We are thankful to Dr Covault for his interest in our paper and for his helpful and constructive comments.

We understand the point of view of the reviewer, two main ideas are indeed present in the paper. After consideration we have decided to keep only one focus for the paper: the topographic impact on the system morphology as it was already the most developed insight. Consequently, the last part of the interpretation section has been removed. Description of the stratigraphic framework, no longer useful, has been also removed. This modification makes the manuscript shorter and its objectives more precise. As suggested, we will certainly keep the removed data and interpretations for another article.

Precisions have been given in the introduction on what we consider "subtle" and why it is important to understand the impact of such changes on deep sea fans morphology. Indeed, our study can help to better constrain the terrestrial sediment routing on topographically complex passive margin and to better trace sand deposits.

We consider as "important' slope gradient changes that are over 0.5°. The gradient changes that have been well documented in the literature are mostly over a degree: Castagnola Formation (4–12°; Felletti, 2002; Southern et al., 2015; Marini et al., 2016); the Laga Formation (6–8°; Marini et al., 2015); and the Grès d'Annot (4–10°; Amy et al., 2007; Salles et al., 2014).

The results and interpretation sections have also been both reorganized in order to shorten them and make a clearer distinction between the two. All the questions raised by the reviewers concerning the interpretations have been answer in the revised manuscript.

The Ogooue Fan (offshore Gabon): a modern example of deep-sea fan on a complex slope profile.

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Abstract. The effects of ~~important~~ changes in slope gradient on ~~turbidity currents velocity~~deposition processes and architecture have been investigated in different deep-sea systems both in modern and ancient environments. However, the impact of subtle gradient changes (<0.~~5~~3°) on sedimentary processes along deep-sea fans still needs to be clarified. The Ogooue Fan, located in the northeastern part of the Gulf of Guinea, extends over more than 550 km westwards of the Gabonese shelf and passes through the Cameroun Volcanic Line. Here, we present the first study of acoustic data (multibeam echosounder and 3.5 kHz, very-high resolution seismic data) and piston cores covering the deep-sea part of this West African system. This study documents the architecture and sedimentary facies distribution along the fan. Detailed mapping ~~and~~of near-seafloor seismic ~~dataset reveal~~ reflection data reveals the influence of subtle slope gradient changes (< 0.2°) along the fan morphology. The overall system corresponds to a well-developed deep-sea fan, fed by the Ogooue River ‘sedimentary load, with tributary canyons, distributary channel-levee complexes and lobes elements. However, variations in the slope gradient due to inherited salt-related structures and the presence of several seamounts, including volcanic islands, result in a ~~more~~topographically complex ~~fan architecture~~slope profile including several ramps and ~~sedimentary facies distribution.~~steps. In particular, turbidity currents derived from the Gabonese shelf deposit ~~across~~cross several interconnected intraslope basins located on the low gradient segments of the margin (<0.3°). ~~The repeated spill-overs of the most energetic turbidity currents have notably led to the incision of a large mid-system valley on~~On a higher

gradient segment of the slope (0.6~~°)~~°), a large mid-system valley developed connecting an intermediate sedimentary basin to the more distal lobe area. Distribution and thickness of turbidite sands is highly variable along the system. However, turbidite sands are preferentially deposited on the floor of the channel and the most proximal depositional areas. Cores description indicates that the upper parts of the turbidity flows, mainly composed of fine-grained sediments, are found in the most distal depocenters. ~~Distribution and thickness of turbidite sand beds is highly variable along the system, however, turbidite sands preferentially deposit on the floor of the channel and the most proximal depositional areas. The most distal depocenters receive only the upper parts of the flows, mainly composed of fine-grained sediments. The Ogooue deep-sea fan is predominantly active during periods of low sea-level because the canyon heads are separated from terrestrial sediment sources by the broad continental shelf. However, the northern part of this system appears active during sea-level highstands. This feature is due to one deeply incised canyon, the Cape Lopez Canyon located on a narrower part of the continental shelf, which receives sediments transported by the longshore drift.~~

Keywords: Ogooue Fan, Gulf of Guinea, complex slope profile, turbidity currents, stepped slope

1 Introduction

Deep-sea fans are depositional sinks that host stratigraphic archives of Earth history and environmental changes (Clift and Gaedicke, 2002; Fildani and Normark, 2004; Covault et al., 2010, 2011), and are also important reservoirs of natural resources (Pettingill and Weimer, 2002). Therefore, considerable attention has been given to the problems of predicting architectures and patterns of sedimentary facies distribution in submarine fans. ~~First~~Early models concerning the morphologies of these systems described submarine fans as cone-liked depositional areas across unconfined basin floors of low relief and gentle slope gradient (Shepard and Emery, 1941; Shepard, 1951; Dill et al., 1954; Menard, 1955; Heezen et al., 1959). However, ~~the development of numerous~~

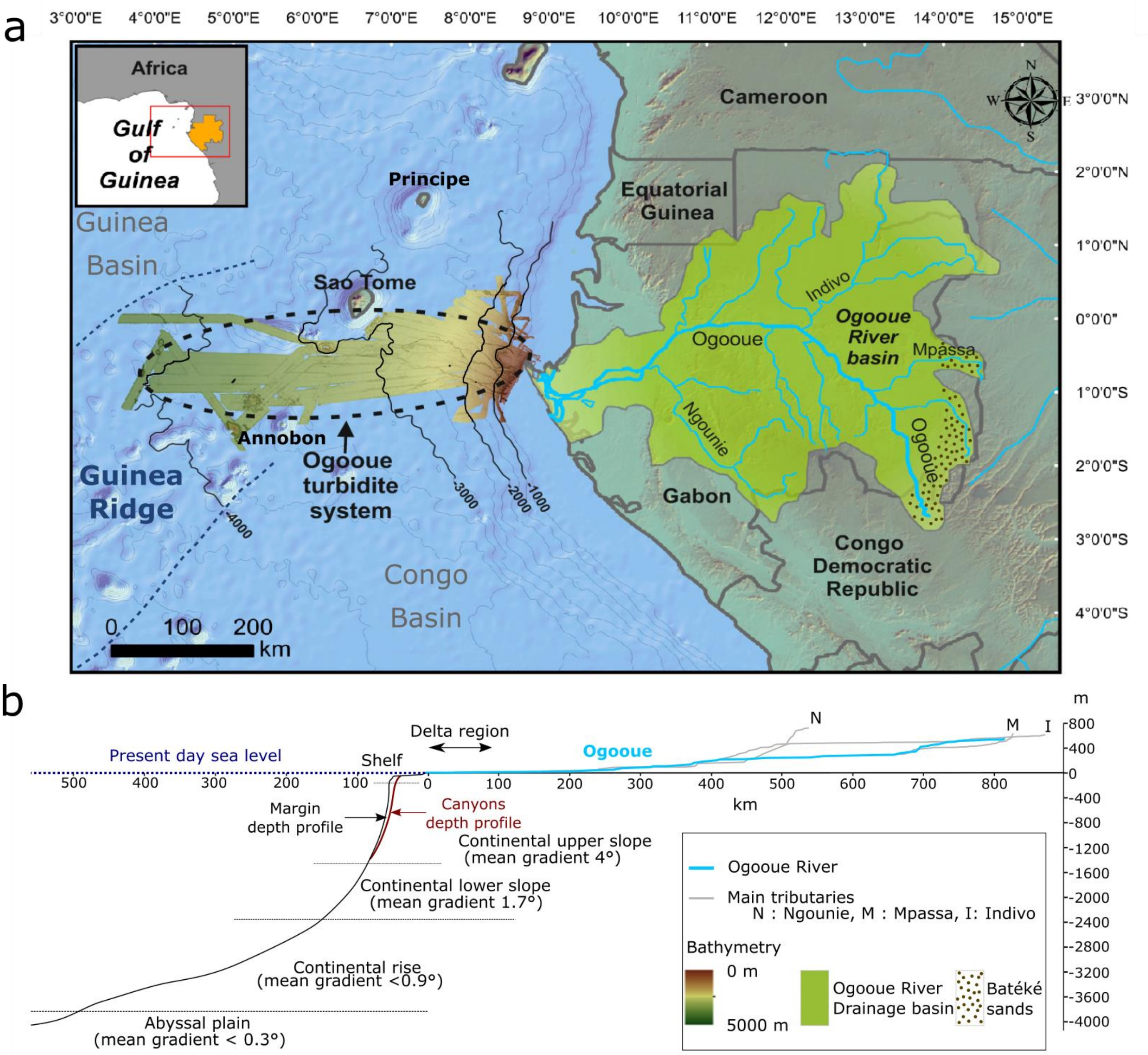
studies ~~realized on both fossil~~of outcrops (Kane et al., 2010) and modern ~~fans~~seabed datasets (Stevenson et al., 2013; Kneller, 1995) showed that topographic complexity across the receiving basin can strongly influence the organization of architectural elements of submarines fans (Normark et al., 1983; Piper and Normark, 2009). A wide range of geometries and architectural features due to topographic obstacles has been described in the literature. Among these features are ponded and intra-slope mini-~~basin~~basins due to three-dimensional confinement (Prather, 2003; Prather et al., 2012, 2017; Sylvester et al., 2015) or tortuous corridors created by topographic barriers (Smith, 2004; Hay, 2012). Spatial changes in slope gradients are also important as they cause gravity flows to accelerate or decelerate along the slope (Normark and Piper, 1991; Mulder and Alexander, 2001) allowing the construction of ~~successive~~several connected depocenters and sediment bypass areas (Smith, 2004; Deptuck et al., 2012; Hay, 2012). These stepped~~-~~ slopes have been described along modern systems such as the Niger Delta (Jobe et al., 2017), the Gulf of Mexico (Prather et al., 1998, 2017) or offshore Angola (Hay, 2012), but also in ancient systems such as the Annot Sandstone Formation (Amy et al., 2007; Salles et al., 2014), the Karoo Basin (Spychala et al., 2015; Brooks et al., 2018) or the Lower Congo basin (Ferry et al., 2005).

On stepped~~-~~ slopes where structural deformation is very slow, sediment erosion and deposition are the dominant processes that control the short-term evolution of slope. In these systems, the slope gradient variations play a key role and studies have shown that subtle gradient changes (<0.3°) can have an important impact on flow velocity and consequently deep-sea fans organization (e.g. Kneller, 1995; Kane et al., 2010; Stevenson et al., 2013). ~~However, despite the growing numbers~~Even though some of ~~studies describing~~ these systems have already been described, the impact of subtle changes in slope gradient on deep-sea ~~fans~~fan organization still needs to be better ~~apprehended~~understood in order to extend our knowledge on terrestrial sediments routing and on the potential for reservoir deposits in stepped slope settings.

The modern Ogooue Fan provides a new large-scale example of the influence of ~~subtle~~ gradient changes on deep-sea sediment routing. This system, which results from the sediment discharge of the Ogooue River, is the third largest system of the Gulf of Guinea after the Congo and the Niger fans (Séranne and Anka, 2005). However, in contrast to these two systems that have been the focus of many studies (Droz et al., 1996, 2003; Babonneau et al., 2002; Deptuck et al., 2003, 2007), the Quaternary sediments of the Gabon passive margin have not been ~~relatively poorly~~ studied, especially in its deepest parts (Bourgoin et al., 1963; Giresse, 1969; Giresse and Odin, 1973). The regional survey of the area by the SHOM (Service Hydrographique et Océanographique de la Marine) in 2005 and 2010, during the OpticCongo and MOCOSED cruises, provided the first extensive dataset on the Ogooue deep-sea fan, from the continental shelf to the abyssal plain.

The objective of this paper is to document the overall fan morphology, and to link its evolution with the local changes in slope gradients or topographic obstacles present in the depositional area. This study contributes to the understanding of the impact of subtle slope gradient changes on a whole deep-water ~~systems and~~ system. This study can be used to develop predictive models of sedimentary facies distribution for systems located on stepped~~-~~ slope with low ~~to very low~~ gradient changes (< ~~1°).~~0.5°) and to better constrain sand deposits.

2 Geological setting



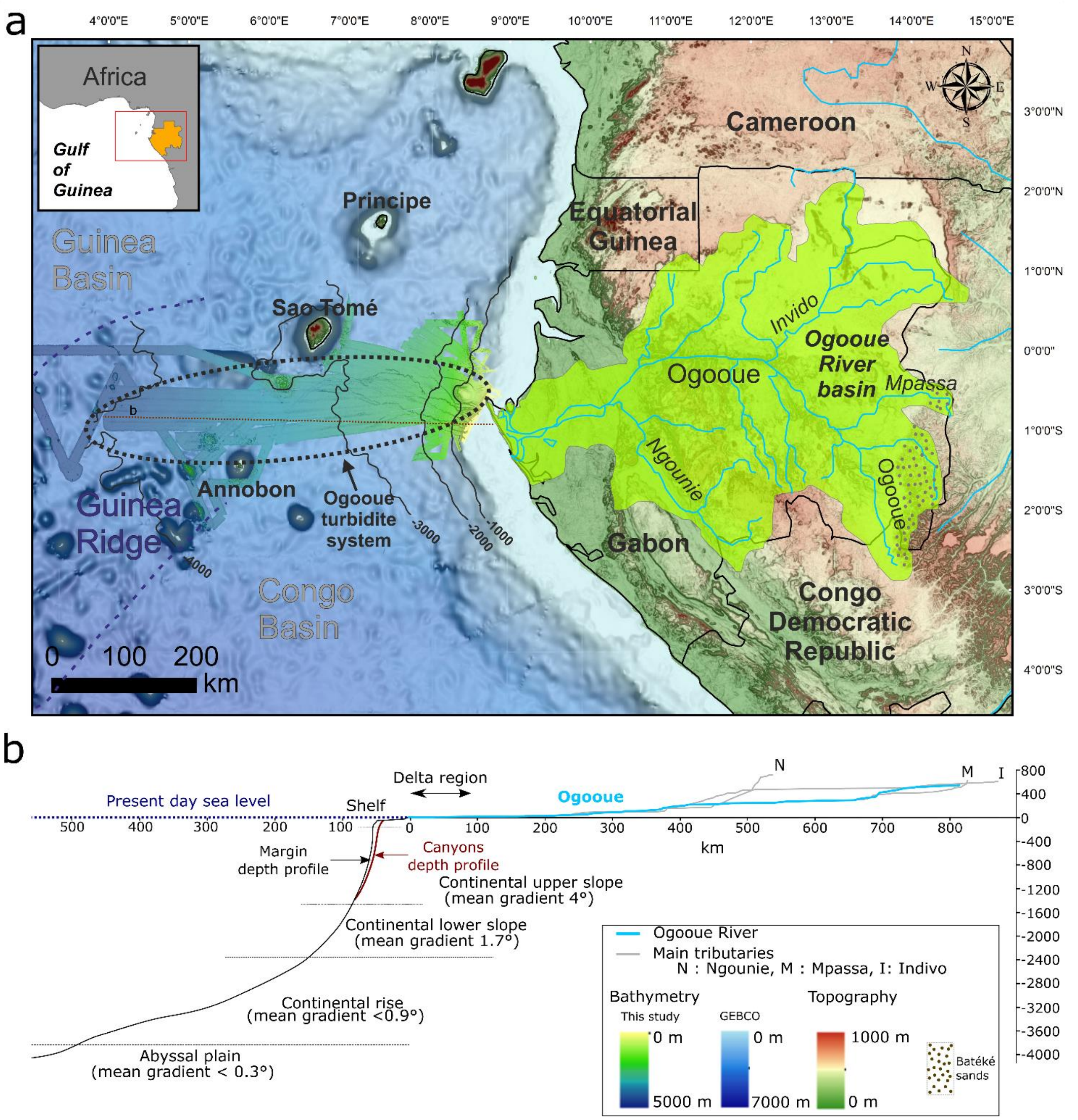


Figure 1: a) The Ogooue sedimentary system from source (river and drainage basin) to sink (Quaternary turbidite fan). b) Channel depth profile of the Ogooue River (blue) and its main tributaries (grey) and mean depth profile along the Gabonese margin.

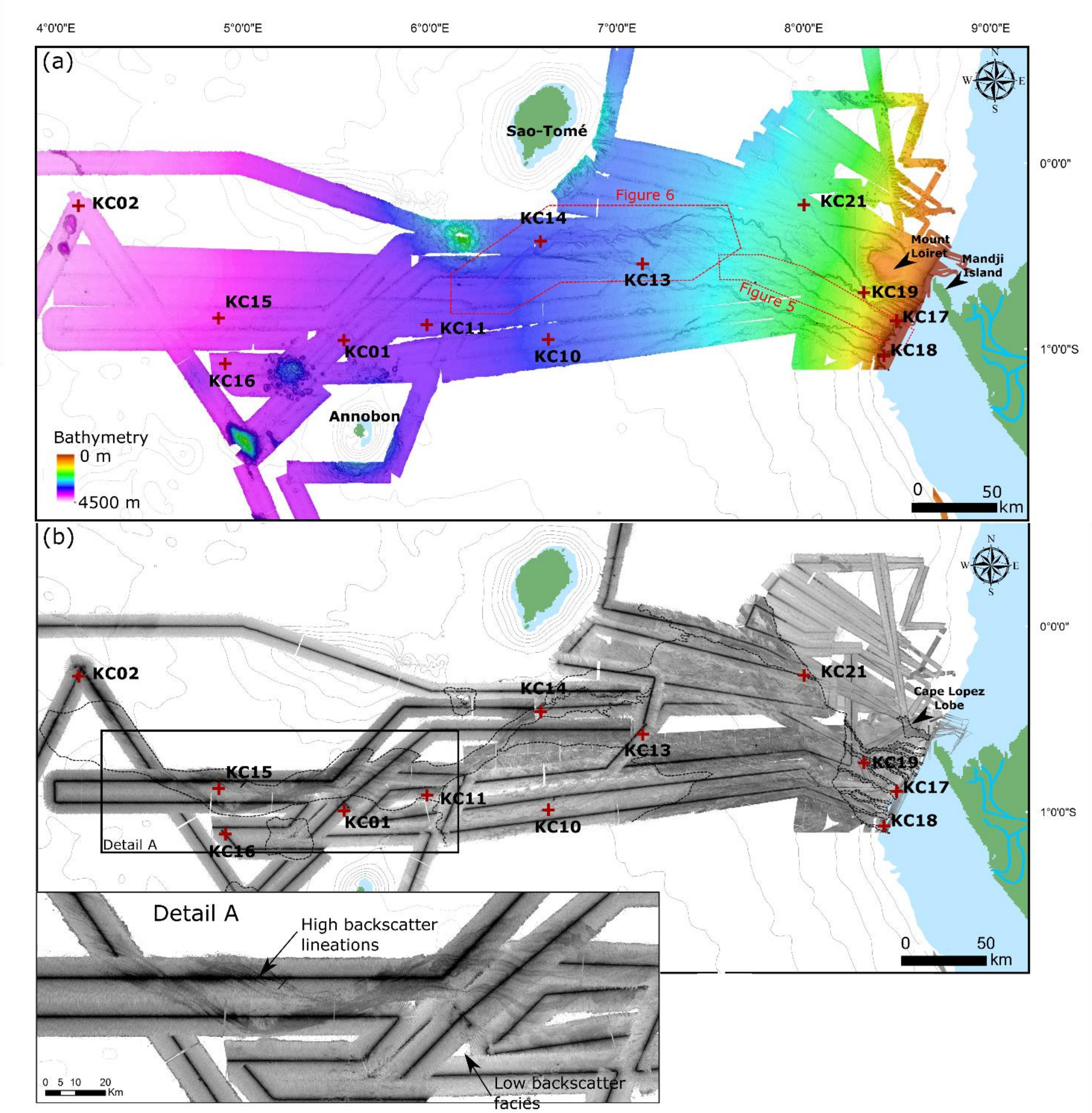
The continental margin of the Gulf of Guinea formed during the rifting that occurred within Gondwana ~~craton~~ in Neocomian to lower Aptian times. Syn-rift deposits are buried by mid-late Cretaceous transgressive ~~sediments~~sedimentary rocks consisting initially of evaporites, which have created salt-related deformations of the margin sediments, followed by platform carbonates (Cameron and White, 1999; Mougamba, 1999; Wonham et al., 2000; Séranne and Anka, 2005). Since the Late Cretaceous, the West African margin has recorded clastic sedimentation fed by the denudation of the African continent (Séranne and Anka, 2005). Different periods of major uplift and ~~canyons incision~~canyon incisions occurred from Eocene to Lower Miocene times (Rasmussen, 1996; Wonham et al., 2000; Séranne and Anka, 2005). The ~~sediments~~sediment depocenters were located basinward of the main rivers, such as the Niger, Congo, Ogooue or Orange River forming vast and thick deep-sea fans (Mougamba, 1999; Séranne and Anka, 2005; Anka et al., 2009).

