Federico Rossetti - Handling Topical Editor

Solid Earth

June 5, 2019

Subject: Response to referee comments – manuscript "Stress Characterization and Temporal Evolution of Borehole Failure at the Rittershoffen Geothermal Project" by Jérôme Azzola et al., se-2019-72

Dear Federico Rossetti,

We thank you and the referee for the extended reviews of our manuscript. We carefully responded to all the comments and revised the manuscript accordingly. Please find below a point by point response to the major and minor comments of the two reviewers (in black the comments and in blue, our response), followed by a copy of the manuscript with tracked changes. The revised manuscript as well as a copy where the changes made to the original version are highlighted, are attached.

Sincerely, on behalf of the authors
Jérôme AZZOLA

Referee n°1

This paper addresses the important issue of evaluating a regional stress field from images of two different failure processes (borehole breakouts and so-called drilling Induced fractures) observed in deep boreholes with different orientations, as well as from results from various water injection tests. The methodology is applied at the Rittershoffen site, located 6km east from the Soultz site, where the stress field is quite well known. This is an important contribution for the understanding of stress field in deep rock masses and the quality of images as well as that of their analysis justify completely its publication.

We thank François Cornet for his careful review. We appreciate his recognition of the importance of our contribution.

1. The GRT-2 borehole is inclined 37° to the vertical so that the axial and tangential stress components at the borehole wall are not principal stresses. Authors must write down the equations they are considering, including the role of pore pressure, and that of thermal stresses. Indeed, the principal directions, at the wellbore wall, of stresses resulting from the far field stresses are not the same as those of the thermal stresses resulting from the cooling of the rock. This issue is completely ignored, and the paper cannot be published before this is properly dealt with. I encourage authors to look at paper by Wileveau et al. that provides good illustrations of en echelon breakouts observed in inclined wells. (Wileveau Y, F.H. Cornet, J. Desroches and P. Blumling, 2007; Complete in situ stress determination in an argillite sedimentary formation; Physics and Chemistry of the Earth (vol. 32, pp 866-878)

The GRT-2 well is strongly inclined: the mean deviation in the section of interest is 37° (as presented in section 2). Equations that describe stress concentration at the borehole wall of a vertical borehole, used for the well GRT-1, are no longer applicable in this case. For the deviated well GRT-2, we used a solution that takes into account the inclined geometry of the borehole (as introduced in section 4.1). The solution is based on equations in which are involved the non-vertical geometry of the well, the orientation of the far field stresses, the thermal stresses and the fluid pressure. We refer to the review of Schmitt et al. (2012) who propose a complete development of the equations, in the general case.

As suggested by the referee, we detailed the used equations and included the computation steps leading to the expression of the effective principal stresses at the borehole wall of the deviated well in the revised version of the manuscript in appendix A. We cite the work of Wileveau et al., as an additional reference to this approach.

2. For their analysis of the width of borehole breakouts, authors refer to three different failure criteria, including the Hoek and Brown criterion. For the parameters to be considered in these criteria, they refer to laboratory work quoted by Rummel, 1991 and by Valley and Evans, 2006. They should also look at the publication by Villeneuve et al. (Villeneuve M.C., M.J. Heap, A.R.L. Kushnir, T. Qin, P. Baud, G. Zhou, and T. Xu, 2018; Estimating in situ rock mass strength and elastic modulus of granite from the Soultz-sous-Forêts geothermal reservoir (France); Geothermal Energy, 6(11), https://doi.org/10.1186/s40517-018-0096-1), which address precisely this issue.

We used all available published data to parametrize our failure criteria, including data provided in Villeneuve et al., 2018, but also from Heap et al. (2019) (which was already cited on lines 90 and 511 of our original manuscript). To clarify this, we added the relevant references in section 5 of our revised manuscript.

3. In their table 3 the density value for the granite is said to be 2570 kg/m3, yet in equation (6) the vertical stress is assumed to be equal to 0.024 z-0.83. These differences should be discussed. In addition, given the vertical stress magnitude is taken into consideration in the three-dimensional failure criteria, authors should show how they determine uncertainties on the vertical stress component evaluation.

The magnitude of the vertical stress *Sv* is obtained from the weight of the overburden, by integrating the density profile from surface to reservoir depth. We apologize for the typo in Eq. (6), which should read *0.0248 z – 0.83*. This misleading rounding will be corrected in the revised manuscript, which leads to a trend in line with density value 2570 kg/m^3 chosen for the granitic zone. Given the fact that the vertical stress is obtained by integrating the density profile from surface to reservoir depth, the uncertainty on density add up and thus the uncertainty on the vertical stress estimation increase with depth. Considering an uncertainty of 50 kg/m^3 on the densities leads to a 2.5 MPa uncertainty on the vertical stress at reservoir depth. This uncertainty is not significant compared to other uncertainties involved in the analysis as for example those related to the mechanical parameters chosen in the inversion of the maximum horizontal stress.

Details about the uncertainty estimation are added in the revised manuscript, in section 8.1.

4. Similarly, equations used for the evaluation of the minimum principal stress magnitude is not described and this should be corrected. Evaluation of associated uncertainty should be discussed.

We follow approaches used in the literature (e.g. Cornet et al., 2007) and estimate the minimum horizontal stress *Sh* from pressure limiting behavior during hydraulic injections. Since we did not have enough information related to the Rittershoffen site to compute a complete *Sh* stress profile, we use measurements carried out at the nearby Soultz-sous-Forêts site. The trend is evaluated following Cornet et al. (2007). In their publication, the uncertainty about the trend is largely discussed but not quantified. Deriving confidence bound through data fits or propagating pressure measurement errors leads uncertainty on *Sh* magnitude of a few megapascals which is irrelevantly low compared to possible misinterpretation of the pressure limiting controlling factors. It is even more insignificant compared to the uncertainty related to the parameterization of the failure criteria that dominates the uncertainty on *SH* magnitude. We rather complete pragmatically our study by discussing the applicability of the trend to the Rittershoffen site. For this purpose, we analyze the wellhead pressure measured during the hydraulic stimulation of GRT-1 and derived an estimate of *Sh* at depth from the pressure reached at a maximum flow rate. However, as the pressure shows a gradual but not definitive stabilization for these maximum flow rates, our measurement is discussed as a lower bound of the minimum horizontal stress *Sh* at depth. We show that our measurement is still consistent with the trend measured at Soultz.

As the second referee also asked for more information on the methodology that we follow, we added details in the manuscript to clarify the steps followed in the estimation of the *Sh* profile.

5. Table 2 indicates values for the Poisson's ratio, but no reference is made to Young's moduli nor to thermal expansion coefficients used in equation

We apologize for not having included the Young's moduli and the thermal expansion coefficient values in the manuscript. They have been added to the revised version. The volumetric thermal expansion coefficient is chosen to be constant for the different layers of our model, $\alpha = 15 \times 10^{-6}\ K^{-1}$, and the Young's moduli are added in Table 2.

6. In equation (2) the stress component τ_{oct} implies the three principal stress components. This should also apply to the mean stress, as opposed to equation written on line 179.

The original derivation of the equations is proposed by Zimmerman & Al-Ajmi (2006). In their review, authors refer to an "effective mean stress", $\sigma_{m,2} = \frac{\sigma_1+\sigma_3}{2}$, for the Mogi-Coulomb criterion. This is not strictly speaking the mean stress, which would also include a contribution of the intermediate stress σ_2. We clarified the terminology and nomenclature in the revised version of the manuscript.

7. In their discussion of results, authors argue that some of the results obtained for the magnitude of the maximum principal stress magnitude do not satisfy the Coulomb stability condition for the rock mass. Interestingly, Cornet (2016) has argued that the large-scale fluid injections conducted at Soultz have generated large scale failure zones

that are changing in orientation with depth, a feature consistent with the Hoek and Brown criterion but not with a Coulomb criterion. This issue should be discussed more carefully (Cornet, F.H., 2016. Seismic and aseismic motions generated by fluid injections; Geomech. Ener. Env., 5, pp 42-54).

We thank the referee for his comment and suggestion to discuss our results at large scale. It addresses an important issue: the upscaling from the borehole scale (centimeter scale) to the reservoir scale (kilometric scale). Our work is based on the stability of the wellbore wall, i.e. processes that are occurring at a centimetric scale. In our study, we use stability criteria and parameters that have been obtained from laboratory experiments at a similar scale.

On the contrary, Cornet (2016) discusses the strength criterion at the scale of rock mass. To discuss our results in terms of stress profiles at larger scale, we can analyze Fig. 11 which shows the evolution with depth of *SH* over 650m. We see a significant trend of a rotation with depth of the *SH* direction which could be related to the observation of Cornet (2016) and support a Hoek-Brown criterion for failure rather than the Mogi-Coulomb criterion. As discussed in section 9.3, we believe that this trend is rather related to the distance to the fault than to the depth and the effect of the fluid pressure on the Hoek and Brown criterion. From the stability criterion computation, we point out that both the Hoek-Brown and Mogi-Coulomb criteria exceed the frictional limit and cannot be used to choose the most relevant criterion. We should also point out that the injection in Rittershoffen was not as "massive" as in Soultz, i.e. the injected volumes and applied well head pressures were much smaller. The transposition of the knowledge from Soultz in this regard may not be directly applicable.

As requested by the reviewer we developed the discussion on this topic in sections 9.2 and 9.3.

8. How are the various parameters measured? How valid are those measurements for in-situ properties? This should be better discussed.

Please note that we are very cautious in describing the criteria and mechanical parameters chosen in the approach. We recognize that the strength parametrization is the main limitation of our approach. We bring this point carefully in the discussion. Given that we do not have "access to direct strength measurements since no cores were collected" (line 582), our results are discussed in the light of the uncertainty on the strength parameters, as stated in section 9.4. In addition to the lack of information to parametrize our criteria, we recognize two other sources of uncertainty:

1) there is no consensus regarding the appropriate failure criterion to asses wellbore wall strength and to be used for borehole breakouts analysis, as mentioned in the manuscript in lines 165-168 (of the original manuscript). In our approach, we used thus multiple criteria and discussed the relevance of the measurements in terms of stress profiles by confronting them to the stability of the rock mass at larger scale.

2) the mechanical and strength parameters that have been selected from core or cutting analyses are not necessarily representative of the *in-situ* conditions.

We added new elements of the discussion into the revised manuscript in section 5 and in section 9.3.

9. Caption of figure 12 has been exchanged with that of fig 13.

We thank the reviewer for pointing out the caption swap between fig. 12 and 13. We fixed this issue in the revised manuscript.

Referee n°2

The authors present a detailed study in orientation and magnitude of the local stress field at the geothermal site of Rittershoffen in France, near the well-known site Soultz-sous-Forêts. The manuscript focuses on the temporal evolution of borehole breakouts and drilling induced tension fractures using acoustic images of two boreholes acquired by Ultrasonic Borehole Imager in 2012, 2013 and 2105. The manuscript is interesting and provides an important contribution for the understanding of the time-dependent deformation.
In this form the manuscript is not ready for publication. Please see my comments.

We thank the reviewer for his careful review. We appreciate his interest in the manuscript and his recognition of the importance of our contribution.

Major comments:

1. The author mentioned in the abstract that they used for their investigation image datasets from two boreholes GRT-1 and GRT-2. In the manuscript the analysis as well as the description and the discussion of the results are mainly focused only on GRT-1. I suggest the authors to show only the analysis on GRT-1 well. In case the author wants to continue keep also the GRT-2, a detailed analysis of the datasets of this borehole is requested. The analysis must be related to the inclined borehole taking into account the orientation of the principal stresses in an inclined borehole.

We acknowledge that the description and the discussion of the results are mainly focused on the data of the GRT-1 well, as the quality of data from GRT-2 is generally lower than for GRT-1 (line 235). The image quality problems with GRT-2 are detailed in section 6.1 of the manuscript and illustrated in figure 3.c. It shows in particular the significant stick-slip effect inducing alternative compression and stretching of the UBI images. Figure 3.d. is an example of an erroneous borehole radius record. Given the extent of the artefacts highlighted in GRT-2, the measurements of the breakout parameters in this borehole are more uncertain than in GRT-1 and no DIFTS have been measured in GRT-2 (line 325). We still analysed the stress tensor in GRT-2 using a proper deviated well approach, which has been clarified in appendix A of the revised manuscript. We feel that it is worth adding the GRT-2 data in the manuscript as the expression of the measurements in TVD enables to compare the results with the measurements performed in GRT-1, even if the data quality doesn't enable to propose an extended analysis of the stress tensor as in GRT-1.

2. The authors show in Figure 15 the magnitude of Sv, Sh and SH from 2000 m to 2500 m of GRT-1. To calculate the Sv magnitude the authors used equation (6). The Sv curve is presented as if it were made using a fixed value of 2440 kg/m3 for the entire well. Can you explain why? At 2300 m using equation (6) as the author wrote the value Sv is 54.37 MPa but using the value of 2570 kg/m3 (Table 3) corresponding to the granite rock at a depth of 2200 m Sv is 58.28 MPa.

The magnitude of the vertical stress *Sv* is obtained from the weight of the overburden. The density profile provided in Table 3, is integrated from surface to maximum depth. The trend provided in equation (6) is obtained from a linear fit to the measurements in the range of depths considered in our study. We apologize for the typo in Eq. (6), which should read *0.0248 z – 0.83.*

This misleading rounding is corrected in the revised manuscript, which leads to a trend in line with density value 2570 kg/m3 chosen for the granitic layer.

I suggest redrawing figure 15 showing the entire section of the GRT-1 between 0 and 2562 m (TVD).

We acknowledge and tested this advice, but by redrawing figure 15 from 0 to 2562 m, we considerably deteriorate the readability of the measurements presented in the figure 15 for greater depths, from 1950 to 2550m, while showing a long wellbore section (from 0 to 1950 m) without data.

After careful consideration, we decided thus not to extend the vertical scale to the entire GRT-1 section.

Furthermore, in line 387 the authors should specify that the density value shown in equation (6) is related to the Jurassic rocks between 1172 and 1447 m of GRT-1 as an example, but that the Sv was calculated taking into account the density values of the different rocks at different depths. No Figure for GRT-2. If the authors want to include this well, they have to show the data and results.

We added details about the procedure followed (after line 394). We measure *Sv* as a function of TVD in order to apply the same measurements in both wells. Figure 17 and 18 are showing results for GRT-2.

Minor comments:

1. I suggest that the figures and tables have the same MD or TVD depths, or that both are reported. For example, Table 3 shows lithologies and densities relative to TVD depths, while if I look at the stratigraphy in Figure 8, the lithologies refer to MD depths.

TVD is the most relevant depth scale to present stress estimate and to compare results across both wells, with GRT-1 almost vertical (TVD and MD are not very different) and GRT-2 being deviated.

We follow thus the advice of the reviewer and made sure to add a TVD depth scale on all our figures and tables.

2. Please include also the fractures distribution (number, dip, dip azimuth) highlighting the main faults or fracture zone to better understand the borehole breakout rotation and/or deviation from the mean of *Sh*.

The major fracture network was observed from acoustic wall imagery in the open-hole sections of GRT-1 and GRT-2 by Vidal (2017). Major continuous fractures (thickness measured on acoustic images higher than 1 cm) are analyzed in both wells. The detailed structural survey is available in Appendix 2 of Vidal's thesis. The fractures are oriented globally in GRT-1 N 15° E to N 20° E with a dip of 80° W. In GRT-2, the main fracture family is oriented N 155° E to N 175° E with a dip of 80° E to 90° E. Fracture density is highest on the roof of the granitic basement. These summary elements are added in section 2 of the revised manuscript, which details the context of the Rittershoffen project. Our analysis doesn't consist in the measurement / discussion of the distribution and orientation of the natural fractures highlighted through the GRT-1 and GRT-2 wells, which has been extensively studied by Jeanne Vidal in her thesis.

We believe thus that adding data regarding the fracture distribution and orientation in the figures doesn't contribute to the discussion of the proposed measurements but would necessitate to analyse data that are not in the focus of our paper.

3. The value from hydraulic test at GRT-1 differs from the data from the boreholes GPK1. Could you explain better the reason? Please add also this Sh- value from GRT- 1 in Figure 15

The approach followed has certainly not been sufficiently clearly explained, and we added details in the section 8.3 to carefully describe the steps in the estimation of the *Sh* profile proposed in the GRT-1 and GRT-2 wells.

The profile of the minimum horizontal stress *Sh* is estimated from pressure limiting behavior during hydraulic injections. Since we did not have enough information related to the Rittershoffen project to compute a complete *Sh* stress profile, we used measurements carried out at the nearby Soultz-sous-Forêts project. The trend is evaluated after Cornet et al. (2007). To complete our analysis, we analyzed the wellhead pressure measured during the hydraulic stimulation of GRT-1 and derived an "estimate at best" of the *Sh* magnitude at depth from the pressure reached at maximum flow rate. The wellhead pressure measured at 1913m in GRT-1 during the hydraulic stimulation (data provided in figure fig. 12) shows a gradual but not definitive stabilization at flow rates up to 80 $L.s^{-1}$. Even if the pressure limiting behavior, related to the creation or reactivation of faults, is not reached, we discuss the measurement as a lower bound for the minimum horizontal stress *Sh* at 1913m. By comparing our measurement in Rittershoffen at 1913m with the trend considered in the stress analysis and measured originally in Soultz-sous-Forêts, we show that both measurements are consistent and that the Rittershoffen measurement is indeed a realistic lower bound for the chosen trend.

