihe Fm in the northeastern NCC is comparable to the lower Beipiao Fm, Xiahuayuan Fm, lower Yaopo Fm, and Zhaogou Fm formations in the northern NCC, as indicated by the presence of similar coal-bearing layers and plant fossil assemblages (JBGMR, 1988; LBGMR, 1989; HBGMR, 1989; BBGMR, 1991; IMBGMR, 1991).

In summary, we suggest that similar late Paleozoic sedimentary successionsformation (North China type) occur in both the northern and northeastern NCC. However, Early Triassic strata are only observed within the northeastern NCC. In contrast, Middle Triassic strata are observed within the northern NCC but are absent in the northeastern NCC. Similar Late Triassic sedimentary formations and contemporaneous volcanics are observed in the northern and northeastern NCC. Earliest Jurassic strata are absent in the northern and northeastern NCC, whereas middle–late Early Jurassic strata are widespread in both regions, and are correlated through thebased on observation of similar coal-bearing layers and plant fossils (JBGMR, 1988; LBGMR, 1989; HBGMR, 1989; BBGMR, 1991; IMBGMR, 1991). Furthermore, regional unconformities are identified between late Permian and Triassic strata, and Late Triassic and Jurassic strata in the northern and northeastern NCC (Fig. 11), respectively.

5.3 Provenance of early Mesozoic strata in the northeastern NCC

A generally accepted method to identify the source of sedimentary units is to compare zircon U–Pb age and Hf isotopic data with areas or units that may have supplied sediment to the region (Dickinson and Gehrels, 2008). Our 502 detrital zircon U–Pb analyses are grouped into three age populations:, namely Neoarchean to Paleoproterozoic, Neoproterozoic, and Phanerozoic (Fig. 12). Magmatic zircon grains from igneous and sedimentary rocks within the northern margin of the NCC generally yield two age populations (of Neoarchean to Paleoproterozoic and late Paleozoic) age populations (Yang et al., 2006; Zhang et al., 2010), whereas igneous rocks and Paleozoic sediments from the southern margin of the XMOB contain

mainly Phanerozoic and minor Neoproterozoic zircon grains (Meng et al., 2010; Wang et al., 2012b~~; Wang et al.~~, 2014). In addition, Phanerozoic zircons from the NCC typically yield negative $\varepsilon_{Hf}(t)$ values, whereas those from the southern margin of the XMOB typically yield positive $\varepsilon_{Hf}(t)$ values (Yang et al., 2006; Cao et al., 2013). These data ~~from 6 samples~~ are inevitably affected by the sampling ~~probability~~strategy, so the given ~~precise~~ percentage data ~~can't~~should be considered qualitatively rather than quantitatively ~~result in the study area~~.

5.3.1 Early Triassic Heisonggou Formation

Approximately 58% of detrital zircon grains from the Early Triassic Heisonggou Fm yield Neoarchean to Paleoproterozoic ages (2.8–1.8 Ga), forming two age peaks at ~2.5 and ~1.8 Ga, which are typical of the NCC (Ma and Wu, 1981; Zhao et al., 2001; Gao et al., 2004). In contrast, the ~42% magmatic zircon grains yield Phanerozoic ages (426–248 Ma). These Phanerozoic magmatic events are not recorded within the interior of the NCC (Yang et al., 2017), but are instead observed within the northern margin of the NCC (Zhang et al., 2004; Zhang et al., 2010; Wu et al., 2011; Cao et al., 2013; Pei et al., 2014; Wang et al., 2015b~~; Wang et al.~~, 2016). These observations, combined with the~~ir~~ negative $\varepsilon_{Hf}(t)$ values (–16.3 to –7.0) and two-stage model (T_{DM2}) ages of 2.3–1.7 Ga, indicate that all sedimentary material within this formation was sourced from the NCC.

In addition, ~~except for some rounded Neoarchean and Paleoproterozoic grains,~~ detrital zircon grains (especially those with Phanerozoic ages) are euhedral–subhedral (with the exception of some rounded Neoarchean and Paleoproterozoic grains), ~~suggesting~~ indicating a proximal source. This interpretation is also supported by the short deposition time (between 252 and 246 Ma) of the Heisonggou Fm.

Based on the present distribution of the NCC basement and Phanerozoic igneous rocks in the northeastern NCC (JBGMR, 1988), we suggest that at least 42% of the sediment was sourced from Phanerozoic igneous rocks along the northern margin of the NCC, to the north of the basin. In contrast, Neoarchean and Paleoproterozoic sediment was likely sourced from regions surrounding the basin.

5.3.2 Late Triassic Xiaoyingzi Formation

Approximately 71% of detrital zircon grains from the Late Triassic Xiaoyingzi Fm yielded Neoarchean to Paleoproterozoic ages (with peaks at ~2.50 and ~1.88 Ga), indicating that they were sourced from the NCC (Gao et al., 2004). The other analyzed detrital zircon grains yielded ~26% Phanerozoic (496–220 Ma) and ~3% Neoproterozoic ages (914–569 Ma). Three zircon grains (with ages ~~at~~ of 569, 496, and 453 Ma) yielded positive $\varepsilon_{Hf}(t)$ values (+2.9 to +8.5), whereas the other Phanerozoic and Neoproterozoic grains yielded negative $\varepsilon_{Hf}(t)$ values (–21.9 to –5.0), ~~suggesting~~ indicating that the former were sourced from the XMOB and the latter from the NCC (Yang et al., 2006; Cao et al., 2013; Pei et al., 2014).

All Phanerozoic zircons are euhedral–subhedral, suggesting the Xiaoyingzi Fm sediments were not transported over long

distances. Combined with the observation of Phanerozoic and minor Neoproterozoic igneous rocks to the north of the basin (JBGMR, 1988), we conclude that ~29% of sediments from the Xiaoyingzi Fm were sourced from an area to the north of the basin, with the remaining ~71% sourced from regions surrounding the basin.

5.3.3 Early Jurassic Yihe Formation

Approximately 91% of detrital zircon grains from the Early Jurassic Yihe Fm yielded Phanerozoic ages (460–177 Ma), with the remaining ~9% yielding Neoarchean to Paleoproterozoic ages (peaks at ~2.5 and ~1.8 Ga). The former are consistent with Phanerozoic magmatism along the northern margin of the NCC and XMOB, whereas the latter are typical of the NCC interior (Gao et al., 2004; Zhang et al., 2004; Yang et al., 2006; Cao et al., 2013; Wang et al., 2015b). ~~The Hf isotopic analyses o~~Of the 17 zircon grains analyzed with ~~the~~ Phanerozoic ages, ~~indicate that~~ two grains (~12%) with ages of 248 Ma yielded $\varepsilon_{Hf}(t)$ values of –5.8 to –1.9, and the remaining 15 grains (~88%) ~~with Phanerozoic ages~~ yielded $\varepsilon_{Hf}(t)$ values of +2.4 to +12.6. We infer that the former, together with Neoarchean and Paleoproterozoic detrital zircon grains, were sourced from the NCC, whereas the latter were derived from the XMOB (Yang et al., 2006). Furthermore, all Phanerozoic zircon grains are euhedral–subhedral, ~~suggesting~~ indicating that the Phanerozoic sediments did not undergo long-distance transport. The age populations and $\varepsilon_{Hf}(t)$ values of the detrital zircon grains indicate that the Yihe Fm was sourced mainly from the XMOB to the north of the basin.

5.4 Implications for the early Mesozoic paleotectonic evolution ~~tectono-paleogeography~~ of the northeastern NCC

The early Mesozoic tectonic evolution of the northeastern NCC was not only ~~was~~ influenced by subduction of the Paleo-Asian oceanic plate beneath the NCC and~~,~~ final closure of the Paleo-Asian Ocean to the north (Zhang et al., 2004; Wu et al., 2011; Cao et al., 2013), but also by subduction and collision between the NCC and the YC to the south (Pei et al., 2008, 2011), as well as by subduction of the Paleo-Pacific Plate beneath Eurasia to the east (Xu et al., 2013; Guo, 2016; Wang et al., 2017). However, the spatio-temporal extents of these influences, and the timing of final closure of the Paleo-Asian Ocean and the onset of subduction of the Paleo-Pacific Plate remain controversial (Sun et al., 2004; Shi, 2006; Ernst et al., 2007; Li et al., 2007; Wu et al., 2011; Xu et al., 2012, 2013; Cao et al., 2013; Sun et al., 2015; ~~Zhou and Li, 2017;~~ Wang et al., 2017; Zhou and Li, 2017). Here we use provenance data and interpretations of changes in provenance within early Mesozoic strata of the northeastern NCC to reconstruct the early Mesozoic paleotectonic evolution~~tectono-paleogeography~~ of the northeastern NCC.

5.4.1 Early Triassic: southward subduction of the Paleo-Asian oceanic plate, and subduction and collision between the NCC and YC

Approximately 42% of the Early Triassic Heisonggou Fm sediments were sourced from the northern margin of the NCC; there is no evidence ~~is seen for the~~of XMOB-sourced grains. We therefore infer that ~~in the Early Triassic,~~ southward subduction of the Paleo-Asian oceanic plate beneath the NCC was ongoing in the Early Triassic (i.e., final closure of the

Paleo-Asian Ocean had not yet occurred) (Cao et al., 2013; Wang et al., 2015b). Subduction resulted in uplift along the northern margin of the NCC, producing a paleogeographic highland that acted as a sediment source during deposition of the Heisonggou Fm (Fig. 13a). The remaining ~58% of sediments of the Heisonggou Fm were sourced from Neoarchean and Paleoproterozoic NCC basement, which is observed in areas surrounding the basin, as well as to the south. We infer that the area to the south of the basin was uplifted in the Early Triassic, consistent with subduction and collision between the NCC and YC at this time (Pei et al., 2008, 2011; Liu et al., 2012; Zheng et al., 2013). Thus, we conclude that southward subduction of the Paleo-Asian oceanic plate and subduction and collision between the NCC and YC resulted in uplift of the northern and southeastern margins of the NCC, respectively. These uplifted highlands provided sediment source areass for the Heisonggou Fm (Figs. 11, 13a).

