

2 **Using Seismic Attributes in seismotectonic research: an application**
3 **to the Norcia's Mw=6.5 earthquake (30th October 2016) in Central**
4 **Italy.**

5 Maurizio Ercoli^{1;4}, Emanuele Forte², Massimiliano Porreca^{1;4}, Ramon Carbonell³, Cristina Pauselli^{1;4},
6 Giorgio Minelli^{1;4}, Massimiliano R. Barchi^{1;4}.

7
8 ¹ Dip. di Fisica e Geologia – Università degli Studi di Perugia (Perugia, Italy).

9 ² Dept. of Mathematics and Geosciences, University of Trieste (Trieste, Italy).

10 ³ Dept. Structure & Dynamics of the Earth, CSIC-Inst. Earth Sciences Jaume Almera (Barcelona, Spain).

11 ⁴ Member of Interuniversity Center for Research on 3D-Seismotectonics (Centro InterUniversitario per l'Analisi
12 SismoTettonica tridimensionale con applicazioni territoriali – CRUST).

13 *Correspondence to:* Maurizio Ercoli (maurizio.ercoli@unipg.it; maurizio.ercoli@gmail.com)

14 **Abstract.** In seismotectonic studies, seismic reflection data are a powerful tool to unravel the complex deep architecture of
15 active faults. Such tectonic structures are usually mapped at the surface through traditional geological surveying whilst
16 seismic reflection data may help to trace their continuation from the near-surface down to hypocentral depth. In this study,
17 we propose the application of the seismic attributes technique, commonly used in seismic reflection exploration by oil
18 industry, to seismotectonic research for the first time. The study area is a geologically complex region of Central Italy,
19 ~~recently~~ struck during the 2016-2017 by a long-lasting seismic sequence, including a Mw 6.5 main-shock. A seismic
20 reflection data-set consisting of three vintage seismic profiles, currently the only ones available at the regional scale across
21 the epicentral zone, ~~constitute~~ represents a singular opportunity to attempt a seismic attribute analysis, by running attributes
22 such as the “Energy” and the “Pseudo Relief”. Our results are critical, because provide information ~~also~~ on the relatively
23 deep structural setting, mapping a prominent, high amplitude regional reflector interpreted as the top of basement, ~~which is~~
24 an important rheological boundary. Complex patterns of high-angle discontinuities crossing the reflectors have also been
25 ~~also~~ identified by seismic attributes. These steep dipping fabrics are interpreted as the expression of fault zones, belonging to
26 the active normal fault systems responsible for the seismicity of the region. Such peculiar seismic signatures of faulting
27 generally well-match with the principal geological and tectonic structures exposed at surface. In addition, we also provide
28 convincing evidence of an important primary tectonic structure currently debated in literature (the Norcia antithetic fault) as
29 well as buried secondary fault splays. This work demonstrates that seismic attribute analysis, even if used on low-quality
30 vintage 2D data, may contribute to improve the subsurface geological interpretation of areas characterized by poor
31 subsurface data availability but high seismic potential.

32 1 Introduction

33 Studying the connections between the earthquakes and the faults to which they are associated is a primary assignment of
34 seismotectonics (Allen et al., 1965; Schwartz and Coppersmith, 1984). Clearly, ~~this is not an easy task; it is in fact generally~~
35 ~~complex~~ to fill the gap between the exposed geology (including the active “geological faults”) ~~mapped by the geologists~~ and
36 the seismological data (e.g. focal mechanisms, earthquake locations, etc...) which are indicators of the geometry and
37 kinematics of the seismic source at hypocentral depth (“seismological faults”, sensu Barchi & Mirabella, 2008), is not an
38 easy task. ~~In case of strong earthquakes ($M_w > 6.5$), impressive~~ important topographic changes and surface ruptures are often
39 reported (e.g. Press and Jackson, 1965; Wyss & Brune, 1967; Jibson et al., 2018; Yi et al., 2018; Civico et al., 2018). While
40 many studies of the surface geology are commonly ~~generally achieved~~ performed, especially after important events,
41 The recovery of deep information on the seismogenic structures at the depth is always more challenging, primarily due to
42 the lack of high-resolution geophysical data and/or wells stratigraphy, generating stratigraphic high degree of uncertainty, and
43 bringing to contrasting geological models and interpretation. -in depth
44 Different geophysical methods (e.g. Gravimetry, Magnetics, Electric and Electromagnetic such as Magnetotellurics and
45 Ground Penetrating Radar) may contribute to define the stratigraphy and structural setting of the upper crust at different
46 scales.- Furthermore, images provided by seismic reflection method are poorly affected by well-known inversion problems
47 typical of the potential methods (Snieder & Trampert, 1999) and are largely the most powerful tool able to produce high-
48 resolution subsurface images. Such type of data, possibly calibrated by deep wells stratigraphy, may provide important
49 constraints to the definition of subsurface geological architecture: these profiles are useful to unveil the deep geometry of
50 active faults from the surface, where they are mapped in the field, down to hypocentral depths. But the
51 Seismic potential fundamental to trace the actual geometry of active faults at surface usually mapped and reconstructed in
52 geological cross sections, from the near surface down to hypocentral depths.
53 Unfortunately, ex-novo acquisition (possibly 3D) of onshore deep-reflection data for research purposes, -is often hampered
54 by high costs, environmental problems and complex logistics (e.g. prohibition of dynamite or vibroseis trucks in Natural
55 Parks or urban areas). -seriously widespread use of for scientific research- Significant exceptions are research projects for
56 deep crustal investigations like BIRPS (Brewer et al., 1983), CoCORP (Cook et al., 1979), ECORS (Roure et al., 1989) and
57 CROP (Barchi et al., 1998; Finetti et al., 2001), IBERSEIS (Simancas et al., 2003), ALCUDIA (Ehsan et al., 2014 and
58 2015). In seismically active regions, Such limitations can be partially overcome by considering old profiles (legacy data)
59 acquired by the ~~exploration~~ industry have been successfully used. ~~When collected in seismically active regions, such data~~
60 ~~may be used~~ to connect the active faults mapped at the surface with the earthquakes seismogenic sources depicted by
61 seismological recordings (Boncio et al., 2000; Bonini et al., 2014; Carvalho et al., 2008; Beidinger et al., 2011; Maesano et
62 al., 2015; Porreca et al., 2018). Legacy seismic lines have in fact some advantages: 1) they are already available from the oil
63 companies 2) they represent a nice source of information in places where new data is difficult to acquire; 3) they can be used
64 to build up and refine geological models. Moreover, such data are often the only available, and are worth to be used in the

65 most appropriate way for constraining the subsurface geological setting and to provide new data on active tectonic structures
66 (see DISS database, Basili et al., 2008). Vintage profiles can therefore significantly contribute to seismo_tectonic researches,
67 even if characterized by intrinsic limitations: i) their location, ~~and~~ orientation and acquisition parameters were not
68 specifically designed with this aim. ~~In addition, ii) they were collected with~~ seismic technologies and acquisition/processing
69 strategies of some decades ago, producing data with ~~both~~ relatively low signal/noise ratio (S/N) and low resolution,
70 especially in comparison to modern data (Manning et al., 2019).
71 ~~In order to~~ improve the data quality and increase the accuracy of the interpretation, two main strategies, ordinarily used by
72 the O&G industry, can be ~~usually applied on legacy data:~~ 1) reprocessing from raw data using modern powerful capabilities
73 processing strategies and ~~developments~~ newly performing algorithms and software; 2) use post-stack analysis processing
74 techniques such as seismic attributes ~~analysis~~.
75 These approaches are ordinarily used by the O&G industry (e.g. in the re-assessment of known reservoirs) and are clearly
76 characterized by variable potential, costs and working time. Some limitations characterize these approaches: the first is
77 particularly demanding in terms of costs and logistic, and not practicable in zones where the use of dynamite or arrays of
78 vibroseis trucks is forbidden or limited (e.g. National Parks or urban areas). The first strategy often requires broad projects
79 encompassing specialized teams, high computation power and generally long processing times (e.g. Pre Stack Depth
80 Migration PSDM strategies); in addition, its efficiency is strictly dependent on the quality of the raw data and survey goals.
81 ~~The second strategy, namely the attribute analysis, exploits a well known and mature technique. It has been used since early~~
82 ~~80s by the O&G exploration industry (Chopra & Marfurt, 2005) for both geometrical and petrophysical characterization of~~
83 ~~reservoirs (Chopra & Marfurt, 2008). An attribute analysis is the easiest, cheapest and fastest strategy to qualitatively~~
84 emphasize the geophysical features and data properties of reflection seismic data sets, producing benefits particularly in
85 complex geological areas.
86 A seismic attribute is a quantity derived from seismic data (pre-stack and/or post-stack) ~~that can be calculated on a single~~
87 ~~trace, on multiple traces, or volumes. This technique is commonly used to extract additional information that may be unclear~~
88 in conventional seismic lines ~~traditional seismic image, therefore leading to a better interpretation of the data. Examples of~~
89 applications on dense 3D seismic volumes produced impressive results, including identification of for instance ancient river
90 channels or sets of faults at variable scales (Chopra & Marfurt, 2005; Chopra & Marfurt, 2007; Chopra & Marfurt, 2008;
91 Marfurt et al., 2011; Hale, 2013; Barnes, 2016, Jacopini et al., 2016; Marfurt, 2018;). ~~Recent developments of approaches~~
92 ~~based on machine learning techniques are currently pushing it further to contribute towards an objective (automatic)~~
93 ~~interpretation of seismic data sets (Wrona et al., 2018; Di & AlRegib, 2019; Nacini & Prindle, 2019). Therefore, among~~
94 between the three two strategies, the attribute analysis is probably the easiest, cheapest and fastest to qualitatively emphasize
95 the geophysical features and data properties of reflection seismic data sets, producing benefits particularly in complex
96 geological areas. ~~Due to different well known~~ limitations and advantages existing between 2D vs 3D seismic data ~~(, these are~~
97 extensively discussed by Torvela et al. (2013) and Hutchinson (2016). ~~For these reasons,~~ 2D post-stack seismic attribute
98 analysis of post stack data may not provide the same quality of information than on 3D, being subjected also to possible

99 pitfalls (Marfurt & Alves, 2015, Ha et al., 2019) and/or ~~may they obviously~~ may not bring so impressive improvements in
100 the seismic images. However, the main point is that in the past inland, it was common most of the sedimentary basins
101 without specific interest for the oil and gas industry, have actually been to sample study areas inland & in the past just by 2D
102 grids of seismic profiles, ~~or at least they have been probed by just a few sparse 2D seismic lines,~~ being the full 3D seismic
103 surveys rare. available only in very few cases. Hence, it is relevant to extract as much information as possible from such
104 data, 2D profiles, which often are the only available data, these 2D surveys in areas not covered by 3D seismic surveys.
105 ~~Whilst in the hydrocarbon industry this process is useful even if mainly driven by a constant necessity to reduce the costs~~
106 (Ha et al., 2019), in seismotectonic researches it is affected by even worse limitations previously aforementioned. Therefore,
107 also slight improvements obtained on vintage 2D data may bring to new and unprecedented subsurface information in complex
108 and active tectonic environments. We think that we might successfully export this approach in a seismotectonic study
109 applying this type of analysis on an active seismic zone, covered only by a very limited number of 2D seismic lines. Based
110 on such considerations, In this work, the selected study area is located in the central Apennines (Central Italy), a region
111 between the southeastern part of the Umbria-Marche Apennines and the Laga Domain, in the outer Northern Apennines
112 (central Italy) (e.g. Barchi et al., 2001). This area presents ideal characteristics to test the application of seismic attributes as
113 proposed a new approach in seismotectonics. ~~In fact,~~ in the past, several seismic profiles were acquired at this location in this
114 region for hydrocarbon exploration, ~~providing and were later used to constrain good constraints for~~ subsurface geological
115 interpretation (Bally et al., 1986; Barchi et al, 1991; Barchi et al., 1998; Ciaccio et al., 2005; Pauselli et al., 2006; Mirabella
116 et al. 2008; Barchi et al., 2009; Bigi et al., 2011). After the 2016-2017 seismic sequence, Porreca et al. (2018) provided an
117 updated regional geological model based on the interpretation of vintage seismic lines, but remarked important differences in
118 the seismic data quality across the region, hampering a straightforward seismic interpretation. After the last 2016-2017
119 seismic sequence, Porreca et al. (2018) have provided a new regional geological model based on the interpretation of vintage
120 2D seismic lines. In such a study, the authors remark important differences in the seismic data quality across the region that
121 hampered the interpretation. Therefore, the present work exploits the use of seismic attributes focuses on three low-quality
122 seismic profiles located close to the Mw 6.5 main-shock of the 2016-2017 seismic sequence, exploit the use of seismic
123 attributes to squeeze additional information. The main goal of this study is to squeeze additional information from the 2D
124 data obtaining as much constraints as possible on the geological structures responsible for the seismicity in the area by
125 defining:

- 126 - geological/structural setting at depth (e.g. depth of the basement and its involvement)
- 127 - trace of potentially seismogenic faults (connection between the active faults mapped at the surface and earthquake's foci).

128 ~~Therefore, Any improvements achievable on the data quality and visualization, for example an increase of the resolution~~
129 ~~and/or an enhancement of the lateral discontinuity of seismic reflectors, would represent a very valuable contribution~~
130 ~~considering the limited data availability in this area. -The manuscript-~~ We think that this innovative approach to
131 seismotectonic research can be extended to other on-shore seismically active areas in the world, especially if covered only by

132 sparse vintage low-quality seismic surveys. In such cases, we think the seismotectonic research may benefit of the potential
133 and improvements generated by the seismic attributes.

134 **2 Geological framework and seismotectonics of the study area**

135 The study area is located in the southeastern part of the Northern Apennines fold and thrust belt, including the Umbria-
136 Marche Domain and the Laga Domain, separated by an important regional tectonic structure, known as the M. Sibillini thrust
137 (MSt) (Fig. 1).

138
139 The Umbria-Marche domain involves the rocks of the sedimentary cover, represented by three main units (top to bottom),
140 characterized by different interval velocities (Bally et al., 1986; Barchi et al., 1998; Porreca et al., 2018):

141 1) on top, the Laga sequence (Late Messinian – Lower Pliocene, up to 3000 m thick, average seismic velocity; $v_{av} = 4000$
142 m/s), consisting of siliciclastic turbidites made by alternating layers of sandstones, marls and evaporites, deposited in
143 marine depositional environment (Milli et al., 2007; Bigi et al., 2011); it is and outcropping in the eastern sector of the study
144 area (i.e. Laga Domain); 2) the carbonate formations (Jurassic-Oligocene, about 2000 m thick, $v_{av} = 5800$ m/s), formed by
145 pelagic limestones (Mirabella et al., 2008) with subordinated marly levels overlying an early Jurassic carbonate platform
146 (Calcare Massiccio Fm.), mainly outcropping in the Umbria-Marche Domain; 3) the Late Triassic evaporites (1500–2500 m
147 thick, $v_{av} = 6400$ m/s), consisting in alternated layers of anhydrites and dolomites (Anidriti di Burano Fm. and and
148 Raethavacula Contorta beds; Martinis & Pieri, 1964), never outcropping and intercepted, only, by deep wells (Porreca et al.,
149 2018 and references therein). For further details on the stratigraphic characteristics of the area, we remind to the works of
150 Centamore et al. (1992) and Pierantoni et al. (2013).

151 These units rest on a basement with variable lithology (Permian-Late Triassic, $v_{av} = 5100$ m/s) that never crops out in the
152 study area (Vai, 2001), but only intercepted by deep wells (Bally et al., 1986; Minelli & Menichetti, 1990; Anelli et al.,
153 1994; Patacca & Scandone, 2001).

154
155 This sedimentary sequence is involved in the Late Miocene fold and thrust belt including a set of N-S trending anticlines,
156 formed at the hangingwall of the W-dipping arc-shaped major thrusts. The most important compressional structure is the M.
157 Sibillini thrust (MSt, Koopman, 1983; Lavecchia, 1985), where the Umbria-Marche Domain is overthrust on the Laga
158 Domain.

159
160 This is a geologically complex region, where in the past the analysis of 2D seismic profiles have produced contrasting
161 interpretations of the upper crust structural setting, i.e. thin- vs. thick-skinned tectonics, fault reactivation/inversion and
162 basement depth (Bally et al., 1986; Barchi, 1991; Barchi et al., 2001; Bigi et al., 2011; Calamita et al., 2012). A review of the
163 geological history of this area has recently been provided by Porreca et al. (2018). These authors propose a tectonic style

164 characterized by coexistence of thick- and thin-skinned tectonics with multiple detachments localized at different structural
165 levels.

167 These compressional structures have been later disrupted by the extensional faults since the Late Pliocene (Fig.1) (Blumetti
168 et al., 1993; Boncio et al., 1998; Brozzetti & Lavecchia, 1994; Calamita & Pizzi, 1994; Pierantoni et al., 2013).

170 The Late Pliocene-Quaternary extensional tectonic phase, characterized by NNW-SSE striking normal faults, consistent with
171 the present-day active strain field as deduced by geodetic data (e.g. Anderlini et al., 2016). The latter have high dip angles
172 (50-70°) and can be synthetic or antithetic structures (WSW or ENE dipping, respectively) dipping normal faults. These
173 faults were also responsible of the tectono-sedimentary evolution of intra-mountain continental basins (Calamita et al., 1994;
174 Cavinato and De Celles, 1999). The most evident Quaternary basins of this part of the Apennines are the Castelluccio di
175 Norcia and Norcia basins (Fig.1), located at 1270 and 700 m a.s.l., here named CNb and Nb respectively. A phase of
176 lacustrine and fluvial sedimentation infilled the two basins with hundred meters of deposits, characterized by fine clayey to
177 coarse grained material (Blumetti et al., 1993; Coltorti and Farabollini, 1995).

179 The area is affected by frequent moderate magnitude earthquakes ($5 < M_w < 7$) and has a high seismogenic potential
180 revealed by both historical and instrumental data (e.g. Barchi et al., 2000; Boncio and Lavecchia, 2000; Basili et al., 2008;
181 Rovida et al., 2016; DISS Working Group, 2018).

183 The major seismogenic structures recognized in the area are the Norcia fault (Nf) and the M. Vettore fault (Vf). The Norcia
184 fault (Nf, Fig.1) is associated to several historical events (Galli et al., 2015; Pauselli et al., 2010; Rovida et al., 2016),
185 probably including the 1979 earthquake (Nottoria-Preci fault, Deschamps et al., 1984; Brozzetti & Lavecchia, 1994; Rovida
186 et al., 2016) and the largest event in 1703 ($M_e = 6.8$, Rovida et al., 2016). The Vettore fault (Vf) in part of the easternmost
187 alignment whose historical and pre-historical activity was recognized by paleoseismological and shallow geophysical
188 surveys (Galadini & Galli, 2003; Galli et al., 2008; Ercoli et al., 2013; Ercoli et al., 2014; Galadini et al., 2018; Galli et al.,
189 2018; Cinti et al., 2019; Galli et al., 2019).. This system was reactivated during the 2016-2017 sequence characterized by
190 multi-fault ruptures occurred within few months (nine $M > 5$ earthquakes at hypocentral depth < 12 km between August 2016
191 – January 2017) having characteristics comparable to previous seismic sequences in Central Italy (e.g. L'Aquila 2009 and
192 Colfiorito 1997-1998, Valoroso et al., 2013 and Chiaraluce et al., 2005).

193 The strongest mainshock of ($M_w 6.5$) occurred on 30th October 2016 (Chiaraluce et al., 2017; Chiarabba et al., 2018;
194 Gruppo di Lavoro Sequenza Centro Italia, 2019; Improta et al., 2019; ISIDe working group, 2019), generating up to 2 m
195 (vertical offset) co-seismic ruptures (Civico et al., 2018; Gori et al., 2018; Villani et al., 2018a; Brozzetti et al., 2019),
196 mainly localized along the Mt. Vettore fault (blue thin lines in Fig. 1).

198 Despite of the large amount of surface data collected (Livio et al., 2016; Pucci et al., 2017; Wilkinson et al., 2017; De Guidi
199 et al., 2017; Brozzetti et al., 2019), the deep extension of the Norcia and Castelluccio antithetic and synthetic faults
200 (particularly Nf and Vf), and the overall complex structure of the area are still debated (Lavecchia et al., 2016; Porreca et al.,
201 2018; Bonini et al., 2019, Cheloni et al., 2018, Improta et al. 2019).

202 **3 Data**

203 We have performed ~~the~~ seismic attributes analysis on three W-E trending 2D seismic reflection data crossing the epicentral
204 area between the Umbria and Marche regions (Central Italy, Fig.1). ~~Such 2D data~~ These seismic profiles are part of a much
205 larger, unpublished dataset including 97 seismic profiles and few boreholes, drilled for hydrocarbon exploration by ENI in
206 the period 1970-1998. The data quality is extremely variable (medium/poor) with limited fold (generally < 60 traces /
207 Common Mid-Point –CMP) mainly due to environmental and logistical factors. Among those, we can list the different
208 acquisition technologies, a limited site access, the complex tectonic setting and especially the different (contrasting)
209 outcropping lithologies (e.g. Mazzotti et al., 2000, Mirabella et al., 2008). The eastern area, showing higher data quality,
210 consists of siliciclastic units of the Laga foredeep sequence, located at the footwall of the MSt. On the contrary, the lowest
211 S/N recordings coincide with outcropping carbonates formations and Quaternary deposits, sediments
212 The analysed lines include seismic reflection profiles include: NOR01 (stack, 14 km long; ~~–) and~~ NOR02 (time-migrated, 20
213 km long, partially parallel to NOR01 on the western sector) located west and east to the Nb, respectively; –and CAS01
214 (stack, 16 km long), located more to the south crossing the Cascia village (Fig. 1). NOR01 and CAS01 were acquired using a
215 Vibroseis source, whilst explosives were used for NOR02; all the lines are displayed in Two-Way-Travel-Time (TWTT)
216 limited to 4.5 s. The amplitude/frequency spectra (computed on the entire time window) of the processed lines show a
217 bandwidth in a range 10-50 MHz, with the NOR02 spectrum displaying a slighter high frequency content (Tab.1). Assuming
218 the average peak frequency of 20 Hz, a vertical resolution of ca. 80 m can be estimated (average carbonate velocity = 6
219 km/s; parameters in Table 1s, ~~supporting information~~). Some processing artefacts are visible in NOR01 as a straight
220 horizontal signal at ca. 1 s (yellow dashed line and label A in Fig. 2a), and ~~two others sub horizontal between 1–2 s another~~
221 in CAS01 (Fig. ~~3a, supporting information~~). ~~However, some seismic events and lineaments, related to geological structures~~
222 ~~of interest, are slightly visible~~ The data ~~display~~ may benefit of potential improvements by selecting potentially improvable
223 with a proper type of choice of seismic attributes to be tested with different calculation parameters type and parameters.
224 Therefore, we loaded the lines into the software OpendTect (OdT, <https://www.dgbes.com/index.php/software#free>)
225 software, setting up a common seismic datum equal to 500 m. Unfortunately, deep borehole stratigraphy is not available for
226 the study area (all details about surrounding deep wells have been already summarized in Porreca et al., 2018). The OdT
227 seismic project was enriched also by some ancillary data, extracted by a complementary GIS project (QGis,
228 <https://www.qgis.org/it/site/>) project. As visible in Fig. 1, we have included a detailed summary of the main normal faults
229 and surface ruptures of the area (Civico et al., 2018; Villani et al., 2018; Brozzetti et al., 2019), obtained after carefully
230 checking the most important regional geological maps and fault patterns (Koopman, 1983; Centamore et al., 1993;

231 Pierantoni et al., 2013; Carta Geologica Regionale 1:10'000 – Regione Marche, 2014; Carta Geologica Regionale 1:10'000 –
232 Regione Umbria, 2016; [Ithaca database, http://www.isprambiente.gov.it/it/progetti/suolo-e-territorio-1/ithaca-catalogo-delle-](http://www.isprambiente.gov.it/it/progetti/suolo-e-territorio-1/ithaca-catalogo-delle-faglie-capaci)
233 [faglie-capaci;](http://www.isprambiente.gov.it/it/progetti/suolo-e-territorio-1/ithaca-catalogo-delle-faglie-capaci)), as well as the most recent works published in literature (e.g. Brozzetti et al., 2019; Porreca et al., 2020). [The](#)
234 [topography was also included using a regional 10 meters resolution DTM \(Tarquini et al., 2007; Tarquini et al., 2012\). The](#)
235 [other important external data-set consists of seismological data, i.e. inferred location and approximated fault geometry as](#)
236 [suggested by the focal mechanisms of the mainshocks and by the distribution of the aftershocks](#) (Iside database,
237 <http://iside.rm.ingv.it/iside/> and Chiaraluce et al., 2017). The integration of such information in a pseudo-3D environment
238 offered us a multidisciplinary platform to clearly display the seismic lines and to link surface data and the deep geologic
239 structures at hypocentral depth.