4. The caption of figure 13 refers to figure 12. Whereas the caption of figure 12 refers to figure 13. Please modify.

The caption of figure 13 has been inverted with caption of figure 12, as mentioned by both referees. We fixed this issue in the revised manuscript.

5. Line 16 GRT-2 instead of GRT2
6. Line 16 2500 m instead of 2500m
7. Line 40 provide an indirect information instead of provide a indirect information
8. Line 90 WSM released in 2016 no in 2008. Please update the reference and cite as: Heidbach, Oliver; Rajabi, Mojtaba; Reiter, Karsten; Ziegler, Moritz; WSM Team (2016): World Stress MapDatabase Release 2016. GFZ Data Services. http://doi.org/10.5880/WSM.2016.001 (http://dataservices.gfz-potsdam.de/wsm/showshort.php?id=escidoc:1680890)

9. Line 120 GRT-1 instead of GRT 1

We corrected the typographical errors referenced in comment #5, #6, #7 and #9 in the revised manuscript. We updated the reference to the WSM (comment #8).

10. Lines 142-143: please specify which failure condition

Details have been added to the manuscript regarding the failure criterion used to the above-mentioned lines.

11. Lines 307-309 Please insert one or more figures to confirm what has been said.

The request of the referee is not very clear, as the lines referred to do not highlight an obvious lack in information. If the referee refers to the deviation of the wells in the open-hole section, this is shown in Figure 1: the trajectories of GRT-1 and GRT-2 show that the deviations are constant in the section of interest and that GRT-1 is quasi-vertical.

12. Line 183: why the authors grouped the Triassic sandstone in a single category? Please add in the manuscript the reason: no alteration, homogeneous lithology, no fractures, etc

The sandstones crossed by the open section of the well are all from the Buntsandstein (section 5 of the manuscript). Heap et al, (2019) studied in detail the strength evolution with depth of the Buntsandstein mechanical properties. As suggested by the referee, they evidenced significant variations of the compressive strength together with elastic modulus changes. They also pointed out the role of the fluid content on the UCS. However, these variations are limited compared to the statistical fluctuations of our measurement. Accordingly, we gathered the Buntsandstein sandstones as a single unit (after line 203). We used typical strength parameters from Hoek and Brown (1997) to characterize the geological unit.

13. Line 387 Sv [MPa] = 0,024 * z [m] – 0,83 or Sv [MPa] = 0.024 * z [m] – 0.83 but no one value as dot and the other a comma. In order not to confuse the reader, I suggest using the asterisk (or an x) as a multiplication sign instead of the point.

To avoid any confusion for the reader about the punctuation used in the equations, we followed the referee suggestions and replaced the dot by an asterisk in the equations proposed in the manuscript.

14. Line 533 please 50 m instead of 50m
15. Line 573 please add a dot after correlation technique
16. Line 579 please add the year of the reservoir stimulation

We modified typographical errors previously referenced (comment #14 and #15) and added details to the manuscript regarding the year of the stimulation to the above-mentioned lines (comment #16).

17. Figure 1: legend: the reference is WSM 2016 not 2006 Helmholtz-Centre Potsdam GFZ. Inset with the sketch of GRT-1 and GRT-2 boreholes: the lithology is not clear, some writings overlap. It would be good if the stratigraphy had the same colours as the geological profile. Highlight the trajectory of the wells on the geological profile. Caption: Heidbach et al., 2016. Cite as: Heidbach, Oliver; Rajabi, Mojtaba; Reiter, Karsten; Ziegler, Moritz; WSM Team (2016): World Stress Map Database Release 2016. GFZ Data Services.

The legend has been updated as well as the caption with the reference proposed by the referee. The writings in the lithology and in the legend have been made clearer. The lithological profile has been set in agreement with the stratigraphy (bottom and left inserts). The geological profile includes the trajectory of the wells even if its scale doesn't enable to distinguish the direction of GRT-1 and GRT-2.

18. Figure 2: Please add two separated scales for radius (mm) and for width (_)

Figure 2 has been updated accordingly to the referee's suggestions.

19. Figure 3: show directly in the figure a, b, c, d, the artefacts (signal loss, stick slip).

In order not to load unnecessarily the figure, we added details in the figure caption regarding to the artefacts. We feel that the mentioned artefacts are now easily recognizable in the images of Fig. 3.

20. Figure 14: please add the fractures as Tadpole related to this section.

In our analysis, we didn't study the distribution and orientation of the fractures, which has been done by Jeanne Vidal in her thesis (Vidal, 2017). We believe that adding data regarding the fracture distribution and orientation in our figures doesn’t contribute to the discussion of the proposed measurements but would necessitate to analyse data that are not in the focus of our paper.

21. Figure 15: Please remove the lithology from inside the figure but add it as litho column to the side of the figure. Please add the fractures as Tadpole related to this section. Is the deviation of the stress values between 2250 and 2380 m, more or less, due to the presence of fractures?

We removed the lithology from the inside of the figure and added it to the side of the figure. The deviation of the stress values is correlated to the increase in the breakout width at the mentioned depths.

22. Figure 18: Please remove the lithology from inside the figure but add it as litho column to the side of the figure. The symbols of Sh and Sv of GRT-2 are not very clear in the figure. Please change the symbol.

We removed the lithology from the inside of the figure and added it to the side of the figure. We modified the symbols related to the stress state estimates in GRT-2 to improve readability.

Stress Characterization and Temporal Evolution of Borehole Failure at the Rittershoffen Geothermal Project

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Abstract. In the Upper Rhine Graben, several innovative projects based on the Enhanced Geothermal System (EGS) technology exploit local deep fractured geothermal reservoirs. The principle underlying this technology consists of increasing the hydraulic performances of the natural fractures using different stimulation methods in order to circulate the natural brine with commercially flow rates. For this purpose, the knowledge of the *in-situ* stress state is of central importance to predict the response of the rock mass to the different stimulation programs. Here, we propose a characterization of the *in-situ* stress state from the analysis of Ultrasonic Borehole Imager (UBI) data acquired at different key moments of the reservoir development using a specific image correlation technique. This unique dataset has been obtained from the open hole sections of the two deep wells (GRT-1 and ~~GRT2, ~2500m~~GRT-2, ~2500 m) at the geothermal site of Rittershoffen, France. We based our analysis on the geometry of breakouts and of drilling induced tension fractures (DITF). A transitional stress regime between strike-slip and normal faulting consistently with the neighbour site of Soultz-sous-Forêts is evidenced. The time lapse dataset enables to analyse both in time and space the evolution of the structures over two years after drilling. The image correlation approach developed for time lapse UBI images shows that breakouts extend along the borehole with time, widen (i.e. angular opening between the edges of the breakouts) but do not deepen (i.e. increase of the maximal radius of the breakouts). The breakout widening is explained by wellbore thermal equilibration. A significant stress rotation at depth is evidenced. It is shown to be controlled by a major fault zone and not by the sediment-basement interface. Our analysis does not reveal any significant change in the stress magnitude in the reservoir.

1 Introduction

Several deep geothermal projects located in the Upper Rhine Graben and based on the Enhanced Geothermal System (EGS) technology exploit local geothermal reservoirs, such as those located in Soultz-sous-Forêts or in Rittershoffen (Baujard et al., 2017; Genter et al., 2010). The principle underlying this technology consists of increasing the hydraulic performance of the reservoir through different types of simulations to achieve commercially interesting flow rates. The stimulation techniques are typically based on high pressure injection (hydraulic stimulation), cold water injection (thermal stimulation) or chemical injection (chemical stimulation). During the injections, a thermo-hydro-chemo-mechanical perturbation induces an increase in permeability due to the reactivation of existing structures or the generation of new ones (Cornet, 2015; Huenges & Ledru, 2011). The *in-situ* stress state is a key parameter controlling rock mass response during stimulation and is required to design stimulation strategies and forecast the response of the reservoir to varying injection schemes.

Despite its importance, the *in-situ* stress state is difficult to assess, particularly in situations where the rock mass is only accessible through a few deep boreholes. In such cases, the assessment of borehole walls using borehole logging imaging is a useful technique to provide information on the type, the orientation and the size of fractures or breakouts which are owed to the stress perturbations related to existence of the well (drilling and fluid boundary conditions). Subsequently, it gives useful constraints on the *in-situ* stress state surrounding the wellbore (Schmitt et al., 2012; Zoback et al., 2003). Borehole breakouts provide ~~a~~an indirect information on the stress orientation that it is difficult to extract in particular for robust quantitative stress magnitudes. Indeed, it relies on the choice of the failure model used to interpret borehole wall images. Indeed, the mechanisms that control the failure evolution of the borehole wall are not well understood both in space and time, and there is no consensus on the most appropriate failure criteria to be used. Parameterizing failure criteria is also a challenge since intact core material is often not available from deep boreholes. Finally, the set of images used to identify borehole failures is typically acquired a few days after drilling completion when it is unclear if the geometry has reached a new stationary state yet. The present analysis addresses these difficulties as we attempt to characterise the stress state at the Rittershoffen geothermal site (France).

We first present in this paper the geological and geodynamical context of the Rittershoffen geothermal site (France). We describe the borehole imaging data acquired in the GRT-1 and GRT-2 wells at the Rittershoffen geothermal project. We then proceed to a brief review of the methods used for UBI analyses with their underlying assumptions. We applied the methodology proposed by Schmitt et al. (2012) and Zoback et al. (2003) in order to assess the stress state at this site. To analyse the three successive images of the wellbore acquired up to two years after drilling completion, we developed an image processing method of the UBI data to compare in time the geometry of breakouts. We deduce from this study, the evolution of breakouts with time and evaluate its impact on our *in-situ* stress state assessment. We finally propose our best estimate of the *in-situ* stress state for the Rittershoffen site, both in orientation and magnitude.

2 Rittershoffen project context

The Rittershoffen geothermal project, also referred as the ECOGI Project is located near the village of Rittershoffen in North-Eastern France (Alsace). It is an EGS geothermal project initiated in 2011 (Baujard et al., 2015, 2017). The doublet has been drilled between Rittershoffen and Betschdorf, 6 km east of the Soultz-sous-Forêts geothermal project, in the Northern Alsace, France (Genter et al., 2010). The aim of the project is to deliver heat through a long pipeline loop to the "Roquette Frères" bio-refinery located 15 km apart. The power plant capacity is 24 MWth, intending to cover up to 25% of the client heat need. Figure 1 gives an overview of the project location and presents in the right insert the trajectory and completion of the two wells GRT-1 and GRT-2 that have been drilled (Baujard et al, 2017). GRT-1 was completed in December 2013. It was drilled to a depth of 2580 m (MD, depth measured along hole) corresponding to a vertical depth (TVD) of 2562 m. The well penetrates the crystalline basement at a depth of 2212 m MD and targets a local complex fault structure (Baujard et al., 2017; Lengliné et al., 2017; Vidal et al., 2016). The 8" 1/2 diameter open-hole section of the well starts at 1922 m MD. The borehole is almost vertical with a maximum deviation of 9° only. The first hydraulic tests concluded in an insufficient injectivity of the injection well GRT-1. Therefore, the well was stimulated in 2013, which resulted in a fivefold increase of the injectivity (Baujard et al., 2017). The target of the production well GRT-2 and its trajectory have been designed benefiting from the results of additional seismic profiles acquired in the meantime. GRT-2 targets the same fault structure but more than one kilometre away from GRT-1. Local complexities of the fault structure as 'in steps' geometry, has been observed *a-posteriori* from the micro-seismic monitoring during GRT-1 stimulation (Lengliné et al, 2017). The GRT-2 borehole was drilled in 2014 to a total depth of 3196 m MD (2708 m TVD) (Baujard et al., 2017). The granite basement is penetrated at a depth of 2493.5 m MD. The 8" 1/2 diameter open-hole section starts at a depth of 2120 m MD. This borehole is strongly deviated with a mean deviation of 37° over the interval of interest. The left insert of Figure 1 shows more specifically the geological units penetrated by the deep boreholes of the geothermal sites in Rittershoffen and Soultz-sous-Forêts. It consists of sedimentary layers from the Cenozoic and Mesozoic that are overlaying a crystalline basement made of altered and fractured granitic rocks (Aichholzer et al., 2016). Natural fractures are well developed in the Vosges sandstones and Annweiler sandstones, as in the granitic basement. The fractures network was observed from acoustic wall imagery in the open-hole sections of GRT-1 and GRT-2 and analysed by Vidal (2017). The analysis of the major continuous natural fractures concluded, in GRT-1, in a global orientation N 15° E to N 20° E with a dip of 80° W. In GRT-2, the main fracture family is oriented N 155° E to N 175° E with a dip of 80° E to 90° E. Fracture density is highest on the roof of the granitic basement (Vidal, 2017). Oil and Gas exploration in the area led to a good knowledge of the regional sub-surface including measures of temperatures at depth. The unusual high geothermal gradient encountered in Soultz-sous-Forêts which is one of the largest described so far in the Upper Rhine graben, encouraged the development of the ECOGI project in this area (Baujard et al, 2017).

The geological context is characterized in the vicinity of the Soultz-sous-Forêts and Rittershoffen sites from numerous studies owing to the extended geophysical exploration in the region (Aichholzer et al., 2016; Cornet et al., 2007; Dezayes et al., 2005; Dorbath et al., 2010; Evans et al., 2009; Genter et al., 2010; Rummel, 1991; Rummel & Baumgartner, 1991). Given that GRT-

1 and GRT-2 wells penetrate geologic units similar to those in Soultz-sous-Forêts, information from Soultz-sous-Forêts site can be used to better characterize the geological units through which the wells in Rittershoffen are drilled (Aichholzer et al., 2016; Vidal et al., 2016). It can be used in particular for the strength and mechanical characteristics of these geological units which are poorly characterized at Rittershoffen site since no coring was made during drilling (Heap et al., 2017; Kushnir et al., 2018; Villeneuve et al., 2018). The World Stress Map (WSM) released in ~~2008~~2016 also compiles the information available on the present-day stress field of the Earth's crust in the vicinity and gives an overview of the values and results which can be expected in Rittershoffen (Cornet et al., 2007; Heidbach et al., 2010; Rummel & Baumgartner, 1991; Valley & Evans, 2007a). The data collected from WSM are presented in Figure 1 and indicate that an orientation of the maximum principal stress close to N169°E and a normal to strike slip faulting regime are expected for our study area.

3. Rittershoffen well data

3.1 GRT-1 data

Several extensive logging programs accompanied the drilling of wells GRT-1 and GRT-2. One was conducted in December 2012 in the open-hole section of GRT-1, few days after drilling (Vidal et al., 2016). UBI acquisitions were carried out (Luthi, 2001). Figure 2 (b) shows the amplitude image acquired in 2012 in GRT-1 and Fig. 2 (c) displays the radius of the borehole computed from the double transit time image. The well logging also included caliper, spectral gamma ray and gamma-gamma acquisitions that enable an estimation of rock alteration and bulk density. The injectivity measured during the first hydraulic test between December 30th, 2012 and January 1st, 2013 showed a low injectivity (Baujard et al., 2017). To enhance the injectivity, the hydraulic connectivity between the well and the natural fracture network has been increased through a multi-step reservoir development strategy. First a thermal stimulation of the well has been performed in April 2013. A cold fluid (12°C) was injected at a maximum rate of 25 $L.s^{-1}$ with a maximum wellhead pressure of 2.8 MPa. The total injected volume was 4230 m^3. Second, a chemical stimulation followed in June 2013. Using open hole packers, a glutamate-based biocide was injected in specific zones of the open hole section of GRT-1 (Baujard et al., 2017). Finally, a hydraulic stimulation of the well has been performed in June 2013 with a large seismic monitoring at the surface (Lengliné et al., 2017; Maurer et al., 2015). During these two last phases, a moderate volume injection, 4400 m^3 were injected in the open hole. The hydraulic stimulation lasted during 30h, with a major phase of stepwise flow rates from 10$L.s^{-1}$ to 80 $L.s^{-1}$ (Baujard et al., 2017). As a result, the injectivity was improved fivefold due to this thermal, chemical and hydraulic (TCH) stimulation program. Two other borehole imaging programs were conducted in December 2013 shortly after stimulation of the well and significantly later in June 2015. The amplitude and travel time (or radius) images used in the analysis are shown respectively in Fig. 2 (e) and Fig. 2 (f) for the logging program of 2013 and in Fig. 2 (h) and Fig. 2 (i) for the logging program of 2015.

This time lapse UBI dataset, whose characteristics are summarized in Table 1, provides the essential information for the present study as it enables to identify evidences of irreversible deformation and failure (natural and induced fractures, breakouts, fault zones, damage zones, etc) along the borehole wall. Vidal et al. (2016) analysed the images acquired in GRT-1 and identified fractured zones impacted by the TCH stimulation, without assessing the stress state and its evolution. Hehn et al. (2016), whose measurements are discussed later in section 9.2, analysed the orientation of DIFTs in GRT-1 in the granitic basement but also in the upper sedimentary layers, investigating the orientation of the stress field with depth.