5.4.2 Late Triassic: final closure of the Paleo-Asian Ocean and post-collisional exhumation of the Su–Lu Orogenic Belt

Zircon Hf isotopic compositions indicate that ~29% of sediments of the Late Triassic Xiaoyingzi Fm were sourced from areas to the north of the basin (~4% from the XMOB and ~25% from the northern margin of the NCC). The presence of the XMOB material indicates that the Paleo-Asian Ocean had closed by the Late Triassic. We therefore suggest that final closure of the ocean occurred in the Middle Triassic, which is also supported by the occurrence of Middle Triassic syn-collisional granitoids in the Yanbian region (Wang et al., 2015b), as well as the absence of Middle Triassic sedimentary strata in the northeast NCC (JBGMR, 1988). Approximately 71% of the Xiaoyingzi Fm zircon grains yield Neoarchean and Paleoproterozoic ages, suggesting indicating that they were sourced from the NCC basement. As the NCC basement dominantly crops out mainly to the south of the basin, we conclude that thise region to the south of the basin, which was still was a highland at the time, provided the primary source during deposition of the Xiaoyingzi Fm. Furthermore, we suggest that Late Triassic rapid exhumation of ultrahigh-pressure metamorphic rocks within the Su–Lu Orogenic Belt in the Late Triassic during the Late Triassic (Zhao et al., 2001; Zheng et al., 2013) produced the inferred uplift in the region to the south of the basin (Figs 11, 13b).

5.4.3 Early Jurassic: rapid uplift of the XMOB and the onset of subduction of the Paleo-Pacific Plate beneath Eurasia

The provenance of the Early Jurassic Yihe Fm changed rapidly from ~71% (the Late Triassic Xiaoyingzi Fm) Neoarchean and Paleoproterozoic basement (the Late Triassic Xiaoyingzi Fm) from an area to the south of the basin in the Late Triassic, to ~91% Phanerozoic rocks from an area to the north of the basin in the Early Jurassic. In addition, Hf isotopic compositions of Phanerozoic detrital zircon grains indicate that >88% of the grains were sourced from the XMOB. In other words, these data show that the depositdepositionthe sediments were was sourced mainly from the NCC in the Early–Late Triassic, and had changed to to the depositdeposition sourced dominantlymainly from the XMOB in the Early Jurassic. The This obvious change in provenance can be directly explained by the change of paleogeographyic changes, consequently suggesting indicating that rapid uplift of the XMOB occurred during the Early Jurassic (Figs. 11, 13c). In aAdditionally, the

paleogeographic changes ~~change of paleogeography could be~~were possibly triggered by a change ~~of~~ in tectonic domains, and the uplift in the northeastern NCC ~~could be also be~~might also have been related to the initial subduction of the Paleo-Pacific Plate beneath Eurasia in the Early Jurassic. This interpretation is also supported by the presence of the Early Jurassic accretionary complex and Early Jurassic calc-alkaline igneous rocks in the eastern Asian continental margin. Firstly, the Early Jurassic igneous rocks ~~occurred~~ in the eastern Asian continental margin (including ~~the~~ eastern Heilongjiang–Jilin provinces, northeastern North Korea, and south Korea) belong ~~chemically~~ to the calc-alkaline series (Xu et al., 2013; Guo et al., 2016; Wang et al., 2017; Tang et al., 2018; Fig. 1), whereas the contemporaneous intracontinental igneous rocks ~~within intracontinent~~ (~~such as~~e.g., ~~within~~those of the Songnen–Zhangguangcai Range Massif) consist of a ~~suite of with~~ bimodal igneous rock association (Yu et al., 2012; Xu et al., 2013). The former ~~reveals~~ is consistent with an active continental margin setting (Gill, 1981), whereas the latter ~~implies~~ indicates an extensional environment (Yu et al., 2012). ~~From continental margin to intracontinent, t~~The polar variations ~~of~~ in K_2O contents ~~in~~ of the continental margin to intracontinental Early Jurassic igneous rocks ~~reveals~~are indicative of the initial subduction of the Paleo-Pacific ~~plate~~ Plate beneath ~~the~~ Eurasia (Xu et al., 2013; Tang et al., 2018). Secondly, the widespread occurrence of ~~the~~ Early Jurassic accretionary complexes in the eastern Asian continental margin (including ~~NE~~ northeastern China, Japan, and the Russian Far East) also ~~reveals~~ indicates ~~the happening of~~ subduction of the Paleo-Pacific ~~plate~~ Plate beneath ~~the~~ Eurasia (Wu et al., 2007; Zhou et al., 2009; Fukuyama et al., 2013; Safonova and Santosh, 2014). Taken together, we conclude that ~~the~~ uplift of the northeastern NCC ~~could be also be~~was related to the onset of subduction of the Paleo-Pacific Plate beneath ~~the~~ Eurasia in the Early Jurassic (Wang et al., 2016).

In summary, Early Triassic deposition in the northeastern NCC was controlled by the southward subduction of the Paleo-Asian oceanic plate, as well as by subduction and collision between the NCC and the YC (Fig. 13a). The absence of Middle Triassic strata in the northeastern NCC ~~suggests~~ indicates the final closure of the Paleo-Asian Ocean ~~at~~ during this time, which is also supported by the provenance of the Late Triassic deposits. The presence of ~~the~~ XMOB material in Late Triassic deposits suggests that the Paleo-Asian Ocean had closed by this time. Rapid exhumation of the Su–Lu Orogenic Belt may have resulted in the formation of a paleogeographic highland within the southeastern margin of the NCC (Fig. 13b). A sudden change in provenance ~~is~~ was triggered by the rapid uplift of the XMOB in the Early Jurassic and~~/or~~ was likely related to the onset of subduction of the Paleo-Pacific Plate beneath Eurasia at this time (Fig. 13c).

6 Conclusions

Based on the U–Pb ages of detrital and magmatic zircons, detrital zircon Hf isotopic data, and biostratigraphic records from early Mesozoic strata of the northeastern NCC, we draw the following conclusions.

(1) The early Mesozoic stratigraphic sequence of the northeastern NCC comprises, from bottom to top, the Early Triassic (252–246 Ma) Heisonggou Formation, Late Triassic Changbai Formation (~227 Ma), Late Triassic Xiaoyingzi Formation, and Early Jurassic Yihe Formation.

(2) The ~~provenance of~~ Early Triassic Heisonggou Formation was ~~all~~ sourced from the NCC, with ~42% of sediment

being derived from an area to the north of the basin and ~58% from the area surrounding the basin. Approximately 88% of sediments of the Late Triassic Xiaoyingzi Fm were sourced from the NCC, whereas the remaining ~12% were derived from the XMOB. In contrast, >88% of sediments of the Early Jurassic Yihe Fm were sourced from the XMOB, with only ~12% being sourced from the NCC.

(3) Early Triassic deposition was controlled by both the southward subduction of the Paleo-Asian oceanic plate beneath the NCC and the northward subduction and collision between the NCC and YC. The Late Triassic deposition ~~could be~~was possibly related to final closure of the Paleo-Asian Ocean and rapid exhumation of the Su–Lu Orogenic Belt between the NCC and YC. The sudden change in provenance recorded by Early Jurassic sediments, together with the coeval calc-alkaline volcanism and accretionary complex, ~~implies~~ indicates the rapid uplift of the XMOB and~~/or~~ the onset of subduction of the Paleo-Pacific Plate beneath Eurasia in the Early Jurassic.

(4) Final closure of the Paleo-Asian Ocean likely occurred in the Middle Triassic, which is consistent with the lack of Middle Triassic strata in the northeastern NCC and the ~~observations~~ occurrence of Middle Triassic syn-collisional granitoids along the Changchun–Yanji Ssuture ~~belt~~Belt.

Data availability. Original data underlying the material presented are available by contacting the authors.

Supplements. Supplementary ~~information~~Information;~~-~~ Table S1; Table S2.

Author contributions. Yini Wang~~,~~ designed the ~~format~~ study and wrote the main content. Wenliang Xu designed the ~~entire~~ overall project and ~~produced~~ wrote the title. Feng Wang participated in ~~the~~ fieldwork and data analysis. Xiaobo Li modified and proofed the figures and captions.

Competing interests. The authors declare that they have no conflicts of interest.

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Figure captions

Figure 1: Geological sketch map of the northeastern North China Craton. JJOB: Jing-Ji Orogenic Belt.

Figure 2: Geological map of the Linjiang area showing ~~the~~ early Mesozoic strata and the three study sections.

Figure 3: Early Mesozoic stratigraphy of the northeastern NCC, sampling sites, and dating results. The ages refer to the youngest age populations of samples.

Figure 4: Stratotype outcrop of the Heisonggou Formation and sampling locations. The location of the stratotype is indicated by the dotted line in Fig. 2 (section 3). The numbers in circles show the sequence of beds in the section. The ages refer to the youngest zircon age populations of samples.

Figure 5: Stratotype outcrop of the Xiaoyingzi Formation and sampling locations. The location of the stratotype is indicated by the dotted line in Fig. 2 (section 1). The numbers in circles ~~show~~ indicate the sequence of beds in the section. The ages refer to the youngest zircon age populations of samples. The dotted line represents vegetation cover.

Figure 6: Stratotype outcrop of the Yihe Formation and sampling locations. The location of the stratotype is indicated by the dotted line in Fig. 2 (section 2). The numbers in circles ~~show~~ indicate the sequence of beds in the section. The ages refer to

the youngest zircon age populations of samples.