240 **4 Methods**

241 The seismic reflection data interpretation is generally accomplished through the definition of specific signal characteristics
242 (seismic signature), supported by the geological knowledge of the study area. A standard seismic interpretation is affected by
243 a certain degree of uncertainty/subjectivity (particularly in case of poor data quality), because [is](#) generally based on a
244 qualitative analysis of ~~reflection~~-amplitude, geometry and lateral continuity [of reflections](#). Over the last years, the
245 introduction of seismic attributes and related automated/semi-automated procedures had an important role in reducing the
246 subjectivity of seismic interpretation. [A seismic attribute is a descriptive and quantifiable parameter that can be calculated on](#)
247 [a single trace, on multiple traces, or 3D volumes and can be displayed at the same scale as the original data. Seismic data can](#)
248 [be therefore considered a composition of constituent attributes \(Barnes, 1999, Taner et al., 1979, Forte et al., 2012\). Their](#)
249 [benefits have been](#) ~~at first appreciated~~ in 2D/3D seismic reflection data (Barnes 1996; Taner et al., 1979; Barnes, 1999; Chen
250 and Sidney, 1997; Taner, 2001; Chopra and Marfurt, 2007; Chopra and Marfurt, 2008; [Iacopini and Butler, 2011; Iacopini et](#)
251 [al., 2012; McArdle et al., 2014; Botter et al., 2014; Hale, 2013 for a review; Marfurt and Alves, 2015; Forte et al., 2016](#)) and,
252 more recently, also in other reflection techniques like the Ground Penetrating Radar (~~GPR~~) (e.g. McClymont et al., 2008;
253 Forte et al., 2012; Ercoli et al., 2015, ~~De~~-Lima et al., 2018). In this work, we have tested several post-stack attributes on three
254 2D vintage seismic lines ([original seismic data in the supplementary material in Fig. 1s](#)), [and starting our analysis by using](#)
255 [first well-known and widely used attributes like the instantaneous amplitude, phase, frequency, and their combinations](#), also
256 using composite multi-attribute ([i.e. simultaneous overlay and display of different attributes e.g. primarily phase, frequency,](#)
257 [envelope, Chopra and Marfurt, 2005; Chopra and Marfurt, 2011](#)). [Later on, we have also tested attributes \(e.g. coherency and](#)
258 [similarity\), generally more efficient on 3D volumes, but without obtaining positive outcomes, due to limited vertical and](#)
259 [spatial resolution of the data](#). Among [tested attributes](#), we selected ~~the~~-three attributes that resulted in the best images
260 ([provided in Figs. s2, s3 and s4 of the supplementary material without any line drawing or labels](#)), making possible to detect
261 peculiar seismic signatures of regional seismogenic layers and fault zones. ~~Details about~~ [The calculated attributes, computed](#)
262 [using OdT software, are: hereafter provided.](#)

264 “**Energy**” (**EN**): one of the RMS amplitude-based attributes, it is defined as the ratio between the squared sum of
265 the sample values in a specified time-gate and the number of samples in the gate (Taner, 1979, [Gersztenkorn,](#)
266 [Marfurt, 1999](#), Chopra & Marfurt, 2005, Chopra & Marfurt, 2007, [for a review of formulas see Appendix A in](#)
267 [Forte et al., 2012](#)). The Energy measures the reflectivity in a specified time-gate, so the higher the Energy, the
268 higher is the reflection amplitude. In comparison to the original seismic amplitude, it is independent of the polarity
269 of the seismic data being always positive, and in turn preventing the zero-crossing problems of the seismic
270 amplitude (Forte et al., 2012, Ercoli et al., 2015, Lima et al., 2018, Zhao et al., 2018). This attribute is useful to
271 emphasize the most reflective zones (e.g. characterization of acoustic properties of rocks). It may also enhance
272 sharp lateral variations in seismic ~~events~~ [reflectors](#), highlighting discontinuities like fractures and faults. In this
273 work, we set a 20 ms time window (i.e. about the mean wavelet length), obtaining considerable improvements in
274 the visualization of higher acoustic impedance contrasts.

275 “**Energy gradient**” (**EG**): it is the first derivative of the energy with respect to time (or depth). The algorithm
276 calculates the derivative in moving windows and returns the variation of the calculated energy as a function of time
277 or depth (Chopra & Marfurt, 2007; Forte et al., 2012). It is a simple and robust attribute, also useful for a detailed
278 semi-automatic mapping of horizons with a relative low level of subjectivity. The attribute acts as an edge detection
279 tool, effective in the mapping of the reflection patterns as well as the continuity of both steep discontinuities like
280 faults and fractures, and channels, particularly in slices of 3D data (Chopra & Marfurt, 2007). In this work, we have
281 used the same time window of the Energy, obtaining considerable improvements in the visualization not only of the
282 strong acoustic impedance reflectors but particularly in the faults imaged in the shallowest part of the seismic
283 sections.

284 **Pseudo-relief (PR)**: it is obtained in two steps: the energy attribute is first computed in a short time window, then
285 followed by the Hilbert transform (phase rotation of -90 degrees). The Pseudo-relief is considered very useful in 2D
286 seismic interpretation to generate “outcrop-like” images allowing an easier detection of both faults and horizons
287 (Bulhões, 1999; Barnes et al., 2011; Vernengo et al. 2017, Lima et al., 2018). In this work, considerable display
288 improvements have been obtained using the Pseudo-relief computed in a window of 20 ms. In comparison to the
289 standard amplitude, it better highlights the reflection patterns and thus the continuity/discontinuity of reflectors,
290 enhancing steep discontinuities and fault zones.

291 5 Results

292 ~~The Figs. 2, 3 and 4 show the comparison between the original seismic lines in~~ [The comparison between the original seismic lines in](#)
293 [amplitude](#) and the images obtained after the attribute analysis, ~~revealing significant~~ [allows to detect considerable](#)
294 ~~improvements~~ [improvements](#) in the visualization and interpretability of the geophysical features. In ~~the~~ [the](#) profiles NOR01, CAS01 and
295 NOR02 ~~(Figs 2, 3 and 4, respectively)~~ [we focus our analysis on three main types](#) [three types](#) of geophysical features

296 highlighted by the attributes: sub-horizontal deep reflectors, low-angle and high-angle discontinuities The main faults
297 mapped at the surface (Fig.1) have been also plotted on top of each seismic line.

298 In the original seismic line NOR01 (Fig. 2a), the overall low S/N ratio hampers the detection of clear and continuous
299 reflectors. At ca. 1 s a horizontal processing artefact is visible (label A, yellow dots), possibly related to a windowed filter.
300 The most prominent sub-horizontal reflections (labelled H) are located in the central portion between 2-3 s (TWT) (strong
301 reflectors in the black box i). Shallower and less continuous reflectors are also visible in the eastern side of the profile,
302 beneath the Nb (black box ii). The EN attribute (Fig. 2b) enhances the reflectivity contrast, better focusing the high-
303 amplitude, gently W-dipping reflector H (blue arrows) and also outlining its lateral extension. In this image most of the
304 reflected energy is concentrated on its top at ca. 2.5 s, so that it is readily apparent that H separates two seismic facies, with
305 higher (top) and lower (bottom) amplitude response, respectively.

306 The EG and PR attributes of NOR01 (Figs. 2c, 2d) better show the geometry of horizon H, characterized by a continuous, ca.
307 8 km long, package of reflectors (ca. 200 ms thick) having common characteristics in terms of reflection strength and period.
308 In the eastern part of the profile, below the Nb, the EG and PR attributes also enhance two major opposite-dipping high-
309 angle geophysical features (red arrows in fig. 2c and 2d), crossing and disrupting the shallower reflectors. The W-dipping
310 lineament propagates down to ca. 2.5 s, intercepting the eastern termination of the reflector H. The two discontinuities border
311 a relatively transparent, shallow seismic facies, corresponding to the area where the Nb crops out. In the same area, the
312 reflectors are pervasively disrupted by other, minor discontinuities.

313 Analysing in detail the line NOR01 (Fig. 2a, line location in the excerpt on the top), the most apparent low angle
314 geophysical features are located in the eastern portion of the line between 2-3 s of the time window. The EN attribute in Fig.
315 2b clearly enhances a high amplitude, gently W-dipping event at about 2.5 s (blue arrows). The EG and R attributes of
316 NOR01 show clearly that this horizon (Figs. 2c, 2d, hereafter H) is characterized by a continuous package of reflectors (ca.
317 200 ms in TWT, ca. 8 km long), with common characteristics in terms of reflection strength and period.

318 A feature showing such a peculiar signature is visible also in CAS01, approximately at the same time interval (Fig. 3a, line
319 location reported on the top insert). But in comparison to NOR01, It appears more discontinuous mainly visible on the
320 westernmost side and beneath the southern termination of Nb (ca. between 11-15 km). all along the seismic profile, and in
321 addition it is partially interfering For those reasons, H is not particularly clear in the standard amplitude line CAS01 (Fig.
322 3a), even if it is mainly visible on the westernmost side and beneath the southern termination of Nb (ca. between 11-15 km).
323 Despite a generalized high frequency noise content, H is better enhanced in fig. 3b by EN attribute (blue arrows), and in
324 particular by the EG and PR attributes (Figs. 3c and 3d), that considerably help to better detect and mark its extension and
325 geometry. A high angle East dipping discontinuity can be noticed in the eastern sector of CAS01 (red arrows in Fig. 3c and
326 3d).

327 Regarding the most visible steep geophysical features detectable in these two seismic profiles, in NOR01 a high angle steep
328 E-dipping lineament in NOR01 is defined by a clear high angle discontinuity of the seismic signal, particularly enhanced in
329 the eastern sector (distance ca. 10 km) below the Nb (red arrows in fig. 2c and 2d). A high angle East dipping discontinuity

330 ~~can be noticed in the eastern sector of CAS01 (red arrows in Fig. 3c and 3d).~~ Another main high-angle W-dipping lineament
331 is enhanced in Figs. 2e-2d of NOR01 (red arrows at the end of the line), that clearly divides two patterns of reflectors
332 showing different dip; this discontinuity propagates down to ca. 2.5 s and intercepts the aforementioned strong reflector H.
333 ~~Between those two main alignments bounding Nb, other minor discontinuities can be also noticed crossing and slightly~~
334 ~~disrupting the shallower reflectors: those high angle features are efficiently displayed by the EG and PR attributes (Fig. 2e,~~
335 ~~2d), whilst in the original line in Fig. 2a cannot be really appreciated.~~
336 The original seismic reflection line CAS01 (Fig. 3a) displays a generalized high-frequency noise content.
337 As in NOR01, a shallow processing artefact (A, yellow dots) is visible and possibly related to a filter. Fragmented packages
338 of high-amplitude reflectors (H) are visible at the same time interval observed in NOR01 (ca. 2.5 s), in both the western
339 (black box i, in Fig. 3a) and, more discontinuous, in eastern part of the line (black box ii, in Fig. 3a). The EN attribute (Fig.
340 3b) emphasizes the presence of the H reflector better focusing the reflectivity (blue arrows). Both the EG and PR attributes
341 (Figs. 3c and 3d) further help to delineate the reflector H. The steeper discontinuities have been analysed mainly in the
342 western part of the profile, closer to the 2016-2017 seismically active area. A major high-angle, east-dipping discontinuity
343 has been traced at about 13 km (alignment of red arrows in Fig. 3c and 3d).
344 The original seismic line NOR02 (Fig. 4a), displays geophysical features similar to the ones detected in NOR01 and CAS01.
345 This seismic profile shows a generalized poor continuity of the reflectors, with the exception of the eastern side, where a set
346 of west-dipping, coherent reflections can be recognized: the higher S/N ratio of this part of the section is due to the
347 outcropping turbidites of Laga sequence, which are known to favour the energy penetration, respect to the carbonates (e.g.
348 Bally et al., 1986; Barchi et al., 1998). The prominent reflection H, gently east-dipping and relatively continuous for more
349 than 8 km (black box in Fig 4a), is located in the centre of the line, at greater depth (3.2–3.5 s TWT), respect to the
350 previously described NOR01 and CAS01 profiles. As in the previous cases, the EN attribute (Fig. 4b) effectively focuses the
351 horizons reflectivity, emphasising the strong amplitude of the reflector H (blue arrows).
352 The EG and PR attributes (Figs. 4c and 4d) improve the overall visualization of the reflection patterns, aiding the detection
353 of the low-angle and high-angle discontinuities.
354 A major westward low-angle discontinuity T (green dots in Figs.4c and 4d) crosses the entire profile, descending from ca. 2
355 s (East) to ca. 4 s (West), where it interrupts marks the continuity of the reflector H. Several high-angle discontinuities have
356 been traced to ca.along the section, marked by alignments of red arrows in Figs. 4c and 4d. The most important alignments
357 have been recognised beneath the two major Quaternary basins (i.e. Nb and CNb) crossed by the profile: in both cases, major
358 W-dipping alignments can be traced from the near surface, where they correspond to the eastern border of the above
359 mentioned basins, down to a depth of ca. 4 s TWT. Other discontinuities, W and E dipping, have been traced in the hanging-
360 wall of these two major alignments. In the seismic volume bounded by these features, many secondary (minor)
361 discontinuities pervasively cross-cut the set of reflectors, producing a densely fragmented pattern. Unfortunately, limited
362 resolution and data quality in the deeper part of the section hampers a univocal interpretation of the cross-cutting

relationships between the low-angle discontinuity T and the major W-dipping high-angle discontinuity: two alternative interpretations are here possible, that will be discussed in detail in the next paragraph 6.

The global improvement in the dataset interpretability can be better appreciated in a 3D visualization of the seismic attributes, also using multi-attribute displays (Fig. 5). Such images better clarify the deep geometry of the main reflectors and the location of the geophysical discontinuities, later interpreted on the light of known and debated tectonic structures on the study area. In Fig. 5a we report a 3D perspective of the seismic line NOR02, after combining in transparency the EN attribute with the PR attribute (EN+PR). The reflectors characteristics and a pattern of discontinuities are clearly visible at different levels of detail, and a first correlation with the surface faults at surface is proposed (red segments on the top). The two boxes (blue and black colours in Fig.5a, respectively) point out the two most representative seismic facies described above. The Fig. 5b and 5c display a comparison of the signature of reflector H in the standard amplitude line (SA) (Fig. 5b) and in a version including PR attribute in transparency with SA itself. The figure 4a display the original seismic line NOR02 characterized by geophysical features. The EN attribute in Fig. 4b again results efficient in enhancing sub-horizontal (blue arrows) and also gently dipping deep events (green dots). On the western sector, the attributes in Figs. 4b and 4e show a pattern of relatively continuous and gently W dipping events between 0-2.5 s (0-5 km along the line). The most evident high amplitude and continuous reflector characterizes the central part of NOR02 at ca. 3.2-3.5 s (blue arrows in Figs. 4b, 4e, 4d), gently East dipping and relatively continuous for more than 8 km. ~~This latter is intercepted by an important and well visible low angle W dipping discontinuity (T, green dots in Figs. 4b, 4e and 4d). It crosses the entire profile, rising from about 4 s (West) to ca. 2 s (East), where it intercepts one of the high amplitude events on the eastern end of the seismic line (18-20 km). Here again the attribute analysis results extremely efficient to clearly detect such geophysical features otherwise poorly visible on the original line NOR02 in Fig. 4a.~~

The most important result provided by the EG and PR attributes is an improved much clear visualization of the reflection patterns of NOR02, aiding an easier detection of high angle primary and secondary (minor) discontinuities, at different scales/levels of detail. ~~This latter is intercepted by an important and well visible low angle W dipping discontinuity (T, green dots in Figs. 4b, 4e and 4d). It crosses the entire profile, rising from about 4 s (West) to ca. 2 s (East), where it intercepts one of the high amplitude events on the eastern end of the seismic line (18-20 km). Here again the attribute analysis results extremely efficient to clearly detect such geophysical features otherwise poorly visible on the original line NOR02 in Fig. 4a.~~The deep continuation of such a main W dipping alignment also truncate and disrupt both the gently dipping discontinuity T and the deep reflector H: at approximately 3.2 s, it appears interrupted laterally on its western side (Figs. 4e and 4d). In fact, a main high angle E dipping discontinuity (red arrows) delimits the NOR02 western sector (ca. 1 km of distance along the line at surface); another steep W dipping alignment (red arrows) that clearly cuts and slightly disrupt the set of reflectors below the Nb (0-2.5 s, ca. 4-5 km). In addition, smaller discontinuities pervasively cross-cut the set of reflectors between 1-4 km bounded by such two main features, producing a densely fragmented reflectors in the middle portion. Another steep E dipping feature is visible at higher depth (red arrows at 1-3 s, ca. 7-9 km) beneath the topographic relief separating Nb by CNb: ~~it end up on the deep surface horizon T and in addition it doesn't reach the shallower portion~~

397 of the seismic line. This discontinuity is subparallel to a similar structure displayed in a more central portion of NOR02
398 (western side of Nb highlighted by red arrows at 10–12 km). The Figs. 4c and 4d show here in this sector sets of reflectors
399 sharply interrupted, fragmented and displaced in a narrow zone. The same seismic pattern is present in the easternwestern
400 side of CNb, but it is due to some west-dipping discontinuities located between 14–16 km. These features highlight a
401 slightly asymmetric “V shape” fabric characterized by very short and fragmented reflectors bounded by those two steep
402 features of opposite dip. The deep continuation of such a main W-dipping alignment also truncate and disrupt both the
403 gently dipping discontinuity T and the deep reflector H: at approximately 3.2 s, it appears interrupted laterally on its western
404 side (Figs. 4c and 4d).

405 The results of this work produced have globally improved the interpretability of the original dataset. In particular, the data
406 integration in a 3D environment and the use of multi-attribute displays clarified the deep geometries of the main reflectors
407 and of the geophysical discontinuities, later interpreted on the light of the known and debated tectonic structures on the study
408 area. This is particularly clear in Fig. 5a, in which we report the seismic line NOR02 after the combined plot of the PR
409 attribute (“similarity” palette) the EG attribute (“energy” palette), overlapped using ODT software (depth conversion with
410 $V_{pav} = 6000$ m/s, vertical scale 2x). The reflectors characteristics and the discontinuities are clearly visible at different
411 levels of detail, and the two boxes (blue and black colours, respectively) highlight on the two most representative seismic
412 facies described before. Fig. 5b and 5c display a comparison of the H in the original line and a of the EN-PR attribute. Again,
413 in the two other inserts in Figs. 5d and 5e, the same data comparison proposed show of the data included in the black box is
414 proposed. Fig. 5d shows the scarce detectability of the dense pattern of steep discontinuities in the original seismic profile
415 (SA). The Fig. 5e displays the enhancement obtained plotting the PR attribute (“similarity palette”) in transparency on the
416 seismic line in amplitude (SA), enhancing well the dense fragmentation of these reflectors.

417 An analogous visualization 3D multi-display of attributes EN and PR is proposed in Fig. 6a for the seismic line NOR01. The
418 comparison between the multi display of attributes PR and ENG (blue box in Fig. 6a), the original line (blue box in Fig. 6b)
419 and the EN+PR plot (Fig. 6c) shows the improved and peculiar signature of the strong reflector H. The black box again
420 reports the original plot vs. line NOR01 and the version PR+SA, which clearly boosting the visualization of the high-angle
421 discontinuities, illustrating a detail on the one beneath aNf.

422 Such results therefore ensure an easier and more accurate interpretation of the subsurface geological structures; some of
423 them those are are apparently connected, whilst others not at all, with the surface geology and related to the hypocentre
424 location of the main seismic events, that will be discussed more in detail within the following chapter.

425 **6 Data Interpretation: New constraints, new elements and insights on the deep geological structure reconstruction of** 426 **the study area.**

427 The comparison between the original seismic data and the images obtained by the attribute analysis ensures an easier and
428 accurate interpretation of the geophysical features, allowing to extend the surface geological data in depth. The geological
429 interpretation of these features requires a thoughtful comparison and calibration with the other data available for the area.

430 e.g. geological, and structural maps, co-seismic ruptures, high-resolution topography and mainshocks hypocentres. The
431 seismic attributes provide a multiple view of the original data through the enhancement of different physical quantities.
432 Therefore, peculiar geophysical signatures have been detected delineating interpretative criteria (e.g. high amplitude
433 reflectors, phase discontinuities, fragmented reflectors patterns etc...). Such geophysical features, after a first order
434 interpretation, fit well with the main outcropping geologic structures. Due to the lack of 3D seismic volumes and of a regular
435 grid of 2D seismic profiles in the area, the geological meaning of the results provided by the attributes analysis have been
436 constrained by integrating all the other available literature data. We have therefore integrated geological, and structural maps
437 (Koopman, 1983; Centamore et al., 1993; Pierantoni et al., 2013), high resolution topographic data (Tarquini et al., 2007 and
438 2012), mainshocks hypocentral data (Chiaraluce et al., 2017) and co-seismic surface ruptures data (Civico et al., 2018;
439 Villani et al., 2018; Brozzetti et al., 2019). Using the same interpretation criteria, other surface-uncorrelated discontinuities,
440 poorly visible in the original amplitude lines, are rising at a more detailed scale after the attribute's analysis. In addition,
441 deep reflectors showing a common signature have been also recognized, revealing a regional character. The geological
442 meaning and the relation of such geophysical features with the surface geology and with the hypocentre location of the main
443 earthquakes are hereafter discussed.

444 ~~In Fig. 7 reports, a global pseudo-3D overview of the study region summarizing all the data analysed across the area,~~
445 ~~together with all the faults mapped at surface (Fig. 7a) and surrounding the location of Mw 6.5 mainshock (30th October~~
446 ~~2016), plotted together with other three strong seismic events in the Northern sector.~~ The two seismic images in Figs. 7b
447 and 7c have been obtained by using again a multi-attributes visualization, in this case overlapping the PR and EN attributes
448 in transparency with the original seismic lines NOR01 and NOR02, following the same procedure used for the images in
449 Figs. 5 and 6. The black boxes centred on the Norcia and Castelluccio di Norcia basins have been magnified above and
450 display the limits of the bounding faults (black dashed lines) and the main important reflectors detected in depth. In the
451 Figs. 7d and 7e, we propose an detailed interpretation of the geophysical features displayed by interpreted on the attribute
452 images, associated to the faults highlighted after an accurate analysis of the discontinuities of attributes signatures, as shown
453 in fig. 5. Regarding the deeper parts of the sections, the together with the location of the focal mechanisms of the principal
454 mainshocks.

455 The deep, high amplitude reflector (H, blue arrows and dashed line) highlighted to the West of Nb in NOR01 (and at 2.5 s,
456 in Figs. 2d and 7d and in Figs. 3d of in CAS01), presents a seismic character and an attribute signature compatible with
457 the deeper reflector one deeper visible in of NOR02 beneath CNb (3.2 s, in Figs. 4b and 7e). This set of reflectors is
458 interpreted as a high acoustic impedance contrast, possibly related to an important velocity inversion occurring between the
459 Triassic Evaporites (anhydrites and dolostones, $V_p \approx 6$ km/s, e.g. Trippetta et al., 2010) and the underlying acoustic
460 Basement (metasedimentary rocks, $V_p \approx 5$ km/s, sensu Bally et al., 1986). Comparable deep and prominent reflections
461 reflectors were detected also in adjacent regions of the Umbria-Marche Apennines (e.g. Barchi et al., 1998; Mirabella et al.,
462 2008) thus confirming its regional importance, particularly because it represents a lithological control marking a seismicity
463 cutoff (Chiaraluce et al., 2017; Mirabella et al., 2008; Porreca et al., 2018; Mancinelli et al., 2019).