The Ogooue Fan is located in the northeastern part of the Gulf of Guinea on the Gabonese continental slope. The fan ~~develops~~developed on the Guinea Ridge, which separates the two deep Congo and Guinea basins. This region is notably characterized by the presence of several volcanic islands belonging to the Cameroon Volcanic Line (CVL) associated with rocky seamounts (Figure 1~~Figure 1~~a). Geophysical studies of the volcanic line suggest that the volcanic alignment is related to a deep-mantle hot line (Déruelle et al., 2007). All the volcanoes of the CVL have been active for at least 65 Ma (Lee et al., 1994; Déruelle et al., 2007). Ar/Ar dates ~~realized~~performed on Sao Tomé and Annobon volcanic rocks ~~evidenced~~proved the activity of theses volcanic island over much of the Pleistocene (Lee et al., 1994; Barfod and Fitton, 2014). The MOCOSED 2010 cruise revealed that numerous mud volcanoes where associated with the toe of the slopes of the volcanic islands (Garlan et al., 2010). They form small topographic highs on the seafloor (< 20 m high and 100 m in diameter) and show active gas venting (Garlan et al., 2010).

The Quaternary Ogooue Fan extends westwards over 550 km through the CVL. Overall, the modern slope profile is concave upward, similar to ~~that of many~~ other passive margins~~.~~, e.g. eastern Canada margin, north Brazilian margin (Covault et al., 2012). The mean slope gradient shallows from 7° on the very upper slope to < 0.3° in the abyssal plain (Figure 1~~Figure 1~~b). The Gabonese continental shelf, which is relatively narrow, can be divided into two sub-parts: the south Gabon margin presenting a SE-NW orientation and the north Gabon margin presenting a SW-NE orientation. The southern part of the margin is characterized by the presence of numerous parallel straight gullies oriented perpendicular to the slope (Séranne and Nzé Abeigne, 1999; Lonergan et al., 2013). On the north Gabon margin, the area located between 1°00 S and the Mandji Island is incised by several canyons that belong to the modern Ogooue Fan (Figure 2~~Figure 2~~a). North of the Mandji Island, the seafloor reveals numerous isolated pockmarks as well as sinuous trains of pockmarks. These features are interpreted as the results of fluid migration from shallow buried channels (Gay et al., 2003; Pilcher and Argent, 2007).

The Ogooue Fan is supplied by the sedimentary load of the Ogooue River, which is third largest African freshwater source in the Atlantic Ocean (Mahé et al., 1990). Despite the relatively small size of the Ogooue River basin (215,000 km²), the river mean annual discharge reaches 4,700 m^3/s due to the wet equatorial climate (Lerique et al., 1983; Mahé et al., 1990). The Ogooue River flows on a low slope gradient in a drainage basin covered essentially with thick lateritic soils that developed over the Congo craton and Proterozoic formations related to Precambrian orogenic belts (Séranne et al., 2008). The estuary area includes several lakes which trap coarse sediments (Figure 1~~Figure 1~~b) (Lerique et al., 1983) ~~that~~and contribute to the dominant muddy composition of the particle load of the Ogooue River that is estimated between 1 and 10 M t/yr. (Syvitski et al., 2005). The limited portion of sand particles in the river originates mainly from the erosion of the poorly lithified Batéké Sands located on a 550-750 m high perched plateau that forms the easternmost boundary of the Ogooue watershed (Séranne et al., 2008) (Figure 1~~Figure 1~~a). On the shelf, recent fluviatile deposits consist of fine-grained sediments deposited at the mouth of the Ogooue River (Giresse and Odin, 1973). The wave ~~regime along~~conditions on the Gabonese coast are characterized by a predominant direction from South to South-West. Reflection of these southwesterly swells causes coastal sediments to be transported northward~~.~~ (Biscara et al., 2013). Sedimentary transport linked to longshore drift ranges between 300,000 m^3/yr. and 400,000 m^3/yr. (Bourgoin et al., 1963) and is responsible for the formation of the Mandji Island, a sandy spit ~~of~~ 50 km long located on the northern end of the Ogooue Delta (Figure 3~~Figure 3~~).

Except for the Cape Lopez Canyon, located just west of the Mandji Island with the canyon head in only 5 m water depth (Biscara et al., 2013), the Ogooue Fan is disconnected from the Ogooue Delta during the present-day high sea-level (Figure 3Figure 3).



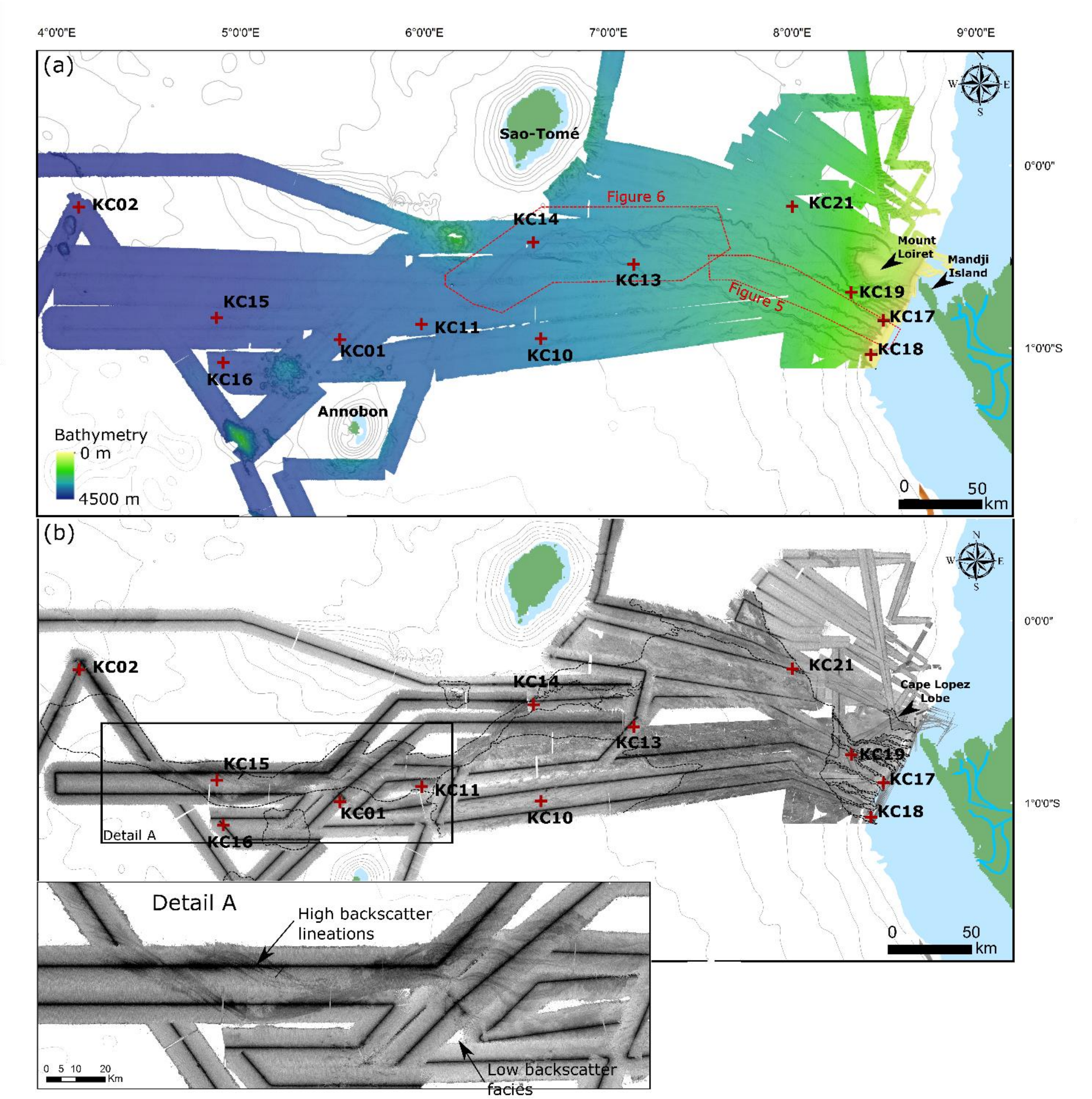
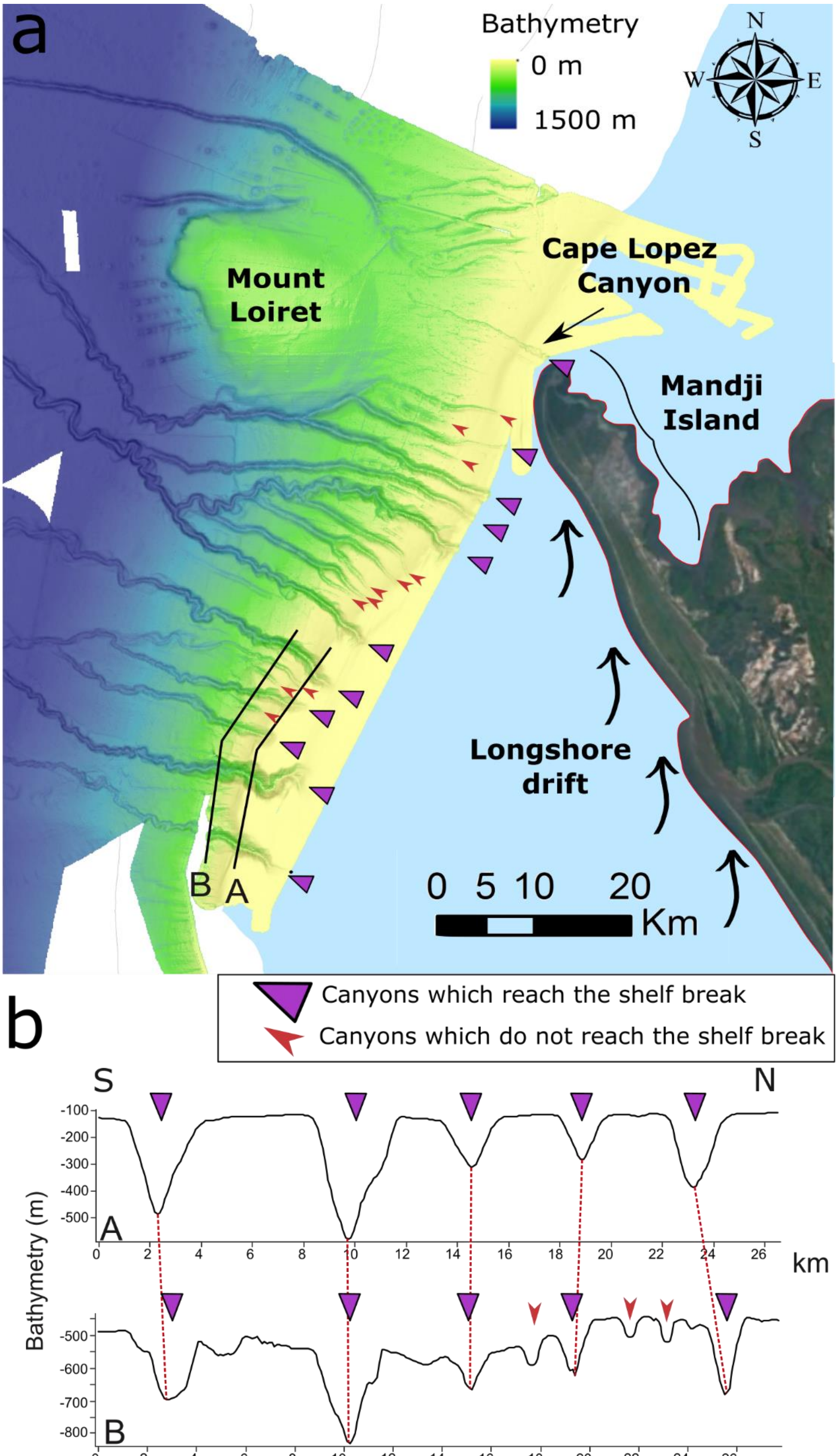


Figure 2: (a) Detailed bathymetric map of the Ogooue Fan, based on the multibeam echosounder data of the Optic Congo2005 and MOCOSED2010 surveys. (b) Acoustic imagery of the Ogooue Fan (high backscatter: dark tones; low backscatter: light tones). Detail A: close-up of the deepest part of the Ogooue Fan. Red crosses: location of the studied cores.


a
Bathymetry
0 m
1500 m
N
S
W
E
Mount Loiret
Cape Lopez Canyon
Mandji Island
Longshore drift
B
A
0 5 10 20
Km
b
Canyons which reach the shelf break
Canyons which do not reach the shelf break
S
N
Bathymetry (m)
km

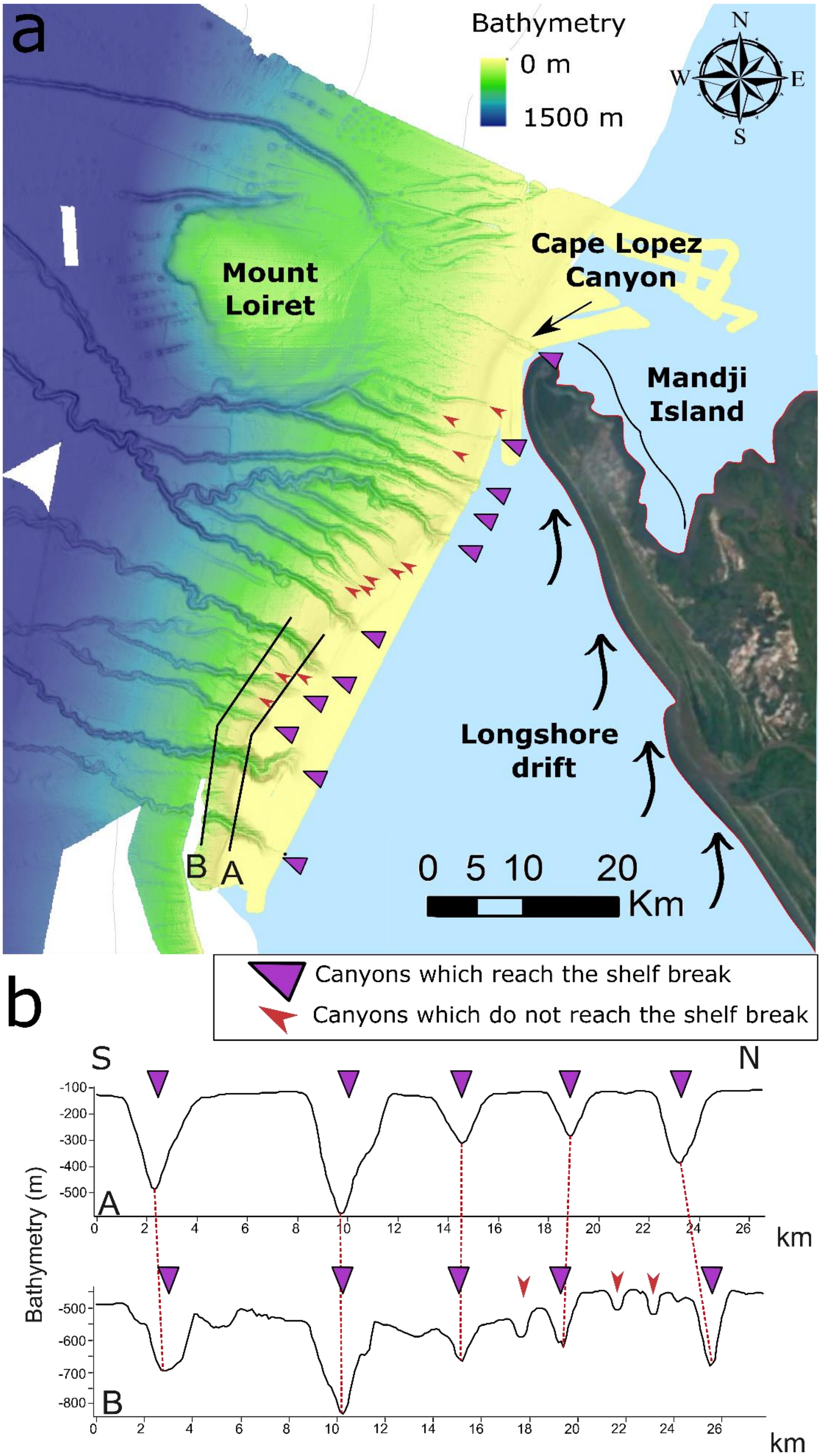
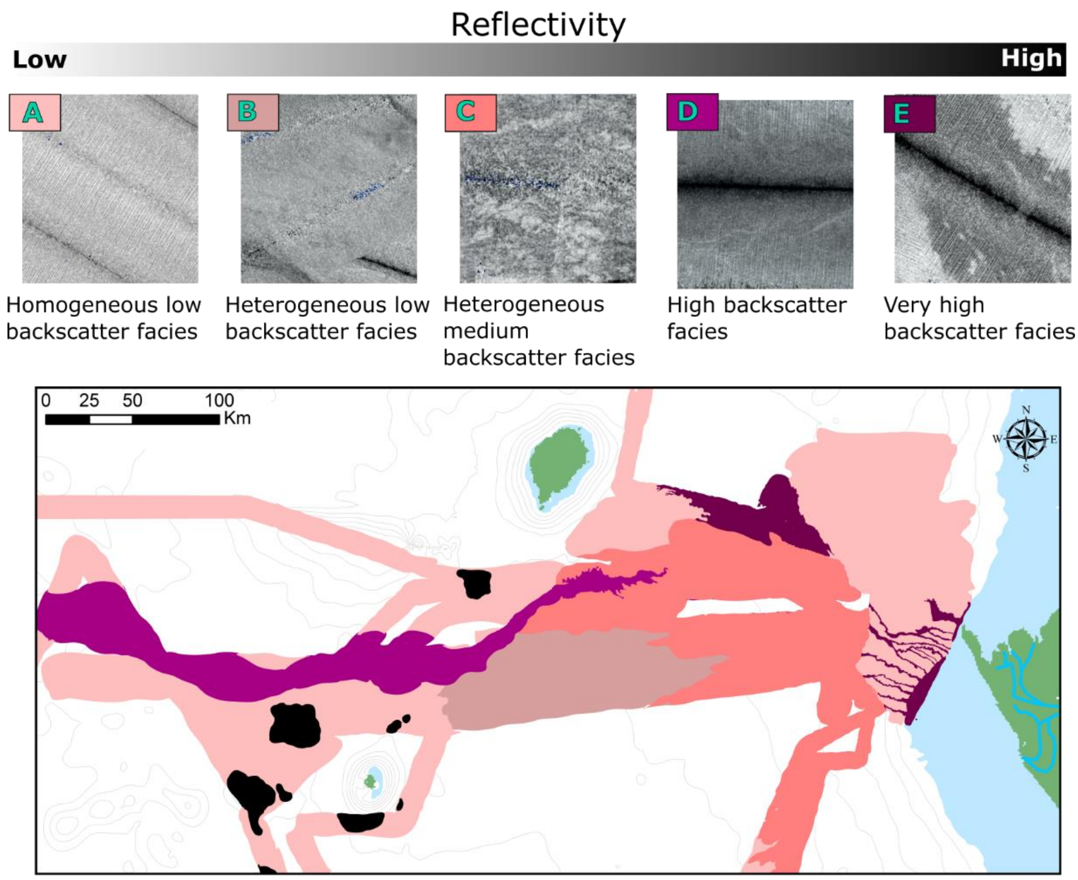