Figure 7: Photomicrographs of representative early Mesozoic samples (cross-polarized light). (a) Sample 16LJ6-1, a medium-grained feldspathic quartz sandstone ~~from~~ of the Heisonggou Fm; (b) Sample 15LJ4-11, a fine-grained feldspathic quartz sandstone ~~from~~ of the Heisonggou Fm; (c) Sample 15LJ4-6, an andesite that intrudes the Heisonggou Fm; (d) Sample 15JFS1-1, a medium-grained feldspathic quartz sandstone ~~from~~ of the Xiaoyingzi Fm; (e) Sample 15JFS2-1, a~~n~~ diabase porphyrite that intrudes the Xiaoyingzi Fm; (f) Sample 15JFS10-1, a pyroxene andesite ~~from~~ of the Guosong Fm; (g) Sample 16LJ1-1, an andesite ~~from~~ of the Changbai Fm; (h) Sample 16LJ3-1, a tuffaceous siltstone ~~from~~ of the Yihe Fm. Af: alkali-feldspar; Pl: plagioclase; Px: pyroxene; Q: quartz.

Figure 8: Cathodoluminescence (CL) images of selected zircon grains from early Mesozoic strata. White circles indicate the locations of U–Pb dating analyses and blue circles show the locations of *in situ* Hf analyses. Values under and ~~upper~~ above the images indicate ~~correspindingcorresponding~~ zircon U~~–~~–Pb ages and measured $\varepsilon_{Hf}(t)$ values, respectively.

Figure 9: U–Pb concordia diagram for zircon grains from the sedimentary and igneous rocks within early Mesozoic strata.

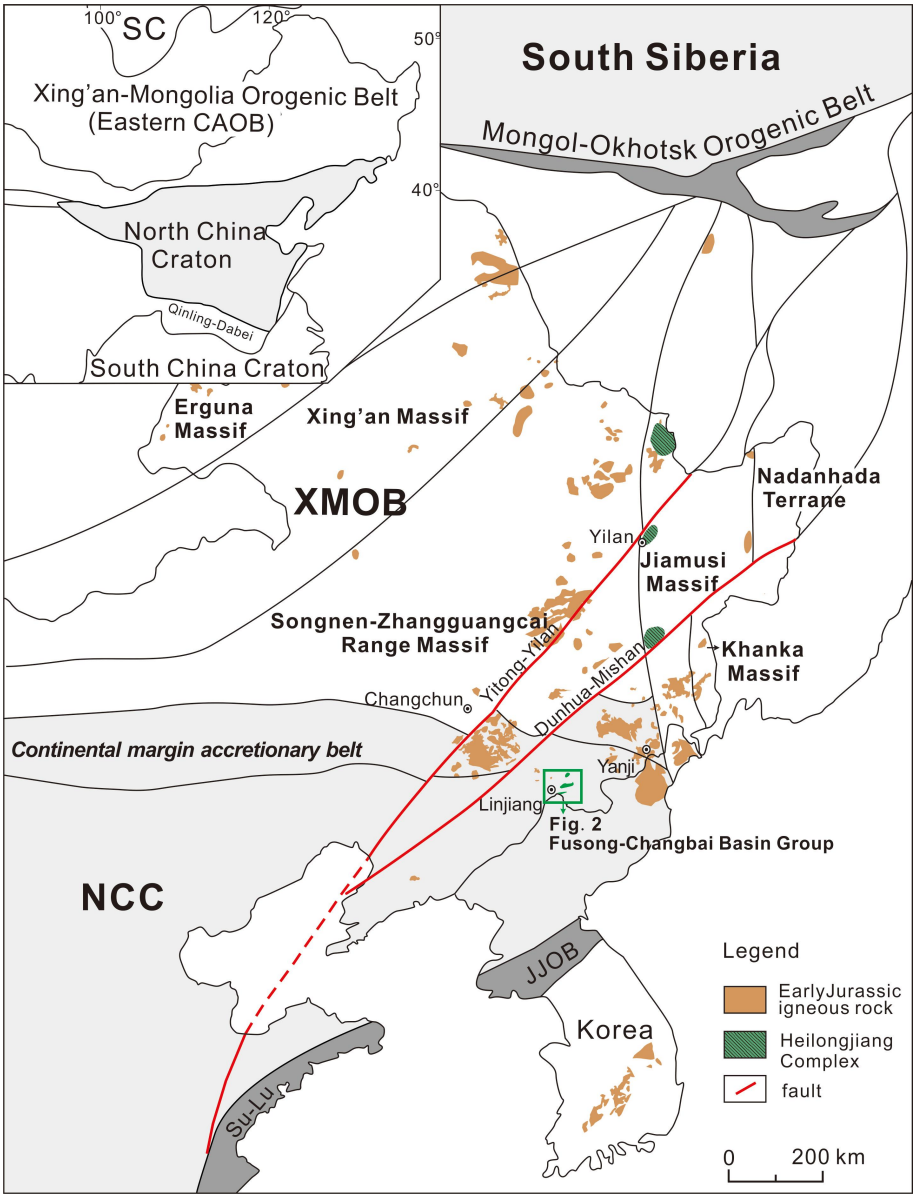
Figure 10: Hf isotopic compositions of detrital zircon grains from early Mesozoic strata of the northeastern NCC. XMOB: Xing'an–Mongolia Orogenic Belt; YFTB: Yanshan Fold-and-Thrust Belt (Yang et al., 2006).

Figure 11: Correlation between the early Mesozoic stratigraphy of the northern and northeastern NCC (the stratigraphic sequence of the northern NCC is from Meng et al., 2018).

Figure 12: Relative probability diagram of detrital zircon grain agess from early Mesozoic strata of the northeastern NCC.

Figure 13: paleotectonic evolution~~Tectono-paleogeography~~ of ~~the~~ northeastern China in the early Mesozoic, showing the paleogeographic evolution and provenance of early Mesozoic basins. (a) Early Triassic: southward subduction of the Paleo-Asian oceanic plate and subduction and collision between the NCC and YC; (b) Late Triassic: ~~the~~ final closure of the Paleo-Asian Ocean and post-collisional exhumation of the Su–Lu Orogenic Belt; (c) Early Jurassic: rapid uplift of the XMOB and the onset of subduction of the Paleo-Pacific Plate beneath Eurasia.

Figure 1


100°
SC
120°
50°
40°
Xing'an-Mongolia Orogenic Belt
(Eastern CAOB)
North China
Craton
Qinling-Dabei
South China Craton
South Siberia
Mongol-Okhotsk Orogenic Belt
Erguna
Massif
Xing'an Massif
XMOB
Nadanhada
Terrane
Yilan
Jiamusi
Massif
Songnen-Zhangguangcai
Range Massif
Yitong-Yilan
Dunhua-Mishan
Khanka
Massif
Changchun
Continental margin accretionary belt
Yanji
Linjiang
Fig. 2
Fusong-Changbai Basin Group
NCC
JJOB
Korea
Su-Lu
Legend
EarlyJurassic
igneous rock
Heilongjiang
Complex
fault
0
200 km