464 As already pointed out in the previous figures, the continuity of the deep reflector H is interrupted in the western edge by the
465 low-angle west-dipping discontinuity T crossing NOR02 (Fig. 7e), and not identified by Porreca et al. (2018). This deep
466 discontinuity can be interpreted as a regional thrust emerging at the footwall of the MSt, in an easternmost sector of the
467 region, and corresponding to the Acquasanta thrust (Centamore et al., 1993).

468 The continuity of the deep reflector H is interrupted in the western edge by the low angle west dipping T discontinuity
469 crossing NOR02 (Figs. 4d and 7e), not identified by Porreca et al. (2018). We interpret this discontinuity as the evidence of a
470 deep thrust emerging in the easternmost sector of the region.

471 The steep discontinuities highlighted by the attribute analysis are here interpreted as the seismic signature at depth of
472 complex normal faults mapped at the surface. More in detail, In NOR01, the most evident high-angle seismic discontinuity is
473 marked by an E-dipping fault, bordering the western area of Nb (Fig. 7d). The location and geometry of this fault, whose
474 presence is still debated in literature, perfectly match its supposed position at surface (Blumetti et al., 1993; Pizzi et al.,
475 2002; Galadini et al., 2018; Galli et al., 2018). Therefore, it may represent the first clear geophysical evidence of the
476 antithetic normal fault of Norcia (aNf), suggested by morphological studies (Blumetti et al., 1990) and paleoseismological
477 records (Borre et al., 2003) and belonging to a conjugate tectonic system (Brozzetti & Lavecchia, 1994; Lavecchia et al.,
478 1994).

479 the most evident seismic discontinuity is marked by an E dipping fault in NOR01, bordering Nb westward to the westthe
480 (Figs. 2d and 7d). The latter does not have a clear surface expression and therefore its presence is still debated in literature
481 (Blumetti et al., 1993; Pizzi et al., 2002; Galadini et al., 2018; Galli et al., 2018): its location and geometry in NOR01
482 perfectly match the supposed position at surface. Therefore, it may represent the evidence of the antithetic normal fault of
483 Norecia (aNf), belonging to a conjugate tectonic system (Brozzetti & Lavecchia, 1994; Lavecchia et al., 1994) and suggested
484 by morphological evidences (Blumetti et al., 1990) and paleoseismological records (Borre et al., 2003).

485 The other principal structure is a synthetic (W-dipping) high-angle, normal fault bordering the eastern flank of Nb
486 ("Nottoria-Preci fault" – Nf, Calamita et al., 1982; Blumetti et al., 1993; Calamita & Pizzi, 1994). The Nf in NOR02 is
487 marked by a downward propagation of a steep alignment (continuous red line in Fig. 7&d). This area is also fragmented by
488 the several minor strands parallel to the main faults (dashed lines in Fig. 7&d). In particular, several west-dipping minor
489 faults are observed in Fig. S5a, where the shallower high-amplitude reflectors of the PR attribute are clearly disrupted.

490 Another discontinuity interpretable as a deep fault is visible slightly eastward, close to the mainshock hypocentral location
491 (Fig. 8e7e). This E-dipping discontinuity, emphasized by the attribute analysis, does not reach the surface, whereas it is clear
492 at depth, as also evidenced by the attribute analysis. The presence of this blind fault has been suggested by several authors in
493 relation to the occurrence of an aftershock (Mw 5.4), which "ruptured a buried antithetic normal fault on eastern side of Nb,
494 parallel to the western bounding fault of CNb" (Chiaraluze et al., 2017, Porreca et al., 2018 and Improta et al., 2019).

495 Athe aNfIt is is a synthetic (W dipping) high angle, normal fault bordering the eastern flank of Nb ("Nottoria Preci fault"–
496 Nf, Calamita et al., 1982; Blumetti et al., 1993; Calamita & Pizzi, 1994). The Nf in NOR02 is evident by a downward
497 propagation of steep alignments (red arrows, Figs. 2e4e, 2d 4d and 7d), which that generates sharp lateral truncations of the

498 gently W-dipping reflectors. This area is also fragmented by several minor strands parallel to the main faults (Figs. 7d). In
499 addition, structure is visible slightly eastward (Figs. 4c, 4d red arrows between 7-9 km, ca. 1-3 s and westernmost dashed
500 black line in Fig. 7e). It is not reaching the shallower portion of the seismic line, but it is clearly visible in depth down to the
501 discontinuity T. This feature might be interpreted as a parallel E-dipping fault, moreover suggested by several authors to be
502 connected with an aftershock (Mw5.4), which that “ruptured a buried antithetic normal fault on eastern side of Nb, parallel to
503 the western bounding fault of CNb” (Chiaraluce et al., 2017, Porreca et al., 2018 and Improta et al., 2019).

504 The central portion of NOR02, corresponding to CNb, shows a peculiar reflection fabric, dominated by high-angle
505 discontinuities, interpreted as two opposite-dipping normal faults bordering the basin, well matching their positions mapped
506 at surface (cfr. Pierantoni et al., 2013).

507 The central sector of NOR02 including CNb, was described as a “triangle-shaped zone” by Porreca et al. (2018), who remark
508 a generalized difficulty to detect the accurate position of the normal faults, thanks to ~~t~~The multi-attribute visualization
509 rendering, shows a clear reflection fabric dominated by high angle discontinuities. Those are interpretable as two opposite
510 dipping normal faults bordering the basin, well matching their positions mapped at surface (cfr. Pierantoni et al., 2013).

511 The main fault is here represented by the W-dipping Vf fault, reactivated during the 2016 earthquake (e.g. Villani et al.,
512 2018a). This structure, which can be traced, from its surface expression downward to hypocentre location, along its deep
513 seismic signature, made by sParallel to the Vf, several high-angle seismic discontinuities representing minor normal faults
514 cross-cutting the gently W-dipping reflectors (Fig. 78e, further details in Fig. s5).

515 Analogous considerations can be extended to a multitude of E-dipping steep discontinuities at the westward side of CNb.
516 These may represent the evidence of an antithetic fault (aVf), actually made by and several minor fault strands characterized
517 by high-angle dip at least in the shallow depths (Villani et al., 2018b). Such a fault appears connected at about 2-3 s to the
518 W-dipping master Vf, producing a geometry of a conjugate system geometry like observed at Nb (Fig. 8e).

519 At depth of 3.2 s, Vf fault clearly interrupt the continuity of the top basement reflector H, whilst the relationships with the
520 Acquasanta thrust (low-angle discontinuity T) is more ambiguous. Two alternative interpretations can be proposed,
521 schematically represented in Fig. 98. In Fig. 9a8a, we propose a model in which Vf merges into the deep Acquasanta thrust,
522 suggesting a negative inversion, as a mechanism proposed by other authors (e.g. Calamita and Pizzi, 1994; Pizzi et al., 2017
523 Scognamiglio et al., 2018). In Fig. 9b8b, Vf cuts and displaces the Acquasanta thrust, following a steeper trajectory (ramp)
524 (Lavecchia et al., 1994 and Porreca et al., 2018).

525 The main fault is here represented by the W-dipping Vf fault, reactivated during the 2016 earthquake (Villani et al., 2018a).
526 It can be traced, from its surface expression downward to hypocentre location along its deep seismic signature, made by
527 several high angle seismic discontinuities cross-cutting the gently W-dipping reflectors (Figs. 4d and 7e). At depth, the Vf
528 displace the Top Basement (H) and the thrust (T) at about 3.2 s.

529 Analogous considerations can be extended to the E-dipping set of steep events at the westward side of CNb. CNb These may
530 represent evidence of an antithetic fault (aVf), made by several minor fault strands (Villani et al., 2018b). Such a fault

531 appears connected at about 2-3 s to the W-dipping master V_f, producing a geometry of a conjugate system Nb (Figs. 4d and
532 7e).

533 For both Norcia and Castelluccio di Norcia basins, the interpreted data suggest two slightly asymmetric fault systems, due to
534 conjugate sets of seismogenic master faults (Ramsay & Huber, 1987) producing a “basin-and-range” morphology (Serva et
535 al., 2002), progressively lowering the topography from east to west, and forming two major topographic steps, corresponding
536 to the CNb and Nb, respectively. Such fault systems control the evolution of the continental basins, and are associated with
537 several complex sets of secondary strands building up complex fault zones. Such fault strands are able to produce surface
538 ruptures in future earthquakes, as occurred in the 2016-2017 seismic swarm, and would require further studies through high-
539 resolution geophysical investigations (e.g. Bohm et al., 2011 and Villani et al. 2019).

540 The results of the seismic interpretation proposed in this work, supported by the attribute images analysis, produced in this
541 work suggests. The attribute images produced in this work suggest that such synthetic and antithetic tectonic structures at the
542 Norcia and Castelluccio di Norcia basins cannot be actually simplified as a unique fault plane, but they could be imaged as
543 complex and fractured fault zones (Fz, in Fig. 7d), like also conceived also by Ferrario & Livio (2018) as “distributed
544 faulting and rupture zones”.

545 Conclusions

546 Taking into account the important role that seismic attributes play in the O&G industry, their usage might be of high interest
547 also for improving the geological interpretation of vintage seismic data, aimed to scientific purposes. When applied to
548 seismically active areas, this analysis may contribute to constrain the buried geological setting and, combined with
549 seismological data (i.e. focal mechanisms and accurate earthquake locations), may have high potential impact for the
550 identification and characterization of the seismogenic sources and eventually on earthquakes hazard assessment.

551 This contribution presents one of the first case studies where the seismic attribute analysis is used for seismotectonic
552 purposes. The analysis is applied to seismic reflection data collected more than 30 years ago in Central Italy. Such industrial
553 data, nowadays irreproducible in regions where the seismic exploration is forbidden, represent, despite the limited quality, a
554 unique source of information on the geological setting at depth. Taking into account the important role that seismic
555 attributes play in the O&G industry their usage might be of interest also for seismotectonic studies and having high potential
556 impact on earthquakes hazard assessment.

557 This contribution presents one of the first case studies where the seismic attribute analysis is used for seismotectonic
558 purposes. The analysis is applied to seismic reflection data collected more than 30 years ago in Central Italy. Such industrial
559 data, nowadays irreproducible in regions where the seismic exploration is forbidden, represent, despite the limited quality, a
560 unique high resolution source of information through high resolution images.

561 This contribution reveals that the use of seismic attributes can greatly improve the interpretation for the subsurface
562 assessment and structural characterization. Certainly, the overall low quality of the data sets did neither allow to extract rock

563 petrophysical parameters, nor more quantitative information. However, the attributes aid the seismic interpretation to better
564 display the reflection patterns of interest and provided new and original details on complex tectonic region in Central Italy.
565 Our attribute analysisWe considerably improved the overall interpretability of the vintage seismic lines crossing the
566 epicentral area of the 2016-2017 Norcia-Amatrice seismic sequence. In particular, we detected peculiar seismic signatures of
567 a deep horizon of regional importance, corresponding, most probably, to the base of the seismogenic layer, and to the
568 location and geometry of the complex active fault zones. Those consists of several secondary synthetic and antithetic splays
569 in both the Quaternary basins, generally consistent with its surface location, but also reinforcing the existence of several
570 faults with no clear surface outcrop, issue currently much debated in the literature.

571 The analysis and integration of the seismic attributes has-allowed the determination of the deep continuation of the (known
572 and supposed) faults and, the recently mapped co-seismic ruptures at surface, providing a pseudo-3D picture of the buried
573 structural setting of the area. The seismic attributes may help to reduce the gap between the surface geology and deep
574 seismological data, also revealing, a high structural complexity at different scales, that-which cannot generally be detected
575 only by using only-traditional interpretation techniques. This- approach has shown the potential of the attributes analysis-,
576 that even when applied on 2D vintage seismic lines, may significantly extend the data value. For all these reasons, we
577 strongly encourage its application for seismotectonic research, aimed to provide new information and additional constraints
578 across other seismically active regions around the world.

579 **Acknowledgments**

580 We are grateful to Eni S.p.A. for providing an inedited set of seismic reflection lines after the 2016-2017 seismic crisis in
581 Central Italy (raw data available in Fig.2 of supporting information). The original seismic reflection lines used in this study
582 are available in the supplementary material, well as the high-resolution Figures 2,3,4,7. The authors are very grateful to dgB
583 Earth Sciences and to QGIS teams for providing the academic software used in this work. We thank Dr. Christian Berndt and
584 Dr. David Jacopini for their valuable comments provided for this paper. We also thank the two anonymous reviewers for
585 their patience in providing useful suggestions and detailed corrections that considerably improved this work.-

- 587 Allen, C. R., St. Amand, P., Richter, C. F., & Nordquist, J.: Relationship between seismicity and geologic structure in the
588 southern California region. *Bulletin of the Seismological Society of America*, 55(4), 753-797, 1965.
- 589 [Anderlini, L., Serpelloni, E., and Belardinelli, M. E.: Creep and locking of a low-angle normal fault: Insights from the
590 Altotiberina fault in the Northern Apennines \(Italy\), *Geophys. Res. Lett.*, 43, 4321– 4329, doi:10.1002/2016GL068604,
591 2016.](#)
- 592 Anelli, L., Gorza, M., Pieri, M., and Riva, M.: Subsurface well data in the Northern Apennines (Italy). *Memorie della
593 Società Geologica Italiana*, 48, 461–471, 1994.
- 594 Bally, A. W., Burbi, L., Cooper, C., & Ghelardoni, R.: Balanced cross-sections and seismic reflection profiles across the
595 central Apennines. *Memorie della Società Geologica Italiana*, 35, 257–310, 1986.
- 596 Barchi, M.: Integration of a seismic profile with surface and subsurface geology in a cross-section through the Umbria-
597 Marche Apennines. *Bollettino della Società Geologica Italiana*, 110, 469–479, 1991.
- 598 Barchi, M. R., Minelli, G. and Pialli, G.: The CROP 03 Profile: a synthesis of results on deep structures of the Northern
599 Apennines, *Mem. Soc. Geol. It.*, 52, 383-400, 1998.
- 600 Barchi M.R., Galadini, F., Lavecchia, G., Messina, P., Michetti, A. M., Peruzza, L., Pizzi, A., Tondi & Vittori, E.: Sintesi
601 delle conoscenze sulle faglie attive in Italia Centrale: parametrizzazione ai fini della caratterizzazione della pericolosità
602 sismica. CNR-Gruppo Nazionale per la Difesa dai Terremoti, Roma, 62 pp., 2000.
- 603 Barchi, M., Landuzzi, A., Minelli, G., & Pialli, G.: Outer northern Apennines. In *anatomy of an orogen: The Apennines and
604 adjacent Mediterranean Basins*. Netherlands, Springer, 215–253, 2001.
- 605 Barchi, M. R., & Mirabella, F.: The 1997-98 Umbria-Marche earthquake sequence: “Geological” vs. “seismological” faults.
606 *Tectonophysics*, 476(1–2), 170–179. <https://doi.org/10.1016/j.tecto.2008.09.013>, ~~2009~~2008.
- 607 Barnes, A. E.: Theory of two-dimensional complex seismic trace analysis. *Geophysics*, 61, 264–272, 1996.
- 608 Barnes, A. E.: Attributes for automating seismic facies analysis. *Seg Technical Program Expanded Abstracts*, 19.
609 doi:10.1190/1.1816121, 1999.
- 610 Barnes, A. E.: "Displaying Seismic Data to Look Like Geology", chapter of: “Attributes: New Views on Seismic Imaging–
611 Their Use in Exploration and Production”, Marfurt, K. J. Gao, D., Barnes, A., Chopra, S., Corrao, A., Hart, B., James, H.,
612 Pacht, J., Rosen, N.C. (2011), *SEPM Society for Sedimentary Geology*, 31, doi: 10.5724/gcs.11.31, 2011.
- 613 Barnes, A., E.: *Handbook of Poststack Seismic Attributes*, Society of Exploration Geophysicists, 21, doi:
614 10.1190/1.9781560803324, 2016.
- 615 Basili, R., Valensise, G., Vannoli, P., Burrato, P., Fracassi, U., Mariano, S., ... & Boschi, E.: The Database of Individual
616 Seismogenic Sources (DISS), version 3: summarizing 20 years of research on Italy's earthquake geology. *Tectonophysics*,
617 453(1-4), 20-43, 2008.

618 Beidinger, A., Decker, K., & Roch, K. H.: The Lasse segment of the Vienna Basin fault system as a potential source of the
619 earthquake of Carnuntum in the fourth century AD. *International Journal of Earth Sciences*, 100(6), 1315-1329, 2011.

620 Bigi, S., Casero, P., & Ciotoli, G.: Seismic interpretation of the Laga basin; constraints on the structural setting and
621 kinematics of the central Apennines. *Journal of the Geological Society*, 168(1), 179–190. doi 10.1144/0016-76492010-084,
622 2011.

623 ~~Borre, K., Cacon, S., Cello, G., Kontny, B., Likke Andersen, H., Moratti, G., Piccardi, L., Stemberk, J., Tondi, E., Vilimek,
624 V.: The COST project in Italy: analysis and monitoring of seismogenic faults in the Gargano and Norcia areas
625 (centralsouthern Apennines, Italy). *J. Geodyn.* 36, 3–18, 2003.~~

626 Blumetti, A.M., Coltorti, M., Dramis, F., Farabollini, P.: Due sezioni stratigrafiche nel Pleistocene medio della conca di
627 Norcia; implicazioni geomorfologiche e neotettoniche. *Rend. Soc. Geol. Ital.* 13, 17–26, 1990.

628 Blumetti, A. M., Dramis, F., & Michetti, A. M.: Fault-generated mountain fronts in the central Apennines (Central Italy):
629 Geomorphological features and seismotectonic implications. *Earth Surface Processes and Landforms*, 18(3), 203–223. doi:
630 10.1002/esp.3290180304, 1993.

631 ~~Bohm, G., Luzi, L., Galadini, F.: Tomographic depth seismic velocity model below the plain of Norcia (Italy) for site effect
632 studies. *Bollettino di geofisica Teorica ed Applicata*, 2011.~~

633 ~~Bonini, L., Toscani, G., & Seno, S.: Three dimensional segmentation and different rupture behavior during the 2012 Emilia
634 seismic sequence (Northern Italy). *Tectonophysics*, 630, 33–42, 2014.~~

635 ~~Bonini, L., Basili, R., Burrato, P., Cannelli, V., Fracassi, U., Maesano, F. E., et al.: Testing different tectonic models for the
636 source of the Mw 6.5, 30 October 2016, Norcia earthquake (central Italy): A youthful normal fault, or negative inversion of
637 an old thrust? *Tectonics*, 38, doi:10.1029/2018TC005185, 2019.~~

638 ~~Boncio, P., Brozzetti, F., Ponziani, F., Barchi, M., Lavecchia, G., & Piali, G.: Seismicity and extensional tectonics in the
639 northern Umbriamarche Apennines. *Memorie della Societa Geologica Italiana*, 52, 539–555, 1998.~~

640 Boncio, P., F. Brozzetti, and G. Lavecchia: Architecture and seismotectonics of a regional low-angle normal fault zone in
641 central Italy, *Tectonics*, 19(6), 1038–1055, doi:10.1029/2000TC900023, 2000.

642 ~~Boncio, P., Lavecchia, G., & Pace, B.: Defining a model of 3D seismogenic sources for seismic hazard assessment
643 applications: The case of central Apennines (Italy). *Journal of Seismology*, 8(3), 407–425, 2004.~~

644 ~~Bonini, L., Toscani, G., & Seno, S.: Three-dimensional segmentation and different rupture behavior during the 2012 Emilia
645 seismic sequence (Northern Italy). *Tectonophysics*, 630, 33-42, 2014.~~

646 ~~Bonini, L., Basili, R., Burrato, P., Cannelli, V., Fracassi, U., Maesano, F. E., et al.: Testing different tectonic models for the
647 source of the Mw 6.5, 30 October 2016, Norcia earthquake (central Italy): A youthful normal fault, or negative inversion of
648 an old thrust? *Tectonics*, 38, doi:10.1029/2018TC005185, 2019.~~

649 ~~Borre, K., Cacon, S., Cello, G., Kontny, B., Likke Andersen, H., Moratti, G., Piccardi, L., Stemberk, J., Tondi, E., Vilimek,
650 V.: The COST project in Italy: analysis and monitoring of seismogenic faults in the Gargano and Norcia areas
651 (centralsouthern Apennines, Italy). *J. Geodyn.* 36, 3–18, 2003.~~

652
653 [Botter, C., Cardozo, N., Hardy, S., Leconte, I., Escalona: From mechanical modeling to seismic imaging of faults: a](#)
654 [synthetic workflow to study the impact of faults on seismic. Mar. Pet. Geol. 57, 187-207, 2014.](#)
655 Brewer, J. A., Matthews, D. H., Warner, M. R., Hall, J., Smythe, D. K., & Whittington, R. J.: BIRPS deep seismic reflection
656 studies of the British Caledonides. *Nature*, 305(5931), 206, 1983.
657 Brozzetti, F., & Lavecchia, G.: Seismicity and related extensional stress field: the case of the Norcia seismic zone. *Annales*
658 *Tectonicae*, 8, 38–57, 1994.
659 Brozzetti, F., Boncio, P., Cirillo, D., Ferrarini, F., de Nardis, R., Testa, A., Liberi, F., & Lavecchia, G.: High resolution field
660 mapping and analysis of the August – October 2016 coseismic surface faulting (Central Italy Earthquakes): slip distribution,
661 parameterization and comparison with global earthquakes. *Tectonics*, 38. <https://doi.org/10.1029/2018TC005305>, 2019.
662 Bulhões, E.M.: Técnica “Volume de Amplitudes”. SBGF/6° Congresso Internacional da Sociedade Brasileira de Geofísica,
663 Rio de Janeiro, Anais (In Portuguese), 1999.
664 Calamita, F., Coltorti, M., Deiana, G., Dramis, F. and Pambianchi, G.: Neotectonic evolution and geomorphology of the
665 Cascia and Norcia depressions (Umbria-Marche Apennines), *Geografia Fisica e Dinamica Quaternaria*, 5, 263-276, 1982.
666 Calamita, F., & Pizzi, A.: Recent and active extensional tectonics in the southern Umbro-Marchean Apennines (Central
667 Italy). *Memorie della Società Geologica Italiana*, 48, 541–548, 1994.
668 Calamita, F., Pace, P., & Satolli, S., Coexistence of fault-propagation and fault-bend folding in curve-shaped foreland
669 fold-and-thrust belts: examples from the Northern Apennines (Italy). *Terra Nova*, 24(5), 396-406, 2012.
670 Carvalho, J., Taha, R., Cabral, J., Carrilho, F. and Miranda, M.: Geophysical characterization of the OtaVila Franca de Xira-
671 Lisbon-Sesimbra fault zone, Portugal. *Geophysical Journal International*, 174, 567-584, 2008.
672 Cavinato, G. P., & De Celles, P. G.: Extensional basins in the tectonically bimodal central Apennines fold-thrust belt, Italy:
673 Response to corner flow above a subducting slab in retrograde motion. *Geology*, 27(10), 955–958, 1999.
674 Centamore, E., Adamoli, L., Berti, D., Bigi, G., Bigi, S., Casnedi, R., et al.: Carta geologica dei bacini della Laga e del
675 Cellino e dei rilievi carbonatici circostanti. In: *Studi Geologici Camerti*, Vol. Spec. Università degli Studi, Dipartimento di
676 Scienze della Terra. SELCA, Firenze, 1992.
677 Cheloni, D., Falcucci, E., & Gori, S.: Half-graben rupture geometry of the 30 October 2016 MW 6.6 Mt. Vettore-Mt. Bove
678 earthquake, central Italy. *Journal of Geophysical Research: Solid Earth*, 124. <https://doi.org/10.1029/2018JB015851>, 2018.
679 Chen, Q. and Sidney, S.: Seismic Attribute Technology for Reservoir Forecasting and Monitoring. *The Leading Edge*, 16
680 (5): 445. <http://dx.doi.org/10.1190/1.1437657>, 1997.
681 Chiarabba, C., De Gori, P., Cattaneo, M., Spallarossa, D., & Segou, M.: Faults geometry and the role of fluids in the 2016–
682 2017 Central Italy seismic sequence. *Geophysical Research Letters*, 45, 6963–6971, 2018.
683 Chiaraluce, L., Barchi, M., Collettini, C., Mirabella, F. & Pucci, S. Connecting seismically active normal faults with
684 Quaternary geological structures in a complex extensional environment: the Colfiorito 1997 case history (northern
685 Apennines, Italy). *Tectonics* 24, TC1002, <https://doi.org/10.1029/2004TC001627>, 2005.