We identify wellbore wall failure and use these observations to characterise the stress state in the reservoir, including its evolution in time. Wellhead pressure measurements of the hydraulic stimulation are also used to estimate a lower bound of the minimum horizontal stress *(Sh)*.

3.2 GRT-2 data

An extended logging program was also conducted in GRT-2, including repeated UBI borehole imaging (see Table 1). Figure 3(c) and 3(d) show respectively the amplitude image acquired in 2014, between 2404 m and 2412 m, and the radius image acquired in 2015 between 2468 m and 2472 m, in GRT-2. No hydraulic stimulation was performed in this well since its initial injectivity was sufficient (Baujard et al, 2017).

4. Stress estimation methodology

The approaches proposed by Zoback et al. (2003) and by Schmitt et al. (2012) are used to fully characterize the *in-situ* stress field at the Rittershoffen geothermal project. In the following, the symbol *S* refers to the total stress when σ refers to the effective stress (Jaeger & Cook, 2009). We suppose that one of the principal stresses of the *in-situ* stress tensor is vertical, which is a common assumption. This hypothesis is justified by the first-order influence of gravity on the *in-situ* stress state, although this assumption may not be valid locally. In the following, we denote the vertical principal stress, *Sv*. The magnitude of the vertical stress *Sv* is obtained from the weight of the overburden. It is calculated by the integration of density logs (see part 8.2). The two other principal stresses act horizontally: *SH*, the maximum horizontal stress and *Sh*, the minimum horizontal stress. The magnitude of the minimum horizontal stress *Sh* is estimated from the wellhead pressure measurements carried out during the hydraulic stimulation of GRT-1 and from the hydraulic tests performed in the reservoir of Soultz-sous-Forêts (see part 8.3). The analysis of the borehole failures is evaluated using televiewer images data (Zemanek et al., 1970; Zoback et al., 1985). The orientation and magnitude of *SH* is assessed using a failure condition at the borehole wall~~.~~: the three common failure criteria considered in our analysis i.e. the Mohr-Coulomb criterion (Jaeger & Cook, 2009), the Mogi-Coulomb criterion (Zimmerman & Al-Ajmi, 2006) and a true triaxial version of the Hoek-Brown criteria (Zhang et al., 2010), are presented in section 4.2.

4.1 Wellbore stress concentration

To express the stress concentration around the quasi-vertical borehole GRT-1 (maximum deviation is only of about 9°), we assumed its shape to be a cylindrical hole, and used the well-known linear elastic solution, often referred to as the Kirsch solution (Kirsch, 1898; Schmitt et al., 2012). For the deviated well GRT-2 where the plane strain approximation is not valid anymore, we used a 3D solution taking into account ~~its deviation (Schmitt et al. 2012).~~the constant deviation of 37° measured along the section of interest. The equations in which are involved the geometry parameters of the well, the far field stresses and the fluid pressure, are well documented in the literature. We refer to the summary proposed in the review from Schmitt et al. (2012) for the general case of a 3D well randomly inclined in regard to the far field stresses. The same methodology has been for example proposed by Wileveau et al., (2007). A summary of the steps leading to the equations used to compute the *SH* stresses for the deviated well GRT-2 is proposed in Appendix A. Note that we included in our solution a thermal stress component that accounts for the thermal perturbation induced by the drilling process. This component is detailed later in section 8.4. We used the formulation of the thermo-elastic stresses arising at a borehole given by Voight & Stephens (1982), also recalled in Schmitt et al. (2012). We computed the effective stress at the borehole wall considering a hydrostatic pore pressure given by $Pp = \rho_f .g.z$, i.e. with the head level located at the surface. The fluid density ρ_f, is taken as 1000 kg.m^{-3} and the gravitational acceleration g, as 9.81 m^2.s^{-1}. z is the vertical depth (TVD) in meter from ground surface.

4.2 Failure criterion

At the scale of the surrounding of borehole (a few decameters), we assume a linear elastic, homogeneous and isotropic rock behaviour prior to failure. When the maximum principal stress exceeds the compressive rock strength, rock fails in compression (Jaeger & Cook, 2009). Failure at the borehole wall is assessed using the elastic stress concentration solutions presented in part 4.1, combined with an adequate failure criterion. There is currently no consensus concerning the appropriate failure criteria to assess wellbore wall strength. Since, in the case where the pore pressure and the internal wellbore pressure are in equilibrium the radial effective stress at the borehole wall is equal to zero, a common assumption is to consider that the Uniaxial Compressive Strength (UCS) is a good estimate of wellbore strength (Barton et al., 1988; Zoback et al., 2003). Others suggest that the strength of borehole walls in low porosity brittle rocks could be less than the UCS, because the failure could be controlled by extensile strains (Barton & Shen, 2018; Walton et al., 2015) or fluid pressure penetration (Chang & Haimson, 2007). The presence of non-zero minimum principal stress and the strengthening effect of the intermediate principal stress however suggest that the borehole wall strength should be larger than UCS (Colmenares & Zoback, 2002; Haimson, 2006; Mogi, 1971). In view of this situation and because stress magnitudes evaluation differs according to the criterion used in the analysis, we compared the estimates obtained using three commonly used failure criteria in borehole breakouts analyses: 1) the Mohr-Coulomb criterion (Jaeger & Cook, 2009), 2) the Mogi-Coulomb criterion (Zimmerman & Al-Ajmi, 2006) and 3) a true triaxial version of the Hoek-Brown criteria (Zhang et al., 2010). The formulation ~~of these criteria~~ is given in ~~the following equations (~~Eq. (1) ~~to (3))~~ for the Mohr-Coulomb criterion in the principal effective stress space $\sigma_1 - \sigma_3$ ~~for the Mohr-Coulomb criterion and in~~. The Mogi-Coulomb and Hoek-Brown criteria include a so-called "effective mean stress" (Zimmerman & Al-Ajmi, 2006) expressed as a function of the principal effective stresses as $\sigma_m = \frac{\sigma_1 + \sigma_3}{2}$ and an octahedral shear ~~vs. mean stress space~~ ~~$\tau_{oct} - \sigma_m$~~ ~~for~~stress, given by $\tau_{oct} = \sqrt{(\sigma_1 + \sigma_2)^2 + (\sigma_2 + \sigma_3)^2 + (\sigma_3 + \sigma_1)^2}$. Eq. (2) and (3) express the Mogi-Coulomb and Hoek-Brown criteria~~:~~ in the space (τ_{oct}, σ_m):

$$\text{Mohr-Coulomb: } \sigma_1 \geq C_0 + q\cancel{.\sigma_3} * \sigma_3 \quad (1)$$

$$\text{Mogi-Coulomb: } \tau_{oct} \geq a + b\cancel{.\sigma_m} * \sigma_m \quad (2)$$

$$\text{Hoek-Brown: } \frac{9}{2.C_0}\cancel{.\tau_{oct}^2} * \tau_{oct}^2 + \frac{3}{2\sqrt{2}}\cancel{.m_i.\tau_{oct}} * m_i * \tau_{oct} - m_i\cancel{.\sigma_m} * \sigma_m \geq C_0 \quad (3)$$

~~Where~~ C_0 is the uniaxial compressive strength and q is a material constant that can be related to the internal friction angle, φ, through $q = \left(\frac{\pi}{4} + \frac{\varphi}{2}\right)$. ~~The octahedral shear stress is given by~~ $\cancel{\tau_{oct} = \sqrt{(\sigma_1 + \sigma_2)^2 + (\sigma_2 + \sigma_3)^2 + (\sigma_3 + \sigma_1)^2}}$ ~~and the mean stress by~~ $\cancel{\sigma_m = \frac{\sigma_1 + \sigma_3}{2}}$~~.~~ The variables a and b in the Mogi-Coulomb criteria and m_i in the Hoek-Brown criteria are parameters that are related to the material friction and cohesion.

5. Strength estimation

Four simplified lithological categories have been used for the strength characterization of the rock at depth in the Rittershoffen reservoir. ~~All the~~The openhole section of GRT-1 and GRT-2 crosses Vosges sandstones and Annweiler sandstones of the Buntsandstein. All the lower Triassic sandstones have been grouped in a single category. The granitic section has been separated in three categories according to the type and intensity of alteration. The simplified lithologic profile for GRT-1 and GRT-2 wells are indicated in Table 2. Considering the methodology used here, the relevance and accuracy of the stress characterization is highly conditioned by the values of the rock strength parameters and by the failure criterion chosen. In Rittershoffen, the drilling was performed exclusively in destructive mode and no sample is available to measure rock moduli and strength characteristics. ~~Thereby, mechanical tests on core samples from the nearby~~GRT-1 and GRT-2 wells penetrate geologic units similar to those in the nearby Soultz-sous-Forêts site. Information from the Soultz-sous-Forêts site are thus used to better characterize the strength and mechanical characteristics of the geological units through which the wells in Rittershoffen are drilled (Heap et al., 2017; Kushnir et al., 2018; Villeneuve et al., 2018, Heap et al, 2019). Mechanical tests that have been carried out on core samples from the Soultz-sous-Forêts site are used to characterize the rock properties (Rummel, 1991; Valley & Evans, 2006). ~~Indeed, boreholes of both sites penetrate the similar lithological units and therefore using Soultz-sous-Forêts mechanical data for an application at the Rittershoffen site is considered acceptable. At~~At the Soultz-sous-Forêts site, EPS-1 borehole was continuously cored from 930 to 2227 m (Genter et al., 2010; Genter & Traineau, 1992, 1996) providing samples of the Sandstones in the Buntsandstein and in the crystalline basement. Some cores have also been obtained in the borehole GPK-1 from various depth sections and were analysed by Rummel (1992). ~~For the Buntsandstein sandstones, because of the high variability of the rocks characteristics within this same geological unit and because only very~~

~~few tests were performed on these sandstones, we rather used typical strength parameters (Hoek & Brown, 1997).~~For the Buntsandstein sandstones, Heap et al, (2019) studied in detail the strength evolution with depth of the Buntsandstein mechanical properties. They evidenced significant variations of the compressive strength together with elastic modulus changes. They also pointed out the role of the fluid content on the UCS. However, these variations are limited compared to the statistical fluctuations of our measurement. Accordingly, we gathered the Buntsandstein sandstones as a single unit. The elastic and strength parameters used for our analyses are summarized in Table 2. The variability range given for elastic parameters, cohesion and UCS reflect natural rock heterogeneities and depict the variability in values encountered. Indeed, we recognize different sources of uncertainty on the mechanical and strength parameters which limit our approach. In addition to the absence of direct strength measurements for the study site, the mechanical and strength parameters are selected from core or cuttings analyses performed in laboratory conditions. The parameters are thus not necessarily representative *in-situ* under large scale conditions, due for example to the presence of core damage.

6. Images processing and borehole failure identification

Stress induced failures are identified and measured from acoustic borehole images. The confidence and accuracy of these determinations depend on the quality of the images. In the following, we describe the original data as well as the processing we applied to improve the quality and comparability of the images. We also explain how we measure borehole failure on these images and the limitations associated with these measurements.

6.1 Quality of the acoustic televiewer images

Several artefacts can deteriorate the quality of acoustic image data (Lofts & Bourke, 1999). The images acquired in Rittershoffen suffer from some of these limitations. The quality of the image depends of the tool specification, the acquisition parameters and logging conditions. All acoustic images at Rittershoffen were acquired by Schlumberger with their UBI (Ultrasonic Borehole Imager) tool. The tool and acquisition parameters were similar between each log, but not identical. For example, the GRT-1 log in 2013 was acquired using a smaller acquisition head (see the changes in transducer diameter detailed in Table 1. The acquisition resolution was the same for every log, i.e. 2° azimuthal resolution and 1 cm depth sampling step. The 2012 log of GRT-1 has the best quality image of the entire suite. The image suffers of signal loss artefact (Lofts & Bourke, 1999) in some limited sections, most commonly related to the presence of breakouts or major fracture zones ~~(Fig. 3 (a)).~~. The zones of signal loss are clearly identified in the radius image presented in Fig. 3 (a) by persisting white patches.

The 2013 log of GRT-1 is of comparable quality than the 2012 log and suffers also of some limited signal loss artefacts. The major issue with the image of GRT-1 acquired in 2013 is that the orientation module was not included in the tool string and thus the image cannot be oriented with magnetometer data as it is usually done for this type of data.

The 2015 log of GRT-1 generally suffers from signal loss issues, not only in areas with major fracture zones and breakouts. In the lower part of the log, wood grain ~~artefact~~textures (Lofts & Bourke, 1999~~) is~~), related to processing noise, are also observable (see Fig. 3 (b)). ~~This is particularly developed~~Wood grain textures are especially encountered below 2431 m MD.

The quality of log data ~~of logs~~ from GRT-2 is generally ~~worse~~lower than ~~the ones of~~for GRT-1. This is due to the deviation of GRT-2 that makes wireline logging more difficult. The 2014 log of GRT-2 suffers from stick-slip artefacts on its entire length ~~(~~. The effects of the alternating compression and stretching on the images and highlighted in Fig. 3 (c~~)).~~), are particularly significant and possibly lead to errors in the recording of the fractures. The 2015 log in GRT-2 does not show any sign of stick-slip but ~~present~~presents an erroneous borehole radius record leading to an incorrect borehole geometry ~~evaluation~~assessment (Fig. 3 (d)).

Despite these difficulties, the images collected in the GRT-1 borehole are of excellent quality. Signal loss is the main problem and it prevents to measure the depth in the radial direction of the breakout in some zones. Given the extent of the artefacts highlighted in GRT-2, the measurements of the breakout parameters in this borehole are much more uncertain.

6.2 Processing of the UBI images

Prior the use of the images for assessing borehole failure, the images went through the following pre-processing steps:

1) Transit time was converted to radius using the fluid velocity recorded during the probe trip down the borehole;
2) Images were filtered to reduce noise;
3) Digital image correlation was applied across the successive logs in order to correct the image misalignment both in azimuth and depth.

The borehole radius was computed from the transit time following Luthi (2001):

$$r = \cancel{\frac{t_{twt} \cdot v_m}{2}}\frac{t_{twt} * v_m}{2} + d \qquad (4)$$

with t_{twt} the two-way travel time, v_m the acoustic wave velocity in the drilling mud, and d the logging tool radius. Images are filtered using a selective despiking algorithm implemented in WellCad™ using a cut-off high level (75%) and a cut-off low level (25%) in a 3x3 pixels window. The goal of this process is to replace outliers by cut-off values when the radius exceeds the cut-off high or low level. Finally, digital image correlation was used to insure proper alignment of the UBI images. This was required for the GRT-1 2013 image because this image was not oriented with a magnetometer/accelerometer tool. The process was also applied to the 2015 GRT-1 data to facilitate comparison between images. For this purpose, we developed a technique based on a Particle Image Velocimetry (PIV) method (Thielicke & Stamhuis, 2014) that relies on optical image correlation but being applied to travel time UBI images. This image alignment process is illustrated in Fig. 4. Figure 4 (a) shows as example the “correlation box” in the travel time UBI image of reference - i.e. 2012 in this case – and the corresponding one in the image to compare with - i.e. the image of 2013 – which it is shifted of a given displacement vector (dX, dY) within the “search box”. The cross-correlation function, which is a measure of the similarity between the thumbnails, is computed between the correlation boxes for each displacement vector (dX, dY). Right panel of Figure 4 (a) shows a map of the cross-correlation function computed for every displacement vector in a given search box. The two-dimensional cross-correlation function is an operator acting on two intensity functions $s(X,Y)$ and $r(X,Y)$, defined as a norm of the colour levels at each position of each thumbnail. C_{sr} is defined at a position (X,Y) and for a shift (dX, dY) by Eq. (5):

$$C_{sr}(dX,dY) = s(X,Y)\cancel{\otimes} \otimes r(X,Y) = \iint_{-\infty}^{+\infty} s(X,Y) r \cancel{*(X-dX,Y-dY)} (X-dX,Y-dY)\, dX dY \qquad (5)$$

The position (dX, dY) within the “search box” with the highest cross correlation correspond to the best alignment (see Fig. 4 (a)). The operation is repeated along the image for each position of the search box. Importantly, the correlation box is taken with an anisotropic shape to account for the rigid rotation of the UBI tool and the linear property of the acoustic camera. The size of the correlation box is 180 x 20 pixels. This configuration is appropriate to identify principally the azimuthal offset while it is less sensitive to the depth mismatch. We investigated offset up to 180 pixels horizontally corresponding for our 2° resolution to a complete 360° rotation. We considered vertical offset of ± 10 pixels corresponding to offsets of about ± 10 cm. Figure 4 (b) gives an example of image realignment and shows the efficiency of the process. This correlation process allows to align finely the successive images and thus to study the borehole shape evolution with time more accurately.