5

Figure 2

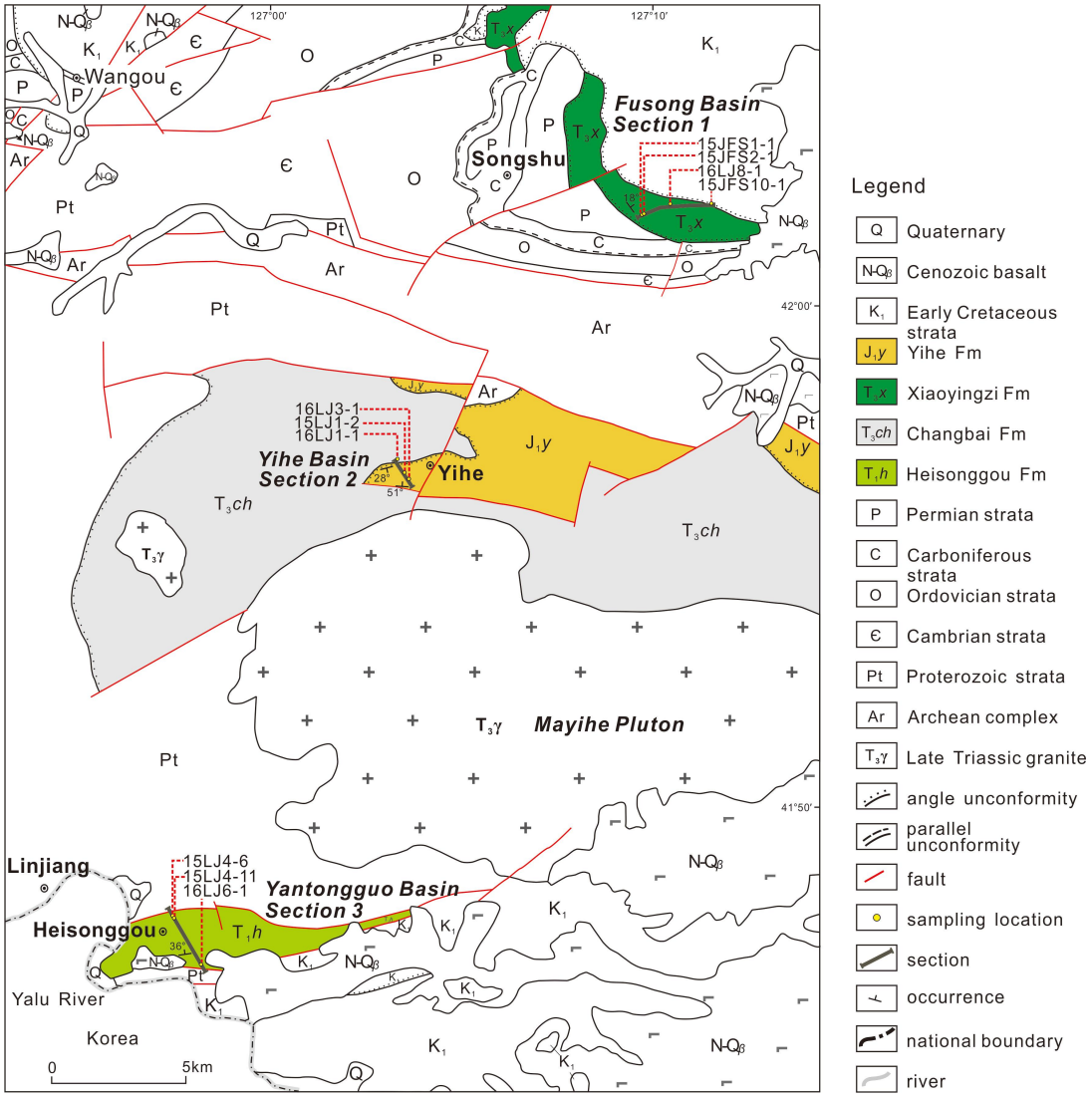

Wangou
Songshu
Fusong Basin
Section 1
15JFS1-1
15JFS2-1
16LJ8-1
15JFS10-1
Yihe Basin
Section 2
16LJ3-1
15LJ1-2
16LJ1-1
Yihe
Mayihe Pluton
Linjiang
Heisonggou
Yantongguo Basin
Section 3
15LJ4-6
15LJ4-11
16LJ6-1
Yalu River
Korea
0
5km
127°00′
127°10′
42°00′
41°50′
Legend
Quaternary
Cenozoic basalt
Early Cretaceous strata
Yihe Fm
Xiaoyingzi Fm
Changbai Fm
Heisonggou Fm
Permian strata
Carboniferous strata
Ordovician strata
Cambrian strata
Proterozoic strata
Archean complex
Late Triassic granite
angle unconformity
parallel unconformity
fault
sampling location
section
occurrence
national boundary
river

Figure 3

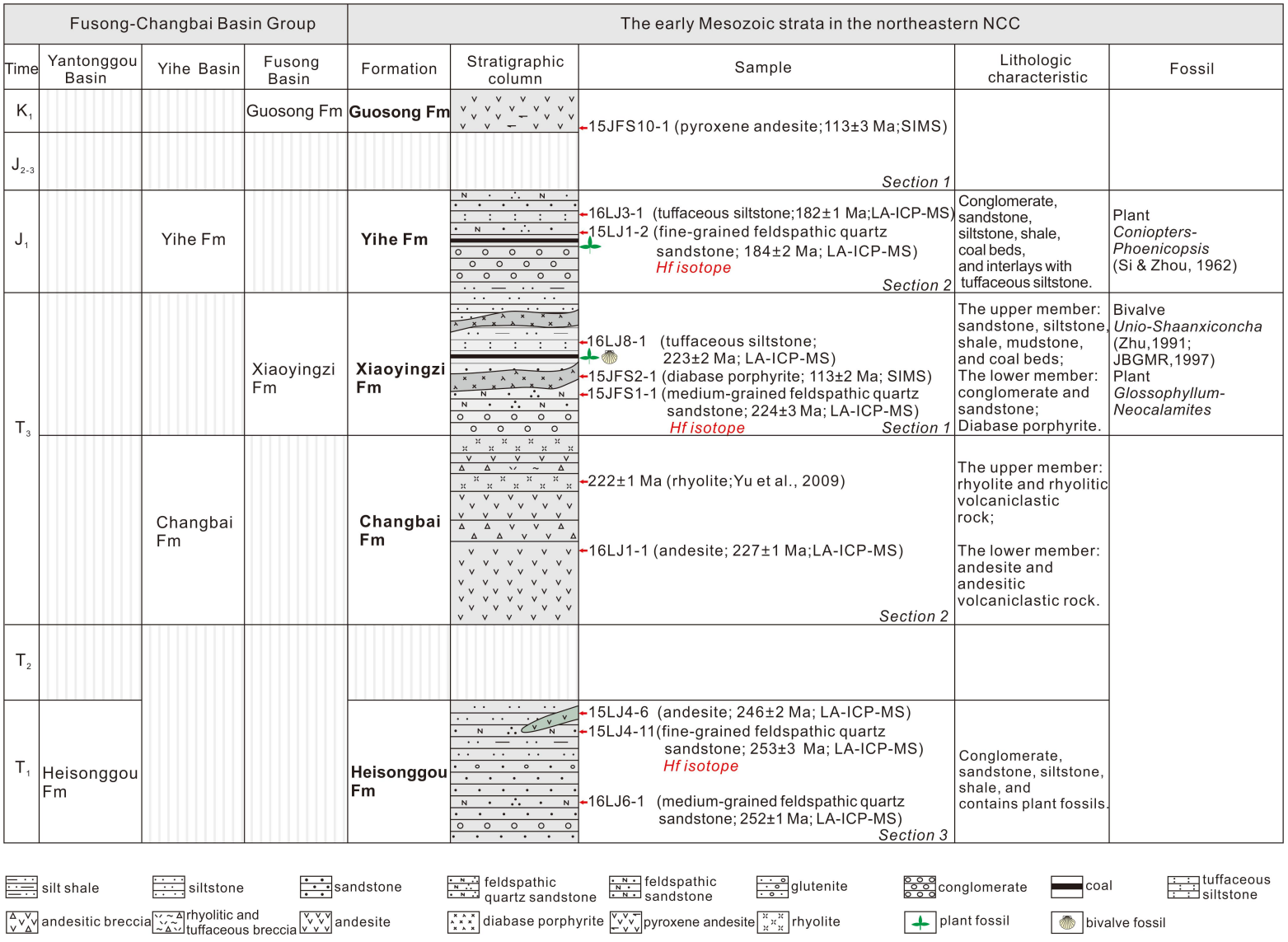

Fusong-Changbai Basin Group
The early Mesozoic strata in the northeastern NCC
Time
Yantonggou Basin
Yihe Basin
Fusong Basin
Formation
Stratigraphic column
Sample
Lithologic characteristic
Fossil
K1
J2-3
J1
T3
T2
T1
Guosong Fm
Yihe Fm
Xiaoyingzi Fm
Changbai Fm
Heisonggou Fm
15JFS10-1 (pyroxene andesite;113±3 Ma;SIMS)
Section 1
16LJ3-1 (tuffaceous siltstone;182±1 Ma;LA-ICP-MS)
15LJ1-2 (fine-grained feldspathic quartz sandstone; 184±2 Ma; LA-ICP-MS)
Hf isotope
Section 2
Conglomerate, sandstone, siltstone, shale, coal beds, and interlays with tuffaceous siltstone.
Plant Coniopters-Phoenicopsis (Si & Zhou, 1962)
16LJ8-1 (tuffaceous siltstone; 223±2 Ma; LA-ICP-MS)
15JFS2-1 (diabase porphyrite; 113±2 Ma; SIMS)
15JFS1-1 (medium-grained feldspathic quartz sandstone; 224±3 Ma; LA-ICP-MS)
Hf isotope
Section 1
The upper member: sandstone, siltstone, shale, mudstone, and coal beds;
The lower member: conglomerate and sandstone;
Diabase porphyrite.
Bivalve Unio-Shaanxiconcha (Zhu,1991; JBGMR,1997)
Plant Glossophyllum-Neocalamites
222±1 Ma (rhyolite;Yu et al., 2009)
16LJ1-1 (andesite; 227±1 Ma;LA-ICP-MS)
Section 2
The upper member: rhyolite and rhyolitic volcaniclastic rock;
The lower member: andesite and andesitic volcaniclastic rock.
15LJ4-6 (andesite; 246±2 Ma; LA-ICP-MS)
15LJ4-11(fine-grained feldspathic quartz sandstone; 253±3 Ma; LA-ICP-MS)
Hf isotope
16LJ6-1 (medium-grained feldspathic quartz sandstone; 252±1 Ma; LA-ICP-MS)
Section 3
Conglomerate, sandstone, siltstone, shale, and contains plant fossils.
silt shale
siltstone
sandstone
feldspathic quartz sandstone
feldspathic sandstone
glutenite
conglomerate
coal
tuffaceous siltstone
andesitic breccia
rhyolitic and tuffaceous breccia
andesite
diabase porphyrite
pyroxene andesite
rhyolite
plant fossil
bivalve fossil