686 Chiaraluce, L., Di Stefano, R., Tinti, E., Scognamiglio, L., Michele, M., Casarotti, E., et al.: The 2016 Central Italy seismic
687 sequence: A first look at the mainshocks, aftershocks, and source models. *Seismological Research Letters*, 88(3), 757–771.
688 <https://doi.org/10.1785/0220160221>, 2017.

689 Chopra, S. & J. Marfurt, K.: Seismic attributes - A Historical Perspective. *Geophysics*. 70(5):3.
690 <https://doi.org/10.1190/1.2098670>, 2005.

691 Chopra, S. and Marfurt, K. J.: Seismic Attributes for Prospect Identification and Reservoir Characterization. SEG
692 Geophysical Developments Series No. 11, Stephen J. Hill, series editor and volume editor. ISBN 978-1-56080-141-2
693 (volume) - ISBN 978-0-931830-41-9 (series), 464 pp, 2007.

694 Chopra, S. & J. Marfurt, K.: Emerging and future trends in seismic attributes. *The Leading Edge*. 27. 298-318.
695 10.1190/1.2896620, 2008.

696 [Chopra, S. and Marfurt, K.J.: Volume co-rendering of seismic attributes — A great aid to seismic interpretation, SEG](#)
697 [Technical Program Expanded Abstracts. January 2011, 1150-1154, 2011.](#)

698 Ciaccio, M., Barchi M. R., Chiarabba, C., Mirabella, F. and Stucchi E.: Seismological, geological and geophysical
699 constraints for the Gualdo Tadino fault, Umbria-Marche Apennines (central Italy), *Tectonophysics*, 406, 233 – 247, 2005.

700 Civico, R., Pucci, S., Villani, F., Pizzimenti, L., De Martini, P. M., Nappi, R. & the Open EMERGE Working Group:
701 Surface ruptures following the 30 October 2016 Mw 6.5 Norcia earthquake, central Italy, *Journal of Maps*, 14:2, 151-160,
702 doi: 10.1080/17445647.2018.1441756, 2018.

703 Coltorti, M., Farabollini, P.: Quaternary evolution of the “Castelluccio di Norcia” basin (Umbro-Marchean Apennines,
704 central Italy). *Il Quaternario* 8(1), 149–166, 1995.

705 Cook, F. A., Albaugh, D. S., Brown, L. D., Kaufman, S., Oliver, J. E., & Hatcher Jr, R. D.: Thin-skinned tectonics in the
706 crystalline southern Appalachians; COCORP seismic-reflection profiling of the Blue Ridge and Piedmont. *Geology*, 7(12),
707 563-567, 1979.

708 De Guidi, G., Vecchio, A., Brighenti, F., Caputo, R., Carnemolla, F., Di Pietro, A., et al.: Co-seismic displacement on
709 October 26 and 30, 2016 (Mw 5.9 and 6.5) earthquakes in central Italy from the analysis of discrete GNSS network. *Natural*
710 *Hazards and Earth System Sciences Discussions*, 2017(May), 1–11. doi: 10.5194/nhess-2017-130, 2017.

711 Deschamps, A., Innaccone, G., & Scarpa, R.: The Umbrian earthquake (Italy) of 19 September 1979. *Annales Geophysicae*,
712 2, 29–36, 1984.

713 [De Lima, R. & Luiz Evangelista Teixeira, W. & Ramos de Albuquerque, F. & Lima Filho, F. P.: Ground Penetrating Radar](#)
714 [digital imaging and modeling of microbialites from the Salitre Formation, Northeast Brazil. *Geologia USP—Serie Científica*,](#)
715 [18, 187-200. doi: 10.11606/issn.2316-9095.v18-146075, 2018.](#)

716 Di, H., and AlRegib, G.: Semi-automatic fault/fracture interpretation based on seismic geometry analysis: Geophysical
717 Prospecting, doi: 10.1111/1365-2478.12769, 2019.

718 DISS Working Group: Database of Individual Seismogenic Sources (DISS), Version 3.2.1: A compilation of potential
719 sources for earthquakes larger than M 5.5 in Italy and surrounding areas. <http://diss.rm.ingv.it/diss/>, Istituto Nazionale di
720 Geofisica e Vulcanologia, doi: 10.6092/INGV.IT-DISS3.2.1, 2018.

721 [Ehsan, S. A., Carbonell, R., Ayarza, P., Martí, D., Pérez-Estaún, A., Martínez-Poyatos, D. J., Simancas, J. F., Azor, A.,](#)
722 [Mansilla, L.: Crustal deformation styles along the reprocessed deep seismic reflection transect of the Central Iberian Zone](#)
723 [\(Iberian Peninsula\). *Tectonophysics*, 621, 159-174, <https://doi.org/10.1016/j.tecto.2014.02.014>, 2014.](#)

724 [Ehsan, S. A., Carbonell, R., Ayarza, P., Martí, D., Martínez Poyatos, D., Simancas, J. F., Azor, A., Ayala, C., Torné, M. and](#)
725 [Pérez-Estaún, A.: Lithospheric velocity model across the Southern Central Iberian Zone \(Variscan Iberian Massif\): The](#)
726 [ALCUDIA wide-angle seismic reflection transect, *Tectonics*, 34\(3\), 535-554, doi: 10.1002/2014TC003661, 2015.](#)

727 Ercoli, M., Pauselli, C., Frigeri, A., Forte, E., & Federico, C.: “Geophysical paleoseismology” through high resolution GPR
728 data: A case of shallow faulting imaging in Central Italy. *Journal of Applied Geophysics*, 90, 27–40.
729 doi.org/10.1016/j.jappgeo.2012.12.001, 2013.

730 Ercoli M., Pauselli C., Frigeri A., Forte E. and Federico C.: 3-D GPR data analysis for high-resolution imaging of shallow
731 subsurface faults: the Mt Vettore case study (Central Apennines, Italy). *Geophysical Journal International*, 198:1(609-621).
732 doi: 10.1093/gji/ggu156, 2014.

733 Ercoli, M., Pauselli, C., Cinti, F.R., Forte, E. and Volpe, R.: Imaging of an active fault: Comparison between 3D GPR data
734 and outcrops at the Castrovillari fault, Calabria, Italy. *Interpretation*, 3(3), pp. SY57-SY66, 2015.

735 Ferrario, M. F., & Livio, F.: Characterizing the distributed faulting during the 30 October 2016, Central Italy earthquake: A
736 reference for fault displacement hazard assessment. *Tectonics*, 37, 1256–1273. <https://doi.org/10.1029/2017TC004935>,
737 2018.

738 Finetti, I. R., Boccaletti, M., Bonini, M., Del Ben, A., Geletti, R., Pipan, M., & Sani, F.: Crustal section based on CROP
739 seismic data across the North Tyrrhenian–Northern Apennines–Adriatic Sea. *Tectonophysics*, 343(3-4), 135-163, 2001.

740 ~~Ferrario, M. F., & Livio, F.: Characterizing the distributed faulting during the 30 October 2016, Central Italy earthquake: A~~
741 ~~reference for fault displacement hazard assessment. *Tectonics*, 37, doi:10.1029/2017TC004935, 2018.~~

742 Forte E., Pipan M., Casabianca D., Di Cuia R., Riva A.: Imaging and characterization of a carbonate hydrocarbon reservoir
743 analogue using GPR attributes. *Journal of Applied Geophysics*, 81, 76–87, 2012.

744 Forte E., Dossi M., Pipan M. and Del Ben A.: Automated phase attribute-based picking applied to reflection seismics,
745 *Geophysics*, 81, 2, V55-V64, doi: 10.1190/GEO2015-0333.1, 2016.

746 Galadini, F., & Galli, P.: Paleoseismology of silent faults in the central Apennines (Italy): The Mt. Vettore and Laga Mts.
747 Faults. *Annals of Geophysics*, 46. <https://doi.org/10.4401/ag-3457>, 2003.

748 Galadini, F., Falcucci, E., Gori, S., Zimmaro, P., Cheloni, D. and Stewart J. P.: Active Faulting in Source Region of 2016–
749 2017 Central Italy Event Sequence. *Earthquake Spectra*, 34, 4, 1557-1583, 2018.

750 Galli, P., Galadini, F., Calzoni, F.: Surface faulting in Norcia (Central Italy): a “paleoseismological perspective”.
751 *Tectonophysics*, 403, 117–130, 2005.

752 Galli, P., Galadini, F. & Pantosti, D.: Twenty years of paleoseismology in Italy, *Earth-Sci. Rev.*, 88(1–2), 89–117, 2008.

753 Galli, P., Galderisi, A., Ilardo, I., Piscitelli, S., Scionti, V., Bellanova, J., Calzoni, F.: Holocene paleoseismology of the
754 Norcia fault system (Central Italy), *Tectonophysics*, 745, 154-169, doi:10.1016/j.tecto.2018.08.008, 2018.

755 [Gersztenkorn, G., Marfurt, K.J.: Eigenstructure-based coherence computations as an aid to 3-D structural and stratigraphic](#)
756 [mapping. *Geophysics*, 64, 1468-1479, 1999.](#)

757 Gruppo di Lavoro Sequenza Centro Italia: Rapporto Bollettino Sismico Italiano sulla revisione dei giorni 24-26 agosto; 26-
758 27 ottobre; 30 ottobre - 1° novembre 2016. Bollettino Sismico Italiano (BSI), 13 pp., 2019.

759 ~~[Jibson, R.W., Allstadt, K.E., Rengers, F.K., and Godt, J.W.: Overview of the geologic effects of the November 14, 2016,](#)~~
760 ~~[Mw 7.8 Kaikoura, New Zealand earthquake, U.S. Geological Survey Scientific Investigations Report, 2017-5146, 39 pp.,](#)~~
761 ~~<https://doi.org/10.3133/sir20175146>, 2018.~~

762 Ha, T. N., Marfurt, K. J. and Wallet B. C., Hutchinson, B.: Pitfalls and implementation of data conditioning, attribute
763 analysis, and self-organizing mapping to 2D data: Application to the Exmouth Plateau, North Carnarvon Basin, Australia,
764 Interpretation, submitted, http://mcee.ou.edu/aaspi/submitted/2019/Ha_et_al_2019_Seismic_attributes_for_2D_data.pdf,
765 2019.

766 Hale, D.: Methods to compute fault images, extract fault surfaces, and estimate fault throws from 3D seismic images.
767 *Geophysics*, 78(2), O33–O43, <https://doi.org/10.1190/geo2012-0331.1>, 2013.

768 Hutchinson, B.: Application and Limitations of Seismic Attributes on 2D Reconnaissance Surveys: Master's thesis,
769 University of Oklahoma, 130 pp., 2016. <https://shareok.org/handle/11244/34658>.

770 [Iacopini, D., Butler, R.W.H.: Imaging deformation in submarine thrust belts using seismic attributes. *Earth Planet. Sci. Lett.*](#)
771 [302, 414-422, 2011.](#)

772 [Iacopini, D., Butler, R.W.H., Purves, S.: Seismic imaging of thrust faults and structural damage: a visualization workflow for](#)
773 [deepwater thrust belts. *First Break* 30, 39-46, 2012.](#)

774 [Iacopini, D., Butler, R. W. H., Purves, S., McArdle, N., & De Freslon, N.: Exploring the seismic expression of fault zones in](#)
775 [3D seismic volumes. *Journal of Structural Geology*, 89, 54-73, 2016.](#)

776 Improta, L., Latorre, D., Margheriti, L., Nardi, A., Marchetti, A., Lombardi, A. M., Castello, B., Villani, F., Ciaccio, M. G.,
777 Mele, F. M., Moretti, M. & the Bollettino sismico Italiano Working Group: Multi-segment rupture of the 2016 Amatrice-
778 Visso-Norcia seismic sequence (central Italy) constrained by the first high-quality catalog of early Aftershocks. *Scientific*
779 *Reports*, 9, 6921, 2019. doi: 10.1038/s41598-019-43393-2

780 ISIDe working group: version 1.0; doi:10.13127/ISIDe, 2016.

781 Ithaca catalogue, Available at: [http://www.isprambiente.gov.it/it/progetti/suolo-e-territorio-1/ithaca-catalogo-delle-faglie-](http://www.isprambiente.gov.it/it/progetti/suolo-e-territorio-1/ithaca-catalogo-delle-faglie-capaci)
782 [capaci](http://www.isprambiente.gov.it/it/progetti/suolo-e-territorio-1/ithaca-catalogo-delle-faglie-capaci), last accessed January 2019.

783 [Jibson, R.W., Allstadt, K.E., Rengers, F.K., and Godt, J.W.: Overview of the geologic effects of the November 14, 2016,](#)
784 [Mw 7.8 Kaikoura, New Zealand earthquake, U.S. Geological Survey Scientific Investigations Report, 2017-5146, 39 pp.,](#)
785 <https://doi.org/10.3133/sir20175146>, 2018.

786 Koopman, A.: Detachment tectonics in the central Apennines, Italy. *Geologica Eltraiectina*, 30, 1–155, 1983.

787 Lavecchia, G.: Il sovrascorrimento dei Monti Sibillini: Analisi cinematica e strutturale. *Bollettino della Società Geologica*
788 *Italiana*, 104, 161–194, 1985.

789 Lavecchia, G., Brozzetti, F., Barchi, M., Keller, J., & Menichetti, M.: Seismotectonic zoning in east-central Italy deduced
790 from the analysis of the Neogene to present deformations and related stress fields. *Geological Society of America Bulletin*,
791 106, 1107–1120, 1994.

792 Lavecchia, G., Castaldo, R., de Nardis, R., De Novellis, V., Ferrarini, F., Pepe, S., Brozzetti, F., Solaro, G., Cirillo, D.,
793 Bonano, M., Boncio, P., Casu, F., De Luca, C., Lanari, R., Manunta, M., Manzo, M., Pepe, A., Zinno, I., Tizzani, P.: Ground
794 deformation and source geometry of the 24 August 2016 Amatrice earthquake (Central Italy) investigated through analytical
795 and numerical modeling of DInSAR measurements and structural-geological data. *Geophysical Research Letters*, 43,
796 12,389–12,398 American Geophysical Union (AGU), 2016.

797 Lima, R. & Teixeira, L. E. W., de Albuquerque, F. R., and Lima-Filho, F. (2018). Ground Penetrating Radar digital imaging
798 and modeling of microbialites from the Salitre Formation, Northeast Brazil. *Geologia USP - Serie Cientifica*. 18. 187-200.
799 [10.11606/issn.2316-9095.v18-146075](https://doi.org/10.11606/issn.2316-9095.v18-146075).

800 Livio, F., Michetti, A. M., Vittori, E., Gregory, L., Wedmore, L., Piccardi, L., et al.: Surface faulting during the August 24,
801 2016, central Italy earthquake (Mw 6.0): Preliminary results. *Annals of Geophysics*, 59. doi: 10.4401/ag-7197, 2016.

802 Maesano, F. E., D'Ambrogi, C., Burrato, P., & Toscani, G.: Slip-rates of blind thrusts in slow deforming areas: examples
803 from the Po Plain (Italy). *Tectonophysics*, 643, 8-25, 2015.

804 Mancinelli, P., Porreca, M., Pauselli, C., Minelli, G., Barchi, M. R., & Speranza, F.: Gravity and magnetic modeling of
805 Central Italy: Insights into the depth extent of the seismogenic layer. *Geochemistry, Geophysics, Geosystems*, 20,
806 <https://doi.org/10.1029/2018GC008002>, 2019.

807 [Manning, T., Ablyazina, D. and Quigley, J.: The nimble node — Million-channel land recording systems have arrived. *The*
808 *Leading Edge*, 38:9, 706-714, doi.org/10.1190/tle38090706.1, 2019.](https://doi.org/10.1190/tle38090706.1)

809 Marfurt, K. J. Gao, D., Barnes, A., Chopra, S., Corrao, A., Hart, B., James, H., Pacht, J., Rosen, N.C.: SEPM Society for
810 Sedimentary Geology, 31, doi: 10.5724/gcs.11.31, 2011.

811 [Marfurt, K.J., Alves, T.M.: Pitfalls and limitations in seismic attribute interpretation of tectonic features. *Interpretation* 3, 5-
812 15. <http://dx.doi.org/10.1190/INT-2014-0122.1>, 2015.](https://doi.org/10.1190/INT-2014-0122.1)

813 Marfurt, K. J.: Seismic Attributes as the Framework for Data Integration throughout the Oilfield Life Cycle, SEG, 508 pp.,
814 2018.

815 Martinis, B., and Pieri, M.: Alcune notizie sulla formazione evaporitica dell'Italia centrale e meridionale. *Bollettino della*
816 *Società Entomologica Italiana*, 4, 649–678, 1964.

817 Mazzotti, A., Stucchi, E., Fradelizio, G., Zanzi, L., Scandone, P.: Seismic exploration in complex terrains: A processing
818 experience in the southern Apennines. *Geophysics*, 65(5), 1402–1417. <https://doi.org/10.1190/1.1444830>, 2000.

819 [McArdle, N.J., Iacopini, D., KunleDare, M.A., Paton, G.S.: The use of geologic expression workflows for basin scale](#)
820 [reconnaissance: a case study from the Exmouth Subbasin, North Carnarvon Basin, northwestern Australia. Interpretation 2,](#)
821 [163-177, 2014.](#)

822 McClymont, A. F., Green, A. G., Villamor, P., Horstmeyer, H., Grass, C. and Nobes, D. C.: Characterization of the shallow
823 structures of active fault zones using 3-D ground-penetrating radar data, *J. Geophys. Res.*, 113, B10315,
824 doi:10.1029/2007JB005402, 2008.

825 Milli, S., Moscatelli, M., Stanzone, O., & Falcini, F.: Sedimentology and physical stratigraphy of the Messinian turbidites
826 deposits of the Laga basin (Central Apennines, Italy). *Bollettino della Società Geologica Italiana*, 126, 37–48, 2007.

827 Minelli, G., and Menichetti, M.: Tectonic evolution of the Perugia massifs area (Central Italy). *Bollettino della Società*
828 *Entomologica Italiana*, 109(5), 445–453, 1990.

829 Mirabella, F., Barchi, M. R. and Lupattelli, A.: Seismic reflection data in the Umbria Marche region: Limits and capabilities
830 to unravel the subsurface structure in a seismically active area. *Annals of Geophysics*, 51(2–3), 383–396.
831 <https://doi.org/10.4401/ag-3032>, 2008.

832 Nacini E. Z. and Prindle, K.: Machine learning and learning from machines, *The Leading Edge*, 37:12, 886-893, 2018.

833 Patacca, E., and Scandone, P.: Late thrust propagation and sedimentary response in the thrust-belt foredeep system of the
834 southern Apennines (Pliocene–Pleistocene). In G. Vai & I. Martini (Eds.), *Anatomy of an Orogen: The Apennines and*
835 *adjacent Mediterranean basins*, 441–454, Norwell, MA: Kluwer Acad., 2001.

836 Pauselli, C., Barchi, M. R., Federico, C., Magnani, M. B. and Minelli, G.: The crustal structure of the northern Apennines
837 (Central Italy): An insight by the CROP03 seismic line. *American Journal of Science*, 306(6), 428–450.
838 <https://doi.org/10.2475/06.2006.02>, 2006.

839 Pauselli, C., Federico, C., Frigeri, A., Orosei, R., Barchi, M.R. & Basile, G.: Ground Penetrating Radar investigations to
840 study active faults in the Norcia Basin (Central Italy), *Journal of Applied Geophysics*, 72, 39-45, 2010.

841 Pierantoni, P. P., Deiana, G., & Galdenzi, S.: Stratigraphic and structural features of the Sibillini Mountains (Umbria–
842 Marche Apennines, Italy). *Italian Journal of Geosciences*, 132, 497–520. <https://doi.org/10.3301/IJG.2013.08>, 2013.

843 Pizzi, A., Calamita, F., Coltorti, M., & Pieruccini, P.: Quaternary normal faults, intramontane basins and seismicity in the
844 Umbria-MarcheAbruzzi Apennine Ridge (Italy): Contribution of neotectonic analysis to seismic hazard assessment.
845 *Bollettino Società Geologica Italiana Special Publication*, 1(January), 923–929, 2002.

846 [Pizzi, A., Di Domenica, A., Gallovič, F., Luzi, L., & Puglia, R.: Fault segmentation as constraint to the occurrence of the](#)
847 [main shocks of the 2016 Central Italy seismic sequence. Tectonics, 36, 2370–2387, doi:10.1002/2017TC004652, 2017.](#)

848 Porreca, M., Minelli, G., Ercoli, M., Brobia, A., Mancinelli, P., Cruciani, F., Giorgetti, C., Carboni, C., Mirabella, F.,
849 Cavinato, G., Cannata, A., Pauselli, C., Barchi, M.R.: Seismic reflection profiles and subsurface geology of the area
850 interested by the 2016–2017 earthquake sequence (Central Italy). *Tectonics*, 37, 1-22, doi: 10.1002/2017TC004915, 2018.

851 Porreca, M., Fabbrizzi, A., Azzaro, S., Pucci, S., Del Rio, L., Pierantoni, P. P., Giorgetti C., Roberts G., Barchi, M. R.: 3D
852 geological reconstruction of the M. Vettore seismogenic fault system (Central Apennines, Italy): Cross-cutting relationship
853 with the M. Sibillini thrust. Journal of Structural Geology, 103938, 2020.
854 Press, F., and D. Jackson: Alaskan earthquake, 27 March 1964: Vertical extent of faulting and elastic strain energy release,
855 Science, 147, 867, 1965.
856 Pucci, S, De Martini, P.M., Civico, R., Villani, F, Nappi, R., Ricci, T., Azzaro, R., Brunori, C. A., Caciagli, M., Cinti, F. R.,
857 Sapia, V., De Ritis, R., Mazzarini, F., Tarquini, S., Gaudiosi, G., Nave, R., Alessio, G., Smedile, A., Alfonsi, L., Cucci, L.,
858 Pantosti, D.: Coseismic ruptures of the 24 August 2016, Mw6.0 Amatrice earthquake (central Italy). Geophysical Research
859 Letters, American Geophysical Union (AGU), 2017.
860 Ramsay, J. G., Huber, M. I.: The Techniques of Modern Structural Geology: Folds and Fractures. Elsevier Science, 391 pp.,
861 1987.
862 Roure, F., P. Choukroune, X. Berastegui, J. A. Munoz, A. Villien, P. Matheron, M. Bareyt, M. Seguret, P. Camara, and J.
863 Deramond: Ecore deep seismic data and balanced cross sections: Geometric constraints on the evolution of the Pyrenees,
864 Tectonics, 8(1), 41–50, doi:10.1029/TC008i001p00041, 1989.
865 Rovida, A., Locati, M., Camassi, R., Lolli, B., & Gasperini P. (Eds.): CPTI15, the 2015 version of the parametric catalogue
866 of Italian earthquakes, Istituto Nazionale di Geofisica e Vulcanologia. <https://doi.org/10.6092/INGV.IT-CPTI15>, 2016.
867 Schwartz, D. P., & Coppersmith, K. J.: Fault behavior and characteristic earthquakes: Examples from the Wasatch and San
868 Andreas fault zones. Journal of Geophysical Research: Solid Earth, 89(B7), 5681-5698, 1984.
869 Scognamiglio, L., Tinti, E., Casarotti, E., Pucci, S., Villani, F., Cocco, M., Magnoni, F., Michelini, A., Dreger, D.: Complex
870 fault geometry and rupture dynamics of the Mw 6.5, 2016, October 30th central Italy earthquake. J. Geophys. Res.: Solid
871 Earth 123, 2943–2964, doi:10.1002/2018jb015603, 2018.
872 Serva L., Blumetti A.M., Guerrieri L. and Michetti A.M.: The Apennine intermountain basins: the result of repeated strong
873 earthquakes over a geological time interval. Boll. Soc. Geol. It., 1, 939-946, 2002.
874 Simancas, J. F., Carbonell .R., González Lodeiro, F., Pérez Estaún, A., Juhlin, C., Ayarza, P., Kashubin, A., Azor, A.,
875 Martínez Poyatos, D., Almodóvar, G.R., Pascual, E., Sáez, R., Expósito, I.: Crustal structure of the transpressional Variscan
876 orogen of SW Iberia: SW Iberia deep seismic reflection profile (IBERSEIS), *Tectonics*, 22, 1062,
877 doi:10.1029/2002TC001479, 6, 2003.
878 Snieder R. and Trampert J.: Inverse Problems in Geophysics. In: Wirgin A. (eds) Wavefield Inversion. International Centre
879 for Mechanical Sciences (Courses and Lectures), vol 398. Springer, Vienna, 1999.
880 Taner, M.T., Koehler, F., and Sheriff, R.E.: Complex Seismic Trace Analysis. Geophysics, 44 (6): 1041.
881 <http://dx.doi.org/10.1190/1.1440994>, 1979.
882 Taner, M.T.: Seismic attributes. Canadian Society of Exploration Geophysicists Recorder, 26. 48-56, 2001.
883 Tarquini, S., Isola, I., Favalli, M., & Boschi, E.: TINITALY/01: a new triangular irregular network of Italy. Annals of
884 Geophysics, 50–53, 2007.