Figure 3: a) Close-up view of the Gabon shelf and canyons ramp. Bathymetry is from the Optic Congo2005 and MOCOSED2010 surveys, satellite view is from Google Earth. b) Two bathymetric profiles across the canyons showing the two types of canyons which are present along the Gabonese slope.

3 Material and method

The bathymetry and acoustic imagery of the studied area result from the multibeam echosounder (Seabat 7150) surveys conducted onboard the R/V “*Pourquoi Pas?*” and “*Beautemps-Beaupré*” during the MOCOSED 2010 and OpticCongo 2005 cruises (Mouscardes, 2005; Guillou, 2010) (Figure 2~~Figure 2~~). The multibeam backscatter data (Figure 2~~Figure 2~~b) have been used to characterize the distribution of sedimentary facies along the margin. Changes in the backscatter values correspond to variations in the nature, the texture and the state of sediments and/or the seafloor morphology (Unterseh, 1999; Hanquiez et al., 2007). On the multibeam echosounder images, lighter areas indicate low acoustic backscatter and darker areas indicate high backscatter. Five main backscatter types are identified on the basis of backscatter values and homogeneity (Figure 4~~Figure 4~~). Facies A is a homogeneous low backscatter facies, Facies B is a low backscatter heterogeneous facies, and Facies C is a medium backscatter facies characterized by the presence of numerous higher backscatter patches. Facies D and E are high and very high backscatter facies, respectively. High backscatter lineations are present within Facies D.


Reflectivity
Low
High
A
B
C
D
E
Homogeneous low backscatter facies
Heterogeneous low backscatter facies
Heterogeneous medium backscatter facies
High backscatter facies
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Km
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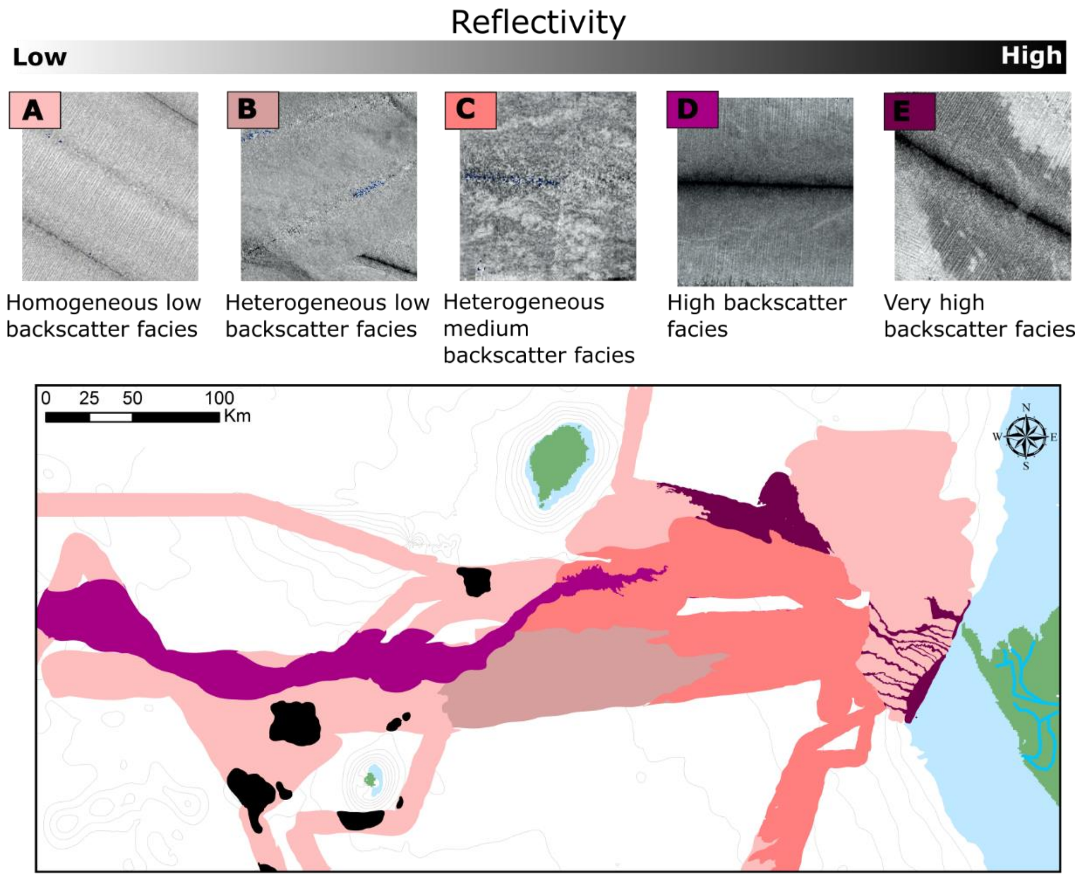
Reflectivity
Low
High
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Homogeneous low backscatter facies
Heterogeneous low backscatter facies
Heterogeneous medium backscatter facies
High backscatter facies
Very high backscatter facies
0 25 50 100
Km
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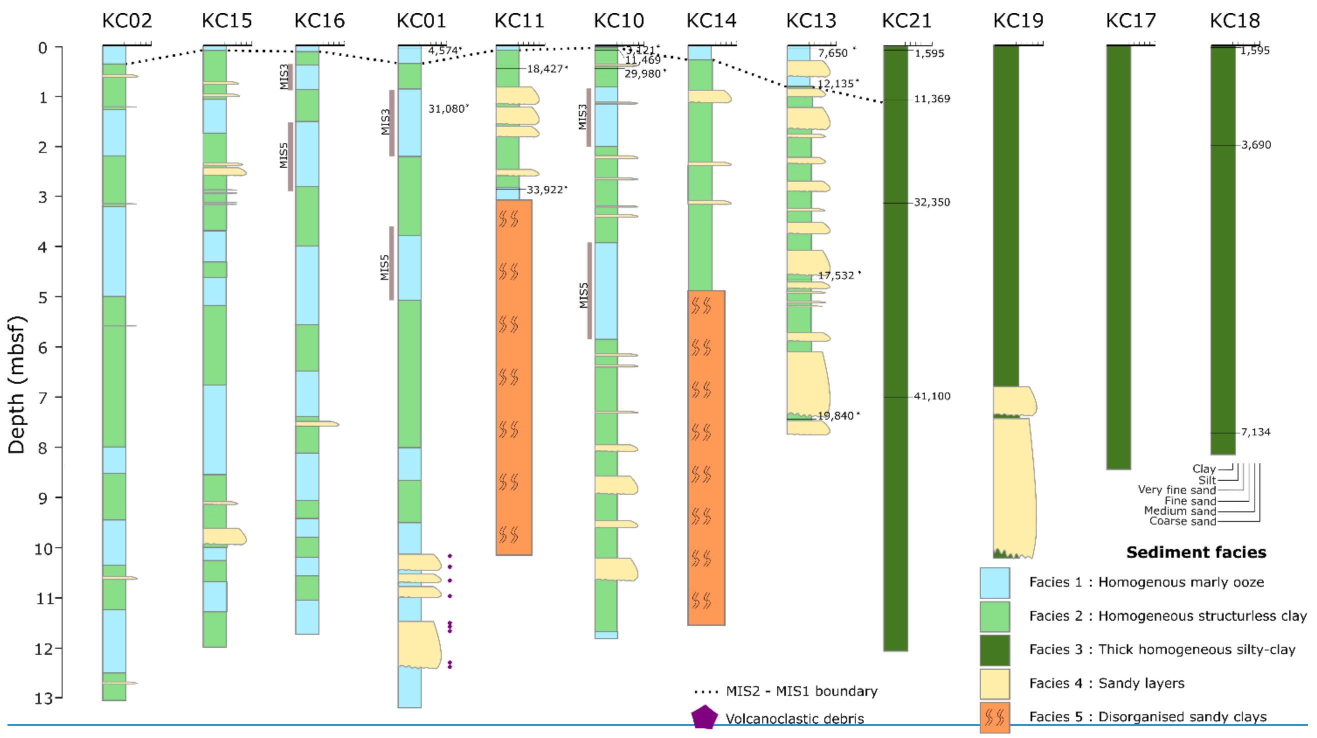
Figure 4: Reflectivity facies map of the Ogooue Fan showing the five main backscatter facies.

A total of four thousand five hundred km of 3.5 kHz seismic lines were collected in the area of the Ogooue Fan during the MOCOSED 2010 cruise and 470 km during the Optic Congo 2005 cruise (iXblue ECHOES 3500 T7). These data were used to analyze the near-surface deposits. The dataset covers the shelf edge, the slope and the abyssal plain. In this study, the 3.5 kHz echofacies ~~has~~have been classified according to Damuth's methodology (Damuth, 1975, 1980a; Damuth and Hayes, 1977) based on acoustic penetration and continuity of bottom and sub-bottom reflection horizons, micro-topography of the seafloor and presence of internal structures.

The twelve Küllenberg cores presented here were collected during the cruise MOCOSED 2010. Five of these cores have already been presented in Mignard et al. (2017) (Table 1). Visual descriptions of the cores distinguished the dominant grain size (clay, silty clay, silt, and fine sand) and vertical successions of sedimentary facies. Thin slabs were collected for each split core section and X-ray radiographed using a SCOPIX digital X-ray imaging system (Migeon et al., 1998). Subsamples were regularly taken in order to measure carbonate content using a gasometric calcimeter and grain size using a Malvern Mastersizer S. ~~The stratigraphic framework is based on the previous work of Mignard et al., (2017), new AMS ^{14}C dating (Table 1) done on core KC21 and KC18 and facies correlation to determine the boundary between Marine Isotopic Stage 1 (MIS1) and Marine Isotopic Stage 2 (MIS2). Indeed, the transition from MIS2 to MIS1 in the south-west Atlantic is marked by an abrupt increase in carbonate content (Volat et al., 1980; Jansen et al., 1984; Olausson, 1984; Zachariasse et al., 1984). This feature is recorded in all the cores of this study collected in the medium and distal part of the system (Figure 5). The new AMS ^{14}C datings were realized on a mixture of different planktonic foraminifer species living in the uppermost water column. Radiocarbon dates have been calibrated using MARINE13 curve (Reimer, 2013) and using a standard reservoir age of 400 years (Table 2; Mignard et al, 2017).~~

Table 1: Characteristics of the twelve studied cores (MOCOSED 2010 cruise).

Core	Depth (m)	Latitude	Longitude	Length (m)
KC01	3504	00°57,010’ S	005°31,806’ E	12,96
KC02	4109	00°13,525’ S	004°07,620’ E	12,76
KC10	3148	00°56,666’ S	006°39,809’ E	11,54
KC11	3372	00°52,008’ S	006°00,008’ E	9,92
KC13	2852	00°32,508’ S	007°08,589’ E	7,62
KC14	3140	00°25,010’ S	006°36,006’ E	11,34
KC15	3850	00°49,996’ S	004°50,009’ E	12,01
KC16	3738	01°05,003’ S	004°52,010’ E	11,48
KC17	565	00°51,188’ S	008°29,377’ E	8,20
KC18	366	01°01,940’ S	008°25,409’ E	7,99
KC19	1610	00°41,593’ S	008°18,592’ E	10,03
KC21	2347	00°13,004’ S	008°00,011’ E	11,81



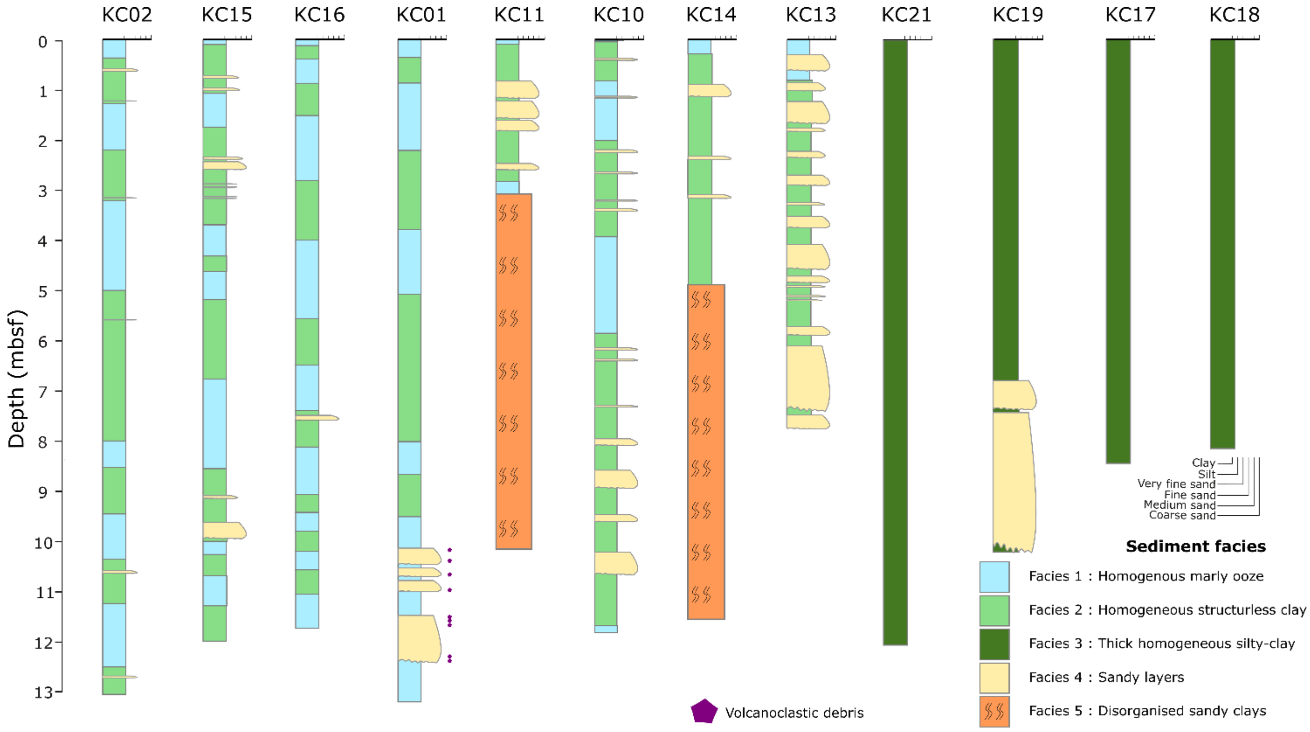


Figure 5: Sedimentological core logs from the Ogooue Fan, showing grain-size variation, lithology and bed thickness (locations of cores are presented in Figure 2~~Figure 2~~). ~~**Ages are from ^{14}C dating (dates with a star are from Mignard et al. (2017), grey bars show MIS3 and 5 sediments for KC16, KC01 and KC10 (Mignard et al., 2017).**~~

~~**Table 2: AMS ^{14}C ages with calendar age correspondences realized for this study (Reimer, 2013).**~~

~~Core number~~	~~Sample depth~~	~~Conventional age (reservoir correction) BP~~	~~Calendar age cal. BP~~
~~KC18~~	~~7~~	~~1,523 ±30~~	~~1,780~~
~~KC18~~	~~197~~	~~3,671±30~~	~~3,690~~
~~KC18~~	~~787~~	~~7056±40~~	~~7,134~~
~~KC21~~	~~12~~	~~1532±30~~	~~1,595~~
~~KC21~~	~~115~~	~~10,654±80~~	~~11,369~~
~~KC21~~	~~327~~	~~30,569±90~~	~~32,350~~
~~KC21~~	~~700~~	~~39,732±120~~	~~41,100~~

~~45~~4 Results

~~45.1~~4.1 Sedimentary facies

The classification in five sedimentary facies used here is based on photography and X-ray imagery, grain size analyses and $CaCO_3$ content (Figure 5~~Figure 5).~~, Table 2). Interpretation of these facies is based on the comparison with previous sedimentary facies classifications such as Stow and Piper~~,~~ (1984); Pickering et al~~.,~~. (1986) and Normark and Damuth~~,~~ (1997).