Figure 4

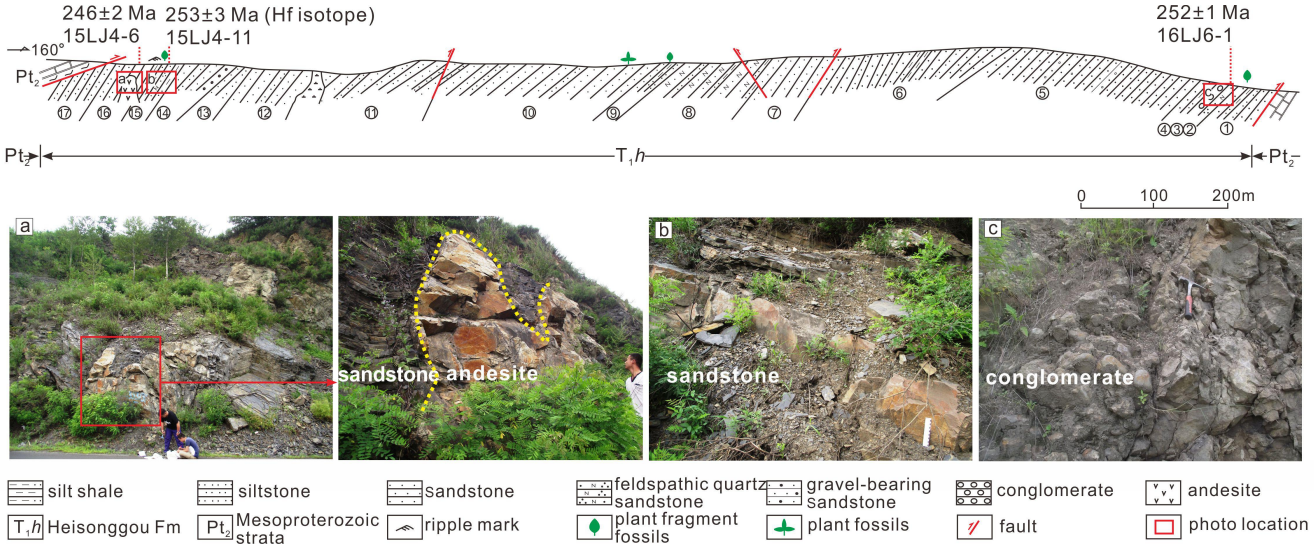

246±2 Ma
15LJ4-6
253±3 Ma (Hf isotope)
15LJ4-11
252±1 Ma
16LJ6-1
160°
Pt2
a
b
T1h
0 100 200m
sandstone andesite
sandstone
conglomerate
silt shale
siltstone
sandstone
feldspathic quartz sandstone
gravel-bearing sandstone
conglomerate
andesite
T1h Heisonggou Fm
Pt2 Mesoproterozoic strata
ripple mark
plant fragment fossils
plant fossils
fault
photo location

Figure 5

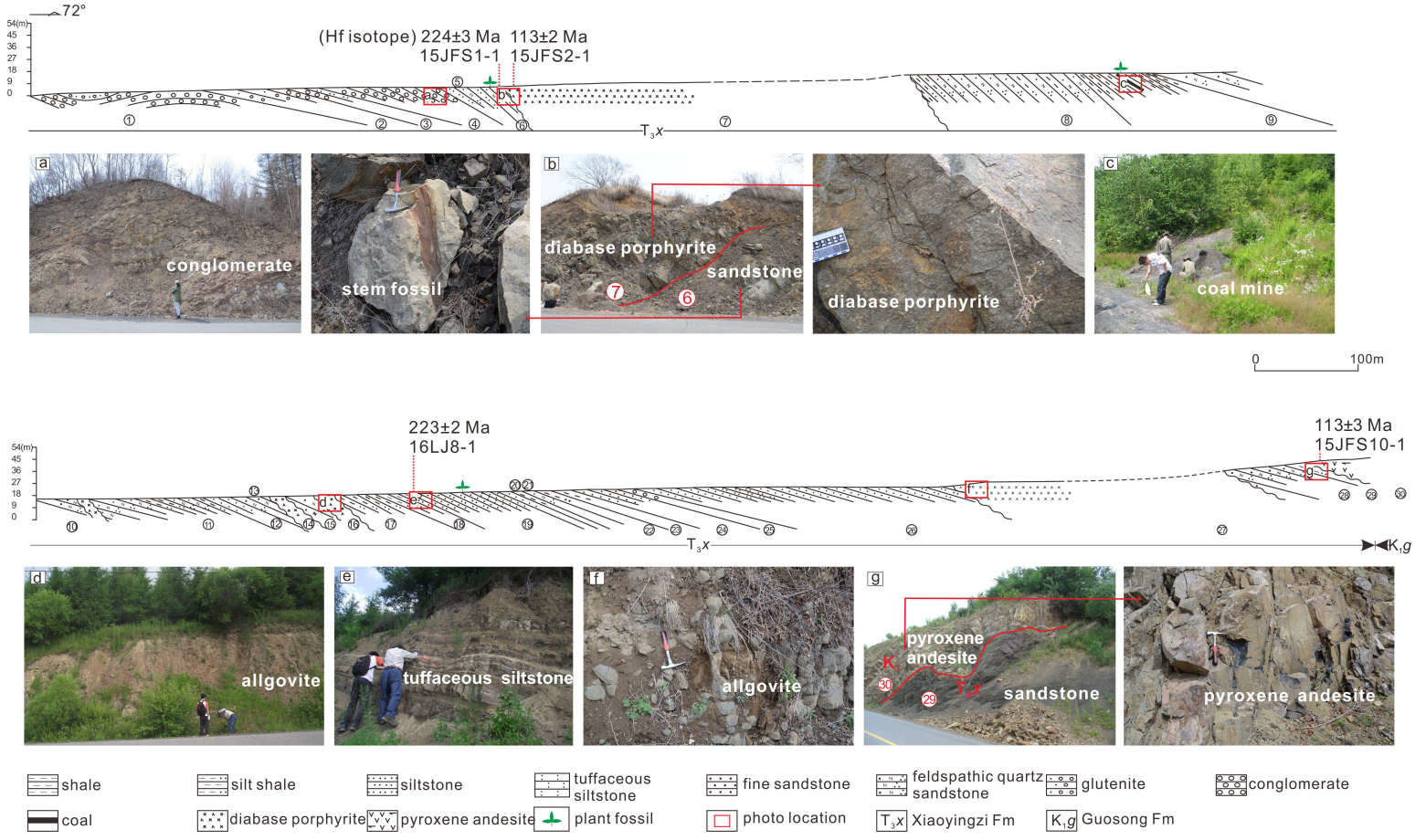
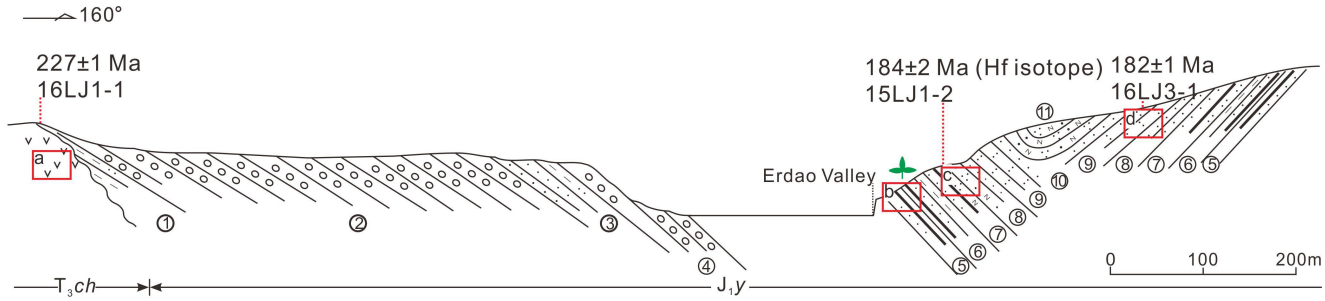
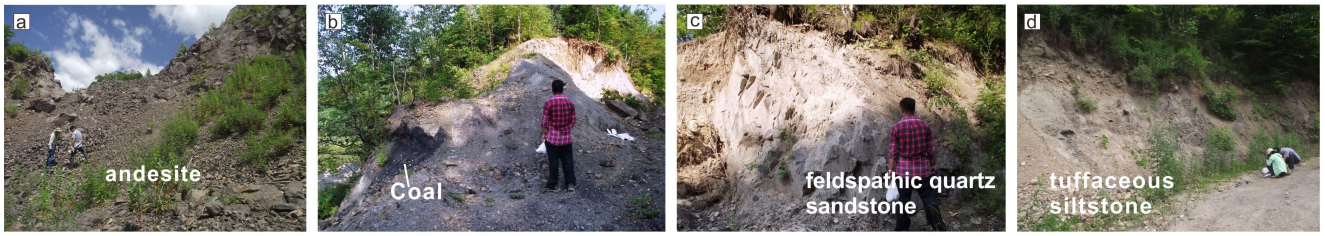

72°
54(m)
45
36
27
18
9
0
(Hf isotope) 224±3 Ma 113±2 Ma
15JFS1-1 15JFS2-1
T_3x
a
conglomerate
stem fossil
b
diabase porphyrite
sandstone
diabase porphyrite
c
coal mine
0 100m
223±2 Ma
16LJ8-1
113±3 Ma
15JFS10-1
K_1g
d
allgovite
e
tuffaceous siltstone
f
allgovite
g
pyroxene andesite
sandstone
pyroxene andesite
shale
silt shale
siltstone
tuffaceous siltstone
fine sandstone
feldspathic quartz sandstone
glutenite
conglomerate
coal
diabase porphyrite
pyroxene andesite
plant fossil
photo location
T_3x Xiaoyingzi Fm
K_1g Guosong Fm