6.3 Determination of the borehole failure

For GRT-1, the breakouts have been determined through a visual analysis of borehole sections computed every 20 cm from 1926 m to 2568 m (MD) from the double transit time data. The borehole sections are computed by stacking (averaging using the median) the data collected every 1 cm over 20 cm borehole interval (with no overlap between two successive sections). The median is thus used because it is less sensitive to extreme values than the mean and thus is efficient at removing local noise from the data. Prior to determining breakout geometrical parameters, the actual borehole center is determined by adjusting the best fitted ellipse to the borehole section. This process corrects for eventual logging probe decentralisation. For each section presenting the characteristic elongated shape of breakouts due to stress induced failure, the azimuthal position of the edges and the center of each limb is determined by visual inspection. Figure 5 gives examples of such determination to depict the process. The breakout edges are defined as the location where the wellbore section departs from a quasi-circular section adjusted by the best fitted ellipse. As it can be seen in Figure 5, this typically spans an azimuthal range much broader than the low amplitude reflections visible as dark bands on the amplitude images and justifies the choice to use the double transit time data. The positions of the breakout edges are not easy to determine in a systematic and indisputable manner, and a significant uncertainty is associated with these measurements. Related to this issue, it is not possible to determine on the images what azimuthal range of the wellbore is enlarged by purely stress redistribution processes and what part is enlarged subsequently by the effects of drill strings wear. These uncertainties about the physical process controlling the enlargement of the breakout could limit the comparisons between the three successive logs acquired in GRT-1. Breakout measurements were thus performed on all three images concomitantly and consistently. We controlled for example that within a tolerance dictated by the uncertainties of the measurements, the width of breakouts only remains identical or increases: no decrease in width is measured between successive logs.

Figure 2 (d), (g) and (j) summarize all the measurements of the breakout's geometry performed in GRT-1, for the images acquired in 2012, 2013 and 2015. Black dots indicate the azimuth at which the radius of the breakout is maximum and red bars link the azimuthal position of the breakout edges used to compute the width of the breakouts. Given the difficulty of measuring breakouts as discussed previously (i.e. artefacts affecting the images, disputable positions of the breakout edges), a confidence ranking has been established for each breakout. This confidence level is presented in Fig. 2 (k). From the geometry of the breakouts, we compute the breakout widths which are obtained from the breakout edge azimuths. The deepest point of the breakout is used to determine the enlargement radius. In some situations, signal loss issues prevent the determination of the enlargement radius, as it is shown in Fig. 5 for the image of GRT-1 acquired in 2015. The measured width (black dots, in degree) and enlargement radius (red dots, in mm) are determined from the GRT-1 data set acquired in 2012 and presented in Fig. 2 (l).

Drilling Induced Tension Fractures (DITFs) are also identified from the GRT-1 borehole images using the same procedure as for the breakout determination. For example, clear DITFs are evidenced in the amplitude image from 2395 m to 2400 m in GRT-1 and presented in Fig. 6. Green crosses show the azimuth of the DITFs that is measured in GRT-1 every 20 cm. Blue dots in Fig. 2 (d), (g) and (j) summarize the azimuth of the DITFs measured in GRT-1, respectively in 2012, 2013 and 2015.

Given the poor quality of the double transit time images acquired in GRT-2, less focus has been given to the analysis of the borehole failure in this well. The data set consists of the acquisitions made in 2014 after completion of the borehole and in 2015. The investigated depths vary from the 2014 to the 2015 dataset. It is from 1950 m (Vertical Depth – 2220 m MD) down to 2125 m (TVD – 2440 m MD) in 2014 when it is down to 2160 m (TVD – 2480 m MD) in 2015. The well is strongly deviated. The concentration of stresses within the borehole wall is expressed under the assumption of a constant deviation of 37° and measurements carried out as a function of the True Vertical Depth, to be comparable with the results obtained in GRT-1 which is considered as vertical. Borehole sections are computed every 50 cm. To this end, borehole sections are stacked using the data collected every 1 cm over 50 cm borehole interval, all along the transit time image. As for GRT-1, the actual borehole centre is determined by adjusting a best fitted ellipse to the borehole section. Breakouts are analysed by visual analysis in a same manner as for GRT-1 data. The difficulties encountered with the identification of breakout geometry are more pronounced for images acquired in GRT-2 as artefacts are more developed. The deviation of this well results on pronounced stick-slip effects. For a more accurate comparison between the measurements carried out on the images acquired in 2014 and 2015, measurements are performed for the two images concomitantly. No DITFs are identified on the GRT-2 borehole images.

7. Analyses of temporal borehole failure evolution

The characterization of the stress tensor derived from the analysis of borehole failures typically relies on a single borehole image data set. From this snapshot in time, stresses are estimated while information on the evolution of breakout shape in time is not available. Interestingly, for the ECOGI project, the acquisition of three successive image logs allows to study this evolution. Here, the time evolution of breakouts, referred as breakout development, is analysed to characterize the time evolution of the borehole failure. A common hypothesis concerning borehole breakout evolution is that their width remains stable and is controlled by the stress state around the well at the initial rupture time. Progressive failure is supposed to lead however to breakout deepening until a stable profile is reached (Zoback et al., 2003).

An example of a time-lapse comparison of breakout shapes is presented in Fig. 7. Images of GRT-1 from 2012, 2013 and 2015 show a clear breakout at a depth of about 2126 m in the "couches de Trifels" in the Buntsandstein. Breakouts can present three types of evolution:

1) They can develop along the well, corresponding to an increase in the vertical length of breakouts. We refer to this process as *breakout extension*;
2) They can widen, corresponding to an apparent opening between the edges of the breakouts. We refer to this process as *breakout widening*;
3) They can deepen, corresponding to an increase of the maximal radius of the breakout (or “depth” of the breakout) measured in the borehole cross section at a given depth. We refer to this process as *breakout deepenin*g.

Figure 7 shows the evolution from 2012 to 2015 of the breakouts, at 2125.6 m. Failure did not occur in 2012 while breakouts are visible in 2013 and 2015. When superposing the 2013/2015 borehole sections, no change in breakout shape is highlighted for the west limb although a slight widening is visible on the east limb. Possible deepening of the east limb is occulted by signal loss issues. The borehole section computed at 2126.2 m shows on the contrary, no modification of the breakout shape from 2012 to 2015 in GRT-1.

Development of borehole failures depends also on the lithology. Breakout extension (longitudinal failure development) is quite common in the Buntsandstein while it is very limited in the basement granites, which is highlighted in Fig. 8. The evolution occurs exclusively between the 2012/2013 data set while no longitudinal extension occurs during 2013 and 2015. In 2012, a total breakout length of 404 m is observed. It increases to 504 m in 2013 and then remains stable in 2015 with a length of 506 m. There is no clear evolution of DITFs along the GRT-1 well despite the hydraulic and thermal stimulation performed between 2012 and 2013.

Figure 9 shows an increase of breakout width. We first compare the data acquired in 2012 and in 2013. 73% of the change of width is within an interval -10° / +10°, i.e. within our measurement uncertainty. For these breakouts no changes of width can be highlighted within our level of uncertainty. However, for 27% of our data, we observe an increase of width larger than 10°. This is reflected by the long tail (with values higher than 10°) of the histogram computed from the width of breakouts (see Fig. 9 (c)). The widening of these breakouts is undisputable. When comparing the data acquired in 2013 and in 2015, very little changes are observed. Indeed, most of the measured changes remain below our uncertainty level of ±10° (red histogram on Fig. 9 (c)).

The evolution of the maximum radial extension (breakout deepening) of the breakout measured in the borehole cross sections is presented in Fig. 10. This parameter is more delicate to track because of signal loss issues (see for example Fig. 3 (a)). In our analysis, we filtered out obvious incorrect depth measurements related to these artefacts, i.e. when the computed radius from transit time image is clearly shorter than the drill bit radius. For both time intervals (2012-13 and 2013-15), the change in the depth of the breakout is symmetrically distributed around 0 mm and spans a variability of about ±15 mm. We interpret this distribution as an indication that if any deepening occurred, it remained within our uncertainty level. Our data analysis does not enable to conclude in a general deepening of the breakouts.

8. Stress characterization

We propose in this section a complete stress characterization at different periods in both the GRT-1 and GRT-2 wells, including a thermal history and thermal stress analyses and discuss the impact of breakout widening in time on stress estimation. To that purpose, we first determine the orientation of the stress tensor. We then detail how we estimate the minimum horizontal stress component *Sh*, the vertical stress component *Sv* and the thermal component. Finally, we propose an estimation of the maximum horizontal stress component *SH* from the measurement of the width of breakouts.

8.1 Maximum horizontal stress *SH* orientation

The orientations of breakouts and DITFs are a direct measure of the principal stress directions in a plane perpendicular to the well. As discussed previously, we assume that *Sv* is in-overall vertical which is a common hypothesis in such an approach and is justified by the first-order effect of gravity on *in-situ* stresses. In GRT-1 which is considered as vertical, DITFs are aligned with the direction of the maximum horizontal stress (*SH*) and breakouts are aligned with the direction of minimum horizontal stress (*Sh*).

Figures 2 (d), (h) and (i) show the orientation of breakouts (black dots) and DITFs (blue dots) measured in GRT-1. The measurements are compiled in Fig. 11 as circular histograms. We chose to only analyse data from the images acquired in 2012 and in 2015. Indeed, data acquired in 2013 were obtained without orientation since the device was not functioning correctly and are reoriented with respect to the 2012 data. Subsequently, the measurements carried out in the 2013 image do not bring additional constraints in terms of stress orientation.

In the Buntsandstein sediments, the failure orientation is stable and indicates that the principle stress *SH* is oriented N15° ±19° (one circular standard deviation). The same failure orientation persists in the upper section of the granite down to about 2270 m. Below this depth borehole failure orientation is much more variable as it seems to be influenced by the presence of major fault zones crossing the GRT-1 borehole at a depth of 2368 m (MD) (Vidal et al., 2016). Below 2420 m, which is the deepest large structure visible on the GRT-1 borehole image, the failure orientation indicates that *SH* is oriented 165° ±14°. This is significantly different from the orientation in the sediments with a 30° counter-clockwise rotation. Such differences in orientation with lithologies have already been noticed by Hehn et al. (2016) from the analysis of the orientation of drilling induced fractures observed on borehole acoustic logs acquired in GRT-1. The orientation of *SH* proposed by Hehn et al. (2016), i.e. globally N155°E in the basement and N20°E in the sedimentary layer, is consistent with our measurements.

The geological study of the cuttings from the drilling of GRT-1 and GRT-2 enabled to determine the rock density profile in both wells (Aichholzer et al., 2016). Thanks to this analysis, we estimate the mean density of each lithological layer. Table 3 shows the rock volumetric mass density as a function of the vertical depth (TVD). The magnitude of the vertical component *Sv* at depth is computed accordingly by integrating the volumetric mass density profile ~~and~~from surface. A linear regression is fitted to the measurements obtained ~~from surface.~~for the depth range studied here, i.e. [1900-2600] m. In the following, the vertical component *Sv* is computed from a linear trend expressed as a function of vertical depth (TVD) *z*:

$$Sv\,[MPa] = 0.\cancel{025\ .}0248 * z\,[m] - 0.83 \qquad (6)$$

As the linear trend is expressed as a function of the vertical depth, we use the same equations in the computation steps leading to the *SH* stress estimates in GRT-1 and GRT-2. As the density profile is integrated from surface to reservoir depth, the uncertainty on density adds up and the uncertainty on the vertical stress increases with depth consequently. Considering an uncertainty of 50 $kg.m^{-3}$ on the densities leads to a 2.5 MPa uncertainty on *Sv* at reservoir depth. This uncertainty is not significant compared to other uncertainties involved in the analysis as for example those related to the mechanical parameters chosen in the inversion of the maximum horizontal stress *SH*.

8.3 Minimum horizontal stress Sh

We take the first order assumption that the minimum horizontal stress *Sh* varies linearly with depth. Usually, the minimum horizontal stress *Sh* is estimated at depth from hydrofracture tests~~,~~ (i.e. ~~Valley & Evans (2007))~~Haimson & Cornet (2003)) but this was not done at Rittershoffen site. ~~If~~As the data available for the ECOGI project ~~limit~~doesn't enable to compute a profile for the *Sh* stresses, our analysis of the minimum stress component~~,~~ is based on the numerous injection tests that were conducted in Soultz-sous-Forêts. We present in Fig. 12 ~~their~~ main trends~~.~~ computed from pressure limiting behavior during hydraulic injections. For large depths, the injection tests performed in the deep wells (GPK-1, GPK-2 and GPK-3 or EPS-1) of Soultz-sous-Forêts (Cornet et al., 2007; Valley & Evans, 2007b) give important constraints for the minimum horizontal stress *Sh* at the Rittershoffen site~~ for large depths.~~. In addition, the study of Rummel & Baumgartner (1991) provides estimates at shallow depth. In our analysis of the stress state in GRT-1 and GRT-2, we compute the horizontal minimum stress *Sh* as a function of the true vertical depth (TVD) *z* from the linear trend proposed by Cornet et al. (2007) for the site of Soultz-sous-Forêts (Figure 15):

$$Sh[MPa] = 0.015\cancel{.} * z\,[m] - 7.3 \qquad (7)$$

~~In order to check~~From the data available for the Rittershoffen site, i.e. the wellhead pressure measured during the hydraulic stimulation of GRT-1 (Baujard et al., 2017), we estimated a lower bound of the minimum horizontal stress *Sh* at 1913 m in Rittershoffen. The measurement enables to verify the applicability of the linear trend inferred from acquisitions in Soultz-sous-Forêts to the Rittershoffen site~~, we estimated a lower bound of the minimum horizontal stress *Sh* at 1913 m in Rittershoffen from the measure of the wellhead pressure during the hydraulic stimulation of GRT-1 (Baujard et al., 2017).~~. Figure 13 shows that the variation of wellhead pressure with the flow is slower during the high rate hydraulic stimulation (above 40 L.s^{-1}) than during the low rate hydraulic stimulation (below 40 L.s^{-1}). ~~This~~The change in ~~behavior~~behaviour highlighted for higher values of the flow rate is interpreted as the beginning of a pressure capping resulting from fractures reactivation. Hydraulic stimulation ~~typically increases~~operations aim at increasing pore pressure, which reduces the effective stress until pressure equals *Sh* in magnitude. In theory, an increase of pressure could activate new fractures which results in the capping of the recorded pressure: in such a case, minimum horizontal stress is inferred at depth from the maximum pressure achieved during ~~such a test.~~the hydraulic operations. Meanwhile, other processes (shearing of existing weak fractures for example) could possibly result in the capping of pressure for lower pressure values. ~~From Figure 13, we assume that wellhead pressure caps at 22.6 MPa at 1913m (TMD) for a flow rate 80 L.s^{-1} (Fig. 12). It provides a lower bound to constrain the minimum horizontal stress *Sh* at depth, which is compared to the Soultz-sous-Forêts trends in Fig. 13 and shows the consistency of the linear trend used in our analysis.~~

The maximum pressure reached at 1913 m (TVD) during the hydraulic test is 22.6 MPa, for a flow rate of 80 L.s^{-1} (Fig. 12). As the measurement is recorded at the end of a gradual but not definitive stabilization of the pressure with the flow rate, the 22.6 MPa stress measured at 1913 m consists in a lower bound for the minimum horizontal stress *Sh* at depth. It is compared to the Soultz-sous-Forêts trends in Fig. 13. and the measurement shows the consistency of the linear trend used in our analysis and inferred from the operations carried out at the Soultz-sous-Forêts site.

8.4 Thermal stresses

The cooling of the well imposed during drilling, results in a thermal stress contribution. Accordingly, the characterization of the stress tensor necessitates to include a thermal stress analysis which requires a good knowledge of the thermal history of the well. We define the thermal contributions in the stress concentration at the borehole wall as: $\sigma^{\Delta T}{}_{r}$, $\sigma^{\Delta T}{}_{z}$ and $\sigma^{\Delta T}{}_{\theta}$ respectively the radial, vertical and tangential components. The thermal stresses resulting from the temperature difference, Δt, between the borehole wall and the so called ambient temperature, i.e. the initial temperature at that depth before the drilling phase or the temperature at a significant distance from the borehole (not influenced by the borehole perturbation), are expressed from Voight & Stephens (1982). These authors adapted the thermo-elastic solutions proposed by Ritchie & Sakakura (1956) for a hollow cylinder to study the stress concentrations at the borehole wall due to the application of a temperature difference. The radial component is null, and the tangential component is expressed as:

$$\sigma^{\Delta T}{}_{\theta} = \sigma^{\Delta T}{}_{z} = \alpha \cancel{.} * E \cancel{.\frac{\Delta T}{(1-\nu)}} * \frac{\Delta T}{(1-\nu)} \quad (8)$$

where ~~α~~ α is the volumetric thermal expansion, E, the Young modulus and ν, the Poisson ratio. The volumetric thermal expansion, which is kept constant in the different layers crossed by the borehole, is α = 14 x10^{-6} K^{-1}. The Young modulus and Poisson ratio values applied at the different layers are indicated in Table 2. Figure 14 (green curve) presents the temperature log acquired in 2015 in GRT-1 (Baujard et al, 2017). It is plotted along with the temperature log acquired in 2013 (red curve). The comparison shows that temperature is close to be stable during that period in GRT-1. As a result, the temperature log acquired in 2015 in GRT-1 is used as an estimate of the ambient temperature since it is considered as in equilibrium with the reservoir. Temperature at the borehole walls at drilling completion is best estimated from the temperature log acquired four days after drilling competition. The temperature log is presented in Fig. 14 (blue curve) and the difference in temperature Δt computed from these logs is presented in the right panel of Fig. 14. Interestingly, these temperature logs show a clear anomaly at 2360m where the wells are crossing the main fault zone associated to a major permeable structure that controls two third of the total flow during flow tests (Baujard et al., 2017).