~~*Facies 1: Homogenous, structureless marly ooze.* This facies is composed of structureless, light beige marly ooze with relatively high concentration of planktonic foraminifers. The mean grain size is around 15 µm and the $CaCO_3$ content ranges between 40 and 60%. This facies is interpreted as a pelagic drape deposit; it forms the modern seafloor of the deepest part of the Ogooue Fan and is observed in most of the core tops corresponding to the MIS 1 interval.~~

~~*Facies 2: Homogenous, structureless clay:* Facies 2 consists of dark brown clay. The mean grain size is less than 15 µm and the $CaCO_3$ content is less than 30%. This facies has been interpreted as hemipelagic drape deposits.~~

~~*Facies 3: Thick, homogeneous silty-clay*: Facies 3 consists of very thick homogeneous dark silt-clay layers containing less than 10% of $CaCO_3$. This facies contains numerous quartz and mica grains and plant debris indicating a continental origin of the sediments. It results from the deposition of the fine-grained suspended load coming from the Ogooue River and flowing down the slope or belonging to the flow tops of the turbidity currents.~~

~~*Facies 4: Silty to sandy layers*: Facies 4 consists of fine- to medium-grained sand beds with a thickness up to several meters. They are either normally-graded or massive and display a variety of bedding structures: ripple cross laminations, parallel laminations. The composition varies from terrigenous (quartz and mica) to biogenic (foraminifers), some sand beds are highly enriched in organic debris (Mignard et al., 2017). They are interpreted as being deposited by turbidity currents initiated on the Gabonese~~

continental shelf. Four beds sampled at the base of core KC01 present a high concentration of volcaniclastic debris, such particles are completely absent in all the other sandy beds (Figure 5) sandy beds. This specific composition and the particular location of the core both suggest that these sequences originate from the nearby Annobon volcanic island.

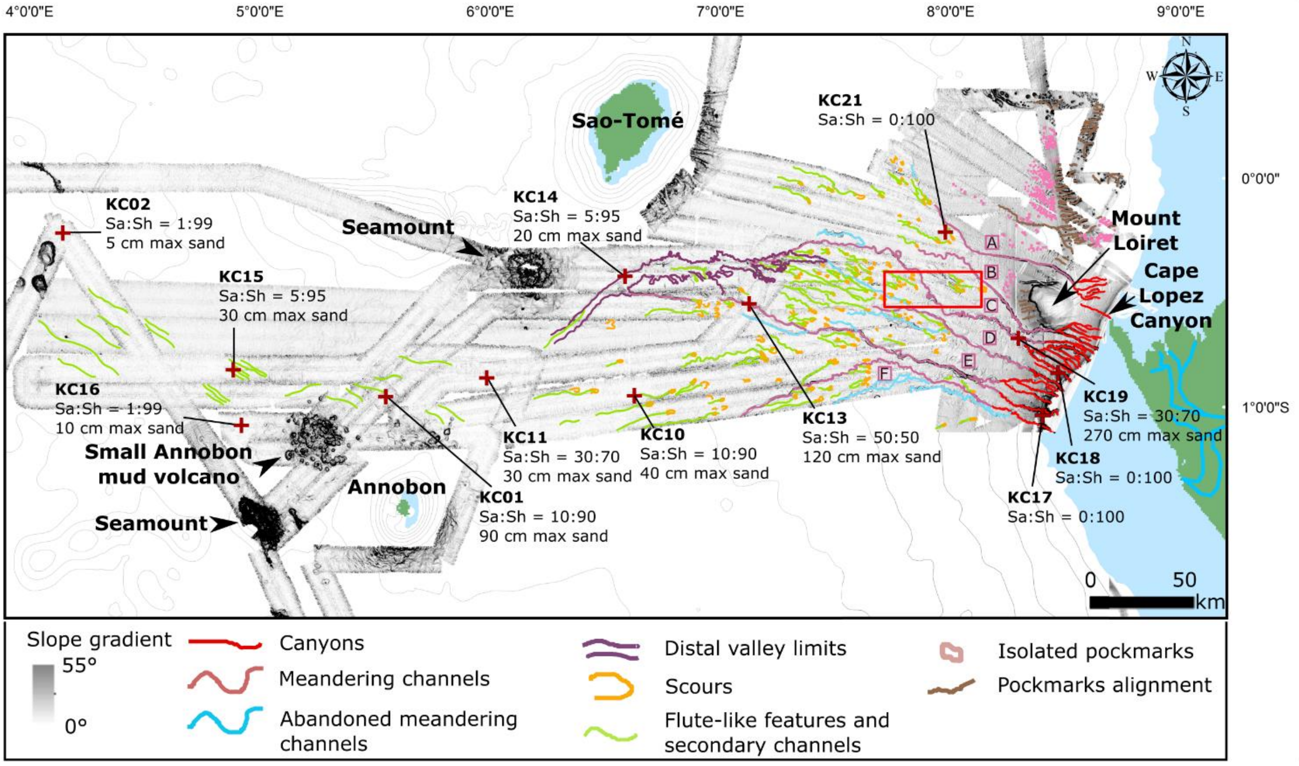
Facies 5: Disorganized sandy clays: Facies 5 consists of thick intervals of deformed or chaotic clay with deformed or folded silty to sandy layers containing mainly quartz grains and rare plant debris. This facies is interpreted as a slump deposit or debrite.

Table 2: Sedimentary facies characteristics.

Facies	Name	Structure	Color	Mean grain size	$CaCO_3$ content	Grains	Deposition process	Remarks
1	*Homogenous, structureless marly ooze*	Massive	Light beige	15 µm	40-60 %	High concentration of planktonic foraminifers	Pelagic drape deposit;	This facies forms the modern seafloor of the deepest part of the Ogooue Fan and is observed in most of the core tops.
2	Homogenous, structureless clay	Massive	Dark brown	15 µm	<30%		Hemipelagic drape deposits	
3	*Thick, homogeneous silty-clay*	Massive	Dark brown	40 µm	<10%	high concentration of quartz and mica grains and plant debris	Deposition of the fine-grained suspended load coming from the Ogooue River and flowing down the slope or belonging to the flow tops of the turbidity currents.	

4	*Silty to sandy layers*	Massive or presenting ripple cross laminations or parallel laminations	Grey to beige	60-120 µm	Highly variable	Composed of quartz and mica grains or foraminifers, some sand beds are highly enriched in organic debris (Mignard et al., 2017)	Deposited by turbidity currents initiated on the Gabonese continental shelf.	Four beds sampled at the base of core KC01 present a high concentration of volcaniclastic debris, such particles are completely absent in all the other sandy beds (Figure 5) sandy beds. This specific composition and the particular location of the core both suggest that these sequences originate from the nearby Annobon volcanic island.
5	*Disorganized sandy clays*	Deformed or chaotic clay with deformed or folded silty to sandy layers			Highly variable	Numerous quartz grains and rare plant debris	Slump deposit or debrite	

45.24.2 Fan morphology



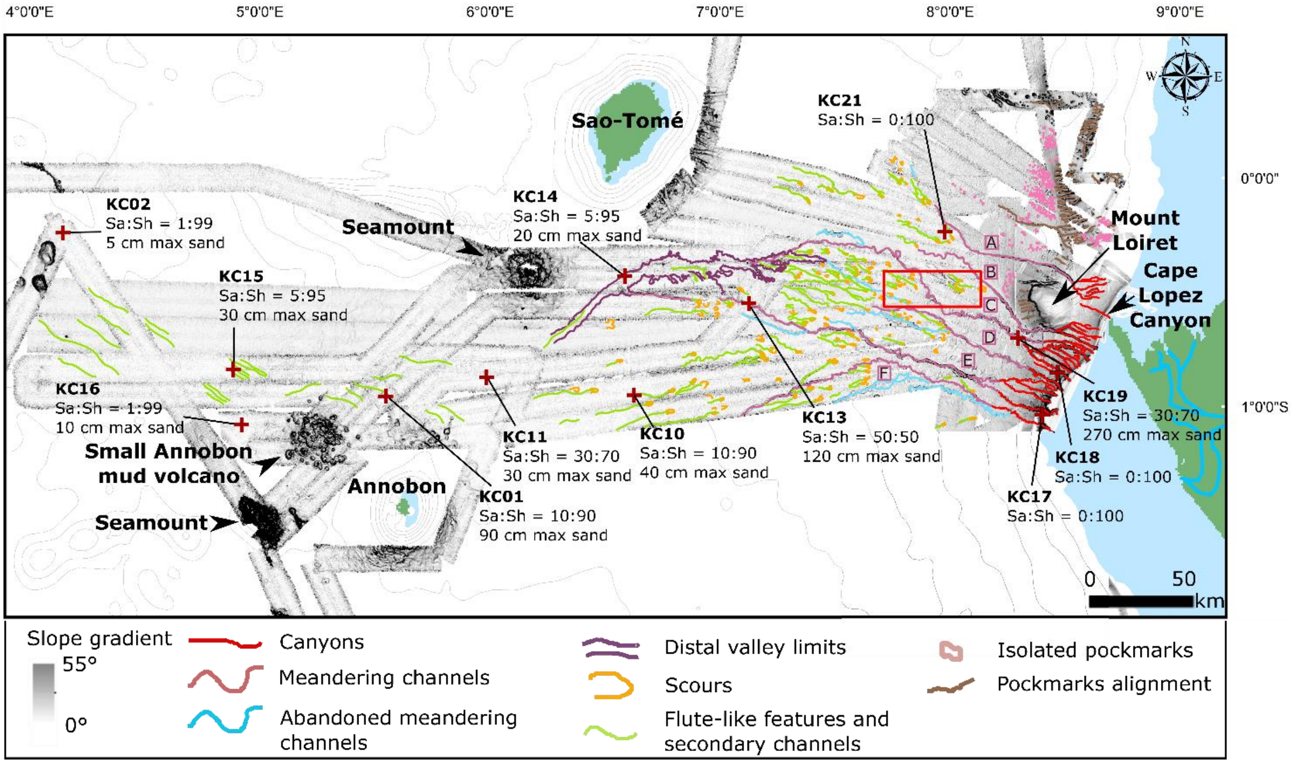


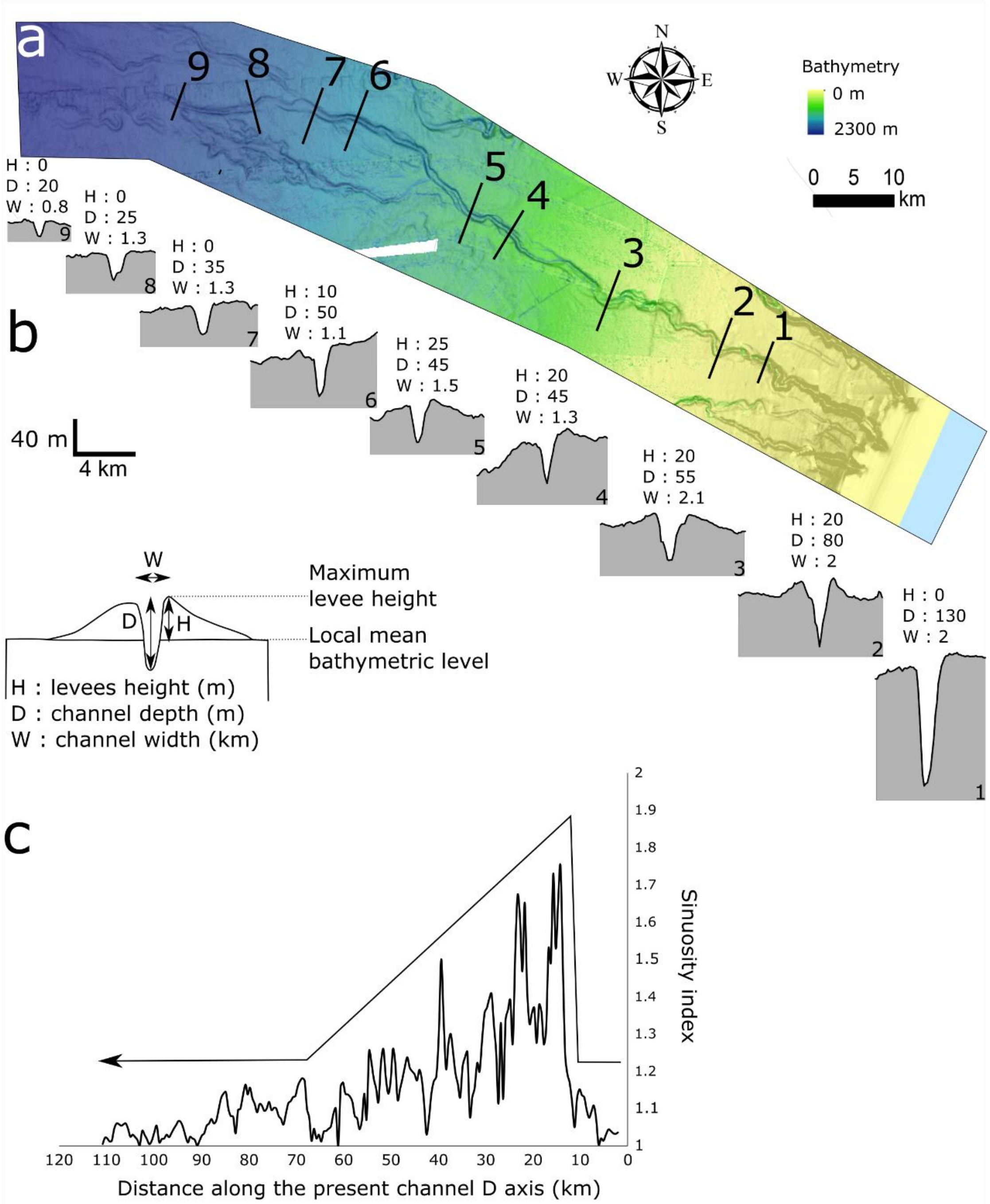
Figure 6: Interpreted gradient-shaded map of the Ogooue Fan showing the main features of the fan. A, B, C, D, E and F are the six main channels discussed in the text. The sand/shale ratio of the cores are shown (Sa:Sh) as well as the maximum sand-bed thickness in each core (max sand). A close-up view of the red rectangle is presented in Figure 8Figure 8.

Analysis of the seafloor data (bathymetry and acoustic imagery) reveals the different domains of the Ogooue sedimentary system and the different architectural features of the Ogooue Fan (Figure 6~~Figure 6~~).

The Gabon continental shelf is relatively narrow, decreasing in width from 60 to 5 km toward the Mandji Island (Figure 3~~Figure 3~~). The slope is characterized by two main topographic features: (1) the Mount Loiret, a guyot located just west of the Manji Island, which forms a bathymetric obstacle on the upper slope and (2) a ramp of several tributary canyons located south of the Mount Loiret (Figure 3~~Figure 3~~). This ramp is composed of several wide and deep canyons (several hundreds of meters deep and 2-3 km wide near the canyons head), with a "V-shape" morphology and which heads reach the shelf break. Several thinner and shallower incisions are located between these deep canyons. They are less than 100 m deep and 1 km wide and their heads are located between 200 and 400 m water depth (Figure 3~~Figure 3~~). The continental shelf and the slope present low backscatter values except for the canyons, ~~that~~which correspond to very high backscatter value (Figure 4~~Figure 4~~).

The transition between the continental slope and the continental rise, between 1,200 and 1,500 m water depth, is marked by a decrease in the slope gradient from a mean value of 2.3° to 0.9°. At this water depth, several canyons merge to form five sinuous channels (B to F in Figure 6~~Figure 6~~). These channels appear with higher backscatter value than the surrounding seafloor (Figure 4~~Figure 4~~). These sinuous subparallel channel-levees complexes extend down to 2,200 m water depth with a general course oriented toward the north-west (Figure 6~~Figure 6~~ and 7). At 2,200 m water depth, the southernmost channel (channel F in Figure 6~~Figure 6~~) deviates its path toward the south-west.