Figure 6


160°
227±1 Ma
16LJ1-1
184±2 Ma (Hf isotope)
15LJ1-2
182±1 Ma
16LJ3-1
Erdao Valley
a
b
c
d
① ② ③ ④ ⑤ ⑥ ⑦ ⑧ ⑨ ⑩ ⑪
0 100 200m
T3ch
J1y


a
andesite
b
Coal
c
feldspathic quartz sandstone
d
tuffaceous siltstone

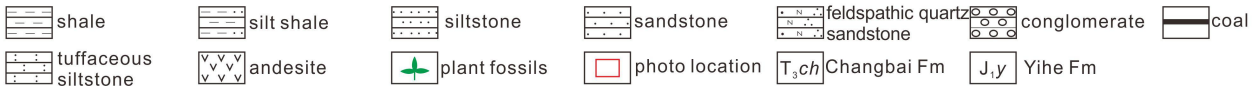
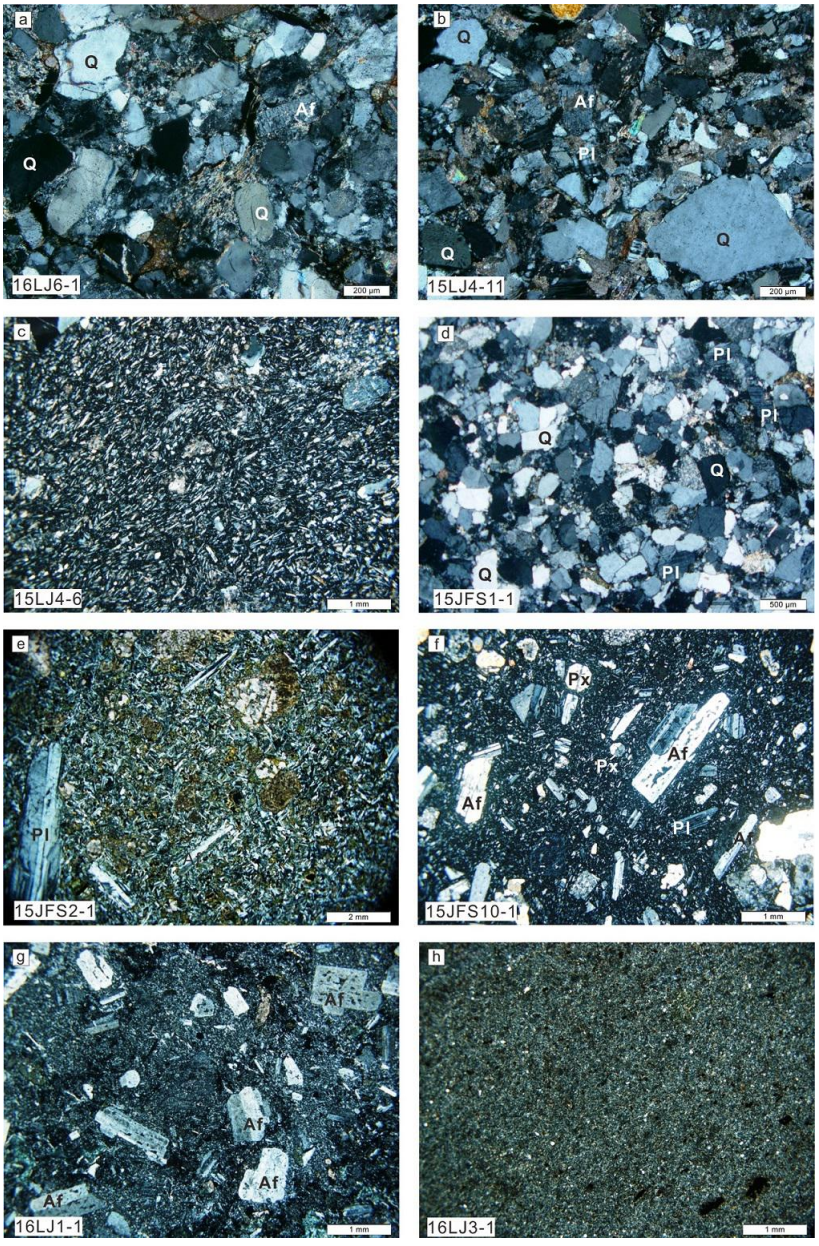

shale
silt shale
siltstone
sandstone
feldspathic quartz sandstone
conglomerate
coal
tuffaceous siltstone
andesite
plant fossils
photo location
T3ch Changbai Fm
J1y Yihe Fm

Figure 7


a
Q
Af
Q
Q
16LJ6-1
200 µm
b
Q
Af
Pl
Q
Q
15LJ4-11
200 µm
c
15LJ4-6
1 mm
d
Pl
Pl
Q
Q
Q
Pl
15JFS1-1
500 µm
e
Pl
15JFS2-1
2 mm
f
Px
Px
Af
Af
Pl
Af
15JFS10-1
1 mm
g
Af
Af
Af
Af
16LJ1-1
1 mm
h
16LJ3-1
1 mm

5

Figure 8

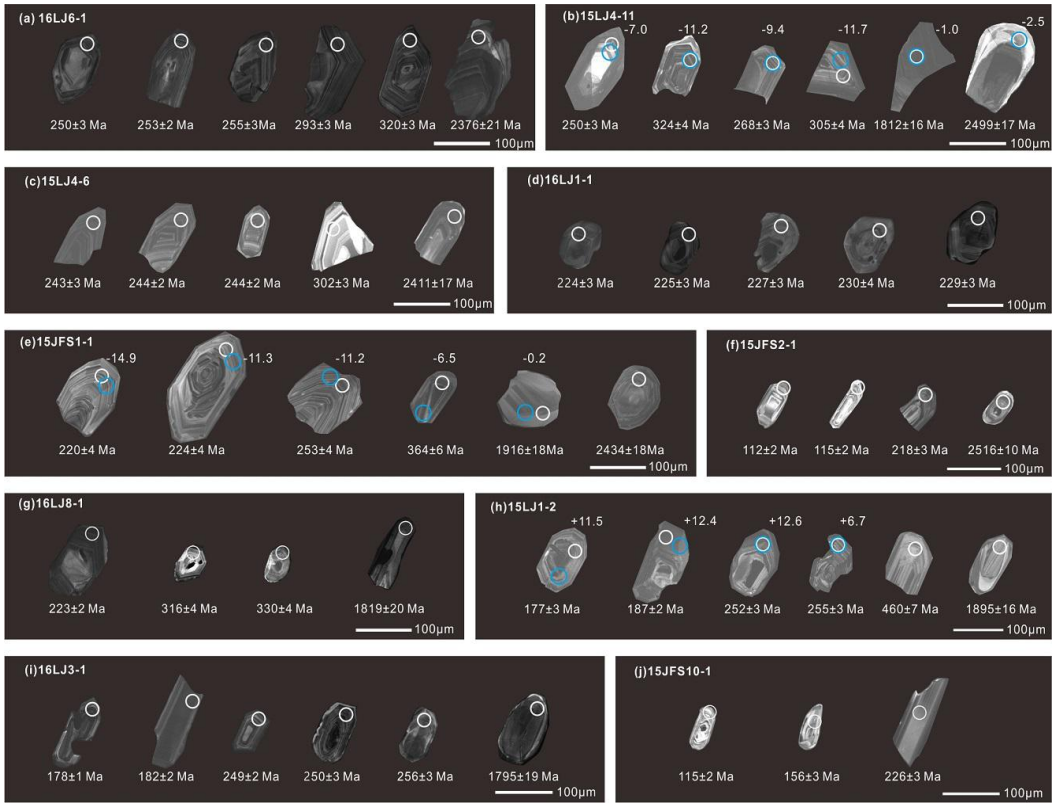

(a) 16LJ6-1
250±3 Ma
253±2 Ma
255±3Ma
293±3 Ma
320±3 Ma
2376±21 Ma
100μm
(b)15LJ4-11
-7.0
-11.2
-9.4
-11.7
-1.0
-2.5
250±3 Ma
324±4 Ma
268±3 Ma
305±4 Ma
1812±16 Ma
2499±17 Ma
100μm
(c)15LJ4-6
243±3 Ma
244±2 Ma
244±2 Ma
302±3 Ma
2411±17 Ma
100μm
(d)16LJ1-1
224±3 Ma
225±3 Ma
227±3 Ma
230±4 Ma
229±3 Ma
100μm
(e)15JFS1-1
-14.9
-11.3
-11.2
-6.5
-0.2
220±4 Ma
224±4 Ma
253±4 Ma
364±6 Ma
1916±18Ma
2434±18Ma
100μm
(f)15JFS2-1
112±2 Ma
115±2 Ma
218±3 Ma
2516±10 Ma
100μm
(g)16LJ8-1
223±2 Ma
316±4 Ma
330±4 Ma
1819±20 Ma
100μm
(h)15LJ1-2
+11.5
+12.4
+12.6
+6.7
177±3 Ma
187±2 Ma
252±3 Ma
255±3 Ma
460±7 Ma
1895±16 Ma
100μm
(i)16LJ3-1
178±1 Ma
182±2 Ma
249±2 Ma
250±3 Ma
256±3 Ma
1795±19 Ma
100μm
(j)15JFS10-1
115±2 Ma
156±3 Ma
226±3 Ma
100μm