8.5 Maximum horizontal stress SH magnitude

The determination of the azimuthal position of the breakout's edges and of their width from the analysis of the UBI images acquired in GRT-1 and GRT-2 enables to estimate the maximum horizontal stress *SH*, and to evaluate its evolution with depth and time. Here, we present the results of our inversion, at multiple dates in GRT-1 and GRT-2.

In GRT-1, we obtain for each UBI log (in 2012, 2013 and 2015), three estimates of the magnitude of *SH*, according to the failure criterion. Figure 15 shows estimates of the magnitude of *SH*. The maximum horizontal stress *SH* in GRT-1 is presented for the 2013 UBI log as a function of the true vertical depth (TVD), along with the *Sh* and *Sv* obtained previously (Eqs. (6) and (7)). The horizontal error bars are calculated from the uncertainty on the elastic parameters, on the *Sh* and *Sv* estimates and on the measurements of the width of the breakouts. The uncertainty ΔSH is obtained by integration, taking into account the uncertainty Δxi on each variable x_i involved in the estimation of *SH*, i.e the strength parameters, the *Sh* and *Sv* trends and the width of the breakouts:

$$\Delta f = \cancel{\sum_i \left|\frac{\partial f}{\partial xi}\right| . \Delta xi} \sum_i \left|\frac{\partial f}{\partial xi}\right| * \Delta xi \quad (9)$$

Figure 15 shows that the *SH* magnitudes vary significantly with the failure criterion. In particular, it shows that the *SH* stresses computed using a criterion that considers the strengthening effect of the intermediate principal stress (i.e. in Mogi-Coulomb or Hoek Brown) are higher than those calculated from a criterion that considers only the minimum and maximum principal stresses (i.e. in Mohr-Coulomb).

To choose the criterion that best describes the failure in the borehole, we use the approach proposed by Zoback et al. (2003) to display the stress state estimates presented in Fig. 15 in the stress polygon whose circumference is defined by a purely frictional, critically-stressed Earth crust. For this purpose, we suppose that crustal strength is limited by a Coulomb friction criterion with a friction coefficient $\mu = 1$. We considered a depth of 2500 m to evaluate the vertical stress and assumed a hydrostatic pore pressure. The possible stress states from 2013 UBI images, are shown in Fig. 16 in a normalized *SH* vs *Sh* space. Because 2500 m is an upper boundary for the investigated depths in our study, the circumference of the polygon sets a maximum value for the maximum and minimum horizontal stresses *SH* and *Sh*. The stresses are normalized by the vertical stress magnitude *Sv* to facilitate the comparison. The maximum principal stresses *SH* measured using both our parametrized Hoek-Brown and Mogi-Coulomb criteria (blue and black dots) exceed the polygon boundaries. With our selection of parameters, the Mohr-Coulomb criterion was therefore retained as the most suitable for characterizing rock failure in our study. The same conclusion was drawn by Valley & Evans (2015) in Basel.

For GRT-2, we calculated the *SH* magnitudes using only the Mohr-Coulomb criterion retained in the previous analysis. GRT-2 is highly deviated and the well has been imaged in 2014 and 2015. The deviation is constant in the section of interest (i.e. the open hole): 37° N355°E. *SH* stresses are shown as a function of the vertical depth (TVD) in Fig. 17 with the according error bars and plotted along with the *Sh* and *Sv* trends in GRT-2.

The impact of breakout widening on stress estimation can be evaluated from our time-lapse characterization of the stress tensor in GRT-1 and GRT-2. For GRT-2, Fig. 17 shows that *SH* magnitude changes are limited between 2014 and 2015, given the uncertainty on the estimates. Figure 18 compares the *SH* stresses estimated using the Mohr-Coulomb criterion at different dates in both GRT-1 and GRT-2 wells. The systematic shift observed between the estimates in both wells suggest that the lower stresses estimated in the deviated well lead to a borehole wall stress concentration closer to the failure condition than in the vertical well. Figure 18 evidences a time evolution of the *SH* stress estimates in GRT-1. Panel b. quantifies the differences in *SH* stress between 2012 and 2015 in GRT-1 in a 1 MPa bins histogram. The confidence in the time-evolution, is discussed in the next section considering the error on *SH*.

9. Discussion

The data set from the Rittershoffen geothermal project and our analyses allow us to discuss both the evolution over time and with depth of the observed borehole failures. The impact of these evolutions on our ability to estimate stress magnitude from borehole failure indicators is important.

9.1 Evolution of breakout geometry with time

Our analysis of the evolution of the breakouts geometry with time proves a development of breakouts along the well GRT-1 during the first year after drilling (Fig. 8). Indeed, we highlighted that sections without breakouts in 2012, four days after drilling, present characteristic breakouts in 2013 and 2015, respectively one year and 2.5 years after drilling. We also observe numerous lengths increases of the 2012 existing breakouts with time in particular in the Buntsandstein. The difficulty is to link this evolution with time with a specific process: time-dependant rheology of the rock (i.e. creep) or the effects of one of the stimulations, thermal, chemical or hydraulic. Moreover, the 2012 data were acquired at a period during which the thermal perturbations due to the drilling operations were still present. The data they are compared with have been collected in 2013 or 2015, after hydraulic, thermal and chemical stimulations of the well. As a result, the observed changes could have taken place during the thermal equilibrium of borehole after drilling or during the simulations operations, i.e. directly after drilling or later. The conclusion brought by our time-evolution analysis of the breakout's geometry contradicts the usual assumption that breakouts deepen (i.e. an increase in the maximum radius measured in the borehole cross sections) but do not widen (i.e. an opening between the edges of the breakouts) with time (Zoback et al. 2003). However, the statistical approach applied in our study along the open-hole of the well GRT-1 must be interpreted with caution. Even if we propose a systematic analysis of a time-evolutive dataset, signal loss artefacts prevent an accurate measurement of borehole radius at some depths. It limits locally our ability to reliably estimate the depth of the breakout, i.e. the extension of the breakout in the radial direction. Given this limitation, we do not totally exclude that breakouts could have deepen with time. Our breakout width evaluation is also affected by uncertainty: the deviation from the nominal cylindrical borehole geometry of the borehole adds complexity to the measurements made considering the disputable positions of breakout edges. Meanwhile, we mitigated this difficulty by proposing a systematic analysis of all dataset to ensure a more consistent measurement and by attributing an uncertainty level on these values. Our study is thus more conclusive concerning this geometric parameter given that measured changes exceed our uncertainty level.

The widening observed in our data set can be explained by the process of thermal stress dissipation. Indeed, the 30 to 35°C of cooling observed at the time of the 2012 logging, are dissipated by the time of the 2013 logging (see Fig. 14). Assuming thermo-elastic properties of the material, the thermal hoop stresses implied by the cooling reaches -17 to -20 MPa (Eq. (8)). This will be sufficient to explain the change in breakout width without including additional time-dependent failure processes.

9.2 Evolution of breakout geometry with depth

The development of breakouts depends on the rock rheology and subsequently on the lithology. For our data set, breakouts are more numerous and extended in the sedimentary cover than in the granitic basement (Fig. 2). Moreover, their development is more pronounced in the sedimentary cover when they develop with time, vertically along the well (Fig. 8). Both observations are consistent with the fact that the sediments have on average a lower strength compared to the granitic rocks (Evans et al., 2009; Heap et al., 2019; Kushnir et al., 2018), i.e. conditions are closer to failure in the sediments.

Another important aspect of the variation of breakout geometry with depth is the evolution of their mean orientation. From the combined measure of the azimuth of maximum radial extension of the breakouts (BOs) and of the azimuth of Drilling Induced Tensile Fractures (DITFs), we analyse in Figure 11 the evolution with depth of the orientation of the maximum principal stress *SH*. The measurements are repeated for the images acquired in GRT-1, in 2012 and in 2015. The consistency between the orientation of our data between the 2012 and the 2015 data set (the 2013 data set was not oriented) builds confidence in the reliability of these indicators.

Figure 11 suggests that the orientation measured in the granitic layers below 2420m in Rittershoffen is consistent with the measurements carried out in the basement of Soultz-sous-Forêts (Valley & Evans, 2007b) and tends to reach the regional orientation. The red line in Fig. 11 is a moving average of the orientation data. It is computed over a 20 m window in depth. The measurement is carried out only if 50 individual measurements or more are present in the averaging window. It shows that the orientation of the maximum principal stress *SH* varies in the studied section. Another important aspect of Figure 11 is the significant rotation of 30° from NNW to NNE highlighted between the bottom and the top of our analysed section. Such rotation with depth has already been evidenced in the Upper Rhine graben area in the Basel geothermal boreholes (Valley & Evans, 2009), in potash mines (Cornet & Röckel, 2012) and at the neighbouring geothermal site of Soultz-sous-Forêts (Valley & Evans, 2007b). Hehn et al. (2016) have also evidenced local stress rotations in the sedimentary section of GRT-1 up to the upper Triassic (Keuper) from the analyses of DITFs. The orientation measured here above the limit set close to 2400m MD (Fig. 11), is also consistent with the measurements of Hehn et al. (2016). They interpreted these variations to be related to mechanical contrasts between stiffer and softer rock layers. Another explanation for the stress rotation has been proposed by Cornet (2016). He suggested that the rotation is the result of the hydrostatic pressure effect on the effective friction angle in the Hoek-Brown failure criterion. In such a case, the rotation would be mainly a depth effect and not link to the presence of the Rittershoffen fault. The particularity of the measurements proposed in Fig. 11 is that the orientation of the maximum principal stress *SH* deviates from the regional trend within the granitic basement, while the measurement in the upper basement aligns with the orientation of the sedimentary cover (Fig. 11). The presence of a major fault crossing the GRT-1 borehole at a depth of 2368 m MD (Vidal et al., 2016) could be the explanation of this rotation. The location of the observed stress rotation, i.e. in the basement and around ~~50m~~50 m above the major fault zone, ~~doesn't~~does not assume that it is related here to the stiffness contrast or decoupling between the sedimentary cover and the underlying basement as typically assumed, but rather to the presence of a neighbouring major fault zone. Considering a high dipping fault geometry for this fault zone, it suggests that the geothermal well tangents the fault zone explaining why breakouts are observed below but also above the major drain of the fault zone located at 2368 m (Fig 11). Moreover, it was clearly demonstrated, based on continuous granite core analyses at Soultz, that fault zone could have a significant thickness due to the presence of a damaged zone characterized by an intense hydrothermal alteration (Genter et al., 2010). Therefore, the absence of breakouts visible in the altered granitic section located just above the main fault drain and the anticipated rotation of the stress field at some distance in the hanging wall and the footwall of the fault zone confirm its major mechanical influence.

9.3 Evaluation of stress magnitude from breakout width

Our study shows the sensitivity ~~to the failure criterion.~~of our approach toward the failure criterion which is chosen to describe the stability of the wellbore wall at a centimetric scale. The absence of consensus regarding the appropriate failure criterion to be used in the analysis of the borehole breakouts is a first limitation in our approach. Our analyses suggest that the Mogi-Coulomb and Hoek-Brown criteria tend to overestimate borehole wall strength because they lead to stress estimates that violate frictional strength limit of the crust (Fig. 16) while the Mohr-Coulomb strength model leads to acceptable results. This conclusion is however dependent of the detailed parameterization of the failure criterion which is in Rittershoffen supported by sparse data. The rock strength is among the main parameters that impact the stress magnitude assessment. ~~At the Rittershoffen project, we have no access to~~ Direct strength measurements are not available for the Rittershoffen project, since no cores were collected. We rely on measurement at the neighbouring Soultz-sous-Forêts site where cores are available. However, even at Soultz-sous-Forêts, a systematic characterization of the rock strength of the various lithologies is not achievable, particularly for the sediments. Also, the mechanical and strength parameters are selected from the analysis of core or cuttings performed at the laboratory scale. The measurements are thus not necessarily representative of the *in-situ* conditions. In addition to the uncertainty on the strength parameterization, the uncertainty on width determination and the evolution of width with time ~~further~~also impact the stress estimation. In the case of the GRT-1, significant changes occur between the 2012 data set (prior to reservoir stimulation operations) and the 2013-15 data sets (after stimulation). Panel (b) of Figure 18 shows that the changes in the *SH* stresses between 2012 and 2015 in GRT-1 are larger than our measurement uncertainty for 15% of the measurements and are showing principally stress increases. This change can be fully explained by the thermal equilibration of the well. ~~The uncertainty on our data doesn't allow to relate stress changes to the reservoir stimulation operations~~The uncertainty on our data does not allow to relate stress changes to the reservoir stimulation operations. Cornet (2016) showed that large-scale fluid injections conducted at the Soultz-sous-Forêts site generated large scale failure zones whose orientation varies with depth. Based on the analyses of borehole failures, considerable stress orientation variations were also highlighted with depth at Rittershoffen (Hehn et al., 2006), at Soultz-sous-Forêts (Valley & Evans, 2007b) and at other sites (e.g. Valley & Evans (2009) or Cornet & Röckel (2012)). In this respect, our measurements at the Rittershoffen site confirm the conclusions drawn at many other sites regarding the change in stress orientation. However, given the difference in the fluid volumes injected into the wells of the two sites during the stimulation processes and in injection pressures, it is difficult to associate the rotation with depth with the hydraulic stimulation of GRT-1 and to apply the conclusions reached by Cornet (2016) in Soultz-sous-Forêts to the Rittershoffen site.

9.4 Stresses magnitude evolution with depth

Stresses estimated in GRT-1 and GRT-2 suggest that *SH*, in regards of their uncertainty, is generally close to the vertical principal stresses *Sv*, consistently with a transitional regime between ~~both~~a strike-slip ~~regime~~ and a normal faulting regime (Anderson, 1951). This result is consistent with the stress characterization of the neighbouring site of Soultz-sous-Forêts, where measurements have highlighted a normal faulting regime in the top granitic layers evolving into a strike slip regime more in depth. The uncertainty about our measurements and about the strength parameterization does not allow, however, for a decision on the faulting regime and its evolution with depth in Rittershoffen. A step in *SH* magnitude is visible on our estimate in Fig. 18 at large depth (below ~~2250m~~2250 m). This step occurs at the interface sediment basement and could be explained by the effect of stiffness contrast between lithologies (Corkum et al., 2018).

10. Conclusion

Thanks to the repeated UBI logging of the geothermal wells GRT-1 and GRT-2 in Rittershoffen (France), this study focuses on the analysis of the evolution with time and depth of the borehole breakouts. The following conclusions are drawn:

(i) Clear evidences of time evolution of the breakout exist in particular in the sedimentary cover.

(ii) The evolution in time of the vertical length and the horizontal width of the breakouts are measured benefiting from the development of a UBI image correlation technique. It is discussed in the limit of the estimated uncertainties. The vertical length of the breakouts is shown to increase with time. No variation in the depth of the breakouts in the radial direction was observed within the limit of the uncertainty of our analysis. However, width increases beyond the uncertainty of our determination were highlighted. This contradict the common assumption that breakouts do not widen but only deepen until the borehole reach a new stable state (Zoback et al. 2003~~);~~).

(iii) The changes in breakout width occur between datasets collected prior and after reservoir stimulation~~.~~, taking place in 2013. However, the most likely effect on breakout width is the thermal equilibration of the wellbore and our data do not evidence stress changes result from reservoir stimulation~~;~~.

In addition to this analysis, the study of the geometry of borehole failures in both wells leads to propose a characterization of the *in-situ* stress tensor at depths including the orientation and the magnitude of the three principal stresses. This detailed stress state analysis includes the estimation of thermal stresses. A Mohr-Coulomb criterion is retained here to estimate the principal stresses magnitude as it is in our parametrization, the most consistent with a frictional strength limit in the crust. The strength parameterization is however uncertain due to the lack of mechanical testing on the Rittershoffen reservoir rocks. Given the uncertainties, we propose the following careful interpretation of our measurements:

(i) Our analyses of the breakout geometry variation with depth suggest a change in mean orientation, with a 30° rotation from NNW to NNE highlighted between the bottom and the top of our analysed section. This observation is robust and independent of the strength parameterisation. The rotation does not occur at the sediment-basement interface but is related to a high steeply dipping major fault zone crossing the GRT-1 borehole at a depth of 2368 m (Vidal et al., 2016).

(ii) Our results suggest also a step in horizontal stress magnitude at the sediment to basement transition that would be consistent with stiffness contrast between these two lithologies. However, such step is determined by the choice of the failure criterion and its parameterization which is uncertain at Rittershoffen.