The sinuosity of these channels decreases ~~Westward~~westward. Channel D sinuosity has been calculated over 2 km long segments (Figure 7~~Figure 7~~C). It is less than 1.1 along the first 13 km corresponding to the canyon part. From 13 to 40 km~~ d~~ the mean sinuosity is 1.4 and then decreases to less than 1.2 between 40 to 90 km from the head. Finally, the most distal part of the channel, from 90 km from the head, is very straight with a sinuosity index lower than 1.1 (Figure 7~~Figure 7~~C).



a
9
8
7
6
5
4
3
2
1
N
W
E
S
Bathymetry
0 m
2300 m
0 5 10
km
b
H : 0
D : 20
W : 0.8
H : 0
D : 25
W : 1.3
H : 0
D : 35
W : 1.3
H : 10
D : 50
W : 1.1
H : 25
D : 45
W : 1.5
H : 20
D : 45
W : 1.3
H : 20
D : 55
W : 2.1
H : 20
D : 80
W : 2
H : 0
D : 130
W : 2
40 m
4 km
W
D
H
Maximum levee height
Local mean bathymetric level
H : levees height (m)
D : channel depth (m)
W : channel width (km)
c
Sinuosity index
Distance along the present channel D axis (km)

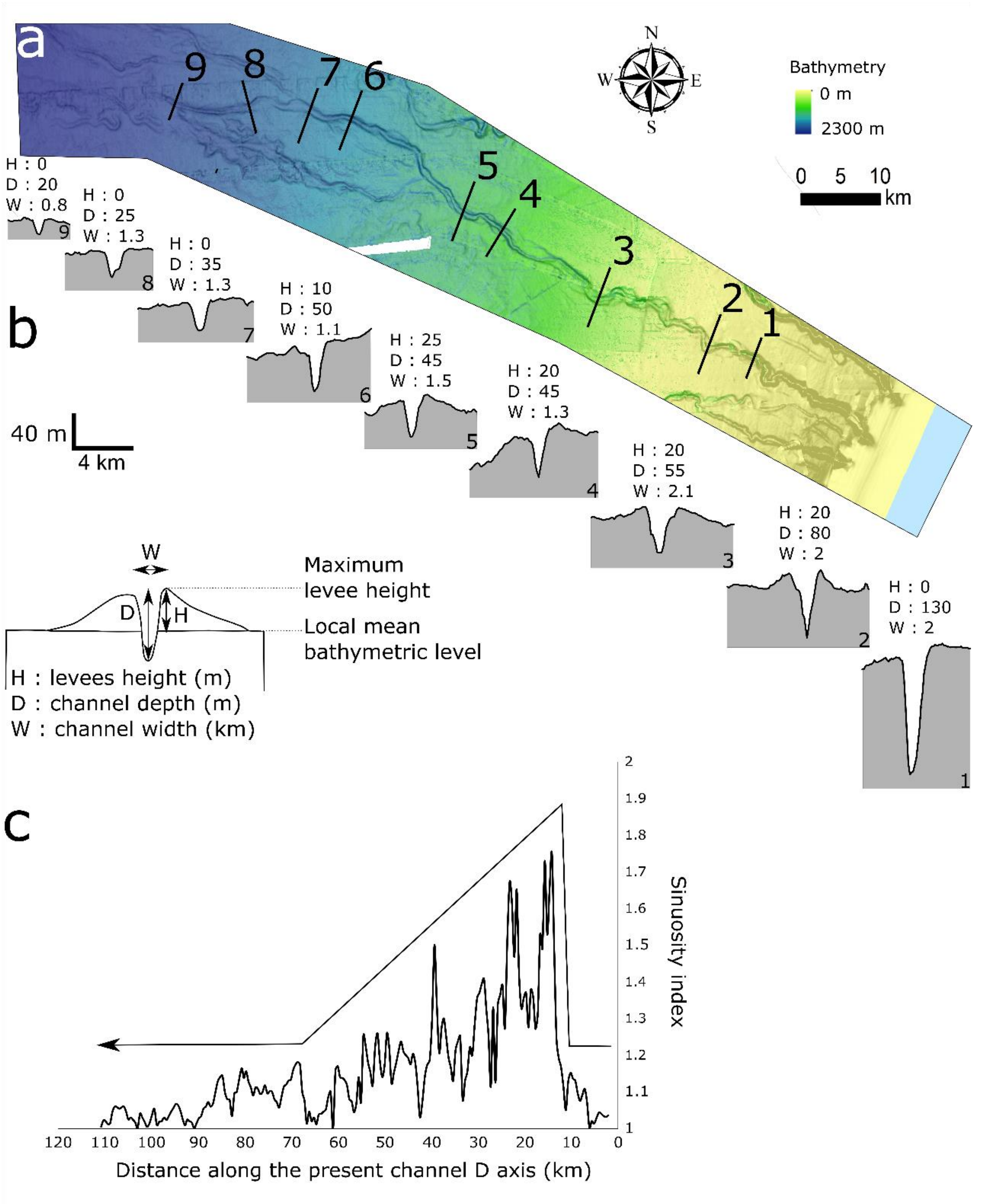
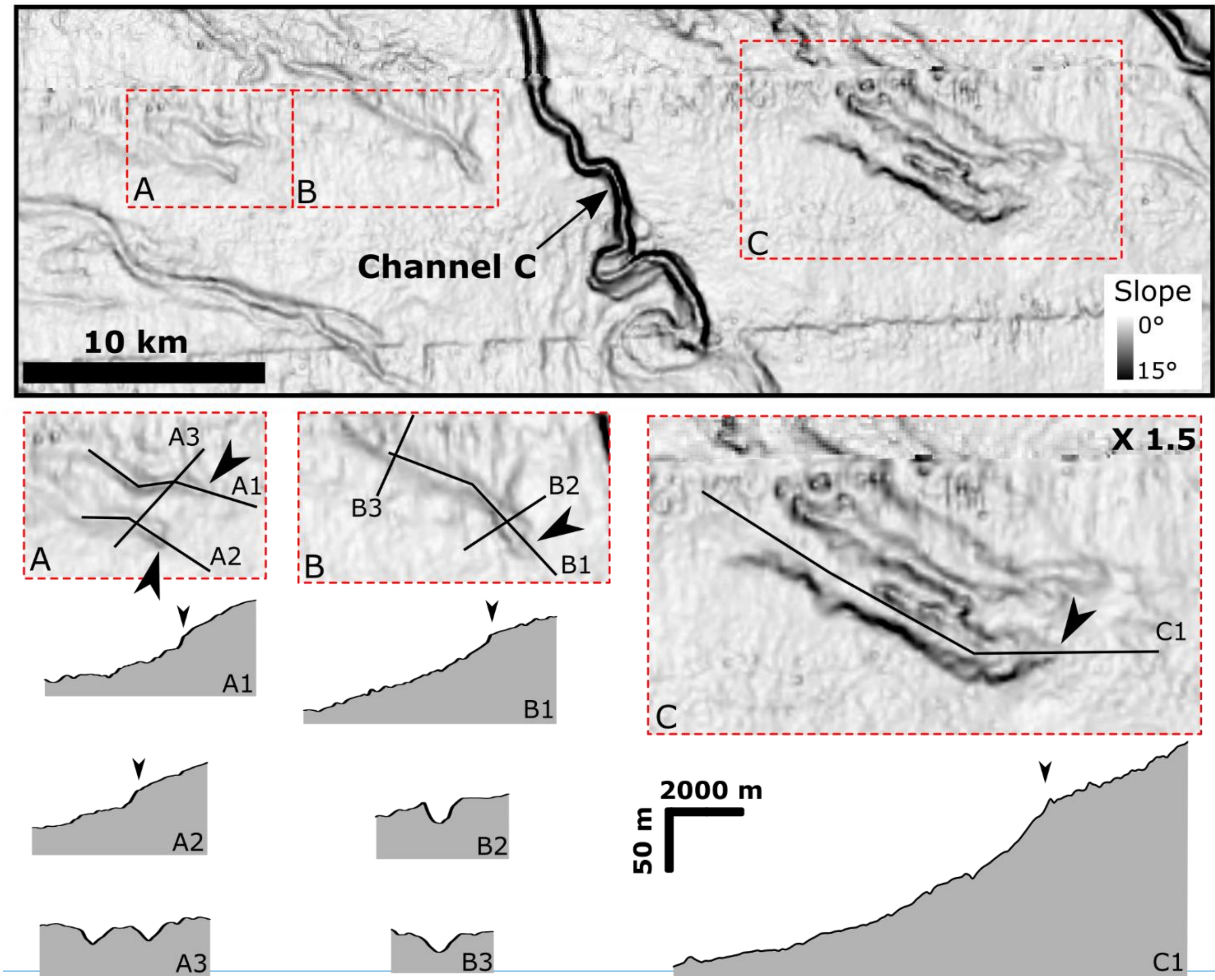


Figure 7: a) Detailed bathymetric map of channel D (location in Figure 2~~Figure 2~~) b) serial bathymetric profiles showing the evolution of the channel-levees along the slope and c) sinuosity down the channel D measured along 2 km channel segments.


A
B
Channel C
C
Slope
0°
15°
10 km
A3
A1
A2
A
B3
B2
B1
B
X 1.5
C1
C
A1
B1
A2
B2
2000 m
50 m
A3
B3
C1

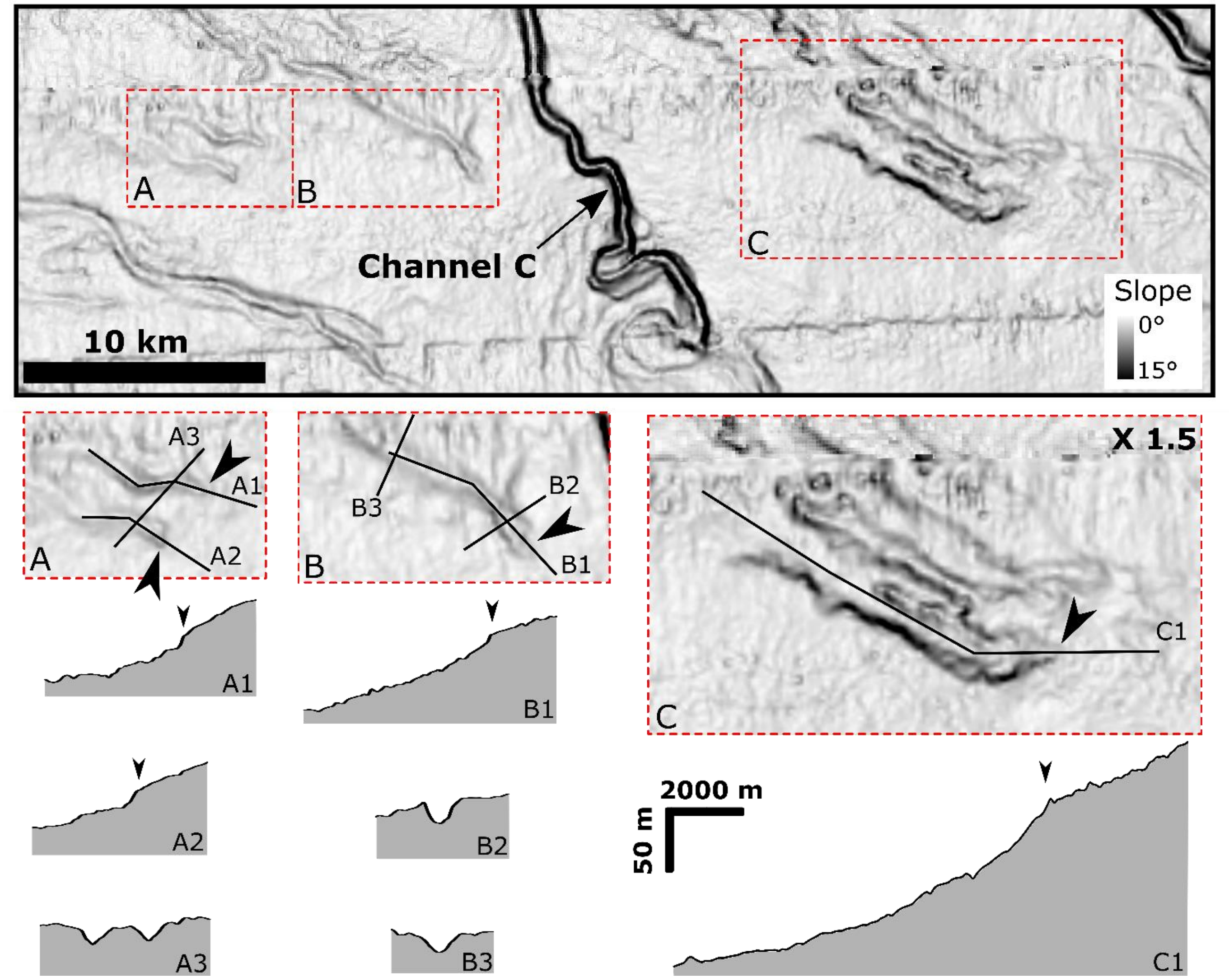


Figure 8: Close-up view of the gradient-shaded map showing erosional lineations (A and B) and amalgamated scours (C) in the central part of the system (location in Figure 6~~Figure 6~~).

Downslope, in the central part of the system, the seafloor located between 2,200 m and 2,500 m water depth~~,~~ presents numerous erosional features including scours, lineations and smaller, subsidiary channels, corresponding to channels with no headward connection with an obvious feeder system according to Masson et al. (1995) (Figure 8~~Figure 8~~). These erosional features appear on a very gentle slope area (0.3°) characterized by a heterogeneous medium backscatter facies (Figure 4~~Figure 4~~). At 2,500 m water depth, just south of the Sao-Tomé Island, the head of a large, 100 km long, mid-system valley appears (Figure 9~~Figure 9~~). This valley can be subdivided in two parts of approximately equal length with two different orientations. The upper part of the valley is oriented E-W, whereas the lower part is oriented NE-SW. This direction change is due to the presence of a rocky seamount located north of the valley and which deflects its course. The upper part of the valley is up to 15 km wide with numerous erosional scars and terraces on its flanks. The valley bottom is characterized by very high backscatter value and small internal erosion channels. Downstream, the valley becomes narrower with a "U" shape (Figure 9~~Figure 9~~, profile 5). Its flanks appear regular with no scar of down-flank mass deposits. The depth of the valley decreases from 60 m in its central part to only 10 m near its mouth. The area located south of the mid-system valley is characterized by a heterogeneous low-backscatter facies. Some erosional features and subsidiary channels are present but scarce.

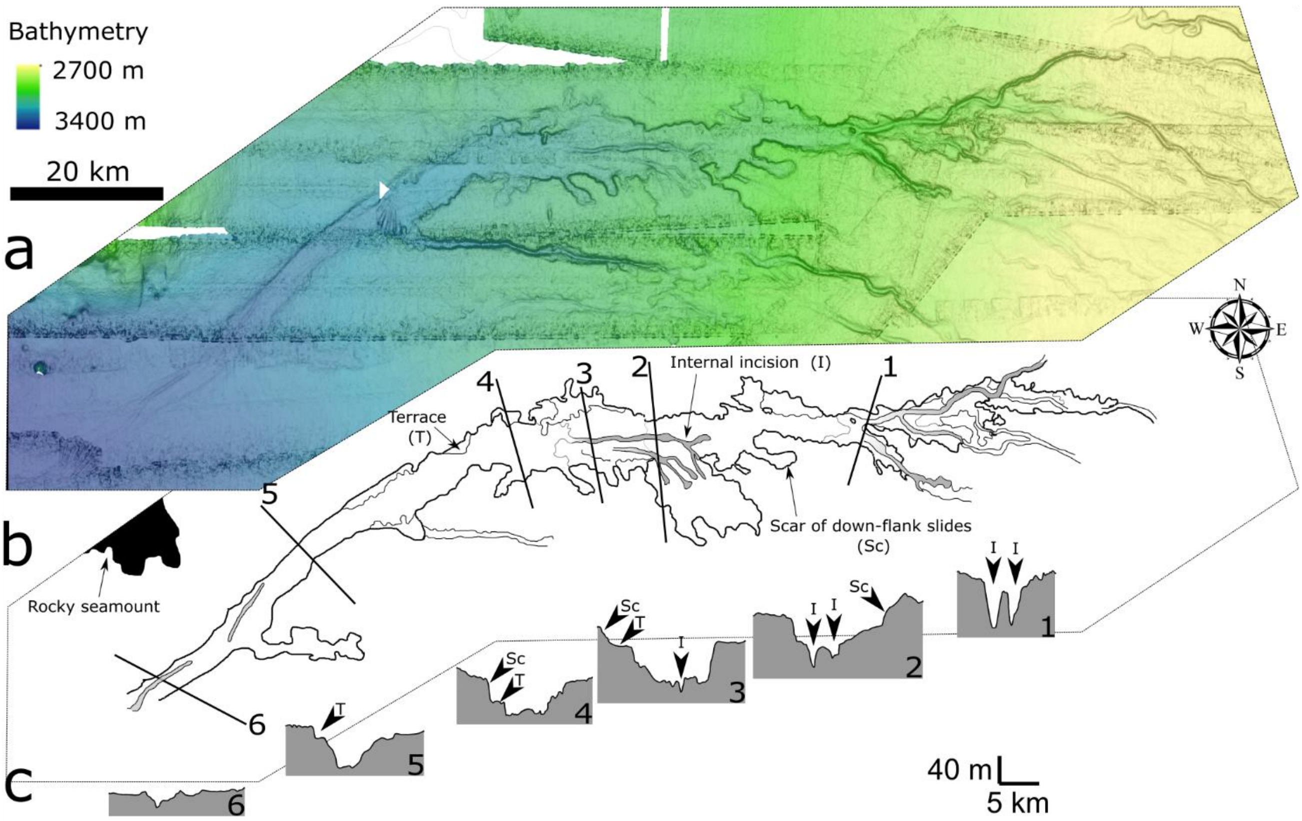
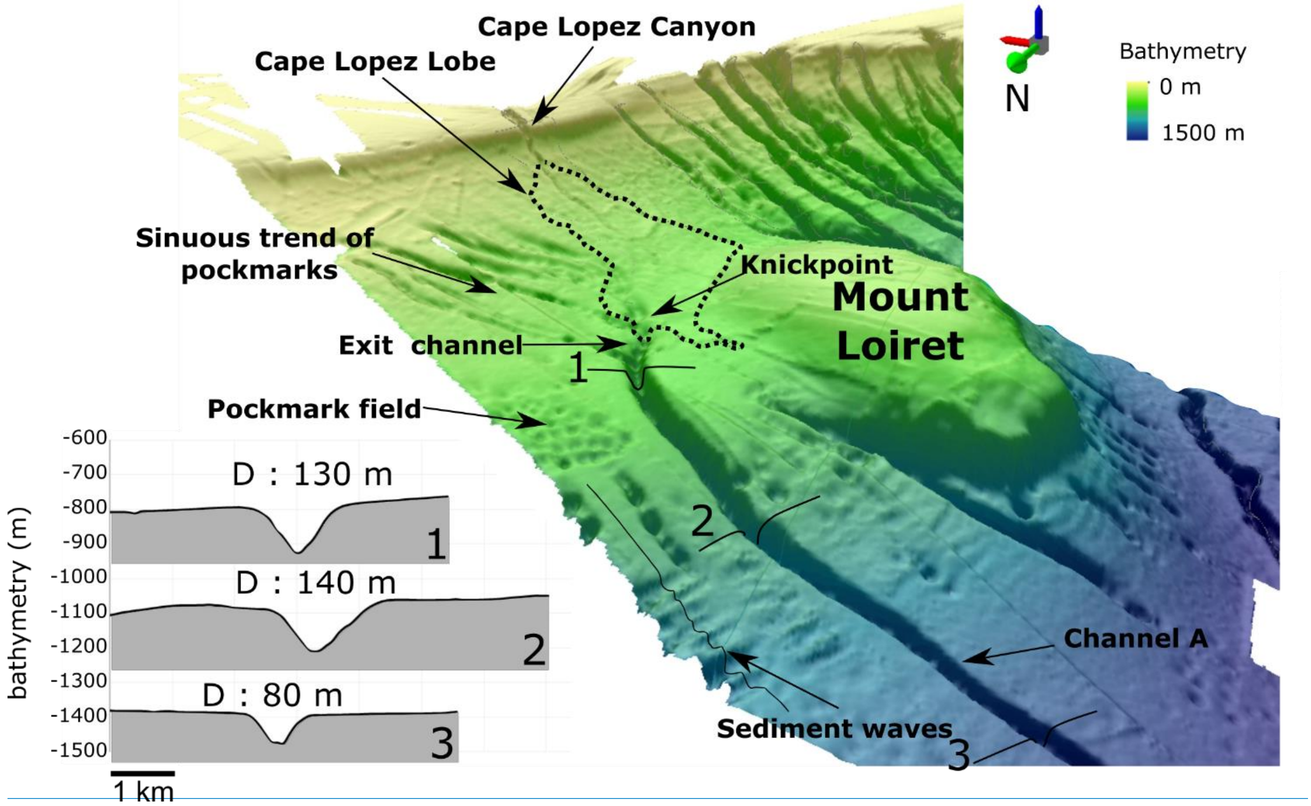


Figure 9: (a) Detailed bathymetric map of the mid-system valley of the Ogooue Fan between 2,700 and 3,400 m water depth; b) Interpretation of the main morphological features of the valley; c) Six transverse profiles of the mid-system valley extracted from the bathymetry data (Sc: scar of down-flank slides, I: internal incision, T; Terrace).