Figure 9

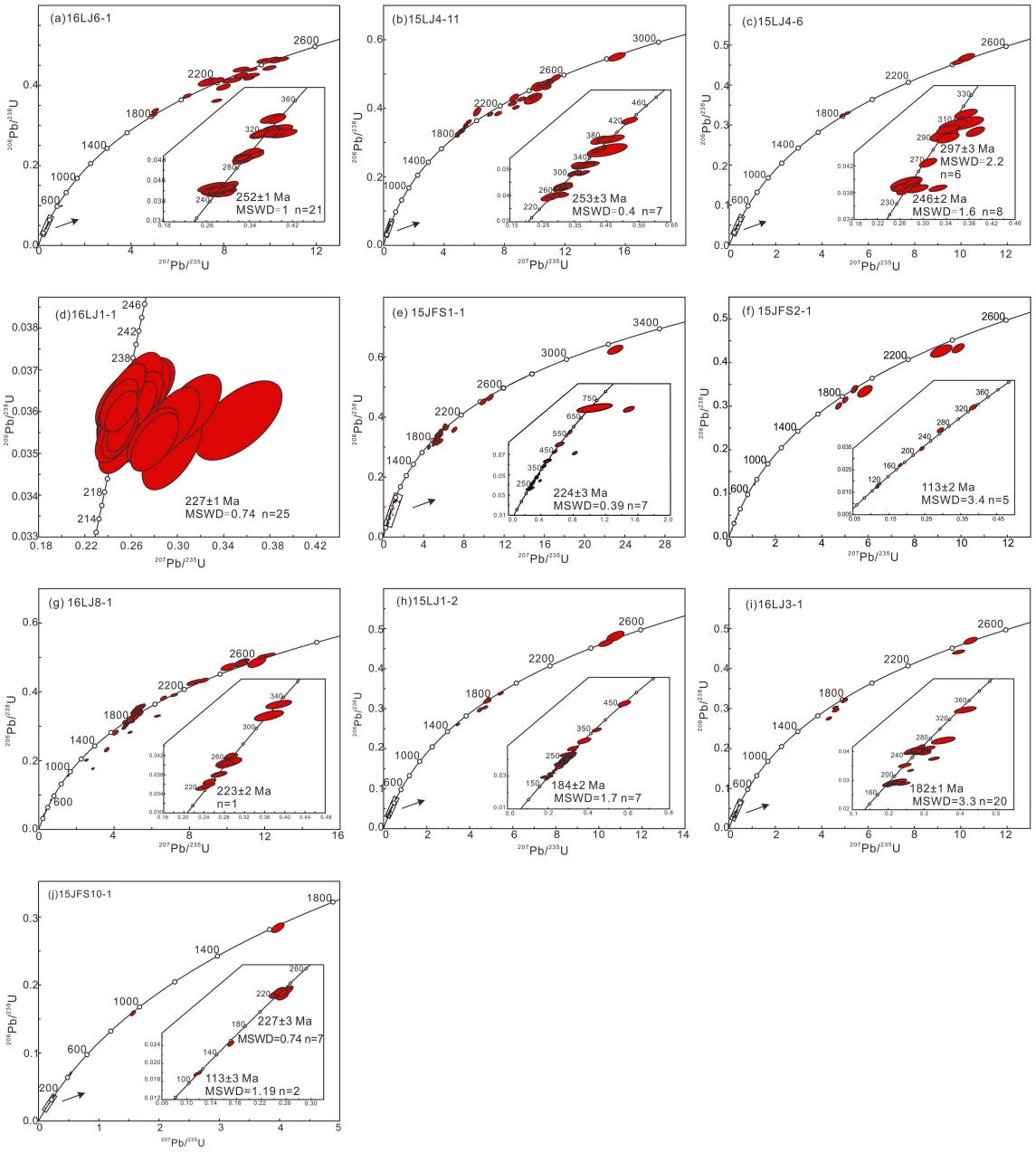

(a)16LJ6-1
252±1 Ma
MSWD=1 n=21
(b)15LJ4-11
253±3 Ma
MSWD=0.4 n=7
(c)15LJ4-6
297±3 Ma
MSWD=2.2
n=6
246±2 Ma
MSWD=1.6 n=8
(d)16LJ1-1
227±1 Ma
MSWD=0.74 n=25
(e) 15JFS1-1
224±3 Ma
MSWD=0.39 n=7
(f) 15JFS2-1
113±2 Ma
MSWD=3.4 n=5
(g) 16LJ8-1
223±2 Ma
n=1
(h)15LJ1-2
184±2 Ma
MSWD=1.7 n=7
(i)16LJ3-1
182±1 Ma
MSWD=3.3 n=20
(j)15JFS10-1
227±3 Ma
MSWD=0.74 n=7
113±3 Ma
MSWD=1.19 n=2
$^{206}Pb/^{238}U$
$^{207}Pb/^{235}U$

Figure 10

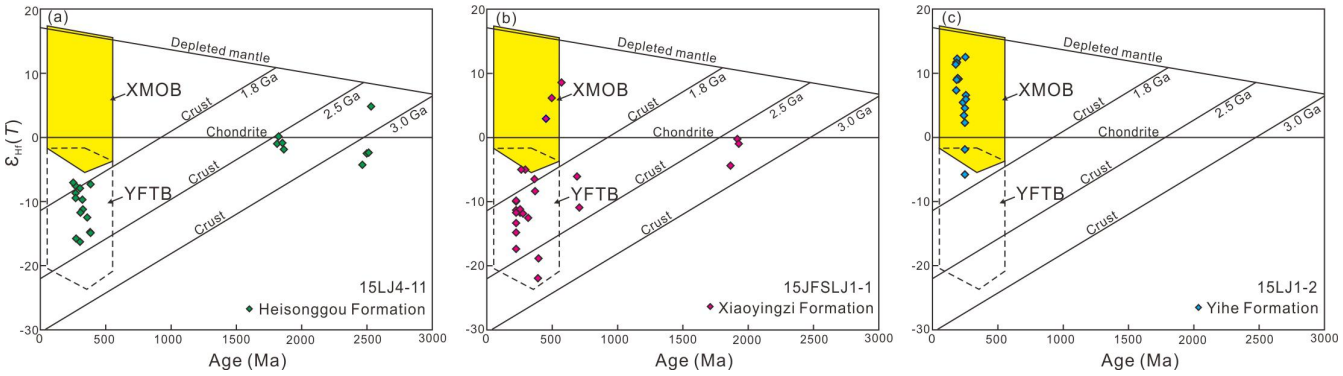

(a)
(b)
(c)
Depleted mantle
XMOB
YFTB
Crust
1.8 Ga
2.5 Ga
3.0 Ga
Chondrite
$\varepsilon_{Hf}(T)$
Age (Ma)
15LJ4-11
Heisonggou Formation
15JFSLJ1-1
Xiaoyingzi Formation
15LJ1-2
Yihe Formation