(iii) *SH* is generally slightly larger than the vertical principal stresses *Sv* consistently with a strike-slip to normal faulting transitional regime. This is consistent with stress characterization at the neighbour site of Soultz-sous-Forêts (Cornet et al., 2007; Klee & Rummel, 1993; Valley & Evans, 2007b)

The Rittershoffen borehole imaging dataset is unique by the fact that repeating logging allowed to study the temporal evolution of borehole breakouts and the possible stress changes induced by reservoir stimulation. Our results change the common view that breakouts mostly deepen but do not widen. Further work is however required to reduce the uncertainties related to stress magnitude estimates from borehole breakouts and to be able to quantify stress changes induced by reservoir stimulation.

Availability of data and materials

Due to the industrial property of the borehole datasets, raw data would remain confidential and would not be shared.

Competing interests

The authors declare no competing financial interest.

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Appendix A:

The Kirsch equations are derived under 2D plane conditions. They provide stress values in a case which is not suited to the one of real deviated boreholes, in which out of plane normal and shear stresses also exist. We consider two Cartesian co-ordinate frames: x–y–z having z aligned with the vertical and x'–y'–z' which is aligned with the three principal stresses noted [$\sigma_{x'x'}$, $\sigma_{y'y'}$, $\sigma_{z'z'}$] respectively. We consider a long cylindrical cavity of radius a. Its axis is arbitrarily oriented with respect to the principal stress state in the Earth. The borehole axis tilts at an angle ϕ relative to the x-axis. The third cylindrical r–θ–ζ co-ordinate frame is borehole centric with the ζ axis which is co-incident with the borehole axis. The azimuth with respect to the borehole axis is noted θ.

The borehole centric stresses are expressed in function of the direction cosines *aij* enabling to transform the principal axes x'–y'–z' to the x–y–z frame, accordingly to Eq. (A1):

$$\sigma' = A \cdot \sigma \cdot A^T \qquad \text{(A1)}$$

where the rotation matrix A is composed of the direction cosines *aij*

$$A = \begin{bmatrix} axx' & axy' & axz' \\ ay'x' & ayy' & ayz' \\ az'x & az'y & azz' \end{bmatrix}$$

Eqs. (A2-A7) express the borehole centric stresses as a function of directional coefficients $\alpha 1$, $\alpha 2$, $\alpha 3$, $\gamma 1$ and $\gamma 2$. They include the contribution of fluid pressure Pf. Indeed, the pressure of the fluid in the mud column increases with depth, which produces tensile hoop stress and compressive radial stress. Eqs. (A2-A7) also include the time-dependant contribution due to temperature changes. The thermal stresses $\sigma^{\Delta T}{}_{\theta}$ and $\sigma^{\Delta T}{}_{r}$ resulting from the temperature difference, Δt, between the temperature applied at the borehole wall and the initial temperature at that depth before perturbation or the temperature at a significant distance from the borehole (not influenced by the borehole perturbation), are expressed from Voight & Stephens (1982). The radial component is null, and the tangential component expressed in Eq. (8) shows that an increase in temperature at $r=a$ effects the compressive hoop stress.

$$\sigma_{rr} = Pf + \sigma^{\Delta T}{}_{r} \qquad \text{(A2)}$$

$$\sigma_{\theta\theta} = 2\,\alpha 1 - 4\,\alpha 2 \cos 2\theta - 4\,\alpha 3 \sin 2\theta - Pf + \sigma^{\Delta T}{}_{\theta} \qquad \text{(A3)}$$

$$\sigma_{\zeta\zeta} = \beta 1 - 4\,\nu\,(\alpha 2 \cos 2\theta + \alpha 3 \sin 2\theta) \qquad \text{(A4)}$$

$$\tau_{\theta\zeta} = 2\,\gamma 1 \cos\theta + 2\,\gamma 2 \sin\theta \qquad \text{(A5)}$$

$$\tau_{r\zeta} = 0 \qquad \text{(A6)}$$

$$\tau_{\theta r} = 0 \qquad \text{(A7)}$$

The geometrical coefficients involved in Eqs. (A2-A7) are expressed as a function of the three far-field principal stress state $[\sigma_{x'x'}, \sigma_{y'y'}, \sigma_{z'z'}]$ and as a function of the geometrical rotations *aij* :

$$\alpha 1 = \tfrac{1}{2}\,[\,(a^2_{x'x} \sin^2\Phi + a^2_{x'y} + a^2_{x'z} \cos^2\Phi - 2\,a^2_{x'z}\,a^2_{x'x} \sin\Phi\cos\Phi)\,\sigma_{x'x'} + (a^2_{y'x} \sin^2\Phi + a^2_{y'y} + a^2_{y'z} \cos^2\Phi - 2\,a^2_{y'z}\,a^2_{y'x} \sin\Phi\cos\Phi)\,\sigma_{y'y'} + (a^2_{z'x} \sin^2\Phi + a^2_{z'y} + a^2_{z'z} \cos^2\Phi - 2\,a^2_{z'z}\,a^2_{z'x} \sin\Phi\cos\Phi)\,\sigma_{z'z'}\,] \qquad \text{(A8)}$$

$$\alpha 2 = \tfrac{1}{2}\,[\,(-a^2_{x'x} \sin^2\Phi + a^2_{x'y} - a^2_{x'z} \cos^2\Phi + 2\,a^2_{x'z}\,a^2_{x'x} \sin\Phi\cos\Phi)\,\sigma_{x'x'} + (-a^2_{y'x} \sin^2\Phi + a^2_{y'y} - a^2_{y'z} \cos^2\Phi + 2\,a^2_{y'z}\,a^2_{y'x} \sin\Phi\cos\Phi)\,\sigma_{y'y'} + (-a^2_{z'x} \sin^2\Phi + a^2_{z'y} - a^2_{z'z} \cos^2\Phi + 2\,a^2_{z'z}\,a^2_{z'x} \sin\Phi\cos\Phi)\,\sigma_{z'z'}\,] \qquad \text{(A9)}$$

$$\alpha 3 = (a_{x'y}\,a_{x'z} \cos\Phi - a_{x'x}\,a_{x'y} \sin\Phi)\,\sigma_{x'x'} + (a_{y'y}\,a_{y'z} \cos\Phi - a_{y'x}\,a_{y'y} \sin\Phi)\,\sigma_{y'y'} + (a_{z'y}\,a_{z'z} \cos\Phi - a_{z'x}\,a_{z'y} \sin\Phi)\,\sigma_{z'z'} \qquad \text{(A10)}$$

$$\gamma 1 = [-a^2_{x'x} \sin\Phi \cos\Phi + a^2_{x'z} \cos\Phi \sin\Phi + a_{x'z}\, a_{x'x} (\cos^2\Phi - \sin^2\Phi)]\, \sigma_{x'x'} + [-a^2_{y'x} \sin\Phi \cos\Phi + a^2_{y'z} \cos\Phi \sin\Phi + a_{y'z}\, a_{y'x} (\cos^2\Phi - \sin^2\Phi)]\, \sigma_{y'y'} + [-a^2_{z'x} \sin\Phi \cos\Phi + a^2_{z'z} \cos\Phi \sin\Phi + a_{z'z}\, a_{z'x} (\cos^2\Phi - \sin^2\Phi)]\, \sigma_{z'z'}\,] \qquad \text{(A11)}$$

$$\gamma 2 = (-a_{x'y}\, a_{x'z} \sin\Phi - a_{x'x}\, a_{x'y} \cos\Phi)\, \sigma_{x'x'} + (-a_{y'y}\, a_{y'z} \sin\Phi - a_{y'x}\, a_{y'y} \cos\Phi)\, \sigma_{y'y'} + (-a_{z'y}\, a_{z'z} \sin\Phi - a_{z'x}\, a_{z'y} \cos\Phi)\, \sigma_{z'z'} \qquad \text{(A12)}$$

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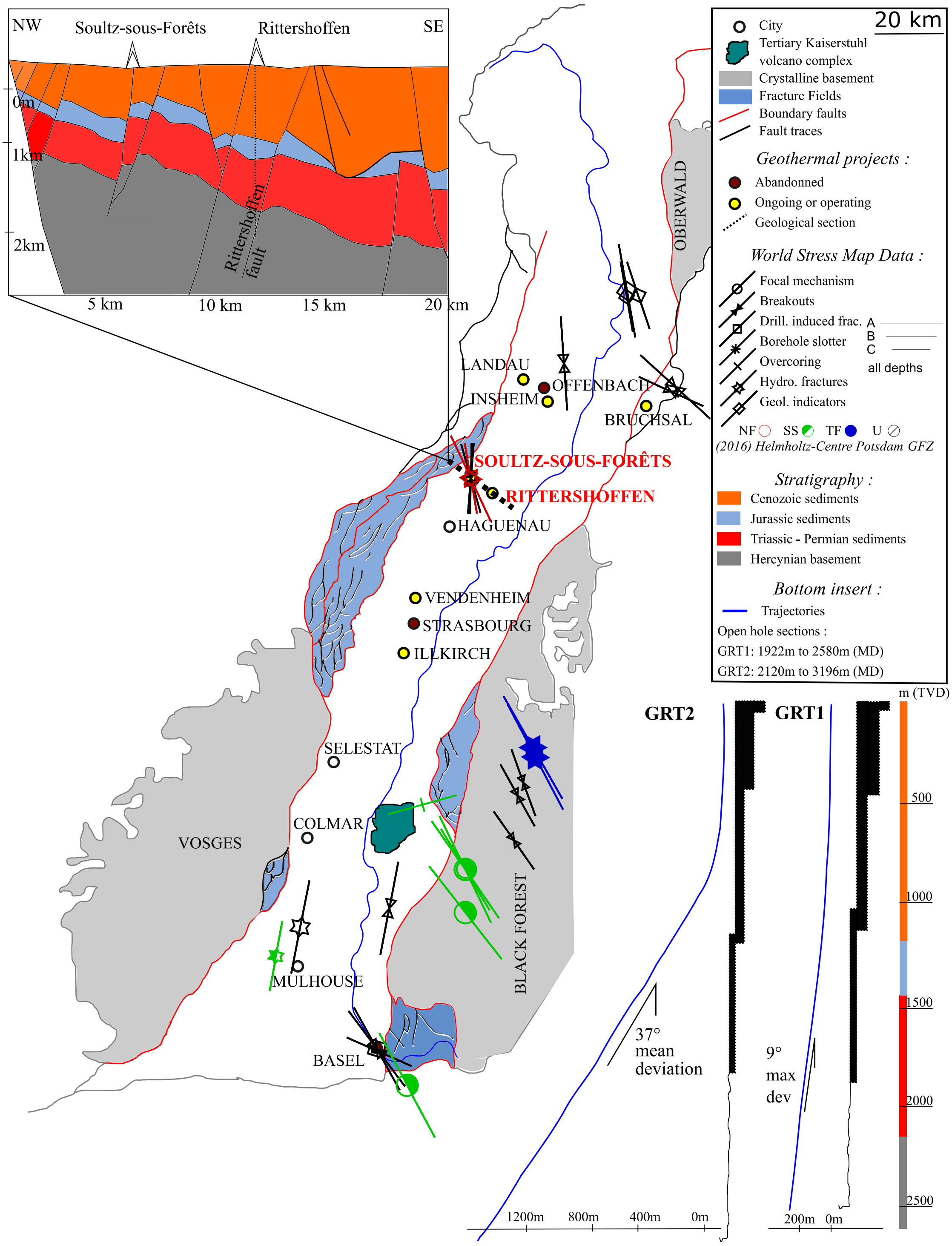
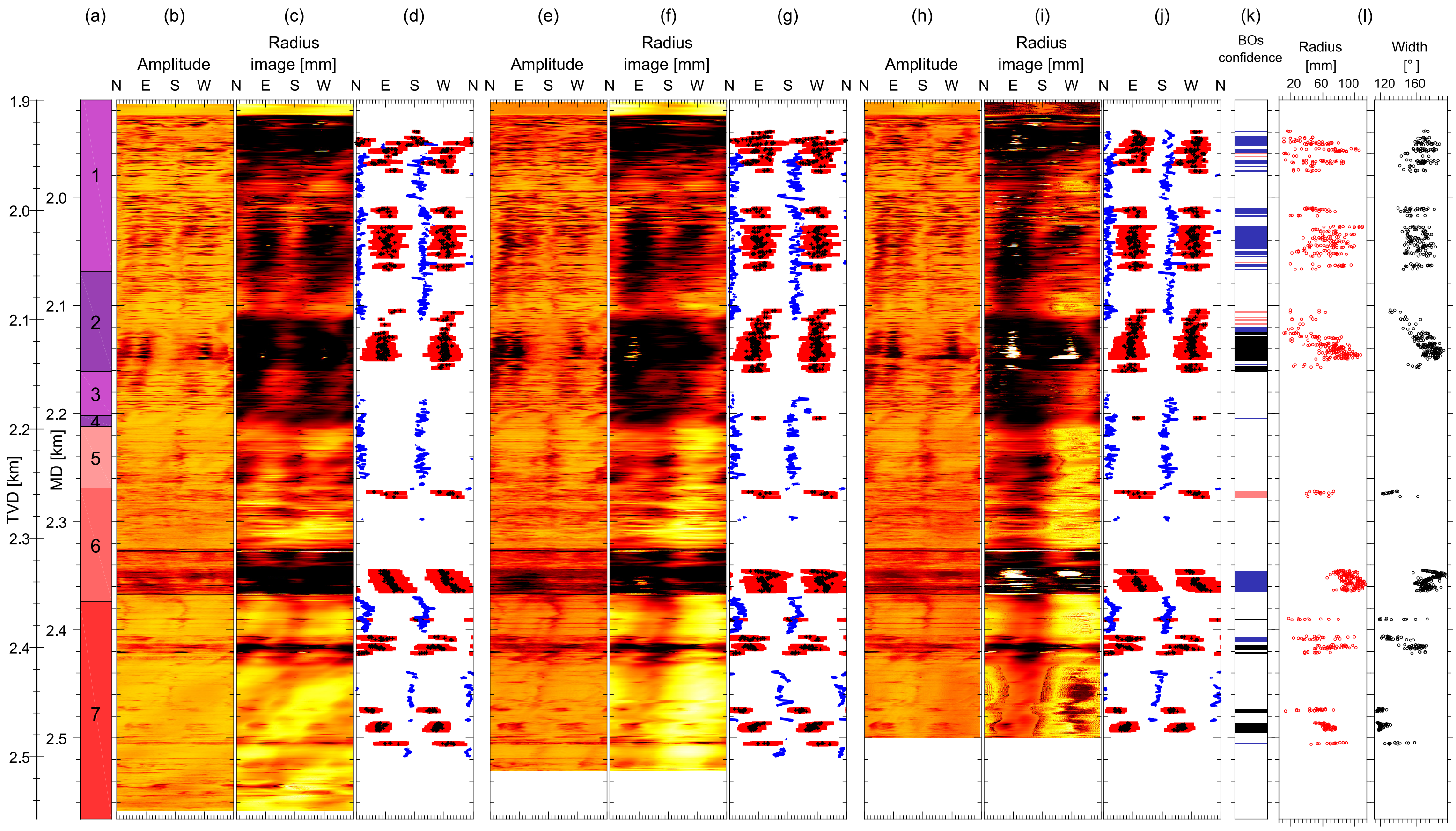


NW
Soultz-sous-Forêts
Rittershoffen
SE
0m
1km
2km
Rittershoffen fault
5 km
10 km
15 km
20 km
20 km
City
Tertiary Kaiserstuhl volcano complex
Crystalline basement
Fracture Fields
Boundary faults
Fault traces
Geothermal projects :
Abandonned
Ongoing or operating
Geological section
World Stress Map Data :
Focal mechanism
Breakouts
Drill. induced frac.
Borehole slotter
Overcoring
Hydro. fractures
Geol. indicators
A
B
C
all depths
NF
SS
TF
U
(2016) Helmholtz-Centre Potsdam GFZ
Stratigraphy :
Cenozoic sediments
Jurassic sediments
Triassic - Permian sediments
Hercynian basement
Bottom insert :
Trajectories
Open hole sections :
GRT1: 1922m to 2580m (MD)
GRT2: 2120m to 3196m (MD)
LANDAU
INSHEIM
OFFENBACH
BRUCHSAL
OBERWALD
SOULTZ-SOUS-FORÊTS
RITTERSHOFFEN
HAGUENAU
VENDENHEIM
STRASBOURG
ILLKIRCH
SELESTAT
COLMAR
VOSGES
MULHOUSE
BASEL
BLACK FOREST
GRT2
GRT1
m (TVD)
500
1000
1500
2000
2500
37° mean deviation
9° max dev
1200m
800m
400m
0m
200m
0m

Figure 1: Geological and structural map of the main of the Upper Rhine Graben with the location of the Rittershoffen and Soultz-sous-Forêts sites. The map shows also the location and status of other neighbouring deep geothermal projects. ~~Stress~~ It includes stress data from World stress map database (Heidbach et al., ~~2010~~2016) ~~are included.~~. Upper left insert shows a geological section highlighting the main units crossed by the wells in Rittershoffen and Soultz-sous-Forêts (Aichholzer et al., 2016; Baujard et al., 2017). Lower right insert is a sketch of wells GRT-1 and GRT-2 drilled in Rittershoffen~~, including~~ which includes their geometry, depths and crossed lithology (after Baujard et al. (2015, 2017)).