West of the mid-system valley outlet, the seafloor is very flat and shows only subtle morphological variations except for local seamounts. Few channel-like, narrow elongated depressions (maximum 10 m deep) presenting high backscatter values can be identified. These lineations are restricted to a long tongue of high backscatter at the mouth of the valley (Figure 2~~Figure 2~~b, Detail A). This tongue is globally oriented E-

W at the exit of the mid-system valley and then deflects toward the NW at 3,700 m water depth, following the steepest slope.

North of Mount Loiret, the upper slope presents a lower slope gradient compared to the south part and is characterized by the presence of numerous linear pockmark trains on the upper part and pockmarks fields on the lower part. This whole area has a very low and homogeneous reflectivity. Trace of active sedimentation on this part of the margin is only visible in association with the Cape Lopez Canyon~~, which is the only canyon located north of the Mount Loiret (~~ (Figure 3~~Figure 3~~). Cape Lopez Canyon terminates at 650 m water depth at an abrupt decrease in slope gradient (from more than 1.7° to 0.6°) caused by the present of Mount Loiret (Figure 10). This canyon is associated with a small intraslope lobe located just north-east of the Mount Loiret and referred as the Cape Lopez Lobe (Figure 10~~Figure 10~~) (Biscara et al., 2011). This northern system continues basinward with Channel A, the head of which is located in the vicinity of the Cape Lopez Lobe. At 2,200 m water depth, Channel A ends and its mouth is associated on the backscatter map with a fan-shaped area of very-high reflectivity, which is associated with some subsidiary channels and erosional marks (Figure 4~~Figure 4~~).


Cape Lopez Canyon
Cape Lopez Lobe
Bathymetry
0 m
1500 m
N
Sinuous trend of pockmarks
Knickpoint
Mount Loiret
Exit channel
1
Pockmark field
2
Channel A
Sediment waves
3
bathymetry (m)
-600
-700
-800
-900
-1000
-1100
-1200
-1300
-1400
-1500
D : 130 m
1
D : 140 m
2
D : 80 m
3
1 km

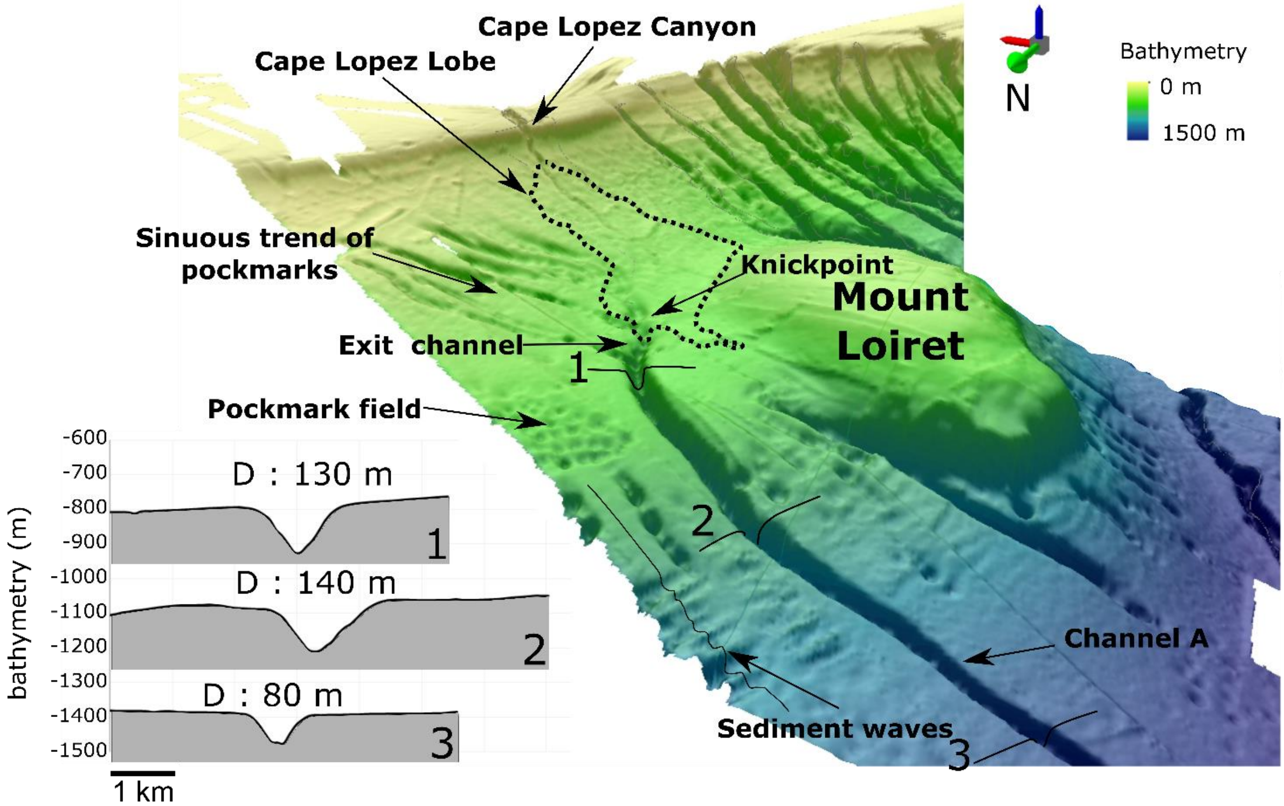
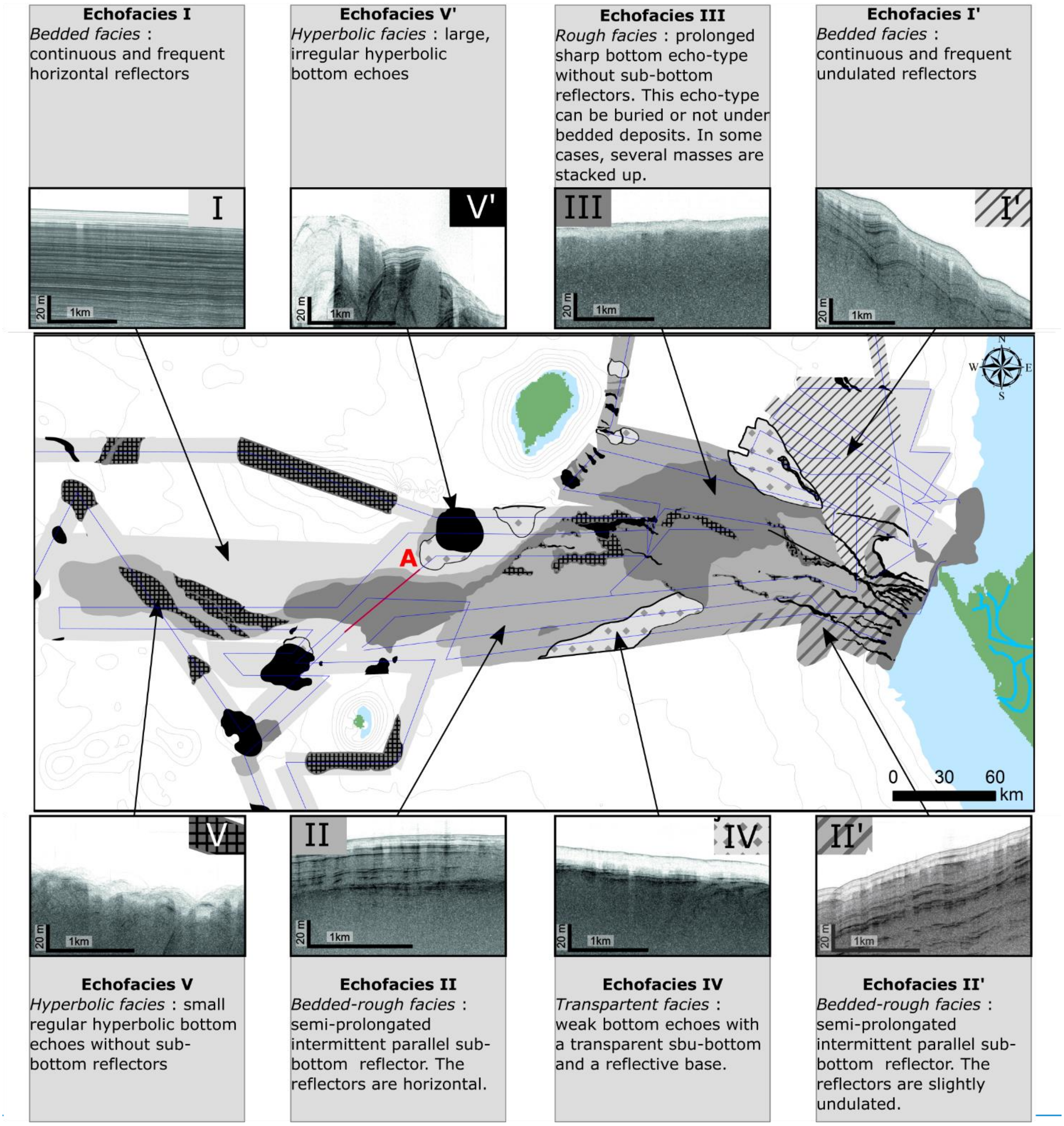


Figure 10: a) Three-dimensional representation of the Cape Lopez, Canyon, Cape Lopez Lobe and Channel A, b) three transverse profiles of Channel A. (Vertical exaggeration: 15).

~~45.3~~4.3 Echofacies ~~analysis and distribution~~classification



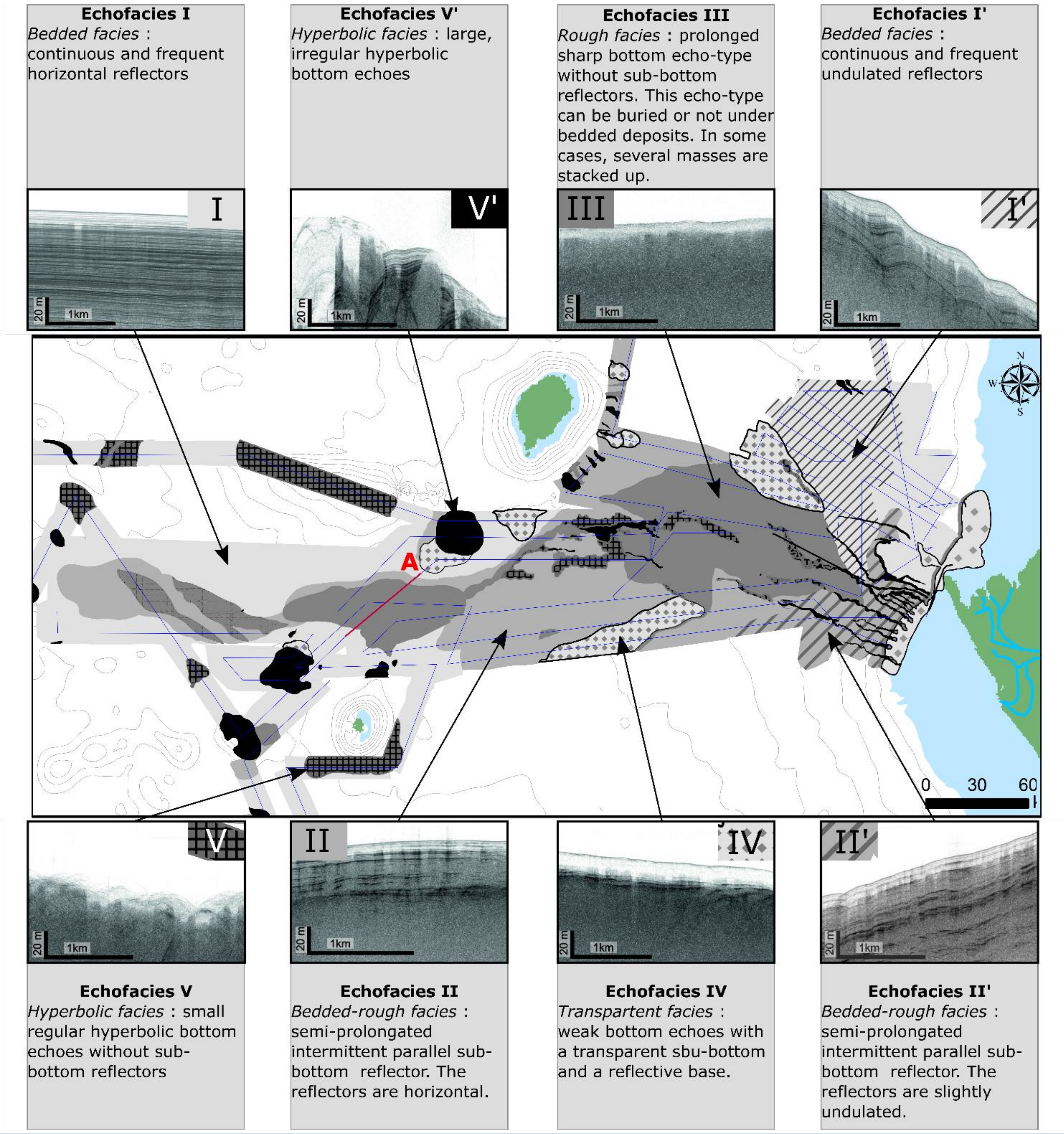


Figure 11: Echofacies map of the Ogooue Fan. Eight shades of grey represent the specific echofacies.

The main echofacies have been discriminated on the profiles based on amplitude, frequency and geometry of the reflections (Figure 11~~Figure 11~~). They have been grouped into five main classes: (I) bedded, (II) bedded-rough, (III) rough, (IV) transparent and (V) hyperbolic. Most transitions between echofacies are gradual.

The echofacies of the edge of the Gabonese shelf consists of ~~rough~~transparent echofacies ~~III~~IV (Figure 11~~Figure 11~~). ~~Core KC18 indicates that this area is dominated by fine-grained, structureless, terrigenous sedimentation~~North of the Mount Loiret, the continental slope presents bedded echofacies I. At 1,500 m, which corresponds to an increase in the slope gradient, echofacies transforms into echofacies I'. South of Mount Loiret, echofacies II and II' dominate on the continental slope.

The echomapping of the continental rise reveals the presence of different facies. The central part, just upstream of the mid-system valley, is characterized by rough echofacies III. Some large channels are marked by hyperbolic facies. South of the mid-system valley, facies II dominates. Echofacies IV is present in two main areas on the continental rise where they respectively form two lobe-shaped zones: one on the northern part, following the limits of the high-reflectivity area located at the mouth of channel A; the second in the southern part of the system in association with channel F. ~~North of the Mount Loiret, the continental slope presents bedded echofacies I, which evolves into echofacies I' down isobath 1,500 m which corresponds to an increase in the slope gradient. Previous studies have shown that bedded echofacies~~

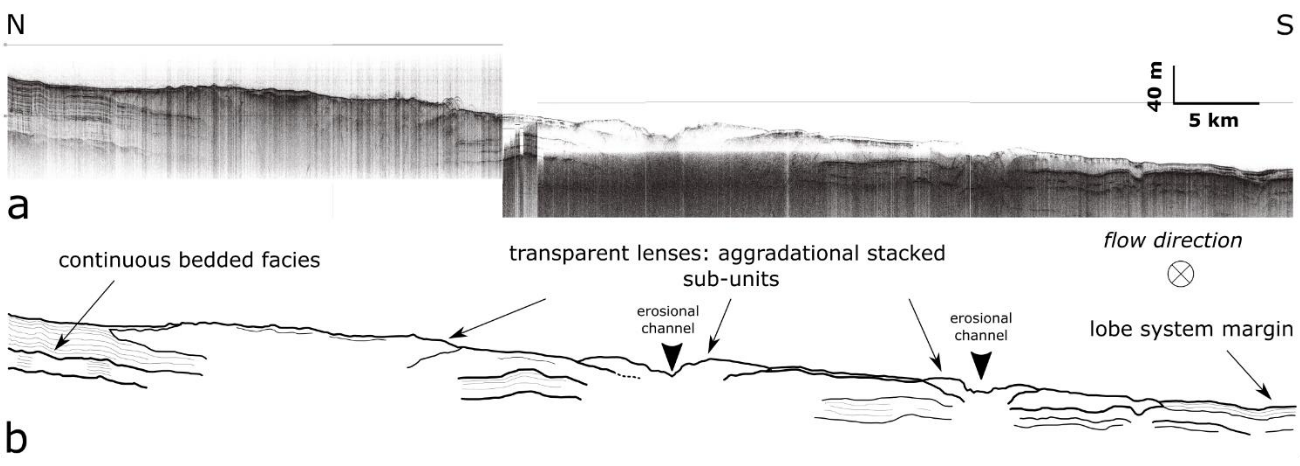


Figure 12: a) Transverse 3.5 kHz very-high resolution seismic line and b) line drawing in the upper distal lobe area, see Figure 11~~Figure 11~~for location of the line.

In the abyssal plain, the area of the elongated tongue noticeable on the backscatter data presents different echofacies. Based on the 3.5 kHz profiles, it can be subdivided into two main domains. The upstream part, at the outlet of the mid-system valley, is characterized by multiple aggradational stacked transparent sub-units from 10 to 30 meters thick are visible on the seismic lines (Figure 12). The downstream part presents is characterized by echofacies (II) associated with hyperbolic echofacies (V).

On the edge of this tongue, high-penetration bedded facies (I) is dominant. Facies V' forms some patches on the seafloor and correspond to seafloor mounts.