885 Tarquini, S., Vinci, S., Favalli, M., Doumaz, F., Fornaciai, A., & Nannipieri, L.: Release of a 10-m-resolution DEM for the
886 Italian territory: Comparison with global-coverage DEMs and anaglyph-mode exploration via the web. *Computers and*
887 *Geosciences*, 38(1), 168–170. <https://doi.org/10.1016/j.cageo.2011.04.018>, 2012.

888 Trippetta, F., Collettini, C., Vinciguerra, S., & Meredith, P. G.: Laboratory measurements of the physical properties of
889 Triassic evaporites from Central Italy and correlation with geophysical data. *Tectonophysics*, 492(1), 121–132, 2010.

890 Torvela T., Moreau, J., Butler, R., W. H, Korja, A. and Heikkinen, P.: The mode of deformation in the orogenic mid-crust
891 revealed by seismic attribute analysis, *Geochem., Geophys., Geosyst.*, 14, 1069–1086, 2013.

892 Vai, G. B.: Basement and early (pre-Alpine) history. In G. B. Vai & I. P. Martini (Eds.), *Anatomy of an orogen: The*
893 *Apennines and adjacent Mediterranean basins*, 121–150, Dordrecht, Netherlands: Kluwer Academic Publisher.
894 https://doi.org/10.1007/978-94-015-9829-3_10, 2001.

895 Valoroso, L. et al. Radiography of a normal fault system by 64,000 high-precision earthquake locations: The 2009 L’Aquila
896 (central Italy) case study. *J. Geophys. Res. - Solid Earth*, 118, 1156–1176, <https://doi.org/10.1002/jgrb.50130>, 2013.

897 Vernengo, L., Trinchero, E., Torrejón, M. G., and Rovira, I.: Amplitude volume technique attributes and multidimensional
898 seismic interpretation. *The Leading Edge*, 36(9), 776–781. <https://doi.org/10.1190/tle36090776.1>, 2017.

899 Villani, F., Pucci, S., Civico, R., De Martini, P. M., Cinti, F. R., & Pantosti, D.: Surface faulting of the 30 October 2016 Mw
900 6.5 central Italy earthquake: Detailed analysis of a complex coseismic rupture. *Tectonics*, 37, 3378–3410.
901 <https://doi.org/10.1029/2018TC005175>, 2018a.

902 Villani, F., Sapia, V., Baccheschi, P., Civico, R., Di Giulio, G., Vassallo, M., et al.: Geometry and structure of a fault
903 bounded extensional basin by integrating geophysical surveys and seismic anisotropy across the 30 October 2016 Mw 6.5
904 earthquake fault (central Italy): The Pian Grande di Castelluccio basin. *Tectonics*, 37.
905 <https://doi.org/10.1029/2018TC005205>, 2018b.

906 [Villani, F., Maraio, S., Bruno, P.P., Improta, L., Wood, K., Civico, R., Baccheschi, P., Sapia, V., Pucci, S., Brunori, C.A.,](#)
907 [De Martini, P.M., Pantosti, D., Conti, P., Doglioni, C.: High-resolution seismic profiling of the Castelluccio basin: new](#)
908 [constraints on the shallow subsurface of the 30 October 2016 Mw 6.5 Norcia earthquake fault \(central Italy\). *Proceeding of*](#)
909 [the 38° Convegno GNGTS, 2019.](#)

910 ~~Yi, S., Wu, C., Li, Y. et al. *J. Mt. Sci.*: Source tectonic dynamics features of Jiuzhaigou Ms 7.0 earthquake in Sichuan~~
911 ~~Province, China, *Journal of Mountain Science*, 15(10): 2266–2275. doi: 10.1007/s11629-017-4703-6, 2018.~~

912 Wilkinson, M. W., McCaffrey, K. J. W., Jones, R. R., Roberts, G. P., Holdsworth, R. E., Gregory, L. C., et al.: Near-field
913 fault slip of the 2016 Vettore Mw 6.6 earthquake (Central Italy) measured using low-cost GNSS. *Scientific Reports*, 7(1),
914 4612, doi:10.1038/s41598-017-04917-w, 2017.

915 ~~Wrona, T., Pan, I., Gawthorpe, R. L. and Fossen, H.: Seismic facies analysis using machine learning, *Geophysics*, 83:5, O83–~~
916 ~~O95, 2018.~~

917 ~~Wyss, M. and Brune, J. N.: The Alaska earthquake of 28 March 1964: A complex multiple rupture, *Bull. Seism. Soc. Amer.*~~
918 ~~57, (5), 1967.~~

919 Wrona, T., Pan, I., Gawthorpe, R. L. and Fossen, H.: Seismic facies analysis using machine learning, Geophysics, 83:5, O83-
920 O95, 2018.

921 Zhao, W., Forte, E., Fontolan, G., Pipan, M.: Advanced GPR imaging of sedimentary features: integrated attribute analysis
922 applied to sand dunes, Geophysical Journal International, 213:1, 147–156, <https://doi.org/10.1093/gji/ggx541>, 2018.

923 ---

924 **Figure 1**

925 **Figure 2**

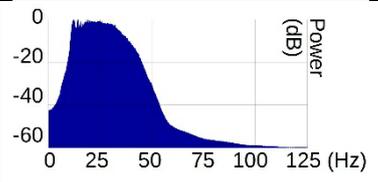
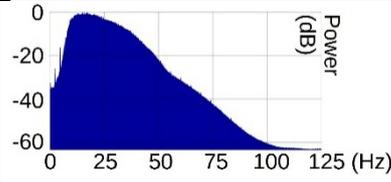
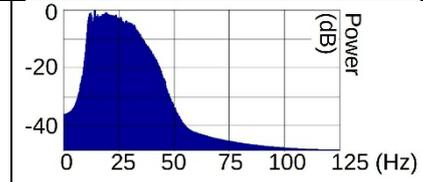
926 **Figure 3**

927 **Figure 4**

928 **Figure 5**

929 **Figure 6**

930 **Figure 7**

Parameters	NOR01	NOR02	CAS01
Source	Vibroseis	<u>Explosive</u>	<u>Vibroseis</u>
Length (km)	14	20	16
Number of traces	938	825	1069
Samples/trace	1600	1750	1600
Time window (ms)	6400	7000	6400
Sampling interval (ms)	4	4	4
Trace interval (m)	15	25	15
Mean Spectral amplitude (dB)			

932

933 **Figures and Tables captions:**

934 **Figure 1:** Simplified geological map of the study area (modified after Porreca et al., 2018), showing the 2D seismic data tracks, the
935 2016-2017 mainshock locations, beachballs with earthquake magnitude ~~beachballs and magnitudes~~, the surface ruptures and the
936 known master faults. Norcia basin (Nb), Castelluccio di Norcia basin (CNb), Monti Sibillini Thrust (MSt), Mt. Vettore fault (Vf),
937 antithetic (aVf), Norcia fault (Nf), antithetic Norcia fault (aNf).

938 **Figure 2:** Stack version of NOR01; a) standard reflection amplitude amplitude line, in the insert on the top the main faults mapped
939 at surface. The yellow dots label A underlines a processing artefact whilst; the boxes i) and ii) indicate the clearest reflectors: b
940 (A); b) Energy attribute enhancing a strong reflectivity contrasts (H, blue arrows); c) Energy Gradient, improving the detection of
941 dipping alignments and continuity of reflectors; d) Pseudo-Relief enhancing the reflection patterns cross-cut by steep
942 discontinuities (red arrows). Nf Norcia fault, aNf antithetic Norcia fault at surface, yellow dots = A, blue arrows = H, red arrows =
943 of the main lineaments and areas with major discontinuities highlighted by the attributes. indication of fault lineaments and fault
944 zones.

945 **Figure 3:** Stack version of CAS01, ~~with same attributes computation:~~ a) standard reflection amplitude line, on the top insert the
946 main faults mapped at surface. The label A underlines a processing artefact, whilst the boxes i) and ii) indicate the main visible
947 reflectors); b) Energy attribute c) Energy Gradient attribute; d) Pseudo-Relief, showing the strong regional reflector H. A high-
948 angle discontinuity on the western margin corresponds with the southern prosecution of aNf inferred at surface. is interpretable as
949 a normal fault, interpreted as aNf. Nf Norcia fault, aNf antithetic Norcia fault at surface, yellow dots = A, blue arrows = H, red
950 arrows = indication of the main fault lineaments and main signal discontinuities enhanced by the attribute's analysis. and fault
951 zones.

952

953 **Figure 4:** Time migrated version of NOR02; a) standard reflection amplitude line, on the top insert the main faults mapped at
954 surface; the box i) points out the most visible reflector; b) Energy attribute displaying the reflector H (blue arrows) and a possible
955 low angle discontinuity (T, green dots); c) Energy Gradient attribute, showing main lineaments detected master faults bounding
956 the basins; the master faults bounding the basins (red arrows); d) Pseudo-Relief, improving the reflectors continuity/discontinuity
957 and the master faults display of the areas with main signal discontinuities (red polygon) after the attribute computation. Nf Norcia
958 fault, aNf antithetic Norcia fault; Vf Mt. Vettore fault, aVf antithetic Mt. Vettore fault at surface, yellow dots = A, blue arrows =
959 H, green dots = T, red arrows = indication of the main lineaments f fault lineaments and fault zones.

960 **Figure 5:** Multi-attribute display of NOR02, displaying the position of the main faults at surface in relation to their deep seismic
961 attribute signature; a) Energy+Pseudo-Relief attributes, the seismic facie in the blue box is compared with the original seismic line
962 (b) and Energy+Pseudo-Relief (c) for comparison; the same plot for the black box is reported in figures d) and e) (original line and
963 Pseudo-Relief+Standard Amplitude, respectively).

964 **Figure 6:** Multi-attribute rendering of NOR01-, displaying the position of the main faults at surface in relation to their deep
965 seismic attribute signature using ODT software (depth conversion with VPav = 6000 m/s, vertical scale 2x); using ODT software
966 (depth conversion with VPav = 6000 m/s, vertical scale 2x); a) Energy+Pseudo-Relief attributes, the seismic facie in the blue box
967 showing a strong set of deep reflectors is compared with the original seismic line in b) and Energy+Pseudo-Relief c). An analogous
968 plot of the black box reports in figures d) and e) the original line and the combination Pseudo-Relief+Standard Amplitude.

969 **Figure 7:** Integration of surface and subsurface data (DTM by Tarquini et al., 2012); a) 3D-view (DTM by Tarquini et al., 2012) of
970 a W-E section crossing the Norcia and Castelluccio di Norcia basins (Nb and CNb), and the mainshock locations (ISIDE working
971 group, 2016). Surface and deep data allow to correlate the master faults and coseismic ruptures mapped at the surface. The multi-
972 attribute display of NOR01 (b) and NOR02 (c), is obtained overlapping the reflection amplitude in transparency with the Pseudo-
973 Relief and Energy attributes (red palette). The black boxes centred on the Norcia and Castelluccio di Norcia basins Nb and CNb
974 have been magnified for displaying the limits of the bounding faults (black dashed lines) and the main important reflectors
975 detected in depth. An important improvement of the subsurface images provides additional details on the seismogenic fault zones:
976 the sketches d) and e) show an interpretation reporting the two conjugate basins, showing master faults along the borders and
977 several some minor synthetic and antithetic splays (see d) and e) sketches).

978 **Figure 8:** The figure proposes two alternative interpretations of the relation between the normal Vf, the deep Acquasanta thrust
979 (T) and the Top- Basement reflector (H). Fig. 8a reports a model in which Vf merges into the deep Acquasanta thrust, suggesting a
980 negative inversion, as a mechanism proposed by some authors (e.g. Calamita and Pizzi, 1994; Pizzi et al., 2017 Scognamiglio et al.,
981 2018). In Fig. 8b, Vf cuts and displaces the Acquasanta thrust, following a steeper trajectory (ramp) as proposed by other authors
982 (Lavecchia et al., 1994 and Porreca et al., 2018; 2020).

983 **Table 1:** List of some parameters extracted from SEG-Y headers and three mean frequency spectra of the three seismic lines. An
984 approximate vertical resolution equal to 7580 m was derived ($v=6$ km/s).

985
986 Fig.s1: Figure summarizing the three original seismic reflection profiles in amplitude used in this work.

987 Fig.s2: Figure 2 reporting the computed seismic attributes without any line drawing and labels.

988 Fig.s3: Figure 3 reporting the computed seismic attributes without any line drawing and labels.

989 Fig.s4: Figure 4 reporting the computed seismic attributes without any line drawing and labels.

990 Fig.s5: The image is a magnification of two portions of NOR01 and NOR02, focused on the two basins of Norcia and Castelluccio
991 di Norcia, aiming to better display the discontinuities enhanced by the Pseudo Relief; a) PR on the Nb and interpretation of the
992 primary (continuous lines) and secondary faults (dashed lines); b) PR on the CNb and interpretation of the primary (continuous
993 lines) and secondary (dashed lines) faults bordering the basin.

994 The continuous red lines are the primary normal faults bounding Nb, whilst the dashed red segments compose a pattern of
995 possible secondary splays within the basin.

998 **Point-to-point authors response to Revision Files, by corresponding author MAURIZIO ERCOLI**
999 **on behalf of all co-authors.**

1000

1001 **Solid Earth Discussion Paper:**

1002 **Using Seismic Attributes in seismotectonic research: an application**
1003 **to the Norcia's Mw=6.5 earthquake (30th October 2016) in Central**
1004 **Italy.**

1005 Maurizio Ercoli^{1,4}, Emanuele Forte², Massimiliano Porreca^{1,4}, Ramon Carbonell³, Cristina Pauselli^{1,4}, Giorgio Minelli^{1,4},
1006 Massimiliano R. Barchi^{1,4}.

1007 ---

1008 **Colour and text code:**

1009 **- original text (first manuscript submission)**

1010 *- Rev1 and Rev2 comments: black italic*

1011 **- Authors replies: blue**

1012 ---

1013 **Manuscript Revision file – Reply to Rev1**

1014 **REV1 General Comments:**

1015 Ercoli et alii discuss the use of seismic attributes, applied to vintage seismic reflection data, for enhancing the structural
1016 interpretation and faults recognition with seismotectonic purposes. They present a case study by analyzing 3 vintage lines
1017 crossing the area interested by the 2016-2017 Central Italy seismic sequence. The study area is provided with updated
1018 geological maps, a dense cloud of earthquake foci and some moment tensor solutions following the 2016-2017 earthquake
1019 sequence and a dataset of earthquake-related surface ruptures, as well. This manuscript is quite well-written and the dataset
1020 worth publication, nevertheless this work needs some major revisions, due to i) **a badly addressed paper scope**, ii) **poor**
1021 **quality of the graphics** in their present form and iii) the somehow confusing way **the data and interpretations** are
1022 reported.

1023 I'm attaching an annotated version of the manuscript with many notes and suggestions; however, the major points of concern
1024 are summarized below:

1025 **- Data and interpretations are presented in a confusing way.** It is really difficult to follow the description of the
1026 recognized seismic features by means of a purely qualitative pattern recognition. Graphics are not helpful in this sense and
1027 the lack of univocal codes for e.g., faults and all the figures is making things worse. See the annotated text.

1028 **- The main point of the paper is that the use of seismic attributes can help in perform a better structural interpretation, in**
1029 **particular if applied to seismotectonic studies.** Some seismic features are here described through a **qualitative** approach and
1030 a possible interpretation is proposed. If the main target of the work is to show the usefulness of the seismic attributes an
1031 external dataset is needed for validation, but this is presently lacking. The use of seismic attributes allowed to identify a
1032 **possible set of secondary structures**, near the surface, in both the Castelluccio and Norcia basins, and to propose the
1033 presence of an antithetic fault bordering the Norcia basin to the west. Such an interpretation is not compared to detailed
1034 geological maps (only the main structures are shown but geology is not discussed (e.g., comparing possible offset from

1035 surface geology with geophysical data). As a result, **the comparison with mapped faults** is only qualitative and quite poor.
1036 Moreover, the **seismotectonic implications of the new interpretation** is totally overlooked in the discussion and/or
1037 conclusions. In this line, I would suggest changing the title: in the present form your focusing the attention on seismotectonic
1038 research it's a really side story in the present form. A possible way to solve the lack of validation would be to make **two**
1039 **different interpretations**, with and without attributes, on the same dataset, basing interpretation on objective and declared
1040 principles (e.g., cutoff, peculiar seismic facies, direct fault detections, axial surfaces dying out: : : etc.) and finally compare
1041 the results with published geological maps and or sections, including the discussion on opposite interpretations in literature.
1042 - Some recent works (see a note in the text – I'm reporting here e.g., Iacopini et al. 2016 - Iacopini, D., Butler, R. W. H.,
1043 Purves, S., McArdle, N., & De Freslon, N. (2016). Exploring the seismic expression of fault zones in 3D seismic volumes.
1044 Journal of Structural Geology, 89, 54-73.) proposed the use of seismic attributes for fault recognition. One of the advantages
1045 of this and other works is that you can produce a **quantitative analysis** of the wavelet, filtering out, on a statistical basis, the
1046 most probable fault plane locations. This could be helpful especially in cases where a direct detection of the seismic features
1047 is problematic. Any quantitative approach is lacking in this work: at least you should discuss the attribute range and
1048 distribution in the areas where you assume the fault should be located. I would strongly suggest **trying a quantitative**
1049 **approach**, at least a descriptive one. In summary, I had the impression that the aim of the work, as presently stated, is only
1050 partially achieved if an external dataset is not used for a detailed validation. Conversely, some interesting observations are
1051 arising from the Authors' interpretations: *the presence of an antithetic fault in the Norcia basin*, **the deep thick-skinned**
1052 **thrust** in NOR-2 section and the amount of possible distributed faults in the two basins. These points would benefit from
1053 more detailed **discussion and comparison with present proposed models in literature**. Finally, you surely have to expand
1054 the seismotectonic implications from your new interpretation. I'm sure the Authors will be able to face these criticisms and I
1055 hope that these notes will be useful to improve the present manuscript.

1056 ---

1057 **Reply to general comments of REV1:**

1058 Dear Rev1,

1059 thank you for your comments and corrections.

1060 Following your suggestions, we have deeply revised the manuscripts, and we hope that we addressed all the main criticisms.
1061 We have also revised all the minor suggested comments, even if in most of the paragraphs have been totally rewritten in this
1062 new revised version as explained below. Regarding your main comments, we have:

1063 i) improved the paper scope, focusing the attention on the use of the seismic attributes for a seismotectonic interpretation of
1064 the complex geological area affected by the recent seismic crisis;

1065 ii) improved the quality of the figures and graphics;

1066 iii) better distinguished the description of the data and their interpretation.

1067 In particular,

1068 - regarding the quality of the figures, we have improved the description of the seismic features and the graphics, that now
1069 have univocal codes (e.g. the faults line drawing and transparent polygons highlighting the interpreted fault zones) to avoid
1070 any confusion. All the main structural elements and discontinuities are now labelled and referred to the text.

1071 - regarding the data validation, we have already remarked in the discussion phase that a validation of the data and
1072 interpretation is basically impossible in this context: wells stratigraphy is available only in the surrounding sectors of the
1073 Apennines and not within or close to the study area. The geological complexity of this sector of the Apennines (involved at
1074 least by three tectonic phases from Jurassic to present day) does not allow to use well data, located far from the study area, to
1075 calibrate our interpretation. We have used all the geological map and stratigraphic information inferred by literature, as
1076 explained in Geological setting and Data chapters. Moreover, we have extensively used the fault patterns at surface
1077 (summarizing those main faults reported in literature) to drive the interpretation, starting from the near-surface, to link such
1078 structures to the hypocentral depth. Of course, we have then made the opposite process, drawing fault splays of fault zones
1079 where the attributes signature suggests their presence.

1080 Using this approach, we detected the presence of antithetic fault (debated in literature) at the Norcia Basin and of a deep
1081 thrust; we also highlight the presence of some secondary faults (unmapped or not outcropping) in both the Norcia (Nb) and
1082 Castelluccio (CNb) basins, characterized by fragmented and differently oriented seismic patterns. In our opinion, the
1083 presence of fault zones makes complex and probably an excessive simplification the drawing of single fault planes, at least at
1084 the resolution provided by these data. However, we have decided to make an additional effort improving the graphics also
1085 drawing, as suggested, some possible faults alignments in a new figure to better explain the interpretation process and
1086 criteria used. Where the high-dipping discontinuity (mainly in phase and/or amplitude) were separating different reflection
1087 patterns and truncating reflectors, we have added a primary fault (continuous red lines and polygons). When similar but
1088 smaller discontinuities between reflectors were particularly evident, parallel or antithetic to the principal faults, we have
1089 added a fault splay/secondary fault. A more quantitative approach as well as an estimation of the offset based on such data is
1090 difficult to achieve, therefore we have rewritten as suggested the description on the attribute performance in the areas where
1091 we think the faults are located.

1092 Regarding the discussion part of the paper, we have improved the seismotectonic implications with respect to the models
1093 debated in the literature. In particular, we have defined the main potential seismogenic faults at depths (e.g. Norcia and
1094 Vettore faults) and discussed the relationships of active normal faults and inherited structures highlighted by attributes
1095 analysis. In this latter case, we have proposed two different interpretations of the cross-cutting relationships between the
1096 seismogenic Vettore fault and a deep thrust (see last part of the chapter 6), as suggested by the Reviewer. We have also
1097 added a new figure (Fig. 8) to describe and compare these two models.

1098 Taking into account all these improved arguments on the seismotectonic features of the area, we have finally decided to keep
1099 the same title, focused on the seismotectonic implications of the seismic active area of the Apennines.

1100

1101 **Manuscript Revision file – Reply to Rev1 supplement:**

1102 ~~Lines 19-20: “This analysis resulted in peculiar seismic signatures which generally correlate with the~~
1103 ~~exposed surface geologic features, and also confirming the presence of other debated structures.”~~

1104 *REV1: Rather than this quite general sentence, insert one sentence summarizing the methods of analysis*
1105 *here adopted.*

1106 Authors: We have added short info on the attributes used, then we move forward the sentence, to
1107 reinforce the outcomes about the detection of faults currently debated in literature (e.g. the Norcia w-
1108 dipping antithetic fault).