(a)
(b)
(c)
(d)
(e)
(f)
(g)
(h)
(i)
(j)
(k)
(l)
Amplitude
Radius image [mm]
Amplitude
Radius image [mm]
Amplitude
Radius image [mm]
BOs confidence
Radius [mm]
Width [°]
N E S W N
20 60 100
120 160
TVD [km]
MD [km]
1.9
2.0
2.1
2.2
2.3
2.4
2.5
1
2
3
4
5
6
7

Figure 2: Synthesis of the data used in this analysis of the borehole GRT-1. The measurements are expressed in function of Measured Depth (MD) and Vertical Depth (TVD). (a) simplified lithologic column. (1) stands for "couches de Rehberg", (2) for "Couches de Trifels", (3) for Annweiler sandstone, (4) for Permian layers older than Annweiler sandstone, (5) for rubefied granite, (6) for hydrothermally altered granite and (7) for low altered granite. The UBI images are presented, as well as the data picked from the visual analysis of the double transit time image for the dataset of 2012 (panel b. - c. - d.), 2013 (e. - f. – g.), and 2015 (h. - i. - j.) collected in GRT-1. The radius of the borehole computed from the double transit time image is displayed in panels b. - e. and h. In panels d. - g. and j., blue dots represent the azimuth of the Drilling Induced Tension Fractures (DITFs), black dots represent the azimuth of the maximal radial depth of the breakouts and red bars represent the extension between the edges of the breakouts. Panel k. informs about the breakouts (BOs) confidence level applied to these results. Panel l. summarizes the width (black dots, in °) and the enlargement radius (red dots, in mm) measured in the 2012, 2013 and 2015 images ~~and panel k. informs about the breakouts (BOs) confidence level applied to these results.~~.

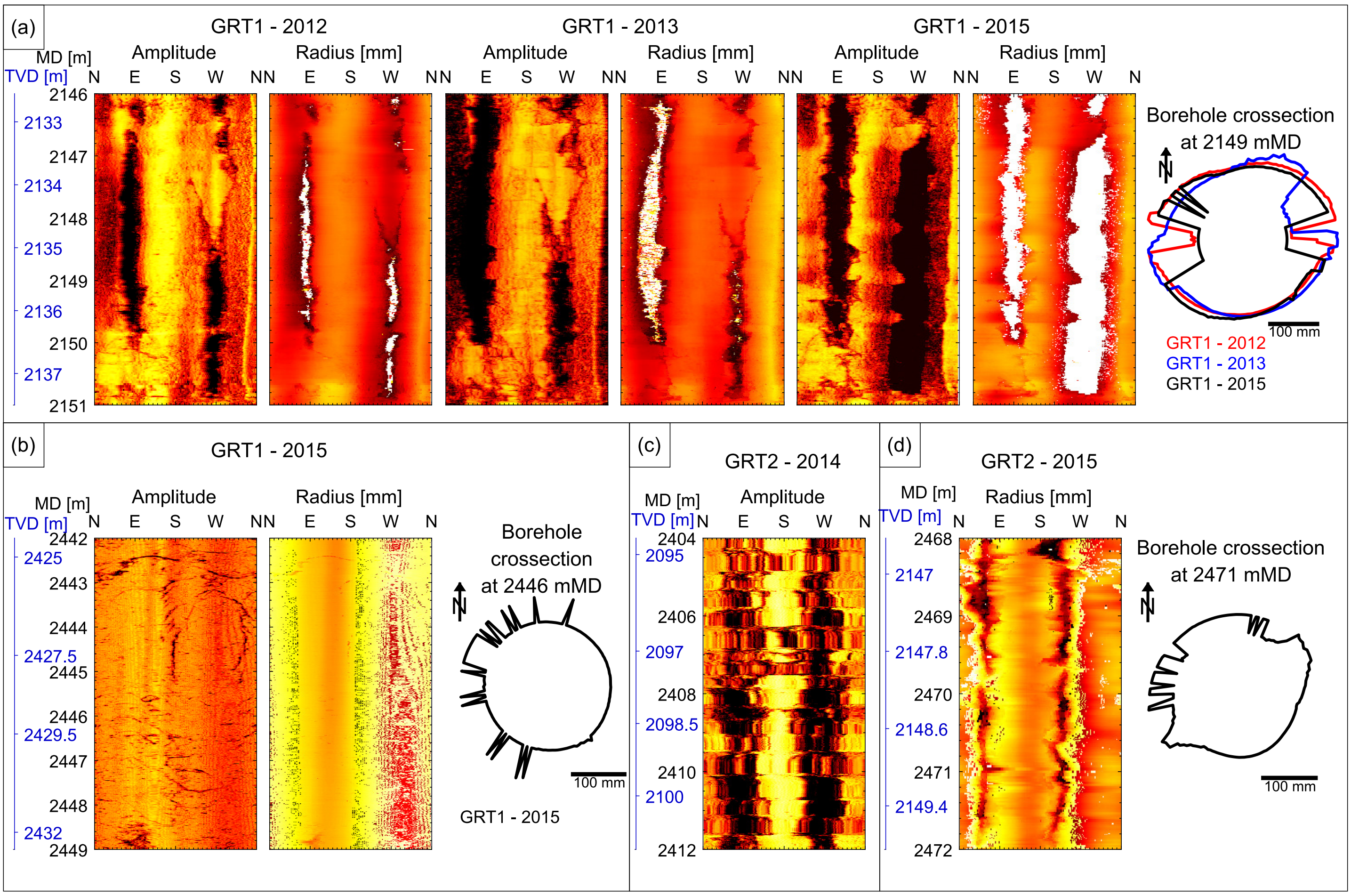

(a)
GRT1 - 2012
GRT1 - 2013
GRT1 - 2015
MD [m]
TVD [m]
Amplitude
Radius [mm]
N E S W N
Borehole crossection at 2149 mMD
100 mm
GRT1 - 2012
GRT1 - 2013
GRT1 - 2015
(b)
GRT1 - 2015
Borehole crossection at 2446 mMD
(c)
GRT2 - 2014
(d)
GRT2 - 2015
Borehole crossection at 2471 mMD

Figure 3: Example of image artefact observed on the GRT-1 and GRT-2 data set. a) Comparison of data from 2012, 2013 and 2015 collected in GRT-1 presenting a signal loss artefact in sandstones~~.~~, clearly highlighted by persisting white patches in the radius signal. b) Processing noise resembling to wood grain ~~artefact~~textures, visible on the 2015 GRT-1 image, both on the amplitude and radius image in granite. ~~This leads to noisy borehole section. c)~~c) Alternating compression and stretching of the image characteristic of stick-slip ~~artefact present~~artefacts, highlighted along the entire GRT-2-2014 image. d) Erroneous radius record observable on the GRT-2-2015 image in granite, possibly related to tool decentralization.

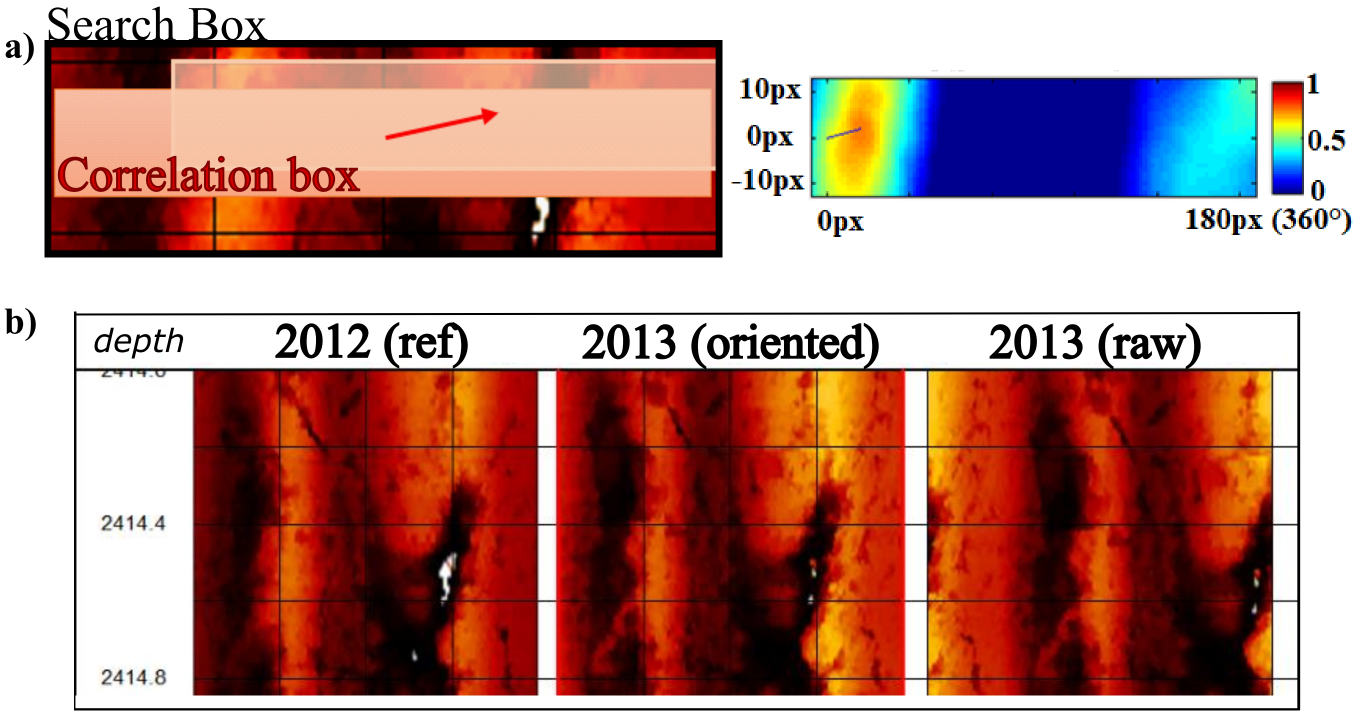
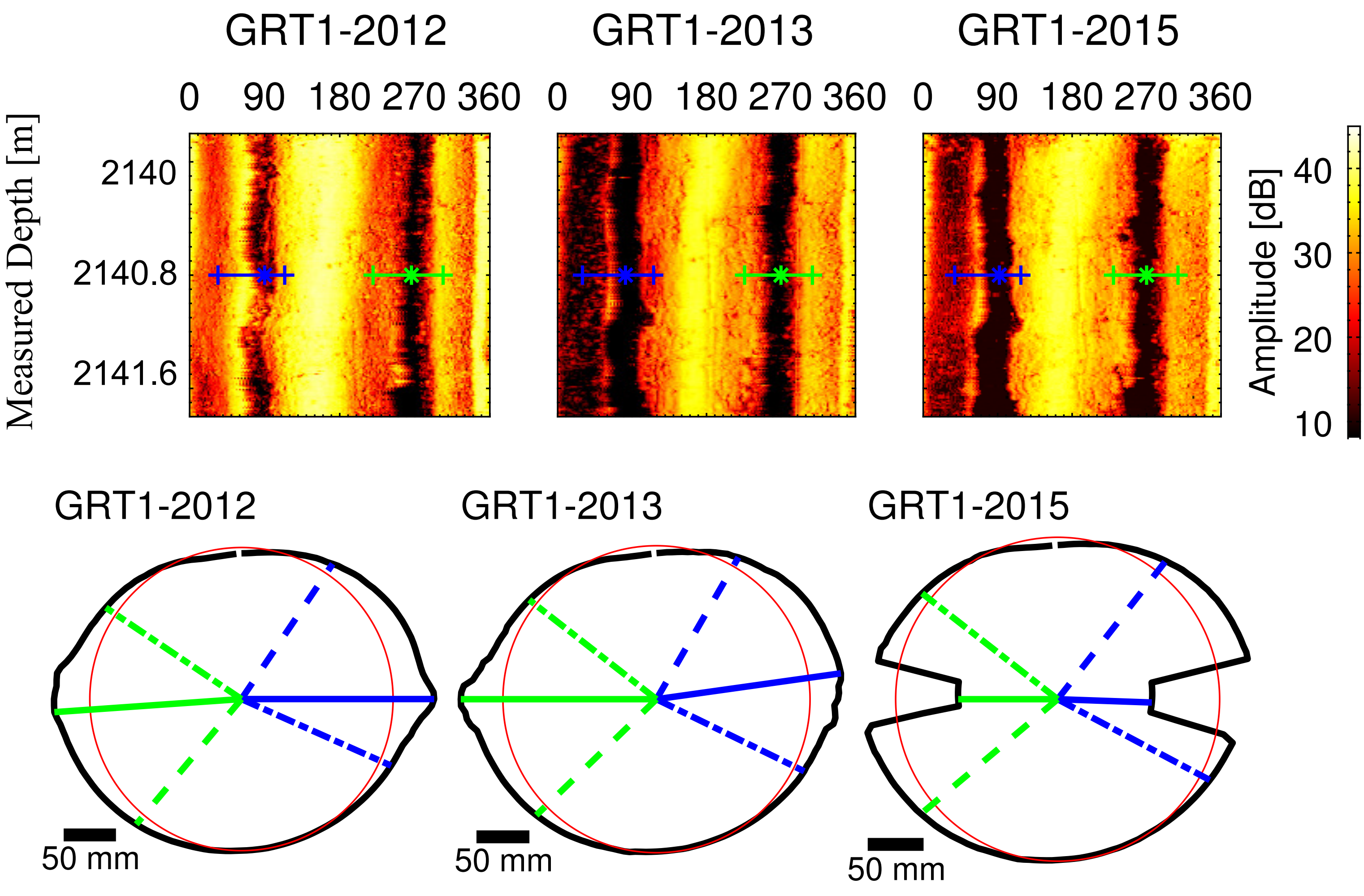
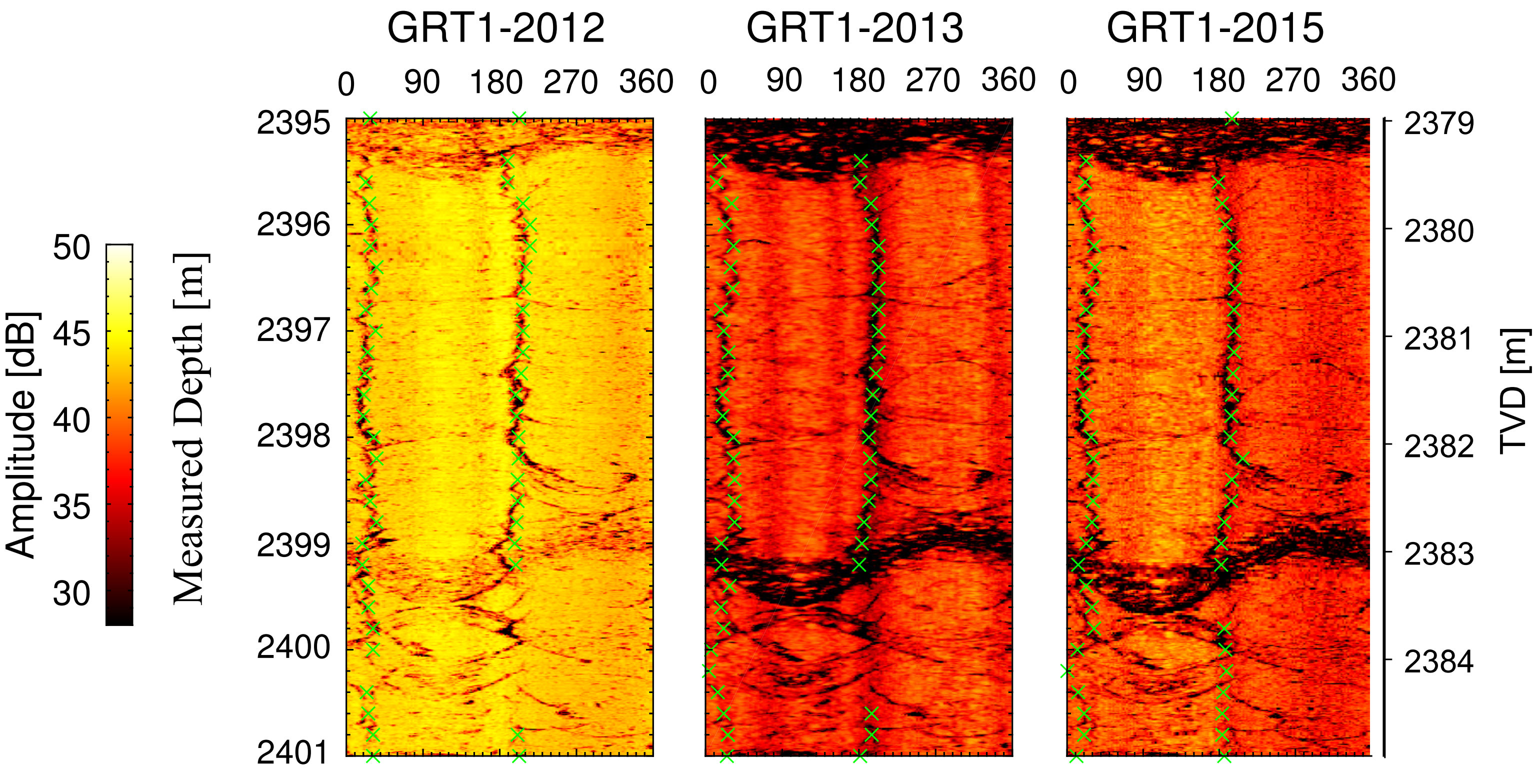

a)
Search Box
Correlation box
10px
0px
-10px
0px
180px (360°)
1
0.5
0
b)
depth
2012 (ref)
2013 (oriented)
2013 (raw)
2414.4
2414.8

Figure 4: a) Sketch presenting the process used to orientate the images of GRT-1. A correlation box is defined in the double transit time image of reference (acquired in 2012) and is progressively shifted in the image to compare with (red windows) within the limits of the search box (black window). We compute the correlation between the correlation box in its initial position in the image of reference and the shifted correlation box in the image to compare with for each position (right insert). The displacement maximizing the correlation factor enables, at a given depth, to rotate and adapt the image of 2013 and 2015 according to the image of 2012. b) example of original and reoriented time transit images of 2013, at a depth of 2414m (TVD) in GRT-1.