Figure 11

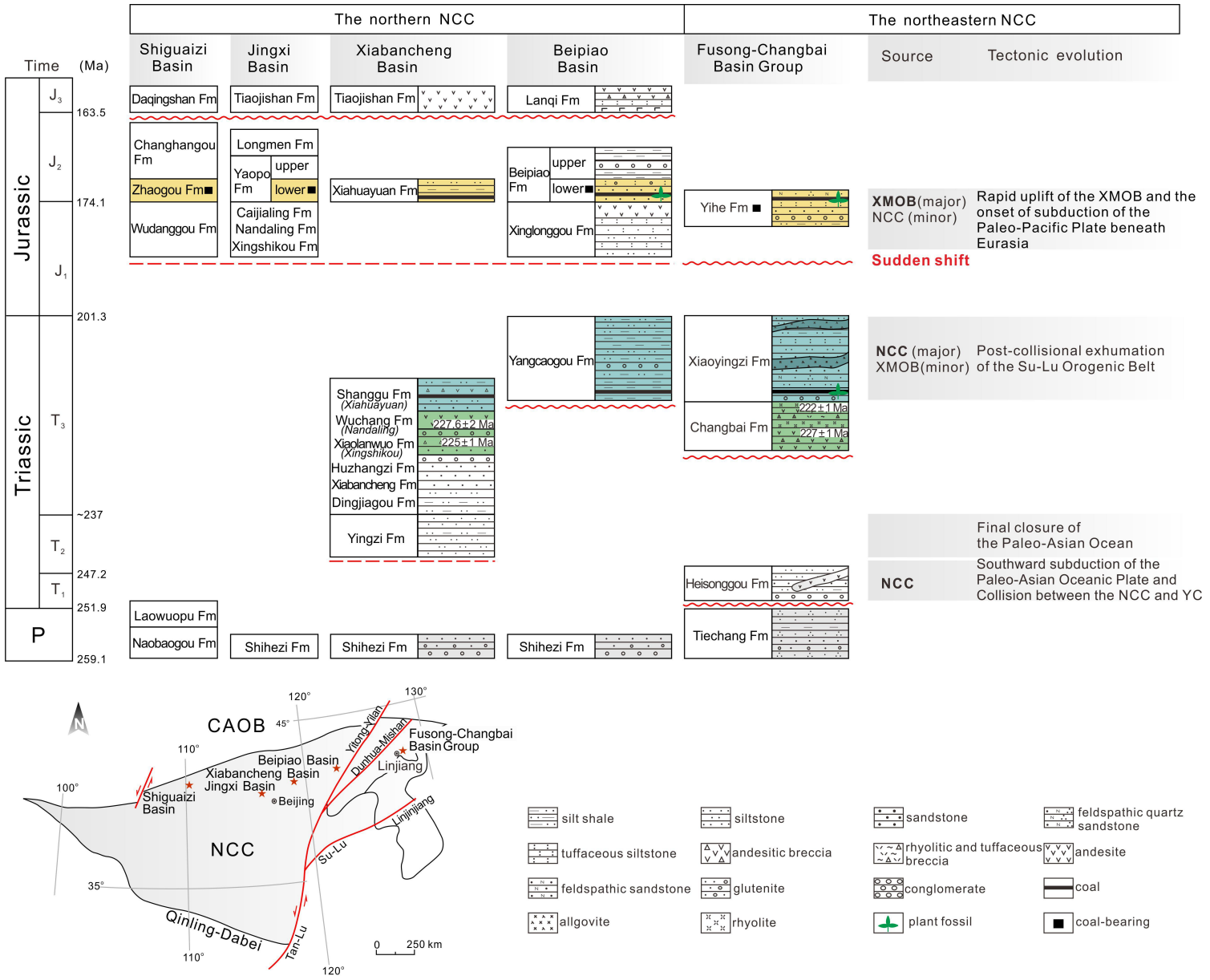

The northern NCC
The northeastern NCC
Time (Ma)
Shiguaizi Basin
Jingxi Basin
Xiabancheng Basin
Beipiao Basin
Fusong-Changbai Basin Group
Source
Tectonic evolution
Jurassic
Triassic
P
J3
J2
J1
T3
T2
T1
163.5
174.1
201.3
~237
247.2
251.9
259.1
Daqingshan Fm
Tiaojishan Fm
Tiaojishan Fm
Lanqi Fm
Changhangou Fm
Zhaogou Fm
Wudanggou Fm
Longmen Fm
Yaopo Fm
upper
lower
Caijialing Fm
Nandaling Fm
Xingshikou Fm
Xiahuayuan Fm
Beipiao Fm
upper
lower
Xinglonggou Fm
Yihe Fm
XMOB(major)
NCC (minor)
Rapid uplift of the XMOB and the onset of subduction of the Paleo-Pacific Plate beneath Eurasia
Sudden shift
Yangcaogou Fm
Xiaoyingzi Fm
Changbai Fm
222±1 Ma
227±1 Ma
NCC (major)
XMOB(minor)
Post-collisional exhumation of the Su-Lu Orogenic Belt
Shanggu Fm
(Xiahuayuan)
Wuchang Fm
(Nandaling)
227.6±2 Ma
Xiaolanwuo Fm
(Xingshikou)
225±1 Ma
Huzhangzi Fm
Xiabancheng Fm
Dingjiagou Fm
Yingzi Fm
Final closure of the Paleo-Asian Ocean
NCC
Southward subduction of the Paleo-Asian Oceanic Plate and Collision between the NCC and YC
Heisonggou Fm
Tiechang Fm
Laowuopu Fm
Naobaogou Fm
Shihezi Fm
Shihezi Fm
Shihezi Fm
CAOB
NCC
Beipiao Basin
Xiabancheng Basin
Jingxi Basin
Shiguaizi Basin
Beijing
Linjiang
Fusong-Changbai Basin Group
Yitong-Yilan
Dunhua-Mishan
Linjinjiang
Su-Lu
Tan-Lu
Qinling-Dabei
120°
130°
45°
110°
100°
35°
0 250 km
silt shale
siltstone
sandstone
feldspathic quartz sandstone
tuffaceous siltstone
andesitic breccia
rhyolitic and tuffaceous breccia
andesite
feldspathic sandstone
glutenite
conglomerate
coal
allgovite
rhyolite
plant fossil
coal-bearing

Figure 12

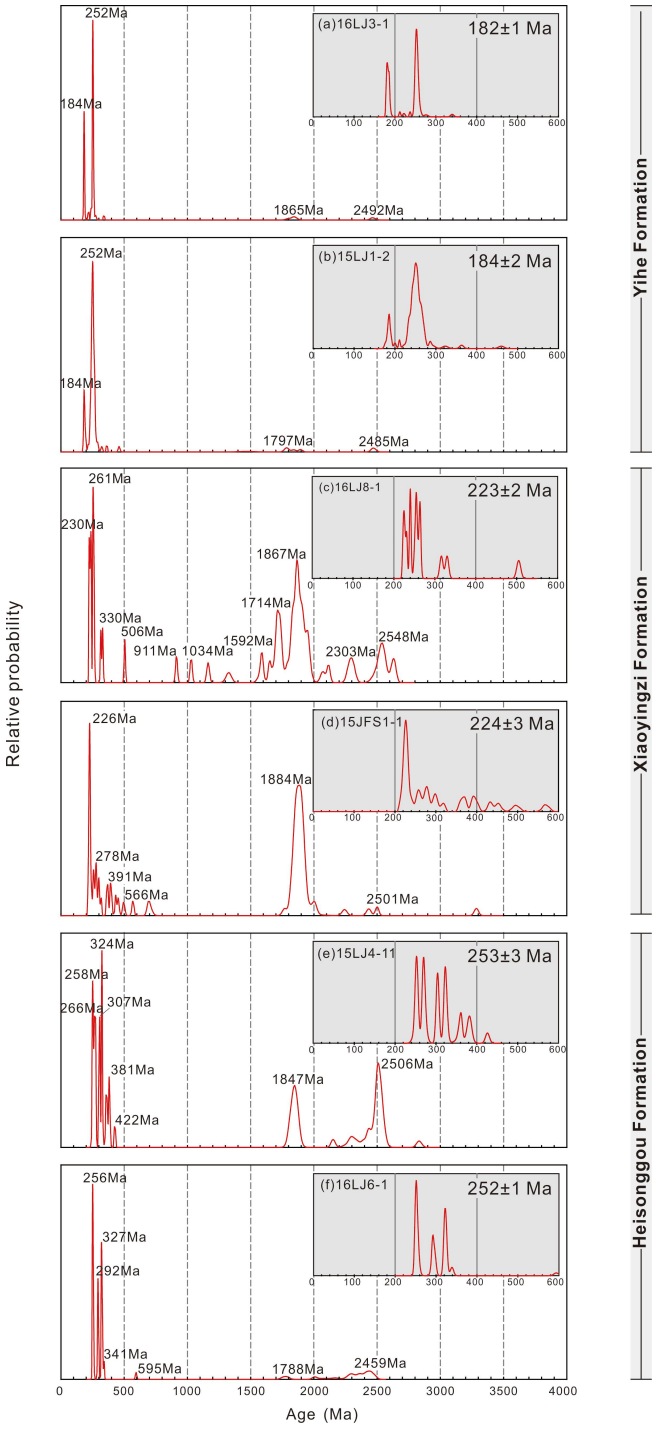
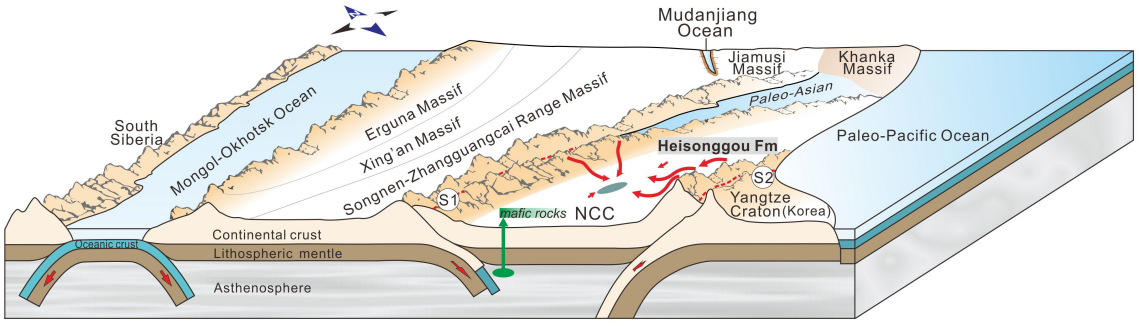
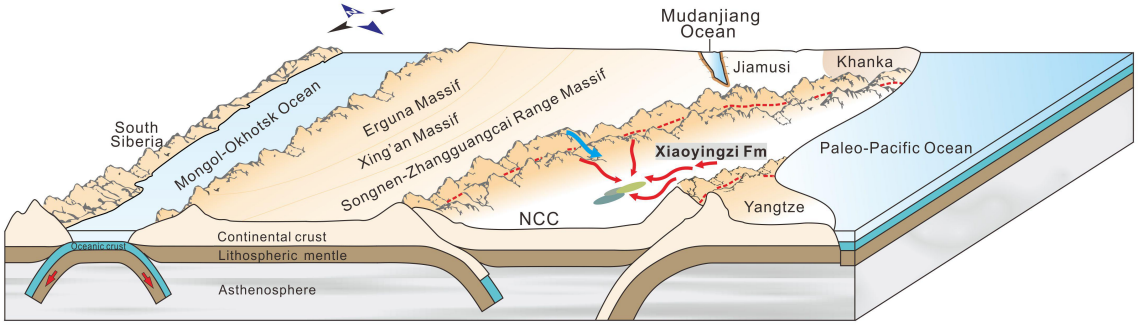
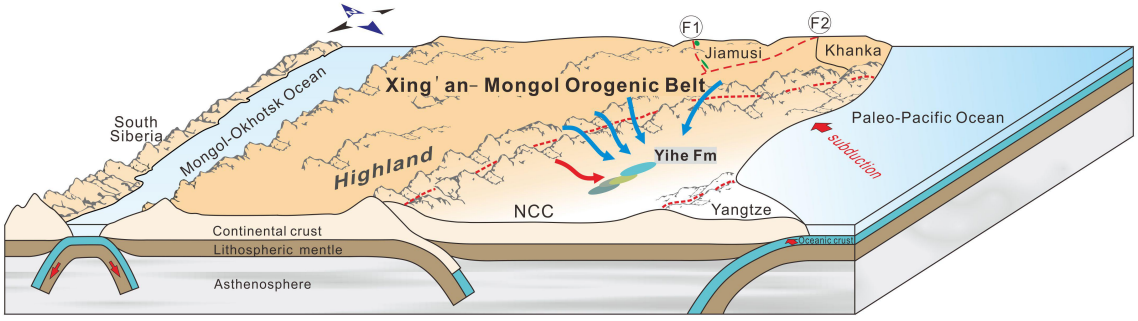

(a)16LJ3-1
182±1 Ma
252Ma
184Ma
1865Ma
2492Ma
(b)15LJ1-2
184±2 Ma
252Ma
184Ma
1797Ma
2485Ma
Yihe Formation
(c)16LJ8-1
223±2 Ma
261Ma
230Ma
1867Ma
1714Ma
330Ma
506Ma
911Ma
1034Ma
1592Ma
2303Ma
2548Ma
(d)15JFS1-1
224±3 Ma
226Ma
1884Ma
278Ma
391Ma
566Ma
2501Ma
Xiaoyingzi Formation
(e)15LJ4-11
253±3 Ma
324Ma
258Ma
266Ma
307Ma
381Ma
422Ma
1847Ma
2506Ma
(f)16LJ6-1
252±1 Ma
256Ma
327Ma
292Ma
341Ma
595Ma
1788Ma
2459Ma
Heisonggou Formation
Relative probability
0
500
1000
1500
2000
2500
3000
3500
4000
Age (Ma)

Figure 13

(a) Early Triassic: southward subduction of the Paleo-Asian oceanic plate, and subduction and collision between the NCC and YC



(b) Late Triassic: final closure of the Paleo-Asian Ocean and post-collisional exhumation of the Su-Lu Orogenic Belt



(c) Early Jurassic: rapid uplift of the XMOB and the onset of subduction of the Paleo-Pacific Plate beneath Eurasia