1109 Line 27: Introduction

1110 Line 27: *shorten up the introduction avoiding repetitions and trying to better focus on the topic of the*
1111 *manuscript.*

1112 Authors: we have rewritten and shorten the introduction chapter, trying to improve the text and better
1113 focusing the main topics, as requested.

1114 Lines 29-31: “Clearly, this is not an easy task: it is in fact generally complex to fill the gap between the
1115 exposed geology including the active “geological faults” mapped by the geologists and the seismic
1116 features describing the geometry”

1117 *REV1: you made a big jump in the logic here. You are already focussing on seismic reflection data*
1118 *while there is a bunch of other techniques. you described some approaches later in the text but you*
1119 *should move that part here, I suppose.*

1120 *Authors: we have corrected and rewritten this sentence, introducing first the other geophysical*
1121 *techniques.*

1122 Lines 38-39: “This fact generates uncertainties that may amplify the scientific debate and the number of
1123 models introduced by the geoscientists. Therefore, this process requires the use of appropriate
1124 geophysical data, aimed at recovering information on the deep geological architecture and, in particular,
1125 on the geometry of active faults.”

1126 *Rev1: This statement is arguable: the aim should not be to obtain a consensus on interpretations but to*
1127 *provide as many constraints to interpretations as possible.*

1128 *Authors: we agree with this comment. We have rewritten this sentence focusing the attention on the use*
1129 *of the seismic attributes to improve the subsurface geological interpretation and to achieve additional*
1130 *information from the 2D data. The final aim is to obtain constraints on the geological structures*
1131 *responsible for the seismicity of the area, and in particular to define geological/structural setting at*
1132 *depth (e.g. depth of the basement and its involvement) and to trace of potentially seismogenic faults.*

1133 Lines 57-58: “To improve the data quality and increase the accuracy of the interpretation, three main
1134 strategies can be usually considered: 1) collection of new reflection seismic data with modern
1135 technologies, optimizing feasibility studies on the base of available vintage datasets;”

1136 *Rev1: this is partly already stated at lines 45-46.*

1137 *Authors: we have deeply reorganized and rewritten the text, removing possible repetitions.*

1138 Lines 63-65: “Some limitations characterize the first two approaches: the first is particularly demanding
1139 in terms of costs and logistic, and not practicable in zones where the use of dynamite or arrays of
1140 vibroseis trucks is forbidden or limited (e.g. National Parks or urban areas)”

1141 *Rev1: also this is already introduced at lines 45-46. try to sum up the three parts.*

1142 *Authors: we have modified this part to avoid repetitions, as requested.*

1143 Lines 92-98: “After the last 2016-2017 seismic sequence, Porreca et al. (2018) have provided a new
1144 regional geological model based on the interpretation of vintage 2D seismic lines. In such a study, the

1145 authors remark important differences in the seismic data quality across the region. In fact, the eastern
1146 area that shows higher overall data quality, is located at the footwall of the Mount Sibillini thrust (MSt)
1147 and, includes (consists of) flyschoid units of the Laga foredeep Domain. It is noteworthy that the Mw
1148 6.5 epicentral zone, is located on the MSt hanging-wall (Lavecchia, 1985). This is characterized by
1149 prevalent carbonate sequence and, its crossed by seismic sections with lower S/N ratio, that hampered
1150 the subsurface interpretation.”

1151 *Rev1: move this part from the introduction to the geologic framework*

1152 *Authors: we have moved this part to the geology chapter. This latter has been extensively re-organized*
1153 *as suggested also by Rev2.*

1154 Lines 100-101: “The main goal of this study is to obtain as much information as possible on the
1155 geological structures responsible for the seismicity.”

1156 *Rev1: try to rephrase. the aim is not clear. could you better explain what characteristics of the*
1157 *seismogenic source are you going to better define thanks to your analysis?*

1158 *Authors: we have rephrased the sentence improving the main aims of the study. See the response above.*

1159 Lines 103-104: “The current manuscript is an example of how can seismic attribute analysis contribute
1160 to seismotectonic research as an innovative approach.”

1161 *Rev1: this should be rephrased. limiting the impact of this work to a simple case study is not promising*
1162 *and adequate to this journal. The importance of this work could be by far better underlined if you*
1163 *clearly state from the very beginning the different interpretations postulated on the Central Italy*
1164 *seismogenic structures and your contribution on this open debate. The introduction should be mostly*
1165 *rewritten in this sense: at the moment there is a general overview on attribute analysis and you end up*
1166 *by proposing a case study.*

1167 *Authors: we totally agree with this comment as it was the aim of this work. We aim to present not only*
1168 *a case history, but we want suggest this approach as a valuable solution for seismotectonic studies*
1169 *around the world. Thus, we have improved and rewritten the introduction, trying to better explain the*
1170 *contribution of this study to seismotectonic interpretation of the area. We refer to Porreca et al. (2018)*
1171 *in the geology chapter for the different interpretations postulated about the Central Italy seismogenic*
1172 *structures.*

1173 Lines 108-109: ~~“Nine earthquakes with $M > 5$ and more than 97'000 events in two years have been~~
1174 ~~recorded at hypocentral depth not exceeding 12 km (Fig.1).”~~

1175 Authors: we have entirely rewritten the chapter. This sentence also has been modified, just to remark
1176 the importance of the 2016-2017 sequence.

1177 Lines 113-114: ”... belt, including the Umbria-Marche thrust and fold belt domain and Laga
1178 Formation.”

1179 *Rev1: add a REF here and introduce to international readers a brief sentence summing up the meaning*
1180 *of Umbria Marche and Laga Fm. significance.*

1181 Authors: we have modified the sentence adding some references.

1182 Line 120: “...faults since the Late Pliocene”

1183 *Rev1: add a ref here*

1184 Authors: done.

1185 Line 122: “sequence”

1186 *Rev1: you were referring to Laga Fm. above. be consistent.*

1187 Authors: done. We now refer to Laga sequence.

1188 Line 124: “velocity ($V_{av} = 4000$ m/s)”

1189 *Rev1: you were referring to Laga Fm. above. be consistent.*

1190 Authors: we have rewritten the text and fixed these issues.

1191 Lines 142-150: “...Norcia (Nb) and Castelluccio di Norcia basins (CNb) (Fig. 1). Nb and CNb are...”

1192 *Rev1: are all these acronimous really necessary? cue them when possible. e.g., Nb and CNb can be*
1193 *probably deleted.*

1194 Authors: we agree that in this section there are many acronyms, but we have decided to maintain in
1195 particular Nb and CNb, also following a Rev2 comment. They help to shorten the document and are
1196 useful to refer them to the figures.

1197 Lines 178-180: “...OpendTect (OdT) software... QGis software... from maps and Ithaca database”

1198 *Rev1: add the project URL, which maps? add the REF and project URL*

1199 Authors: we have added the URL and removed “maps” (already listed in the next raw) rewriting the
1200 sentence.

1201 Lines 192-193: “(Barnes 1996; Taner et al., 1979; Barnes, 1999; Chen and Sidney, 1997; Taner, 2001;
1202 Chopra and Marfurt, 2007; Chopra and Marfurt, 2008; Forte et al., 2016)”
1203 *Rev1: there is some other and more recent literature to be cited... e.g., Iacopini and Butler, 2011;*
1204 *Iacopini et al., 2012; McArdle et al., 2014; Botteret al., 2014; Hale, 2013 for a review; Marfurt and*
1205 *Alves, 2015*

1206 Authors: we have added the recent literature, as requested.

1207 Line 200: “Energy” (E):

1208 *Rev1: it would be better to provide a generalized formula, at least for this attribute.*

1209 Authors: We added a reference in the text, referring to a specific paper of our co-author Emanuele
1210 Forte, in which all the mathematical formulation is already provided within an exhaustive appendix.

1211 Line 206: “...useful to emphasize the most reflective zones...”

1212 *Rev1: provide a reference to the software used for attribute calculations.*

1213 Authors: reference are added in the text.

1214

1215

1216 **5. Results**

1217 Line 226:

1218 *Rev1: the reporting of the results in quite confused. There is excessive use of acronyms, text jumps*
1219 *from continuously from one sector to another making the reading very frustrating. More importantly,*
1220 *the text does not highlight the advantages and limitations of each technique. You should provide a first*
1221 *interpretation of faults, based on geological data and amplitude sections, and then provide a refined*
1222 *interpretation using seismic attributes. This approach would stress the real advantages of using seismic*
1223 *attributes.*

1224 Authors: We agree that there are some acronyms, but we have maintained most of them because the text
1225 would be even worse by repeating the long names of basins and faults (also following the advice of
1226 Rev2 to continue using Nb and CNb once defined). Then, we have reorganized the chapter improving
1227 the text and adding boxes/labels for interpretation of the amplitude section, poorly informative
1228 regarding the faults.

1229 Line 238: seismic profile, and in addition it is partially interfering with suspicious processing artefacts
1230 (highlighted with yellow dots, labelled as “A”, slightly undulated in Fig. 3a whilst horizontal in Fig. 2a
1231 ca. at 1 s)

1232 *Rev1: discuss this artefact. where is it coming from?*

1233 *Authors: the legacy seismic lines have been provided already processed by ENI, so we suspect this is*
1234 *the result of a windowed filter to remove horizontal noise or multiples. We have added this*
1235 *consideration within the text and we have marked these artefacts in the figures.*

1236 Line 243: “...by the EG and PR attributes (Figs. 3c and 3d) ...”

1237 *Rev1: data description is quite confusing: try to label each feature with letters on the seismic lines*
1238 *instead and refer to those codes.*

1239 *Authors: the acronyms for the main faults are provided on the top of the PR attribute, whilst letters are*
1240 *provided for the low angle features (H and T) (and blue/black boxes). We avoided to add extra letters*
1241 *and labels for the secondary faults to within the text and figures that are already dense of items.*

1242 Line 264-266: In fact, a main high-angle E-dipping discontinuity (red arrows) delimits the NOR02
1243 western sector (ca. 1 km of distance along the line at surface); another steep W-dipping alignment (red
1244 arrows) clearly cuts and slightly disrupt the set of reflectors below the Nb (0-2.5 s, ca. 4-5 km).

1245 *Rev1: there is a plenty of red arrows in Fig. 4 c and d. It is really hard to follow such a description.*
1246 *Maybe provide a letter for each element whose you are referring to in the text.*

1247 *Authors: as remarked in the comment above, we avoided to add more letters, the arrows indicate the*
1248 *main areas in which the discontinuities are visible. We have added transparent red polygons the help the*
1249 *readers to focus on the main discontinuity areas. We have also improved the quality and brightness of*
1250 *all the figures, and added an extra figure (s5) with two zooms on two areas to better show the*
1251 *alignments and better clarify the interpretative strategy and criteria.*

1252
1253 Line 268: fragmented reflectors pattern in the middle portion.

1254 *Rev1: there is no line drawing of these secondary elements in Fig. 4 c and d. Instead, in the Norcia*
1255 *basin (kms 0 to 5) some gently W-dipping reflectors can be traced, probably indicating backtilting to*

1256 *the west of this crustal sector. If this is true, the backtilting could possibly indicate that the main fault is*
1257 *the E-dipping one (see also Fig. 7). could you discuss this observation or discard this hypothesis?*

1258 *Authors:* In our opinion the W-dipping reflectors (we agree that these are the most evident features in
1259 the seismic profile) derive from the SW-dipping tectonic units, so they are mostly related to
1260 compressional tectonics. But in particular, in this sentence we wanted to highlight the fragmentation of
1261 these reflectors created by a dense pattern of subvertical discontinuities suggesting the presence of a
1262 fault zone (shown by a peculiar signature of faulting on these seismic lines). Instead of a single fault
1263 lineament we prefer the concepts of fault zone made by many steep secondary discontinuities and
1264 fragmented fabric concentrated in a narrow area. Following this consideration, we used first only some
1265 aligned red arrows to drive the readers' attention on the main discontinuity zones. Then have introduced
1266 also an additional figure (s5) as requested, with two magnifications on representative areas illustrating a
1267 simple interpretation of the most visible faults.

1268 *Line 276: "seems"*

1269 *Rev1: try to avoid the term "seem" and similar. It gives the feedback that your new imaging is not*
1270 *reducing the uncertainties. Moreover, in the data and results section, only objective information should*
1271 *be given.*

1272 *Authors:* ok, removed in the entire part.

1273 *Line 282: ...combined plot of the PR attribute...*

1274 *Rev1: "the multi-attribute rendering method should be introduced in "Methods"."*

1275 *Authors:* "multi-attribute display" was already in "Methods", but we have changed it now with
1276 "rendering" as requested and added a specific reference.

1277 *Line 282: ("similarity" palette) with superimposed the EG attribute ("energy" palette)"*

1278 *Rev1: this is not clear, what do you mean with superimposed? a transparency? or a multi-band false*
1279 *color rendering? the first I suppose.*

1280 *Authors:* transparency, corrected.

1281 *Line 283: "(depth conversion with $V_{pav} = 6000$ m/s, vertical scale 2x)."*

1282 *Rev1: sorry but I'm missing this point... could you be clearer?*

1283 *Authors:* deleted, it was a mistake.

1284 Line 285: “The blue box of Fig.5a is reported in Fig. 5b and 5c by...”

1285 *Rev1: this should go in the figure caption.*

1286 *Authors: text has been changed according to this.*

1287

1288 Lines 294-296: “Such results therefore ensure an easier and more accurate interpretation of the
1289 subsurface geological structures; those are connected with the surface geology and related to the
1290 hypocentre location of the main seismic events, that will be discussed more in detail within the
1291 following chapter.

1292 *Rev1: (divided in some points)*

1293 *1) no interpretations are given for these figures. In order to demonstrate the supposed enhancement you
1294 should provide a line drawing with horizons, cutoffs etc. on each rendering, demonstrating and
1295 discussing which seismic features are better imaged through each rendering.*

1296 *2) how can you assess that you are correctly interpreting the signal?*

1297 *3) Is there any geologic evidence such as the 2016 ground breaks?*

1298 *4) can you compare your sections with detailed fault strand traces after recently published geologic
1299 maps? e.g., Pierantoni et al.?*

1300 *Authors:*

1301 *1) Instead of using a standard line drawing we have used boxes, dashed lines and arrows to leave the
1302 sections cleaner for readers to see the improvements. But we have also added the figures 2s,3s,4s
1303 displaying the attributes without any interpretation labels as requested by the second reviewer, and also
1304 adding a new figure (5s) displaying the interpretation of two representative basins. In addition, our final
1305 interpretation has been summarized in figure 7, in a discussion considering all the other data available
1306 for the area including outcropping geological units (carbonate substrate vs. quaternary basins), the main
1307 faults and the surface ruptures (point 3). Finally, we have improved the figures drawing some boxes,
1308 lines, labels etc ... enhancing the features displayed by the attribute analysis.*

1309 *2) We remarked that the only constraints available are at the surface (geology and traced faults), so our
1310 seismic interpretation is clearly based on our experience and knowledge of the Central Apennines, from
1311 the geologic and geophysical point of views. On the other hands, the geophysical features are*

1312 interpreted using common and well-known principles available in literature, particularly regarding the
1313 signature of faulting. However, for the interpretation of the deepest (less-constrained) part of the
1314 seismic images, we have produced a new figure (Fig. 8) reporting two different interpretations as
1315 suggested by the Reviewer.

1316 3) Surface evidences can be observed in the field and there is a wide literature cited in this work, like
1317 co-seismic ruptures (e.g. Civico et al. 2018, Villani et al., 2018a, Brozzetti e al., 2019 and many others).
1318 Not only geomorphological and geological evidences, but also paleoseismological data (citations in the
1319 text). Surface ruptures have been observed in the Central Apennines area, also in the past, only after
1320 earthquakes of $M_w > 6$.

1321

1322 4) this is what we aimed to do in this work, but probably unclear in the first manuscript version.
1323 However, in this revision we have better separated in the text and figures the surface data (including
1324 known faults and surface ruptures, detailed in Fig.1) by our fault interpretation.

1325 We basically have started our seismic interpretation using the surface data, therefore “driving” our
1326 workflow using the location of the known faults and ruptures at surface. Secondly, by considering
1327 “peculiar signature of faulting” obtained by attributes computation, we interpreted other buried faults,
1328 fault zones or secondary splays. The best example, among our results, is the detection of a primary fault
1329 still debated in literature due to scarce surface evidences: it is the Norcia antithetic fault, that in our
1330 opinion is “seismically” very clear in our attribute sections.

1331

1332

1333

1334 Lines 310-312: “The deep, high-amplitude reflector (H, blue arrows and dashed line) highlighted to the
1335 West of Nb in NOR01 (at 2.5 s, in Figs. 2d and 7d and in Figs. 3d of CAS01), presents an attribute
1336 signature similar to the one deeper visible in NOR02 beneath CNb (3.2 s, in Figs. 4b and 7e).”

1337 *Rev1: this is a repetition...*”

1338 Authors: the aim of this sentence was to correlate and group the observation done for the H reflection
1339 visible in NOR01(and CAS01) with NOR02, that was not the objective of the previous chapter.
1340 However, we have rewritten the text, particularly focusing on the interpretation aspect.

1341 Lines 313-315: “This set of reflectors are interpreted as a high acoustic impedance contrast, possibly
1342 related to an important velocity inversion occurring between the Triassic Evaporites (anhydrites and
1343 dolostones, $V_p \approx 6$ km/s, e.g. Trippetta et al., 2010) and the underlying acoustic Basement
1344 (metasedimentary rocks, $V_p \approx 5$ km/s, sensu Bally et al., 1986).”

1345 *Rev1: this interpretation implies that the Sibillini thrust is thick-skinned. this is an important*
1346 *consequence of your interpretation. try to stress it in the discussion.*

1347 Authors: We have discussed to possible scenario for the interpretation of the deeper discontinuities and
1348 reflectors. In both cases we are not able to resolve the duality between thick- and thin-skinned tectonics.
1349 We have described this in the Discussion chapter.

1350 Line 323: “(Figs. 2d and 7d)”

1351 *Rev1: why are you not using the codes in Figs? this paragraph is really confusing. try to rewrite it with*
1352 *the help of univocal codes for surface geology and seismic sections...*

1353 Authors: we have rewritten the paragraph, using the codes introduced for the surface geology/faults and
1354 seismic sections.

1355 Line 351: “Those”

1356 *Rev1: ???*

1357 Authors: corrected

1358 FIGURES

1359 REV1

1360 Figure 1:

1361 *Rev1: Provide the codes for the faults reported in sections. Are these all the potentially active faults*
1362 *reported in geological maps or a selection of?*

1363 Authors: We have added the codes for the main faults bounding the basins. We provide, after a
1364 comprehensive literature review, a summary of all the main faults and secondary splays mapped on the
1365 area.

1366

1367 Figure 2:

1368 *Rev1: 2C -> these features in red are not well detectable. maybe you should use a more quantitative*
1369 *approach to characterize them. e.g., semblance coherence or other quantitative measures of attribute*
1370 *similarity...*

1371 *blue on green is not a good choice for the readability indicate H also here.*

1372 Authors: due to the nature of the data, we have declared that this study has a qualitative approach, being
1373 such results the best we are able to provide. The tests performed with other attributes like the similarity
1374 didn't perform well (see our reply the reply to Prof. Iacopini during the discussion).

1375 However, the Norcia antithetic fault looks clearer in comparison to the Norcia fault. The position of Nf
1376 is constrained by surface outcrops, but also looking at all the three attributes (particularly the PR, better
1377 showing the changes in the reflection patterns) it is plausible in this position and with this geometry,
1378 and suggesting a deformation spread in a narrow fault zone.

1379 We have updated the figure as requested, modifying the arrows for better visibility.

1380 Figure 3:

1381 *Rev1: 3A -> CHANGE THE COLORS IN ORDER TO INCREASE CONTRAST. provide a colorscale*
1382 *for the use palette: what is the range of values of each attribute?*

1383 Authors: We have increased the images contrast and added the colour bars as requested.

1384 Figures 5-6:

1385 *Rev1: in the main section report the letters for the insets... the fault from surface geology, in red, and*
1386 *their codes are not readable....*

1387 Authors: we have used a colour code (blue and black) thicker on the boxes. We have improved the fault
1388 labels at surface and the overall quality of the attribute images.

1389 Figure 7

1390 *Rev1: these beachballs have not been projected onto the 3D perspective. it could be misleading... you*
1391 *can simply report them in sections as done for the Mw 6.5 event.*

1392 Authors: apart the Mw 5.3 event very close to the line, the other events a too far for a reliable re-
1393 projection on the sections. So, we left only the beachball of the mainshock (rotated considering the
1394 perspective).

1395 Line 693: “EN+PR”

1396 *Rev1: expand the codes in the caption....*

1397 Authors: fixed

1398

1399 Final comment:

1400 We have produced a new figure in the main text (Fig.8) proposing two possible interpretations of the
1401 cross-cutting relations between deep reflectors, normal faults and thrust. We have also improved and
1402 added new figures in the Supplementary as requested during the first revision.

1403

1405 **REV2 General Comments:**

1406 The manuscript “Using Seismic Attributes in seismotectonic research: an application to the Norcia’s Mw=6.5 earthquake
1407 (30th October 2016) in Central Italy” by Maurizio Ercoli et al. submitted to Solid Earth proposes the use of seismic attribute
1408 analysis approach on three vintage reflection seismic profiles across the Norcia and Castelluccio di Norcia basins to
1409 determine the extension and geometry of the geological structures. This region was the epicentral area of the 2016-2017
1410 seismic crisis in central Italy. This manuscript could be of interest to geologists and geophysicists working in active
1411 tectonics and using reflection seismic data. However, in my opinion, it needs still some work in the structure of the writing
1412 and, most important, more work in the interpretation of the data or, at least, it needs to show more clearly all the
1413 interpretations the authors are doing. I am not an expert in the analysis of this type of data (onshore seismic data across rocky
1414 regions) but I have many difficulties to identify the same structures the authors are interpreting. At the end, I have had the
1415 impression that the authors have extended the surface map structures in depth following some possible alignments. My
1416 question is, would have they interpreted the same structures without the surface information? To me, there is a **high**
1417 **uncertainty in the interpretation of the alignments** in the seismic profiles that, then, I have problems to believe the final
1418 structural model proposed in the manuscript. Following there are some general comments on the different sections. I also
1419 provide a commented manuscript that hope will help to improve the quality of the manuscript and the presented results.
1420 Despite my criticism, to be intended solely as constructive, I warmly encourage the **authors to make any effort for the**
1421 **publication of this manuscript**, because of the relevance of the proposed approach and objectives.

1422 1. **Introduction** I think that in general the introduction needs to be restructured to emphasize the main aspects of what
1423 authors wants to expose. It is a very confusing introduction. I am not a native English speaker and I have found some errors,
1424 so I think that a native English speaker should review the final version of the manuscript. Some specific comments:
1425 Paragraph from lines 69 to 104 is a long paragraph that jumps from one idea to another and then back on. It is confusing and
1426 needs to be rewritten. Why mention 2D data vs 3D data various times? Just need to stress the differences and then stress the
1427 information and advantages of using 2D dataset, mainly which it is available and ready to work on. In addition, sentences
1428 like the one in lines 82-84 are out of sense in that paragraph. The stated between lines 85 and 98 is confusing. This may be
1429 rewritten, but also, I think that it makes no sense to explain all this in the introduction.

1430 2. **Geological framework** This section of the manuscript is a little bit confusing and difficult to follow. The authors jump
1431 from one topic to another in some paragraphs and is difficult to understand the geological structure of the area. I think it is
1432 necessary some organization. Begin for the big geological units, as done. Then, explain the structures, the fault systems in
1433 the area. Continue with the basins object of study. Finally talk about the seismicity in the area and the recent earthquakes and
1434 the faults that show surface rupture. In addition, I recommend the authors to be consistent with the names of the units, faults,
1435 for example, the Laga foredeep domain is referred in three or four different ways, and that is confusing.