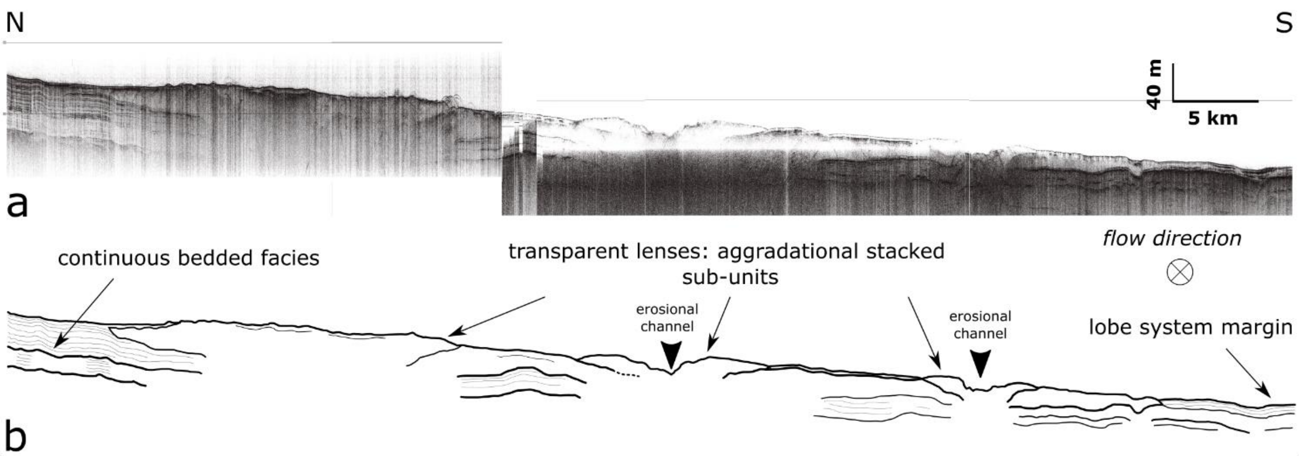
Facies V and IV are also present and form lenses around the island of Sao-Tomé and Annobon.

Based on previous studies and core samples, we speculate the following links between echofacies, type of sediments and associated depositional processes:

-Bedded facies (I, I') are commonly associated with alternating sandy and silty beds (Damuth, 1975, 1980a; Pratson and Laine, 1989; Pratson and Coakley, 1996; Loncke et al., 2009) or with hemipelagic sedimentation ~~(Gaullier and Bellaiche, 1998). The very low reflectivity of the area and the absence of any channel suggest that only hemipelagic sedimentation occurs in this area. The wavy aspect of echofacies I' is certainly due to the post-deposition deformation of the hemipelagic sediments (Bouma and Treadwell, 1975; Jacobi, 1976; Damuth and Embley, 1979; Damuth, 1980b).~~ when associated with very low reflectivity this is confirmed by facies description of cores KC16 and KC02 (Gaullier and Bellaiche, 1998).

~~South of Mount Loiret, echofacies II~~ - Rough and bedded-rough facies (II, II' ~~dominate on the continental slope. Despite the lack of sampling, the presence of discontinuous seismic reflectors can indicate the presence of coarse~~, III), as described in Loncke et al. 2009, are attributed to coarse-grained ~~sediment interpreted as turbidites~~turbidite (Damuth, 1975; Damuth and Hayes, 1977). ~~The echo-mapping of the continental rise reveals the presence of different facies. The central part, just upstream of the mid-system valley, is~~ Damuth and Hayes (1977) have shown that a quantitative relationship exists between the relative abundance of coarse sediment in the upper few meters of the seafloor and the roughness of the echo-types. Rough echofacies characterized ~~by rough echofacies III~~areas that ~~suggests~~contain the ~~presence of a high proportion~~highest concentrations of coarse~~-grained~~ grains, like lobe areas, whereas bedded-rough facies contain little coarse sediments. ~~Some large channels are marked by hyperbolic facies certainly due to the irregular and steep seafloor. South of the mid-system valley, facies II dominates.~~ Core KC10 and KC15, collected in ~~this~~an area of facies II, indicates the alternation of clayey and sandy layers but with a predominance of fine-grained sediments (Figure 5~~Figure 5~~).

- Transparent ~~Echofacies IV is present in two main areas on the continental rise where they respectively form two lobe-shaped zones: one on the northern part, following the limits of the high-reflectivity area located at the mouth of channel A; the second in the southern part of the system in association with channel F. This echo-~~facies (IV) commonly corresponds to structureless deposits ~~without internal organization~~ due to mass-flow processes such as debris flows (Embley, 1976; Jacobi, 1976; Damuth, 1980a, 1980b, 1994) but it can also characterize basinal fine-grained turbidites (Cita et al., 1984; Tripsanas et al., 2002). ~~Core KC21, collected in the northern area indicates homogeneous silty-clay sediments with numerous detrital debris similar to those collected near the continental shelf.~~In this study transparent facies is also associated with fine-grained, structureless, terrigenous sedimentation of the shelf (Core KC18).



- Hyperbolic facies (V, V') is linked to the degree of roughness of the seafloor topography. Large, irregular hyperbolae (V') are generally associated with abrupt topographies such as seamounts or canyons and deep channels. Small regular hyperbolae (V) are commonly associated with deposits generated by debris-flow (Damuth, 1980a, b, 1994).

Figure 12: a) Transverse 3.5 kHz very-high resolution seismic line and b) line drawing in the upper distal lobe area, see Figure 11for location of the line.

In the abyssal plain, the area of the elongated tongue noticeable on the backscatter data presents different echofacies. Based on the 3.5 kHz profiles, it can be subdivided into two main domains. The upstream part, at the outlet of the mid-system valley, is characterized by rough echo character but with a specific organization: multiple aggradational stacked transparent sub-units from 10 to 30 meters thick are visible on the seismic lines (Figure 12). This organization is characteristic of sandy lobes deposits (Kenyon et al., 1995; Piper and Normark, 2001). Core KC11, collected in this environment, presents several decimeters-thick sandy layers and a several meter-thick disorganized sandy-clay units interpreted as a slump. The downstream part presents bedded-rough echofacies (II) associated with hyperbolic echofacies (V). Core KC15 intersected fine-grained sediments and several silty layers corresponding to the distalmost turbidites.

On the edge of this tongue, high-penetration bedded facies (I) is dominant. The highly continuous parallel bedding indicates hemipelagic sedimentation with no coarse-grained fraction, which is confirmed by core KC16 and core KC02 both composed of alternating carbonate-rich and carbonate-poor clay sediments. Facies V' forms some patches on the seafloor and correspond to seafloor mounts. The hyperbolic facies is due to the steep slopes and the irregular topography.

Facies V and IV are also present and form lenses around the island of Sao-Tomé and Annobon. These features indicate some downslope sedimentary transfer from these islands. The limited area covered by these facies suggests short transport by sliding.

~~49~~5 Interpretation and discussion

~~49.1~~5.1 Sedimentary processes along the fan

The Ogooue Fan ~~is~~could be classified as a delta-fed passive margin deep-sea ~~mud/sand-rich~~ submarine fan according to ~~the classification of~~ Reading and Richards (1994). However, analysis of sub-surface data (bathymetry, acoustic imagery and 3.5 kHz ~~echo-characters~~echocharacters) reveals a great variability of sediment processes in the different domains of the margin, controlled by variations in slope gradient and the presence of seamounts (Figure 13a).

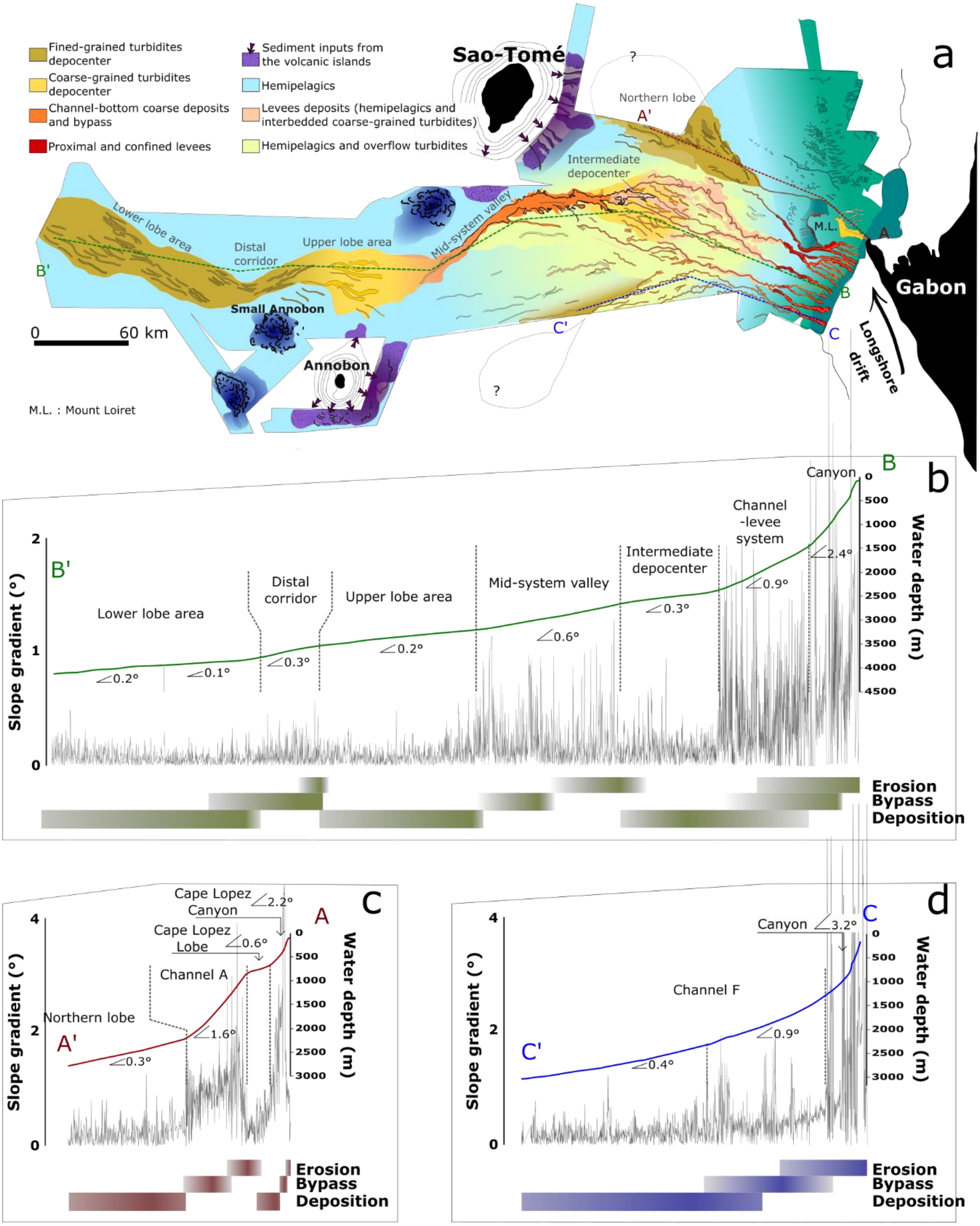


Figure 13: a) Synthetic map showing the architecture and the recent sedimentary processes of the Ogooue Fan determined by imagery and echofacies mapping; b) c) and d) Longitudinal profiles from the bathymetric data along the central, northern and southern part of the Ogooue Fan and slope gradient (in degree, measured every 100m). The differences in slope gradient along the transects are associated with the main sedimentary processes encountered along the slope.

~~49.1.1~~5.1.1 ~~Canyons~~Upslope area and canyons system

Cores collected in the upslope area (KC18 and KC17) show mostly hemipelagic sediments with a very low carbonate content. This reflects significant detrital flux associated with proximity to the Ogooue platform and the influence of the Ogooue river plume. Erosional processes ~~predominate~~are also active on the upper part of the slope as indicated by the presence of numerous tributary canyons (Figure 3~~Figure 3~~). Based on the comparison of the canyon depths, widths and head positions, we observe the existence of two types of canyons as described in Jobe et al. (2011) along the Equatorial Guinea margin. The canyons presenting a deep (> hundreds of meters deep) "V" shape and which indent the shelf edge are type I canyons (*sensu* Jobe et al., 2011), whereas the shallower canyons (<100 m deep) with a "U" shape and which do not indent the shelf are type II canyons (*sensu* Jobe et al., 2011). The difference between these two types of canyons indicates different initiation and depositional processes. Type I are commonly associated with high sediment supply and the canyons initiation and morphology are controlled by frequent sand-rich erosive turbidity currents~~ and mass-wasting processes~~ (Field and Gardner, 1990; Pratson et al., 1994; Pratson and Coakley, 1996; Weaver et al., 2000; Bertoni and Cartwright, 2005; Jobe et al., 2011). Core KC19 collected down of a type I canyon shows two several meters-thick sandy successions corresponding to top-cut-out Bouma sequences (Ta) interbedded with the upper slope hemipelagites. These sandy turbidites, which are the thickest sand beds recorded in all the cores (Figure 6), indicate the occurrence of high-density turbidity currents flowing down this canyon. In contrast, Type II canyons are found in areas of low sediment supply. Their initiation is attributed to retrogressive sediment failures and subsequent headward erosion (Shepard, 1981; Twichell and Roberts, 1982; Stanley and Moore, 1983). ~~Their~~The evolution of these canyons is controlled by ~~depositional processes involving~~ fine-grained ~~sediments -~~sedimentation: hemipelagic deposition and dilute turbidity currents ~~-~~ that can be carried over the shelf ~~and upper slope~~ into the canyon heads~~ but without~~. These sedimentary processes do not cause significant erosion in the canyons (Thornton, 1984).

North of the Mount Loiret, the fine-grained sedimentation has completely infilled several type II canyons~~ creating~~. The fluid migration from the previously deposited coarse-grained sediments inside the paleo-canyons has created sinuous trains of pockmarks. These pockmarks have been previously described in Pilcher and Argent (2007). Variations in the localisation of coarse-grained sediment supplies play a key role on the development of the two types of canyons. Along the central Gabonese shelf, the very recent development of the Mandji Island 3,000 years~~-~~ BP (Giresse and Odin, 1973; Lebigre, 1983) concentrated most of the coarse sediments near the Cape Lopez and favoured the construction of the presently active Cape Lopez Type I canyon~~, which is presently active~~ (Biscara et al., 2013).

~~49.1.2~~5.1.2 Channels system

The transition from deep canyons to sinuous channels with ~~external~~ levees ~~(*sensu* Kane and Hodgson, 2011)~~ is related to a decrease in slope gradient from the continental slope (> 2°) to the continental rise (< 1~~°). The sinuous channel-levees systems develop on a relatively gentle slope (0.9°) from 1,500 to 2,200 m water depth. These channels are mainly erosive in their axial part (Normark et al., 1993) while deposition occurs on low-developed external levees (25 m maximum levees height for channel D; Figure 7).~~°) that slows turbidity currents and reduces their erosional power. The external levees of the four central channels (B, C, D and E in Figure 2~~Figure 2~~) show high reflectivity ~~that suggests frequent turbidity currents overspill. These channels are deeply incised in the seafloor~~ (compared to the surrounding seafloor which indicates a different sedimentological nature. This suggests that deposition occurs on the low-developed external levees (25 m maximum levees height for channel D; Figure 7) due to turbidity currents overspills. External levee deposits have been sampled by core KC13, which shows numerous turbidites made up of centimeter-thick, normally graded, parallel or ripple cross-laminated of silt and fine sands (Figure 5). In their axial part, these channels are mainly erosive (Normark et al., 1993) as indicated by their deep incision in the seafloor: average 70 m deep for channel D and 90 m deep for channel A; (Figure 7~~Figure~~

~~7~~ and Figure 10~~Figure 10~~) below the associated levees, when present. This feature is similar to the modern Congo Channel (Babonneau et al., 2002) and is opposed to the morphology of aggrading channels (such as the Amazon Channel) where the thalweg is perched above the base of the levees system (Damuth, 1995). This entrenched morphology prevents extensive overflow of turbidity currents and is the probable cause of low development of external levees and ~~inhibits~~limits channel~~ bifurcation~~ by avulsion. It has been proposed for the Congo Channel that the entrenched morphology of the channel confines the flow and ~~keeps the energy~~ maintains a high ~~enough~~velocity. The high velocity of the flow enables the sediments to ~~allow a transport of sediment~~be transported to very distant areas (Babonneau et al., 2002).

Several studies have documented that sinuosity of submarine channels increases with time (Peakall et al., 2000; Babonneau et al., 2002; Deptuck et al., 2003, 2007; Kolla, 2007). The sinuous upper parts of the channels (1.3 < sinuosity < 1.75 for channel D; Figure 7~~Figure 7~~C) have consequently undergone a long history whereas the distal straighter parts of the channels are in a more immature stage. Moreover, the height of the external levees and the depth of the channels both decrease in the lower parts of the ~~channels~~channel system (Figure 7~~Figure 7~~). These morphological changes are due to a slope gradient decrease (< 0.5° from transect 6 along channel D; Figure 7~~Figure 7~~) that progressively slows down the flow velocity and reduces ~~its~~the erosional power of the turbidity current. Simultaneously, deposition of fine particles by spilling of the upper part of the flow on the external levees leads to a progressive decrease of the fine-grained fraction transported by the channelized flows (Normark et al., 1993; Peakall et al., 2000).

At 2,200 m water depth, the appearance of numerous erosional features such as isolated and amalgamated spoon–shaped scours (Figure 8~~Figure 8~~ C1), erosional lineations and subsidiary channels with limited surface expression (10-20 m deep, Figure 8~~Figure 8~~ B2, B3) are characteristic of the channel lobe transition zone (Figure 8~~Figure 8~~) (Kenyon et al., 1995; Wynn et al., 2007; Jegou et al., 2008; Mulder and Etienne, 2010). The appearance of these features correlates with a second abrupt decrease in slope gradient (from 0.9° to 0.3°) and with the transition from bedded~~ rough~~ echofacies with low penetration to rough ~~echo-facies~~echofacies indicating a change in the sedimentary process and suggest a high sand/mud ratio. This area corresponds to deposition by spreading flows in an unchanneled area referred as the intermediate depocenter in Figure 13 and covering area surface of ca. 4,250 km². However, the low penetration of the 3.5 kHz echosounder and the limited number of seismic lines in this area did not allow a more detailed interpretation of the sedimentary processes in this part of the system.