1436 3. **Data** The authors mention a couple of times **the supporting information**, but in fact the information is provided in tables
1437 and figures in the manuscript. Also, the figures in the supporting information are not correctly identified and some errors of
1438 profiles identifications are present and must be corrected.

1439 4. **Methods** Authors comments that they have tested **several post-stack attributes**, but it is not clear at all why they select
1440 ones and not others. Maybe it is not necessary to explain this? I am not an expert in seismic attribute analysis.

1441 5. **Results** To me it is necessary to **include in the supplementary information the profiles** (original and attribute analysis)
1442 without any interpretation and each one on one page at a bigger scale. The profiles on the manuscript show arrows pointing
1443 to specific features that attract the attention towards the author’s interpretation. For example, in Fig2c the authors points with
1444 red arrows to some discontinuity (?) but at the same time the arrows mask reflectors around. I could point to similar features
1445 (orange arrow in the corresponding figure on my commented manuscript) that could point to a normal fault dipping to the
1446 W? That suggests me that the authors are just looking for structures that have been recognized at surface and not for all the
1447 other possible structures in the area/profiles. But again, without the un-interpreted profiles it is difficult to compare
1448 observations. I would recommend to describe each profile independently pointing to the observations done in each attribute
1449 profile and follow the same structure from one profile to the other. Begin with the seismic section and describe what you see

1450 and what is or could correspond the observed artefacts, then, the EN section with the specific observations, after, the EG
1451 section and, finally, the PR section. This makes things easy to the reader and not necessary to jump from one profile to the
1452 other and return. I suggest to identify the different high-dipping lineaments in the figures with letters (e.g., L1, L2) and then
1453 refer to them in the text. It would be much easier for the reader to understand to which lineament the authors are referring.
1454 In profile NOR02 the relationship between horizons T, H and the west-dipping lineament interpreted as bounding the CNb is
1455 not clear. In lines 256-259 it is said that horizon H is interrupted by horizon T, which crosses all the profile from east to west
1456 and dipping to the west. Later on, in lines 275-276 it is said that a west-dipping lineament truncates and disrupts horizons
1457 (discontinuities) T and H. In general, to me is very difficult to interpret the lineaments in all the profiles (as pointed in a
1458 number of comments in the manuscript) but in that case I think that the authors are proposing different interpretations for the
1459 same observations. This needs to be clarified.

1460 **6. Discussion and conclusions as said in various comments I have problems to interpret the steep discontinuities** on the
1461 different seismic profiles (amplitude and attributes). All the discussion is based on the authors interpretation and since I
1462 cannot interpret the same things, I cannot support it. But I am not a specialist in this type of seismic interpretations.

1463 ---

1464 Manuscript Revision file – Reply to Rev2 supplement:

1465 Dear REV2,

1466 thank you for all your detailed comments. In this new revised manuscript, we have improved, as
1467 suggested, the data interpretation in the text and the figures to show more clearly all the interpretations
1468 that we propose. We agree that an attribute analysis done for seismotectonics is a new and complex
1469 approach for non-experts, particularly on onshore vintage data like the ones reported in this work. But
1470 we aim to give some slight improvements (possibly not fantastic like in offshore 3D seismic volumes)
1471 supporting the data interpretation. We aim to suggest to scientists working on such topic a new
1472 approach able to achieve better constraints on seismic areas characterized by scarcity of deep data. To
1473 do this, we have declared at the beginning of the manuscript that our strategy is based on the extension
1474 of the surface map structures in depth by following some possible seismic alignments, as the geologic
1475 data at surface are the only constraints available (absence of deep wells stratigraphy).

1476 Regarding the main points:

1477 1) The introduction has been completely rewritten following all the suggestions and the correction of
1478 both reviewers. In particular, we have shortened it and better focused the aims of this work as
1479 explained above (please see also responses to Rev1).

1480 2) The geological framework has been totally reorganized and rewritten in a more logic way, using the
1481 scheme proposed by Rev2.

1482 3) The supporting material contained the raw seismic lines plus the high resolution (pdf) images of the
1483 attributes, effectively with some possible mistakes in the filenames. However, we have entirely

1484 reorganized the material. Now we have added 5 figures to the Supplementary material: fig.1
1485 summarizes the original lines, the figs. 2s, 3s, 4s reports the attributes without labels as requested for
1486 better comparison, fig.5 is finally another figure regarding the details of the PR attributes and their
1487 interpretation, related to the two tectonic-controlled Quaternary basins.

1488 4) We have improved this paragraph briefly describing the workflow done to select the attributes.
1489 Further details have been provided during the discussion phase in the reply to Iacopini, but later we
1490 have inserted in the manuscript only a summary. This is to avoid an excessive technical description
1491 which in our opinion would have distracted the reader from the main theme of the work.

1492 5-6) We have included in the supplementary information the original amplitude profiles as well as the
1493 attribute analysis without interpretations (point 3). We have also remarked in the text that we looked for
1494 structures that have been recognized at surface. We started our interpretation using this constraint at
1495 surface, but then we extended the interpretation to the geophysical signature of faulting also belonging
1496 to possible structures not outcropping in the area/profiles (mainly the two basins of Norcia and
1497 Castelluccio di Norcia). We have rewritten the text following the Rev2 advice, even without grouping
1498 similar observations to avoid boring repetitions. We have better labelled at least the main structures
1499 (aNf, aVf, Nf, Vf), even if without labelling each secondary splay to avoid an excessive use of the
1500 acronyms/labels in the text (note by Rev1).

1501
1502 **Manuscript Revision file – Reply to Rev2 supplement:**
1503 Lines 16: ...recently...

1504 *Rev2: Recently is an ambiguous term. Instead, you could include the time range of the seismic*
1505 *sequence.*

1506 *Authors: we agree with this comment, we have added the time range 2016-2017, as requested.*

1507 Lines 18: ... ~~currently the only available across the epicentral zone...~~

1508 *Authors: we decided to maintain this sentence but adding “at the regional scale” because such data are*
1509 *the only available, so we’d like to remark their importance.*

1510 Lines 34: ... impressive topographic changes...

1511 *Rev2: Consider to delete.*

1512 Authors: removed and changed with “important”

1513

1514

1515 Lines 36: ... While many studies on the surface geology are generally performed, especially after
1516 important events ...

1517 *Rev2: I do not agree with this. There have been studies of active faults around the world before the*
1518 *occurrence of a large earthquake, not just after. In fact, I would say that is on the contrary, a lot of*
1519 *faults have been studied that do not have produced an earthquake nor in recent or historical times.*

1520 Authors: we have entirely rewritten the introduction. See main comments.

1521 Lines 38-40: ... This fact generates uncertainties that may amplify the scientific debate and the number
1522 of models introduced by the geoscientists. Therefore, this process requires the use of appropriate
1523 geophysical data, aimed at recovering information on the deep geological architecture and, in particular,
1524 on the geometry of active faults.

1525 *Rev2: I have understand what authors want to express with this sentence after read it few times.*
1526 *Recommend to rewrite. Which process? Obtaining? Adquireing?*

1527 Authors: we have entirely rewritten the introduction.

1528 Lines 42-49: ... Different geophysical methods (e.g. Gravimetry, Magnetics, Electric and
1529 Magnetotellurics, Ground Penetrating Radar) may contribute to define the stratigraphy and structural
1530 setting of the upper crust at different scales. But the seismic reflection is largely the most powerful tool
1531 producing high-resolution images fundamental to trace the actual geometry of active faults at surface
1532 (usually mapped and reconstructed in geological cross-sections), from the near surface down to
1533 hypocentral depths. However, the ex-novo acquisition of onshore deep reflection data, possibly 3D, is
1534 often hampered by environmental problems, complex logistics, and high costs. These issues seriously
1535 limit the possible, widespread use of this technique for scientific research. Significant exceptions are
1536 research projects for deep crustal investigations like BIRPS (Brewer et al., 1983), CoCORP (Cook et
1537 al., 1979), ECORS (Roure et al., 1989) and CROP (Barchi et al., 1998; Finetti et al., 2001), IBERSEIS
1538 (Simancas et al., 2003).

1539 *Rev2 (grouped questions): is the method that provides...? Confusing, rewrite. Is this necessary?*
1540 *Nowadays seismic acquisition is extensively used, although I agree that it is being more difficult to*
1541 *acquire deep seismic data, but it is still possible. Is that necessary? I know that some research groups in*
1542 *France and Spain have acquired deep seismics (reflection and refractions) in the Mediterranean in the*
1543 *last decade, so in more recent times that all these other datasets.*

1544 *Authors: the entire paragraph was rewritten and recent references updated as requested.*

1545 *Lines 50-51: ...Such limitations can be partially overcome by considering old profiles (legacy data)*
1546 *acquired by the exploration industry. When collected in seismically active regions, such data may be*
1547 *used to connect the active faults mapped at the surface...*

1548 *Rev2: I am not in agreement with this statement, I think that even a little more difficult it is not*
1549 *impossible to acquire new seismic data. I think that the use of legacy data could be a nice source of*
1550 *data in places that new data is difficult to acquire due to lack of funding or that could provide new*
1551 *information to improve the geological models. I think that you try to justify the use of legacy data*
1552 *pointing to limitations instead of pointing to advantages, as would be the already availability of these*
1553 *data. I would consider to rewrite this part.*

1554 *Authors: We actually agree that it is not impossible, we have just remarked that currently it is not*
1555 *common to see research projects including acquisition of regional seismic reflection data for*
1556 *seismotectonic purposes. More common is the acquisition of high-resolution seismic at the scale of*
1557 *single basins. We appreciate the advice regarding a justification for using legacy data considering the*
1558 *advantages instead only the limitations. So, we have rewritten the introduction following this indication*
1559 *as requested.*

1560 *Line 51: ...such data may be used to connect the active faults...*

1561 *Rev2: to improve geological models... Usually researchers working in seismotectonics has tried to do*
1562 *that link between surface geology and earthquakes proposing different fault models, isn't it? The Italian*
1563 *active faults database localize active faults provide fault dip, seismogenic depth, so it defines a fault*
1564 *model for each source. Your data may improve the determination of the fault geometry and other*
1565 *characteristics.*

1566 Authors: We totally agree with this comment. We have specified that the results of this approach can be
1567 useful for constraining the subsurface geological setting and to provide new data on active tectonic
1568 structures. We have also cited the DISS database (Basili et al., 2008) as an example of database of
1569 active faults in Italy.

1570 Line 57: ... three main strategies can be usually considered...

1571 *Rev2: Where? In seismic processing?*

1572 Authors: we have rephrased the sentence: "In order to improve the data quality and increase the
1573 accuracy of the interpretation, two main strategies, ordinarily used by the O&G industry, can be applied
1574 on legacy data: 1) reprocessing from raw data using modern powerful capabilities, processing strategies
1575 and developments of newly performing algorithms and software; 2) use post-stack analysis techniques
1576 such as seismic attributes."

1577 Lines 66-67: ... The second requires broad projects encompassing specialized teams, high-computation
1578 power and generally long processing times, the latter is dependent on the quality of the raw data. The
1579 third strategy, in the case of the attribute analysis exploits a well-known and mature technique...

1580 *Rev2 (grouped questions): That is not true. I agree that it is a time consuming task and maybe you may
1581 need a dedicated workstation, but reprocessing seismic data does not requires a broad project and
1582 large teams.*

1583 *I do not understand this sentence, at the beginning I thought you were describing the third type of
1584 strategy, just after I have seen I was wrong. Rewrite*

1585 Authors: we have rewritten and simplified the entire paragraph, following the advices of both reviewers.
1586 Regarding the costs, time, and team availability, the problem is wide and complex to be fully described
1587 here. However, with "modern processing techniques" we meant specific type of workflows e.g.
1588 including Pre-Stack Depth Migration (PSDM), that may require high computational power, long time
1589 and teamwork if performed on densely sampled 2D lines and/or 3D data in a short time period.
1590 Currently, only the oil companies or their contractors have such possibilities, whilst clearly, it's less
1591 easy, even if not impossible, in academic environments. Of course, we agree that more conventional
1592 workflows, depending on the survey goals, can be accomplished with more limited efforts.

1593 Line 72: ... seismic volumes produced spectacular results...

1594 *Rev2: This is ambiguous. Could you describe very briefly these results or give a couple of examples?*
1595 *For example: "...volumes allow identifying ancient river channel and ..." in agreement with your*
1596 *citations.*

1597 *Authors: thank you for the suggestion. We have integrated the text as requested.*

1598 *Line 77: ... in complex geological areas ...*

1599 *Rev2: Just in complex geological areas? Consider to delete.*

1600 *Authors: We agree with your comment. We have modified the sentence as: "... the attribute analysis is*
1601 *probably the easiest, cheapest and fastest to qualitatively emphasize the geophysical features and data*
1602 *properties of reflection seismic data sets, producing benefits particularly in complex geological areas."*

1603 *Line 79: ... may not bring so impressive improvements ...*

1604 *Rev2: may not provide the same quality of information than on 3D*

1605 *Authors: corrected*

1606 *Lines 79-81: ... However, the main point is that inland, most of the sedimentary basins have actually*
1607 *been sampled by 2D grids of seismic profiles, or at least they have been probed by a few sparse 2D*
1608 *seismic lines.*

1609 *Rev2: Maybe that is your case, but it could be not the same thing in other areas. I would rewrite this*
1610 *sentence pointing that you use this data because it is the available data.*

1611 *Authors: We rephrase as follow: "However, the main point is that in the past, it was common to sample*
1612 *study areas inland by 2D grids of seismic profiles, being the full 3D seismic surveys rare"*

1613 *Lines 82-84: ... Whilst in the hydrocarbon industry this process is useful even if mainly driven by a*
1614 *constant necessity to reduce the costs (Ha et al., 2019), in seismotectonic researches it is affected by*
1615 *even worse limitations previously aforementioned ...*

1616 *Rev2: Consider to delete.*

1617 *Authors: we have cancelled this sentence as requested*

1618 *Line 87: ... Based on such considerations,...*

1619 *Rev2: Which ones?*

1620 *Authors: Deleted*

1621 *Line 90: ... proposed new approach ...*

1622 *Rev2: Which new approach?*

1623 *Authors: we have rewritten the text explaining better which approach we propose.*

1624

1625 Line 96: ... Mount Sibillini thrust (MSt) ...

1626 *Rev2: Indicate it in figure 1*

1627 *Authors: MSt added in Fig.1*

1628 Lines 103-104: ... The current manuscript is an example of how can seismic attribute analysis

1629 contribute to seismotectonic research as an innovative approach

1630 *Rev2: This is a conclusion.*

1631 *Authors: we have rewritten and integrated the sentence following the comments of both reviewers.*

1632 Line 108: ... L'Aquila and Colfiorito, ...

1633 *Rev2: Indicate the years of the events*

1634 *Authors: years added in the text.*

1635 Line 109: ... 97'000 events ...

1636 *Rev2: ?*

1637 *Authors: it was the total number of earthquakes recorded in two years. However, we have rewritten the*

1638 *sentence following the comments of both reviewers.*

1639 Lines 111-112: ... generating impressive co-seismic ruptures (Civico et al., 2018; Brozzetti et al.,

1640 2019)...

1641 *Rev2: Necessary? Where? Along the Mt Vettore fault? Also point to Fig1*

1642 *Authors: corrected.*

1643 Lines 113-128: The study area is located in the easternmost part of the Northern Apennines fold and

1644 thrust belt, including the Umbria-Marche thrust and fold belt domain and Laga Formation. This is a

1645 geologically complex region, where in the past the analysis of 2D seismic profiles have produced

1646 contrasting interpretation of the upper crust structural setting, e.g. thin vs. thick skinned tectonics, fault

1647 reactivation/inversion, basement depth (Bally et al., 1986; Barchi, 1991; Barchi et al., 2001; Bigi et al.,

1648 2011; Calamita et al., 2012; Porreca et al., 2018). The Umbria-Marche fold and thrust belt was formed

1649 during the Miocene compressive phase, and overthrusts the Laga foredeep sequence, through arc-

1650 shaped major thrusts, namely the M. Sibillini thrust (MSt, Koopman, 1983; Lavecchia, 1985), with
1651 eastward convexity. The compressional structures were later disrupted by the extensional faults since
1652 the Late Pliocene. The Umbria-Marche domain involves the rocks of the sedimentary cover, represented
1653 by three main units:

1654 1) on top, the Laga sequence consisting of siliciclastic turbidites belonging to the Laga foredeep and
1655 foreland Formation (Milli et al., 2007; Bigi et al., 2011); it is made by alternating layers of sandstones,
1656 marls and evaporites (Late Messinian – Lower Pliocene, up to 3000 m thick, average seismic velocity
1657 (v_{av}) = 4000 m/s), mainly outcropping in the eastern sector of the study area (i.e. at the footwall of the
1658 MSt).

1659 2) in the middle, carbonate formations (Jurassic-Oligocene, about 2000 m thick, v_{av} = 5800 m/s) formed
1660 by pelagic limestones (Mirabella et al., 2008) with subordinated marly levels overlying an early Jurassic
1661 carbonate platform (Calcare Massiccio Fm.)

1662 *Rev2 (grouped questions): Identify in Fig.1 “Umbria-Marche thrust and fold belt domain”. You*
1663 *identify it as Laga foredeep domain in Fig1, later as Laga foredeep sequence and here as Laga*
1664 *Formation. Be consistent and use the same terminology along the manuscript and figures.*

1665 Lines 131-132: representing the main and deeper detachment of the region.

1666 Line 133: An underlying basement of variable lithology (V_{av} = 5100 m/s)

1667 Line 135: aforementioned units by the aforementioned important regional decollement.

1668 Lines 136: ... complex ...

1669 *Rev2: represents or is where the detachments are localized?*

1670 *Rev2: Rewrite*

1671 *Rev2: Repetitive and ambiguous. Rewrite.*

1672 *Rev2: I wouldn't say complex, is just a quite simple thrust and fold system, isn't it?*

1673 *Authors: all these corrections have been considered and this paragraph has been totally rewritten.*

1674 Line 137: ... produced NNW-SSE striking WSW-dipping normal faults ...

1675 *Rev2: All the faults are dipping to the WSW? Also that bounds to the west the Norcia basin?*

1676 *Authors: we agree with this comment. Among the steep normal faults, the WSW dipping faults are not*
1677 *the unique characterizing the area, but the ones that generally produce the stronger earthquakes and*

1678 therefore are better known with respect to the ENE dipping faults. However, the antithetic ENE dipping
1679 faults are also important in this structural context, because they seem to be able to produce moderate
1680 earthquakes (as highlighted e.g. by Chiaraluce et al., 2017 in this seismic sequence). The fault that
1681 bounds the west side of the Norcia basin, that we clearly recognize in this work, belongs to the second
1682 type (ENE dip). We have improved the text following the Rev2 suggestion.

1683 Line 139: ... are the Castelluccio di Norcia (CNb) and Norcia (Nb) basins ...

1684 *Rev2: Refer to fig 1*

1685 *Authors: reference to Fig.1 added.*

1686 Lines 140-141: ... They have been subjected to a lacustrine and fluvial sedimentation of hundreds of
1687 meters ...

1688 *Rev2: rewrite*

1689 *Authors: we have rewritten the paragraph.*

1690

1691 Lines 143-149: ... The recent 2016-2017 seismic sequence has been caused by the activation of a
1692 complex NNW-SSE trending fault system, characterized by prevalent high-angle WSW-dipping normal
1693 faults (Lavecchia et al., 2016). More in detail, the easternmost fault system of the region recently
1694 activated is the NNW-SSE trending "Monte Vettore fault system" (Vf). This was the responsible of the
1695 mainshock nucleation between the continental Norcia (Nb) and Castelluccio di Norcia basins (CNb)
1696 (Fig. 1). Nb and CNb are two asymmetrical grabens, bordered by high-angle WSW-dipping normal
1697 faults located on their eastern flanks. Both fault systems are thought to have high seismogenic potential
1698 and able to generate earthquakes up to Mw 7.0 ...

1699 *Rev2 (Grouped questions): Rewrite. You repeat the same idea in different ways. You have defined*
1700 *acronyms in line 139, use them. This must be located after line 142, when you are describing the basins.*

1701 *Authors: corrected and rewritten.*

1702 Line 151: ... The Nb master fault (Nottoria-Preci fault, Nf) ...

1703 *Rev2: Localize in Fig1*

1704 *Authors: corrected*

1705 Line 158: ... Norcia and Castelluccio faults ...

1706 *Rev2: Which fault is this one? Localize in Fig1*

1707 Authors: we refer always to the same fault systems mentioned so far, including the synthetic and
1708 antithetic ones and their secondary splays. Therefore, we have integrated the text and figures.

1709 Line 171: ... whilst explosive was used for NOR02; ...

1710 *Rev2: In Table 1 you mention that CAS01 was acquired with explosives. Which ones is correct?*

1711 Authors: text ok, we have updated the table.

1712 Line 174: ... parameters in Table 1s, supporting information ...

1713 *Rev2: This is in Table 1. There is no table in supporting information*

1714 Authors: corrected, it is in the manuscript.

1715 Line 175: ... Some processing artefacts (A) are visible ...

1716 *Rev2: In the corresponding figures, put the A on top of the line identifying the artefact.*

1717 Authors: There is already in the figure a label A on the top of the artefact (yellow dashed line). We have
1718 improved the figures and text.

1719 Line 176: ... CAS01 (Fig. 1s-a, supporting information) ...

1720 *Rev2: There is no Fig 1s-a in supporting information. This profile is not well identified in on of the*
1721 *figures available in the supporting information section. Also the figures in the supporting information*
1722 *section are not identified. Finally, most of the figures in the supporting information section are*
1723 *repetitions of the figures provided in the manuscript and must be deleted if not used for anything.*

1724 Authors: corrected, the figure is the Fig. 3a. The figures in the supporting material are effectively the
1725 same. But we added here the high-resolution (PDF) version of each figure, because we noticed an
1726 excessive compression and quality reduction of the journal printed-pdf after its creation. So, HR figs
1727 were added only to help the reviewers during the revision. In addition, after this revision, we'll use the
1728 supporting material to add the attributes images (Figs. 2s, 3s, 4s) without any interpretation and line
1729 drawing for comparison with the fig.s 2, 3 and 4.

1730 Line 176: ... some seismic events and lineaments ...

1731 *Rev2: ? Events? Earthquakes? What do you mean by events?*

1732 Authors: "events" removed and replaced in the text as requested.

1733 Lines 177-178: ... seems potentially improvable with a proper choice of seismic attributes type and
1734 parameters ...
1735 *Rev2: ? Rewrite*
1736 *Authors: we have rewritten the sentence as requested.*
1737 Line 180: ... Ithaca database ...
1738 *Rev2: reference*
1739 *Authors: reference added*
1740 Line 183: ...Inside database ...
1741 *Rev2: reference*
1742 *Authors: reference added*
1743 Line 190: ... Over the last years, ...
1744 *Rev2: Maybe explain briefly what is a seismic attribute?*
1745 *Authors: we have integrated the text adding also new references*
1746 Line 196: ... also using composite multi-attribute displays ...
1747 *Rev2: Maybe explain briefly what this means?*
1748 *Authors: ... we have integrated the sentence... as requested*
1749 Line 200: “Energy” (E):
1750 *Rev2: I think that is identified as EN in Figures 2, 3 and 4. Be consistent.*
1751 *Authors: corrected*
1752 Line 207: lateral variations in seismic events,
1753 *Rev2: What do you mean by seismic events? I have done the same question before.*
1754 *Authors: we agree there is confusion with events as earthquake. We have modified the text.*
1755
1756 **5. Results**
1757 Line 228: ... considerable improvements ...
1758 *Rev2: I wouldn't say considerable, at least just looking at figures 2, 3 and 4.*
1759 *Authors: We do not agree. Surely the low-quality images in the revision pdf don't show efficiently the*
1760 *improvements, but in comparison to the standard lines displayed in amplitude, there are many details*

1761 and signal characteristics that are enhanced and that in our opinion improve the data interpretability.
1762 However, we have attempted to improve the figures to show the benefits provided by attributes.

1763 Line 235: ...200 ms in TWT...

1764 *Rev2: Thickness?*

1765 *Authors: the thickness in TWT it's about 200 ms. Considering an average velocity of 6000 m/s for the*
1766 *carbonates, the thickness in meters would be about 600 m.*

1767 Lines 236-239: A similar feature showing such a peculiar signature is visible also in CAS01,
1768 approximately at the same time interval (Fig. 3a, line location reported on the top insert). But in
1769 comparison to NOR01, it appears more discontinuous all along the seismic profile, and in addition it is
1770 partially interfering with suspicious processing artefacts (highlighted with yellow dots, labelled as "A",
1771 slightly undulated in Fig. 3a whilst horizontal in Fig. 2a ca. at 1 s).

1772 *Rev2: I agree, but you should mentioned that it is masked by the artefact (yellow dotted line). In some*
1773 *places seems that it could be directly related to this artefact. I would point that the most clear area is*
1774 *close to the western end of the profile at about 3s TWT. Al the seismic facies in this area are similar to*
1775 *those shown in profile NOR01.*

1776 *Authors: We agree, thank you. We left only the shallower artefact, that is very sharp and clear (see in*
1777 *particular EN and PR attributes) and it seems a copy of the topography.*

1778 Line 241: ... and beneath the southern termination of Nb (ca. between 11-15 km)...

1779 *Rev2: I agree that is clear in the western part of the profile, but not that clear in the eastern. Needs to*
1780 *indicate fault Nb somewhere in the figure, maybe the upper geological map?*

1781 *Authors: Nb is already indicated in Figure above the line in standard amplitude, reported as Norcia*
1782 *basin. We have preferred to use in the figures the entire names of the basin, whilst in the text the*
1783 *acronyms to facilitate the reading. However, we have added Nb on the geological map on the top.*

1784 Lines 242-243: H is better enhanced in fig. 3b by EN attribute (blue arrows), and in particular by the EG
1785 and PR attributes (Figs. 3c and 3d), that considerably help to better detect and mark its extension and
1786 geometry.

1787 *Rev2: To me some of the characteristics that you attribute to horizon H are also related to the observed*
1788 *artefact. If you compare the signal of the upper artefact with the signal of the lower artefact it is not*

1789 *very different. Clearly, in the western end of the profile it is more similar to the results in NOR01, but in*
1790 *the other areas is more arguable. Maybe in the places marked by the blue arrow, but I could point to*
1791 *places related to the upper artefact that have a similar signature (yellow arrows in fig 3 show zones*
1792 *with similar characteristics on the upper artefact than those identified as H in the lower part with a*
1793 *blue arrow). To me it is not clear that you can clearly mark its extension and geometry.*

1794 *Authors: As described in the comment above, we left only the shallower artefact.*

1795 *Line 246: ... Nb ...*

1796 *Rev2: No Nb in the figure*

1797 *Authors: we have added Nb on the geological map on the top.*

1798 *Line 249: this discontinuity propagates down to ca. 2.5 s and intercepts the aforementioned strong*
1799 *reflector H.*

1800 *Rev2: To me that is not evident at all. Below the artefact it is almost impossible to distinguish any west*
1801 *dipping lineament.*

1802 *Authors: We have improved the red arrows to suggest the W-dipping discontinuity visible in fig.2 EG*
1803 *and even better in PR. There is a different reflectors pattern beneath the basin in comparison to the*
1804 *external (east) part and the high-angle discontinuity is in our opinion clearly visible: it propagates down*
1805 *to the depth level in which we find the reflector H. We have made many efforts to improve the figures*
1806 *in the text and also added a new figure with magnifications of the PR attribute (see Fig. 5s).*

1807 *Lines 250-252: other similar but minor discontinuities can be also noticed crossing and slightly*
1808 *disrupting the shallower reflectors: those high angle features are efficiently displayed by the EG and PR*
1809 *attributes (Fig. 2c, 2d), whilst in the original line in Fig. 2a cannot be really appreciated.*

1810 *Rev2: I would not say efficiently displayed. In fact it is difficult to see anything in that zone, even in Fig*
1811 *6a,d,e. I question that you would have interpreted anything there without the knowledge of surface*
1812 *geology.*

1813 *Authors: we have better declared in the text that the surface geology and structural information has been*
1814 *used to “drive” a first phase of interpretation, at least to detect the extension of the basins and the main*
1815 *faults. However, most of the faults interpreted later on the base of the attributes signature (see the new*

1816 image in Fig. s5) don't have evidences at surface, apart a couple of splays detected by the
1817 paleoseismologists close to the Norcia centre (Galli et al, 2005 and 2019).

1818 In our opinion the aNf is clearly visible thanks to the seismic attributes and it has been detected for the
1819 first time in a geophysical data across this basin (it is still debated in literature, as it does not have clear
1820 surface evidences). Then, the seismic attributes enhanced in particular the secondary splays close to the
1821 surface, visible by following the lateral discontinuity of the quaternary deposits at shallow depth. We
1822 hope to have provided here useful elements for better illustrate our interpretation. In addition, we have
1823 revised and improved all the images in the manuscript.

1824

1825 Line 253: ... by similar geophysical features ...

1826 *Rev2: Similar to what?*

1827 *Authors: we have added "to ones detected in NOR02 and CAS01"*

1828 Line 255: ... in Figs. 4b and 4c ...

1829 *Rev2: Do you mean Figs. 4c and 4d?*

1830 *Authors: yes, thank you. We have rewritten the sentence and in general all the paragraph.*

1831 Line 256: W-dipping

1832 *Rev2: Do you mean the west dipping or the east dipping? In the range you indicate there is just the east*
1833 *dipping, the west dipping may begin around km 6 and end at 3-4? I could agree that there is something*
1834 *corresponding to the west dipping lineament but I have more difficulties to interpret the east dipping*
1835 *lineament, mainly in 4c, maybe 4d shows a change in general facies east and west of this lineament. On*
1836 *4c (HR image) seems that the red arrows are pointing to arbitrary places not to places with the same*
1837 *characteristics, and that is confusing.*

1838 *Authors: We have rewritten the text better separating within the description the reflectors and the*
1839 *alignments (discontinuities). We have improved the figures adding more accurately other smaller*
1840 *arrows and red semi-transparent polygons to attract the readers' attention on the main lineaments thus*
1841 *simplifying the text comprehension.*

1842 Lines 259-260: It crosses the entire profile, rising from about 4 s (West) to ca. 2 s (East), where it
1843 intercepts one of the high amplitude events on the eastern end of the seismic line (18-20 km).

1844 *Rev2: I agree that there is a west dipping lineament T that cuts H, but I am not sure it is possible to*
1845 *follow that lineament from km 10-11 to the east as the authors interpret. It would be necessary an un-*
1846 *interpreted section.*

1847 Authors: Thank you. Yes, we agree that T cuts H and this discontinuity is basically not visible in the
1848 original line, but well enhanced for example in PR attribute. In our opinion T can be traced along almost
1849 all the line. To convince the reviewer we have added the figure without green dots in the supplementary
1850 material as requested (see figs. 2s, 3s and 4s).

1851 Lines 261: ... original line ...

1852 *Rev2: Amplitude data?*

1853 Authors: yes thank you, it was a repetition so we have decided to remove the sentence.

1854 Lines 263: ...is a much clear visualization of the reflection patterns...

1855 *Rev2: ... Much clear? Maybe there are some improvements for some horizons (H and T) but I am not*
1856 *sure about the ones pointed with red arrows or just for some of them. ...*

1857 Authors: we clearly agree about the improvements for the horizons like H and T. Also, the overall
1858 pattern of reflectors is much better. To avoid misunderstanding, with the red arrows we don't indicate
1859 reflectors, but only the steep discontinuities of phase/amplitude separating the reflectors, that later we
1860 interpret as attributes evidence of fault zones, however usually simplified in seismic interpretation with
1861 only one red line. We improved the figures to better help the readers in the interpretation.

1862

1863 Lines 264-265: ... a main high-angle E-dipping discontinuity (red arrows) delimits the NOR02 western
1864 sector (ca. 1 km of distance along the line at surface); a ...

1865 *Rev2: Specifically, this is one of the high-angle discontinuities that I am not sure about. I could agree*
1866 *that to the east the seismic facies changes, but I cannot identify a clear lineament in any of the attribute*
1867 *profiles.*

1868 Authors: We think the fact that there is a clear and sharp change of the reflection pattern, as the
1869 reviewer also noticed, is already an indication of a lateral discontinuity (or sets of discontinuities). We
1870 have introduced the concept that we should rethink the concept of single fault planes with distributed
1871 "fault zones", made by many secondary splays and discontinuities at different scales concentrated in a

1872 relatively narrow area. This was one of the reasons on the base of our initial seismic interpretation with
1873 the arrows and without a conventional line drawing. Seismic attributes like the Pseudo-Relief are able to
1874 clearly enhance also small-scale discontinuities providing an outcrop-like seismic line. We think that
1875 regarding this fault, the different reflection patterns as well as the phase discontinuities and truncations
1876 of some reflectors, suggest the presence of distributed fault zones. Again, we remark that to better
1877 support the readers, we have also added an additional image to the Supporting material (Fig. 5s). We
1878 made an additional effort refining the arrows, and introducing a simpler line drawing on the main
1879 visible discontinuities better magnified by this figure.

1880 Line 266: (red arrows)

1881 *Rev2: Why you do not identify the different lineaments with a letter and a number? E.g., L1, L2... That*
1882 *would be more easy to localize them in the figures and to refer to them.*

1883 *Authors: Thank you for the suggestion, but the main faults are already labelled with Nf, aNf, Vf and*
1884 *aVf on the PR figures. Regarding the minor faults, we have preferred to do not add other labels: as*
1885 *remarked also by Rev1 there are already many acronyms and this may make the reading of the*
1886 *document fragmentary. However, we have improved also the text.*

1887 *Lines: smaller discontinuities pervasively cross-cut the set of reflectors between 1-4 km bounded by*
1888 *such two main features, producing a densely fragmented reflectors pattern in the middle portion.*

1889 *Rev2: With the resolution and quality of the data this is very difficult to see. I could point to zones with*
1890 *similar characteristics (just to the east of the profile). Maybe there is some over-interpretation based on*
1891 *surface geology.*

1892 *Authors: As (we hope) better visible in the new image, there are many secondary steep discontinuities,*
1893 *that we have traced giving to the data this peculiar fragmented pattern across the fault zones. We have*
1894 *drawn on the top of the PR images only few faults reported on the geological maps by literature. There*
1895 *are many others interpreted not on the base of the surface geology, but we interpreted just the main*
1896 *secondary ones avoiding possible over-interpretations.*

1897 *Lines 268-269 and 271: ... Another steep E -dipping feature is visible at higher depth (red arrows at 1-3*
1898 *s, ca. 7-9 km) beneath ... to a similar structure displayed in a more central portion of NOR02 ...*

1899 *Rev2: Again this lineament it is not clear to me. In EG I could say that maybe the zone between both*
1900 *lineaments, the one east and the other west dipping, shows a different facies, but there is not a clear*
1901 *lineament. In fact arrows 1 and 2 mark zones with different lineaments to me (see my annotations on the*
1902 *figure and orange dotted lines).*

1903 *Identify the different lineaments with letters in the figures and refer to them in the text. The way you are*
1904 *doing this is confusing. I thought that you were describing the lineament dipping east in the center of*
1905 *the profile. My previous comment was referring to this lineament.*

1906 *Authors: we agree that is less clear than other discontinuities, and in addition, here our interpretation*
1907 *has not been driven by the surface geology, because this discontinuity is quite deep. However, as better*
1908 *shown in the new figure, we cannot avoid to notice the E-dipping lineament highlighted in*
1909 *correspondence to the arrowheads in the area at 2 seconds. About the aVf fault, we have already replied*
1910 *above and suggested to see the new figure 5s in Supplementary. We have reorganized the labels of the*
1911 *main discontinuities in the figures as requested.*

1912 *Line 272: ... here ...*

1913 *Rev2: Here? Where?*

1914 *Authors: we have rewritten the text.*

1915 *Lines 274: characterized by very short and fragmented reflectors bounded by those two steep features of*
1916 *opposite dip.*

1917 *Rev2: As said in one of my comments before when I thought you were describing this area, I have*
1918 *difficulties to interpret both lineaments.*

1919 *Authors: we hope the revised paper has been improved as well as the images easy to interpret.*

1920 *Lines 275-277: ... of such a main W-dipping alignment also seems to truncate and disrupt both the*
1921 *gently-dipping discontinuity T and the deep reflector H: at approximately 3.2 s, it appears interrupted*
1922 *laterally on its western side (Figs. 4c and 4d) ...*

1923 *Rev2: I cannot interpret this lineament so far, but according to your interpretation this lineament*
1924 *dipping to the west seems to die on horizon T (Fig4c). As you have mentioned before, seems that is*
1925 *horizon T that truncates horizon H. Profiles does not have the lateral scale to allow the easily*

1926 *identification of the place the authors are mentioning. In fact I am not sure what reflector/discontinuity*
1927 *they are referring here.*

1928 Authors: The lateral scale is reported on the top profile of Fig.s 2a, 3a, 4a) and now we have modified
1929 also adding the scale to the bottom of each one. We also suggest that there is a grid of thin black lines in
1930 all the images, vertical for the distance (intervals of 5 km) and horizontal for the travel time (every 1
1931 second).

1932 Regarding the deep area between 5-10 km, it is much complex in the data. It seems that T (low angle)
1933 intercept and interrupt H, but we prefer to present and discuss possible interpretations on which tectonic
1934 structure cuts the horizon H. As also suggested by the Reviewer 1, we discuss the relationships with the
1935 Acquasanta thrust (low-angle discontinuity T) is more ambiguous. We propose two alternative
1936 interpretations can be proposed, schematically represented in Fig. 9: a model in which Vf merges into
1937 the deep Acquasanta thrust (T), suggesting a negative inversion, and another in Vf cuts and displaces
1938 the Acquasanta thrust, following a steeper trajectory (ramp). In both cases the H horizon is truncated,
1939 but the relations between the Vettore fault, Vf and the Acquasanta thrust, T, are different.

1940 Lines 279-280: displays clarified the deep geometries of the main reflectors and of the geophysical
1941 discontinuities, later

1942 *Rev2: Some of the reflectors/discontinuities may have been highlighted by the attribute processing (H*
1943 *and T) but in general I have some problems to identify the lineaments interpreted by the authors.*

1944 Authors: we are glad that some discontinuities have been detected by the reviewer. Regarding the
1945 others, we hope our revision have improved the figures, allowing a better visualization of the reflectors
1946 mentioned in the text.

1947 Lines 282-283: overlapped using ODT software (depth conversion with $VP_{av} = 6000$ m/s, vertical
1948 scale 2x).

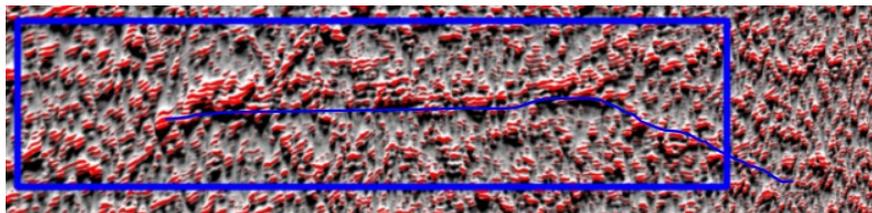
1949 *Rev2: This must go on the figure caption*

1950 Authors: moved to caption, as requested.

1951 Lines 285-286: The blue box of Fig.5a is reported in Fig. 5b and 5c

1952 *Rev2: ...corresponds... ...to figures ... To me the reflector corresponding to H is very clear in the*
1953 *original seismic image 5b, maybe more clear than in 5c. Then, I am not sure what the attribute analysis*
1954 *is providing. ...*

1955 Authors: H is relatively clear in the original line only for the short portion in which it shows higher
1956 amplitude. Looking more globally the lines, the attributes are not only giving a peculiar signature
1957 recognizable also on other lines (like CAS01 and NOR1) contributing to give a better idea of its
1958 regional extension, but also its lateral extend in NOR02 is better appreciable enhancing its continuation
1959 to the east with a gentle W-dip in comparison to the original line (see figure below).



1960
1961 Lines 288-289: The Fig.5e displays the enhancement obtained plotting the PR attribute (“similarity
1962 palette”) in transparency on the seismic line in amplitude (SA).

1963 *Rev2: I am not an expert in interpreting this type of datasets, but I am having difficulties to see any*
1964 *enhancement in the data in 5e. I cannot see any lineament.*

1965 Authors: we have improved the figure. The dense steep lineaments (discontinuities) highlighted in 5e
1966 produce peculiar seismic facies, that in our opinion can be used to interpret the area as a fault zone in
1967 which there is a strongly deformed associated with a main fault.

1968
1969 Lines 291-293: ... The comparison between the multi-display of attributes PR and EG (blue box in Fig.
1970 6a), the original line (Fig. 6b) and the EN+PR plot (Fig.6c) shows the improved signature of the strong
1971 reflector H. The black box again reports the original line NOR01 and the version PR+SA, clearly
1972 ...boosting the visualization of the high-angle discontinuities. 293

1973 *Rev2: ... I think the attributes maybe highlight the reflector H but I also think that in the original dataset*
1974 *is also quite clear, so talking about improving...*

1975 Authors: we have partially already replied above. We think that the improved images are often self-
1976 explicating, improving the interpretability of the data. The alternative way is to not use these seismic

1977 data. Clearly, the outcomes here provided may be not dramatic as usually happens in modern high-
1978 resolution 3D survey, but any improvements, even if only on some reflectors or on limited area, are
1979 welcome considering the uniqueness of these seismic lines and the importance of the study area.

1980 Lines 294-295: ... those are connected with the surface geology and related to the hypocentre location of
1981 the main seismic events, that will be discussed more in detail ...

1982 *Rev2: ... Not sure at all about this ... In fact, it seems that the authors have been interpreting high-angle*
1983 *lineaments based on surface geology as I have commented before. Some of this interpreted lineaments*
1984 *are quite questionable, at least I have difficulties to interpret them, but, as mentioned before, I am not*
1985 *an expert in this kind of interpretations.*

1986 Authors: as already replied in other comments, we admit that our interpretation was driven by, but not
1987 limited to the surface geology. As already pointed out in comments above, some faults have been
1988 interpreted in this way, but then for many others, we used typical elements of a standard seismic
1989 interpretation (phase discontinuities, lateral variation in amplitude, offset between reflectors etc..).

1990 Lines 306-307:

1991 *Rev2: This must be in the figure caption.*

1992 Authors: Thank you, part of this sentence has been moved to the figure caption.

1993 Lines 312-322: The steep discontinuities highlighted by the attribute analysis are here interpreted as the
1994 seismic signature at depth of complex normal faults mapped at the surface.

1995 *Rev2: As said in various comments I have problems to interpret these steep discontinuities. All this*
1996 *discussion is based on the authors interpretation and since I cannot interpret the same things I cannot*
1997 *support it. But again, I am not a specialist in this type of seismic interpretations.*

1998 Authors: we are confident that after the revision and integrations provided, the interpretability of the
1999 seismic features will be now more clear for the readers.

2000 Lines 326-329: belonging to a conjugate tectonic system (Brozzetti & Lavecchia, 1994; Lavecchia et al.,
2001 1994) and suggested by morphological evidences (Blumetti et al., 1990) and paleoseismological records
2002 (Borre et al., 2003). It is a synthetic (W -dipping) high-angle, normal fault bordering the eastern flank of
2003 Nb (“Nottoria-Preci fault” – Nf, Calamita et al., 1982; Blumetti et al., 1993; Calamita & Pizzi, 1994).

2004 *Rev2: This is referred to the interpreted discontinuity/fault or to the Nf? Which one? Ok to Nf but the*
2005 *previous sentence it is not clear, so this "It is" is also no clear what you are referring to. Rewrite both*
2006 *sentences.*

2007 Authors: yes, it is referred to the E-dipping (antithetic Norcia Fault -aNf-), currently still debated in
2008 literature because not clearly visible in outcrops and only inferred, before this study, by
2009 geomorphological evidences and by paleoseismological studies (e.g. Galli et al, 2018; Borre et al.,
2010 2003) .We have rewritten the text.

2011 Line 330 and 332: ... red arrows, Figs. 2c, 2 d ... and red arrows between 7-9 km, ca. 1-3 s

2012 *Rev2 (grouped comments): Do you mean 4c and 4d? Red arrows, which ones? There are a lot of red*
2013 *arrows. Again, it is necessary to identify the lineaments by names in the figures and in the text. See my*
2014 *previous comments about lineaments identification.*

2015 Authors: yes, thank you for this correction. We have updated the text and the figures and already replied
2016 in previous comments.

2017 Lines 345: and the thrust (T) at about 3.2 s.

2018 *Rev2: Seems strange that the normal fault is cutting the thrust plane. Usually in inversion tectonics, the*
2019 *"new" faults use the slip planes of the previous faults, since it requires less effort to slide along a*
2020 *preexisting plane than to generate a new one. In that case, it seems more plausible that the normal*
2021 *faults would be using the thrust detachments at depth as fault planes and not rupturing them and*
2022 *generating new ones.*

2023 Authors: there are currently different interpretations and models available in literature. Our data do not
2024 allow to clarify this point in detail, so, following also the suggestions of the Rev. 1, we have provided
2025 two possible interpretations on the relations between the thrust and normal fault. In one case the normal
2026 fault cuts the thrust and another case characterized by negative inversion of the pre-existing thrust. We
2027 have compared and discussed these two models in the Discussion (chapter 6) and produced a new figure
2028 (Fig. 8)

2029 Line 363: high-resolution

2030 *Rev2:?*

2031 Authors: high-resolution images. Corrected.

2032 Lines: However, the attributes aid the seismic interpretation to better display the reflection patterns of
2033 interest and provided new and original details on complex tectonic region in Central Italy.

2034 *Rev2: Arguable*

2035

2036

2037

2038

2039 FIGURE 1.

2040 REV2:

2041 For each earthquake indicate the date in which occurred and the depth. Could be also possible to plot
2042 seismicity? Above 3.0 or 3.5, to show where the earthquakes are localized. I would suggest to plot more
2043 clearly the surface rupture traces on the map.

2044 *Authors:*

2045 *We have added and updated all the information in the Fig.1 as requested.*

2046 FIGURE 2.

2047 REV2:

2048 Consider to put the A on top of the line.

2049 *Authors:*

2050 *We have moved A to the left, in a place where it doesn't obscure reflections.*

2051 REV2:

2052 Identify with a name each possible lineament (L1, L2,...). The same lineament in two different profiles
2053 could have the same name (NOR01 and NOR02).

2054 *Authors:*

2055 *Thank you, we have updated and enhanced the labels for the main faults (aNf, aVf, Nf, Vf) using a*
2056 *continuous line for each one. We didn't add more labels on the interpreted secondary splays but we*
2057 *have added a new figure (fig. 5s) as supplementary material to provide further details on the shallow*
2058 *part of NOR01 and NOR02.*

2059 FIGURE 3.

2060 REV2:

2061 Consider to put the A on top of the line.

2062 *Authors:*

2063 We have moved A to the left in a place where it doesn't obscure reflections, we think it's preferable to
2064 maintain the label close to the yellow dots to aid the readers.

2065

2066 FIGURE 4.
2067 REV2:
2068 Identify with a name each possible lineament (L1, L2,...). The same lineament in two different profiles
2069 could have the same name (NOR01 and NOR02).
2070 *Authors:*
2071 *Please, see our replies in previous comments.*
2072 FIGURE CAPTIONS:
2073 Line 681: Figure 2
2074 *Rev2: For the different figures containing seismic profiles I suggest to explain at the end what means*
2075 *each arrow, dotted line,... Then you avoid repetition or not mentioning in one of the subfigures, as for*
2076 *example not mentioning the red arrows in 2c and mentioning in 2d.*
2077 *Authors: fixed following the comment and almost all the captions have been considerably rewritten.*
2078 Line 685: ..., with same attributes computation ...
2079 *Rev2: Same as what? A figure caption has to be self explained.*
2080 *Authors: deleted*