~~49.1.3~~5.1.3 Mid-system valley and distal lobe complexes

The presence of a steeper slope downslope of the intermediate depocenter (0.6°) led to the incision of the multi-sourced mid-system valley, which acts as an outlet channel for turbidity currents that are energetic enough to travel through the flatter depositional area (Figure 13b). The numerous erosional scars present in the upstream part of the valley ~~is multi-sourced and~~suggest that this section has migrated upstream by retrogressive erosion, whereas the downstream part appears more stable with a straighter pathway and steeper flanks, these features being similar to the Tanzania Channel described by Bourget et al. (2008). According to the available bathymetric data, the volume of sediment removed from the mid-system valley is between 8 and 10 km^3. The pathway of the valley seems to be controlled by the seafloor topography as the valley deviates near the rocky seamount located west of Sao-Tomé. This large mid-system valley ~~corresponds to a single feeding “source” for the lower fan and, consequently, the final depositional area is located downstream of the valley.~~delivers sediments to the lower fan.

At the outlet of the mid-system valley, the echofacies shows an area mainly characterized by rough echofacies (III) forming stacked lenses. This organization is characteristic of sandy lobes deposits (Kenyon et al., 1995; Piper and Normark, 2001). This area, referred as the upper lobe area in Figure 13, constitutes the main lobe complex (*sensu* Prélat and Hodgson, 2013) of the Ogooue Fan. Core KC11 shows that coarse-grained turbidity currents are deposited in the proximal part of the lobe complex. The abrupt transitions between erosional/bypass and depositional behavior observed notably at the mouth of the mid-system valley is the result of hydraulic jumps affecting flows when they become unconfined between channel sides and spread laterally (Komar, 1971; Garcia and Parker, 1989). According to the seismic data, the depositional area of the lobe complex is ~ 100 km long, reaches ~ 40 km in width, spreads over 2,860 km² and reaches up to 40 m in thickness. The transparent lenses are interpreted as ~~lobe elements and~~ lobes: they seem to be bounded by erosive bases and separated vertically by fine-grained units (Mulder and Etienne, 2010; Prélat and Hodgson, 2013). Some incisions (< 15 m deep) are imaged on the top surface of the lobes; two of them are visible in Figure 12~~Figure 12~~. The area where incisions are present is interpreted as the channelized part of the lobe complex. This lobe area presents a gentle slope (0.3°) oriented north-south, suggesting that topographic compensation would shift future lobe ~~element~~ deposition southward. However, the few numbers of seismic lines do not allow the precise internal geometry and the timing of the construction of the different lobe units.

This depositional area is not the distalmost part of the Ogooue Fan. West of this lobe, evidences of active sedimentation are visible on the reflectivity map (Figure 2~~Figure 2~~, Figure 4~~Figure 4~~). The reflectivity map shows high-backscatter finger-shape structures suggesting pathways of gravity flows (Figure 2~~Figure 2~~b, detail A). These lineations (< 10 m deep) are concentrated in a 20 km wide corridor just west of the lobe area and then form a wider area extending up to 550 km offshore the Ogooue Delta. This part of the system follows the same ~~scheme~~pattern as the one previously described between the intermediate depocenter and the upper lobe area (Figure 13b). The corridor appears on a segment of steeper slope (0.3°) just at the downslope end of the upper lobe area (0.2°). This corridor, which disappears when the slope becomes gentler (0.1°), ~~was certainly~~is certainly dominated by sediment bypass (*sensu* Stevenson et al., 2015). Core KC15, located downstream of this corridor in the lower lobe area, is composed of very thin silty turbidites corresponding to the upper parts of the Bouma sequence interbedded with hemipelagic deposits. The upper lobe acts as a trap for the basal sand-rich parts of gravity flows and the lower lobe area receive only the upper part of the flows, which is composed of fine-grained sediments. The spatial distribution of facies suggests a filling of successive depocenters with a downslope decrease of the coarse-grained sediment proportion (Figure 6).

Considering the sedimentary facies of core KC15 located downstream this corridor, we can assume that this corridor was formed by the repeated spill-over of the fine-grained top of turbidity currents over the upper lobe area. This architecture suggests that this corridor ~~is dominated by sediment bypass (*sensu* Stevenson et al., 2015).~~. On the most distal segment with a very low slope gradient (0.1-0.2°) sediment deposition dominates.

~~49.1.4~~5.1.4 Isolated systems

On the northern part of the slope, the isolated system composed of the Cape Lopez Canyon, Cape Lopez intraslope lobe, channel A and northern lobe follows the same ~~scheme (Figure 13c). Cape Lopez Canyon terminates at 650 m water depth at an abrupt decrease in slope gradient (from more than 1.7° to 0.6°) caused by the present of the Mount Loiret (Figure 10).~~pattern (Figure 13c). The Cape Lopez intraslope lobe occupies a small confined basin, 6 km wide and 16 km long and covers an area of 106 km². This lobe appears very similar with the "X fan" described in Jobe et al. (2017) on the Niger Delta slope (8 km x 8 km, 76 km²) and is in the same size range as the intraslope complexes studied in the Karoo Basin by Spychala et al. (2015) (6-10 km wide and 15-25 km). The two successive depositional areas, composed by the Cape Lopez lobe and the northern lobe, are located on areas with a low slope gradient (0.6-0.3°) whereas erosion and sediment bypass dominate on segments of steeper slope gradient (1.6°). The high slope gradient between the two depositional areas favored the construction of a straight deeply entrenched channel (>100 m deep near the knickpoints) without levee (Figure 7b) instead of a large valley similar to the central mid-system valley.

In the southern part of the fan, channel F transports sediments southward (Figure 13d). At 2,200 m water depth, a transparent echofacies appears associated with the pathway of this channel. This echofacies suggests that sediment transported by this channel might be partly deposited in this area by turbidity current overflow. This channel might also be associated with a depositional lobe; however, the area covered by the MOCOSED survey does not allow us to image it.

~~49.2~~5.2 The Ogooue Fan among other complex slope fans

The Ogooue Fan develops on a stepped~~-~~ slope (Prather, 2003) which creates a succession of depositional areas on segments with gentle slope (referred as ‘steps’ in Smith~~,~~ (2004)) and segments of steeper slope (“ramps” in Smith, 2004) associated with erosion or sediment bypass (Figure 13) (Demyttenaere et al., 2000; Deptuck et al., 2012; O’Byrne et al., 2004; Smith, 2004). The depositional behavior in these systems is guided by an equilibrium profile of the system that forms preferential areas of sedimentation or erosion (Komar, 1971; Ferry et al., 2005). As described in the conceptual model of O’Byrne et al. (2004), erosion is favored where local gradient increases, the eroded sediments being delivered downstream resulting in a local increase in sediment load (O’Byrne et al., 2004; Gee and Gawthorpe, 2006; Deptuck et al, 2012). This kind of fan geometry is common along the West African margin where abrupt changes in slope gradient and complex seafloor morphology are inherited from salt tectonic movement (Pirmez et al., 2000; Ferry et al., 2005; Gee and Gawthorpe, 2006; Gee et al., 2007). Deptuck et al. (2012) has described the influence of stepped~~-~~ slope on sedimentary processes along the western Niger Delta. ~~He~~They showed that differences of slope gradient between ramps (0.8° to 2.1°) and steps (0.3° to 1.1°) induce the transition from vertical incision and ~~sediments~~sediment removal to preferential ~~sediments~~sediment accumulation (Deptuck et al., 2007; Deptuck, 2012). Gradient changes along the Gabonese margin are however lower than the ones reported in Deptuck et al.~~,~~ (2012) and variation in slope gradient of 0.2° appears to be enough to modify sedimentary processes. The impact of subtle changes of slope gradients has already been highlighted by studies of the Karoo basin (Van der Merwe et al., 2014; Spychala et al., 2015; Brooks et al., 2018) and Moroccan margin where sedimentary processes are controlled by very subtle gradient changes (<0.1°) (Stevenson et al., 2013; Wynn et al., 2012).

Moreover~~,~~ in the ~~case of the~~ modern Ogooue Fan~~, and conversely to what is observed in the Congo and Niger systems~~, the presence of several bathymetric highs including the volcanic islands of the CVL and the Mount Loiret ~~constitutes additional stresses~~acts as obstacles for the flows and creates a more complex slope profile. ~~These~~Such topographic highs are not present in the Congo and Niger systems. The bathymetric highs on the Ogooue fan area induce a lateral shift of the pathways of different channels as well as the pathway of the mid-system valley and form several downslope depositional lobes such as the Cape Lopez lobe that is constrained by the presence of the Mount Loiret. Several complex-slope systems have already been described in the literature with slope complexity due to salt-related deformations (e.g. Gulf of Mexico (Prather et al., 1998; Beaubouef and Friedmann, 2000), offshore Angola (Hay, 2012) or basin thrusting (offshore Brunei; McGilvery and Cook, 2003, Markan margin; Bourget et al., 2010). For these systems, the slope evolves rapidly, and sedimentation and erosion are unlikely to establish an equilibrium profile. In contrast, the Gabonese margin reached a mature evolutionary stage with ~~only weak and slow salt tectonic activity (Chen et al., 2007), and sedimentation~~salt diapir piercement rate much lower than deposition rate and thus no conspicuous effect of salt tectonics on the deposition of overburden sediment (Chen et al., 2007). Sedimentation and erosion certainly dominate the short-term evolution of the slope. The Ogooue Fan appears to be much more similar to the morphology of the Northwest African margin where the Madeira, the Canary and the Cape Verde islands create a complex slope morphology along the Moroccan and Mauritanian margin (Masson, 1994; Wynn et al., 2000, 2002, 2012).

~~50.0 Sedimentary facies distribution~~

~~The main processes involved in the deposition of the Upper Quaternary sediments of the Ogooue Fan are pelagic and hemipelagic suspension fall-out together with turbidity currents. Fine-grained pelagic/hemipelagic 'background' sedimentation is dominant across a large area of the margin, particularly on the lower rise and the adjacent basin plains. These sediments are then overprinted by downslope gravity flows such as~~

turbidity currents. However, the previously described fan organization implies a specific distribution of the sedimentary facies and grain-size distribution within the system (Figure 6).

Cores collected in the upslope area (KC18 and KC17) show mostly hemipelagic sediments with a very low carbonate content. This reflects significant detrital flux associated with proximity to the Ogooue platform and the influence of the Ogooue river plume. Core KC19 collected down the slope just at the transition from canyon to channel-levee complexes show two several meters-thick sandy successions corresponding to top-cut-out Bouma sequences (Ta) interbedded with the upper slope hemipelagites. These sandy turbidites, which are the thickest sand beds recorded in all the cores (Figure 6), indicate the occurrence of high-density turbidity currents flowing down the canyons. The lack of the upper parts of the turbidite is consistent with deposition in the canyons of coarse-grains located at the base of the turbidity currents, whilst the finer upper part of the current is transported downstream and/or spills over the external levees. External levee deposits have been sampled by core KC13, which shows numerous turbidites made up of centimeter-thick, fining upwards parallel or ripple cross-laminated of silt and fine sands (Figure 5). Unfortunately, no core has been collected directly in the intermediate depocenter. However, the rough echofacies III found in this area associated with various erosional features both suggest a high sand/mud ratio.

The mid-system valley acts as a conduit for the sediments coming from the upper part of the system, transporting them further downstream. However, the sediments resulting from the erosion of this valley constitute certainly a part of the sediments deposited in the lobe complex area. According to the available bathymetric data, the volume of sediment removed from the mid-system valley is between 8 and 10 km^3. We assume that these sediments are mainly fine-grained due to the deep location of the valley. Core KC14, collected on an internal terrace of the valley, shows that this valley is also an area of active sedimentation notably due to down-flank sliding. The bottom of the valley comprises slump deposits and coarse-grained sediments deposited by gravity flows coming from the upper part of the system.

Downstream of the mid-system valley, core KC11 show that coarse-grained turbidity currents are deposited in the proximal part of the lobe complex. The abrupt transitions between erosional/bypass and depositional behavior observed notably at the mouth of the mid-system valley is the result of hydraulic jumps affecting flows when they become unconfined between channel sides and spread laterally (Komar, 1971; Garcia and Parker, 1989). Core KC15, located in the lower lobe area, is composed of very thin silty turbidites corresponding to the upper parts of the Bouma sequence interbedded with hemipelagic deposits. The upper lobe acts as a trap for the basal sand-rich parts of gravity flows. Consequently, only the upper part of the flows, which is composed of fine-grained sediments, travels beyond this area. The spatial distribution of facies suggests a filling of successive depocenters with a downslope decrease of the coarse-grained sediment proportion (Figure 6). The same facies distribution can be observed in the northern system. No sandy turbidites are recorded in KC21 located in the Northern lobe, only fine-grained sedimentation, whereas the study of cores taken in the Cape Lopez lobe shows the presence of numerous sandy turbidites (Biscara et al., 2011). The Northern lobe is thus fed by the downslope flow stripped suspended fines transported at the top of turbidity currents flowing through the Cape Lopez Canyon, similarly to intraslope lobes observed in other locations (e.g. Spychala et al., 2015; Jobe et al., 2017). Whatever the current pathways are, the deposited material has a continental origin as suggested by the abundance of quartz, micas and plant debris in the coarse-grained fraction. The important proportion of planktic foraminifers in the coarse-grained fraction of turbidites located in the distal part of the system (core KC10- KC11- KC15) suggests that turbidity currents previously entrained pelagic and hemipelagic deposited upslope where such deposits cover large areas (Viana and Faugères, 1998). The presence of volcanoclastic debris in a sandy layer found at the base of core KC01 suggests that sedimentary input may also come from the volcanic islands of Sao Tomé or Annobon. However, acoustic data indicate that these inputs are limited to the close vicinity of the Sao-Tomé and Annobon islands. In contrast to the model proposed by Wynn et al., (2000) for the Northwest African slope, the volcanic islands and other seamounts present on the Ogooue Fan act mainly as obstacles for the flow pathway but are not important sediment sources for the fan.

56.0 Palaeoceanographic control on the fan activity

The results of Mignard et al. (2017) concerning the study of five cores located along the central part of the Ogooue Fan showed that the fluvial system fed the fan with sediments almost only during times of relative low sea-level. This relative sea-level control on turbidite activity (switch on/off behavior) is classical for mid and low latitude passive margin fans where canyon heads are detached from terrestrial sediment sources (e.g. Mississippi Fan (Bouma et al., 1989), Amazon Fan (Flood and Piper, 1997), Rhone Fan (Lombo Tombo et al., 2015), Indus Fan (Kolla and Coumes, 1987). Conversely, sedimentation during periods of relative high sea-level such as the Holocene, is dominated by hemipelagic to pelagic fall-out with a low part of fine terrigenous particles. Therefore, all cores collected in the central part of the system are capped by 8 to 20 cm of light-brown nannofossil ooze corresponding to Holocene hemipelagites (Figure 5).

However, the northern part of the system appears to have a different behavior. Biscara et al., (2011) showed that the Cape Lopez lobe is currently recording both hemipelagic and turbidity current sedimentation despite the present-day high sea-level. This lobe is fed with sediment from the Cape Lopez Canyon, which incises the shelf to the edge of the Mandji Island (Biscara et al., 2013). The deep incision of the continental shelf up to the coast combined with the longshore sediment transport along the Mandji Island and the narrow shelf in this area (4 km wide) favor the capture of sediment by this canyon during time of high sea-level (Reyre, 1984; Biscara et al., 2013). The northern lobe, which is directly connected to the Cape Lopez lobe by Channel A, appears to be also fed by terrigenous sediments during the Holocene. Core KC21, located at the entrance of the northern lobe, is entirely composed of *facies 3*, even for sediments deposited during MIS1 (Figure 5).

In the Ogooue Fan, the shelf width between the littoral area and the canyon heads is the main control factor on the fan activity. During periods of relative low sea-level, the canyons of the central part of the system receive sediment from the river system that extended across the subaerially exposed continental shelf. During periods of relative high sea-level, river sediments are unable to reach the canyon heads south of the Manji Island and accumulate on the continental shelf close to the Ogooue Delta. However, part of these sediments mixed with sediments coming from the south Gabon margin are drift-transported and contribute to supply the Cape Lopez Canyon and consequently the Cape Lopez and Northern Lobe. Due to their specific location and favorable hydrodynamic conditions on the shelf, sedimentation on the Cape Lopez and the Northern lobes is active during relative sea-level highstands, in contrast to the rest of the Ogooue Fan. Examples of this type of supply have already been described along the California margin where the La Jolla canyon is fed by drift-transported sediments during highstand (Covault et al., 2007, 2011) but also on the southeast Australian coast near the Fraser Island (Boyd et al., 2008), which appears very similar to the Mandji Island.

6 Conclusions

This study provides the first data on the morphology of the recent Ogooue Deep-sea fan and interpretations on sedimentary processes occurring in this environment. The Gabonese margin presents a pelagic/hemipelagic background sedimentation overprinted by downslope gravity flows. The fan is made up of various architectural elements and consists of both constructional and erosional sections. The pattern of sedimentation on the margin is controlled by subtle slope gradient changes (< 0.3°). The long-term interaction between gravity flows and the seafloor topography has induced the construction of successive depocenters and sediment bypass areas. The gravity flows have modified the topography according to a theoretical equilibrium profile, eroding the seafloor where slopes are steeper than the theoretical equilibrium profiles and depositing sediments when slopes are gentler than the theoretical equilibrium profile. Three successive main sediment depocenters have been identified along a longitudinal profile. They are associated with three areas of low slope gradient (0.3°-0.2°). The two updip deposition areas – the intermediate depocenter and the upper lobe area – have recorded coarse-grained sedimentation and are connected by a well-developed large mid-system valley measuring 100 km long and located on a steeper slope segment (0.6°). The distalmost depocenter – the lower lobe area - receive only the fine-grained portion of the sediment load that has bypassed the more proximal deposit areas. Sedimentation on this margin is made more complex by the presence of several volcanic islands and seamounts that constrain the gravity flows. The presence on the slope of the Mount Loiret has caused the formation of an isolated system composed of the Cape Lopez Canyon and lobe, which continues downstream by the Northern Lobe area. ~~The Ogooue Fan is currently in a low activity period since the recent Holocene rise of sea-level. Nowadays, the sedimentation is mostly located on the shelf, in the Ogooue Delta and on the upper slope. The fan was more active during the last glacial lowstand. Nonetheless, the northern part of the system appears to have an asynchronous activity with the rest of the fan as this part is fed by the drift-transported sediments during time of relative high sea-level when the activity in the rest of the system is shut-down.~~

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