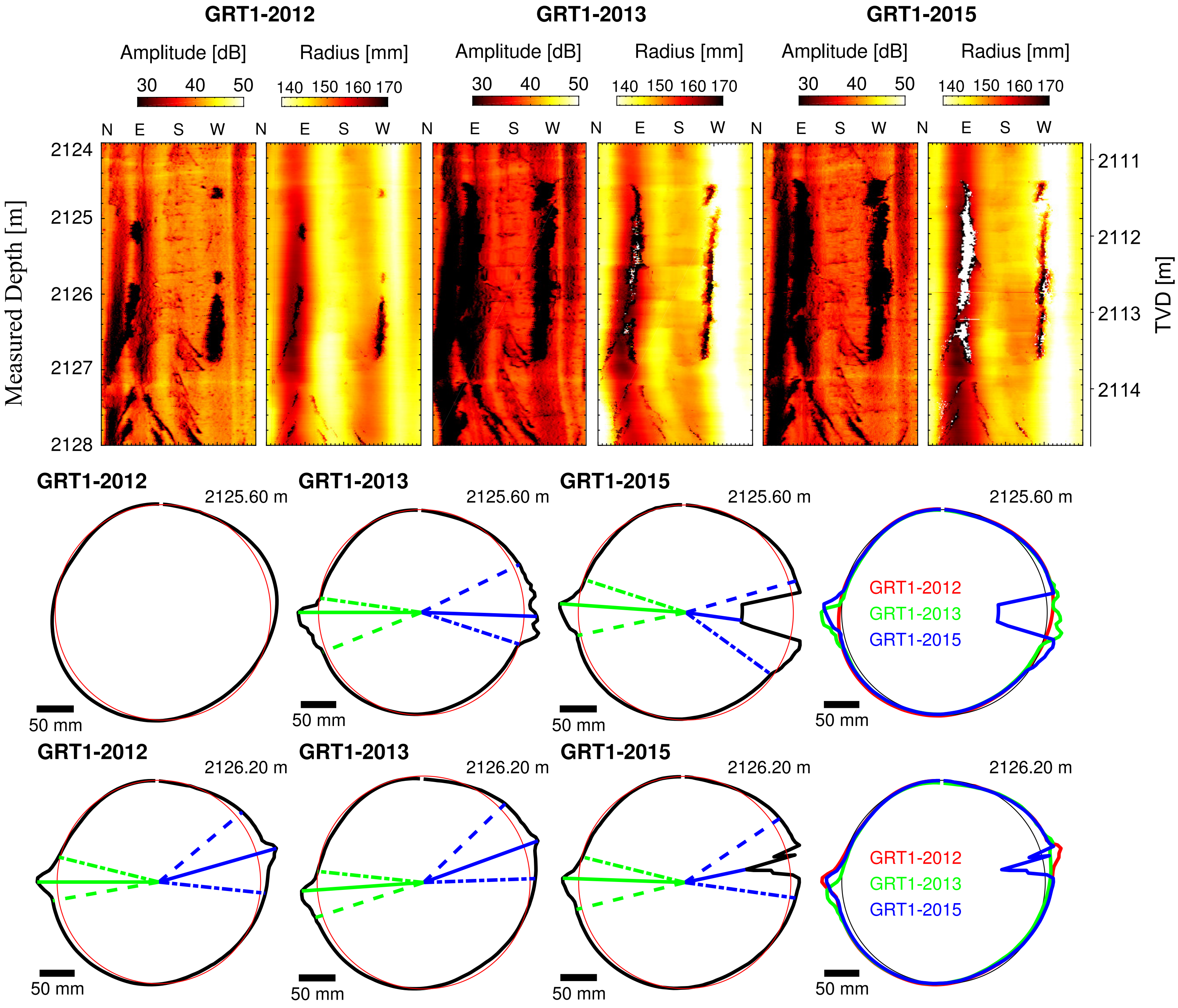
GRT1-2012
GRT1-2013
GRT1-2015
0 90 180 270 360
Measured Depth [m]
2140
2140.8
2141.6
Amplitude [dB]
40
30
20
10
GRT1-2012
GRT1-2013
GRT1-2015
50 mm

Figure 5: Example of breakout geometry determination in sandstones. Upper figures: amplitude images for GRT-1 at 2140.8 m for the logs from 2012, 2013 and 2015. Lower figures: wellbore section at 2140.8 m computed from the transit time images from the 2012, 2013 and 2015 logs respectively. The breakout extent is determined on the wellbore section. The blue and green dashed lines represent the extent of the breakout when the plain lines represent the azimuth of maximum radial extension of the breakout.



GRT1-2012
GRT1-2013
GRT1-2015
0 90 180 270 360
2395
2396
2397
2398
2399
2400
2401
2379
2380
2381
2382
2383
2384
Amplitude [dB]
50
45
40
35
30
Measured Depth [m]
TVD [m]

Figure 6: Examples of Drilling Induced Tension Fractures (DITFs), observed in the granitic ~~setion~~section of GRT-1 in the amplitude images acquired in 2012, 2013 and 2015. The azimuth of the DITFs is measured every 20 cm (green crosses).



GRT1-2012
GRT1-2013
GRT1-2015
Amplitude [dB]
Radius [mm]
30 40 50
140 150 160 170
N E S W N
Measured Depth [m]
2124
2125
2126
2127
2128
TVD [m]
2111
2112
2113
2114
2125.60 m
2126.20 m
50 mm
GRT1-2012
GRT1-2013
GRT1-2015

Figure 7: Examples of breakout shape evolution between the three successive images collected in GRT-1 in sandstones. Upper figures show the amplitude images and the radius computed from the time transit images for a section of GRT-1 from 2124 to 2128m (~~TVD~~MD) in 2012, 2013 and 2015. Lower figures show the mean section computed at 2125.6 and 2126.2m (~~TVD~~MD) from the time transit images averaged over 60cm intervals. The sections are represented along with an 8.5 inch radius circle representing the unaltered open hole section. The sections from the image of 2012, 2013 and 2015 are superposed in the right panel.

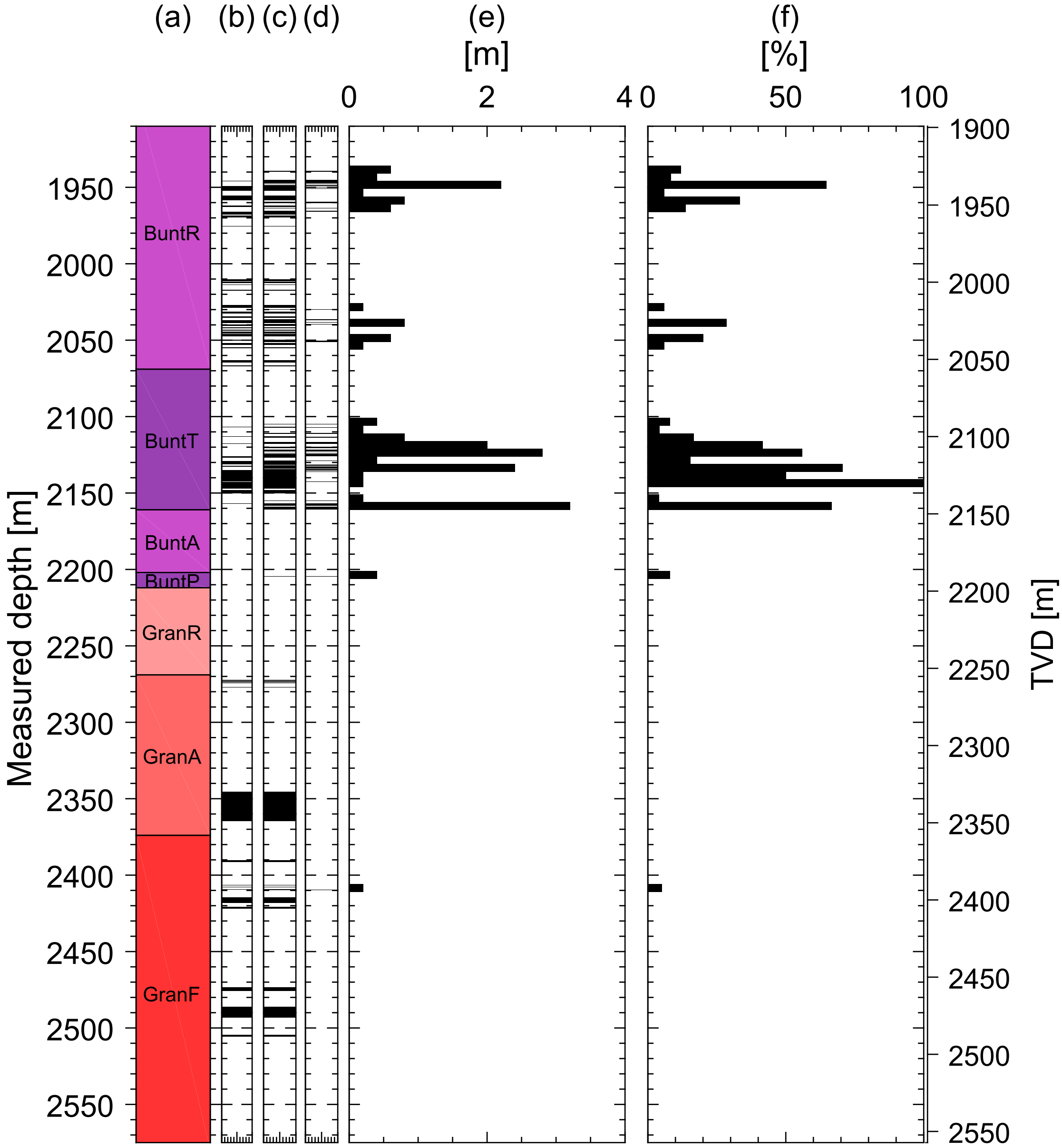
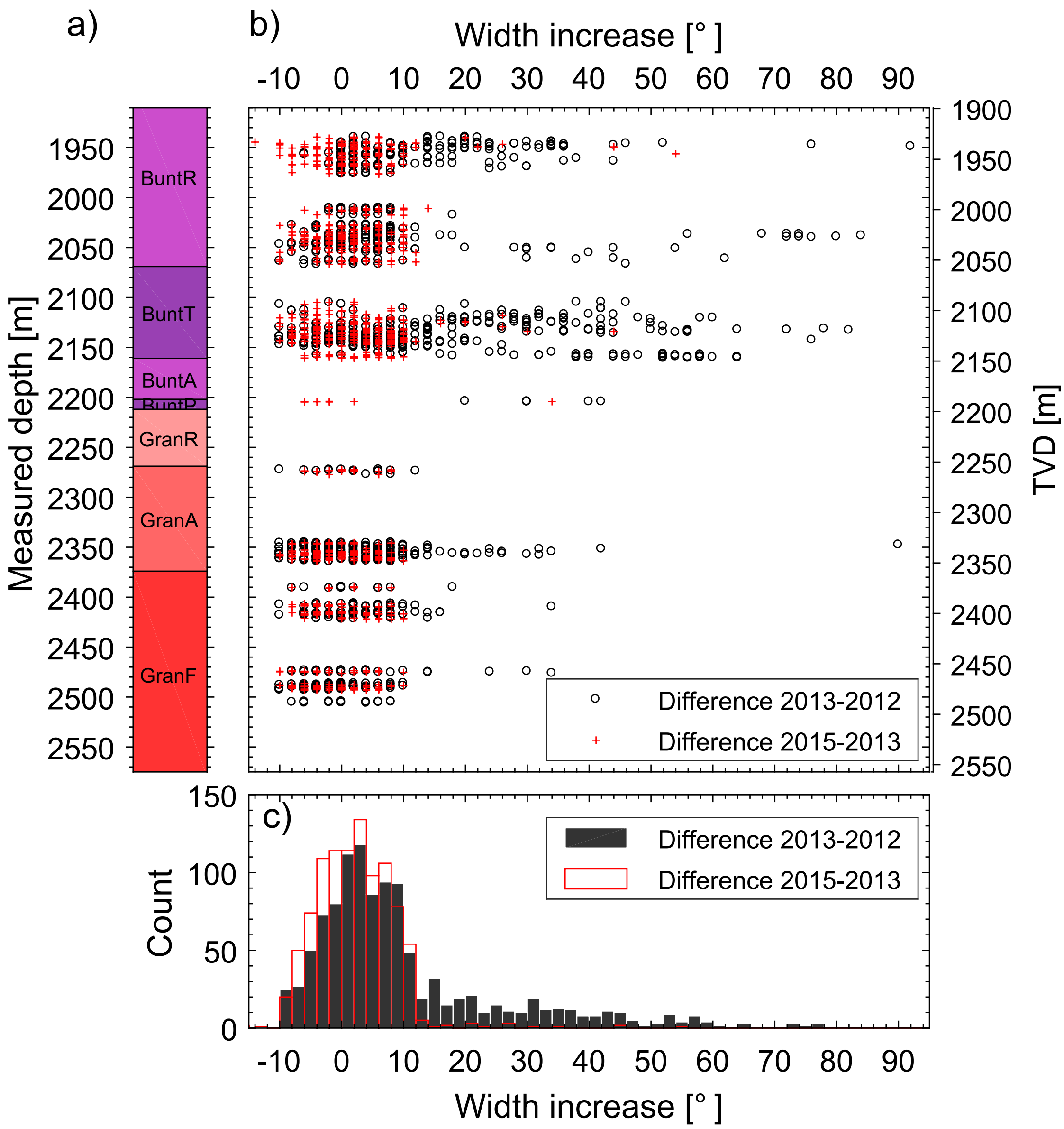

(a)
(b)
(c)
(d)
(e)
[m]
(f)
[%]
0
2
4
0
50
100
Measured depth [m]
1950
2000
2050
2100
2150
2200
2250
2300
2350
2400
2450
2500
2550
BuntR
BuntT
BuntA
BuntP
GranR
GranA
GranF
TVD [m]
1900
1950
2000
2050
2100
2150
2200
2250
2300
2350
2400
2450
2500
2550

Figure 8: Development of breakouts along GRT-1 borehole between 2012 and 2013. a) Simplified lithologies along GRT-1 borehole ~~with~~in function of Measured Depth (MD) ~~:~~or Vertical Depth (TVD). BuntR stands for "couches de Rehberg", BuntT for "Couches de Trifels", BuntA for Annweiler sandstone, BuntP for Permian layers older than Annweiler sandstone, GranR for rubefied granites, GranA for hydrothermally altered granite and GranF for low altered granite. The major fault zone crossing GRT-1 at 2368m is represented as a black band. b) Breakouts positions in GRT-1 in 2012. c) Breakouts positions in GRT-1 in 2013. d) Intervals where breakouts are present in 2013 but not in 2012. e) Breakout length increase in [m] along the borehole between 2012 and 2013 in 5 m bins. f) fraction in [%] of wellbore length that was free of breakout in 2012 that is presenting breakout on the 2013 image, computed in 5 m bins.

Figure 9: Evolution of breakout width in GRT-1 borehole ~~with~~in function of Measured Depth (MD) or Vertical Depth (TVD). a) Simplified lithologies along GRT-1 borehole (see Fig. 8 for the legend). b) Width increase between the 2012-13 time interval (black circles) and the 2013-15 time interval (red crosses) presented as a function of the vertical depth. c) histograms in 2° classes of breakout width changes for the 2012-13 interval (black) and the 2013-15 interval (red).



a)
b)
Width increase [°]
-10
0
10
20
30
40
50
60
70
80
90
1900
1950
2000
2050
2100
2150
2200
2250
2300
2350
2400
2450
2500
2550
TVD [m]
Measured depth [m]
BuntR
BuntT
BuntA
GranR
GranA
GranF
Difference 2013-2012
Difference 2015-2013
c)
Count
150
100
50
0
Width increase [°]

Figure 10: Evolution of the depth of the breakouts in the GRT-1 borehole ~~with~~in function of Measured Depth (MD) or Vertical Depth (TVD). a) Simplified lithologies along GRT-1 borehole (see Fig. 8 for the legend). b) Increase of the maximum radial extension between the 2012-13 time interval (black circles) and 2013-15 time interval (red crosses) presented in function of depth. c) histograms in 2 mm classes of breakout with changes for the 2012-13 interval (black) and 2013-15 interval (red).

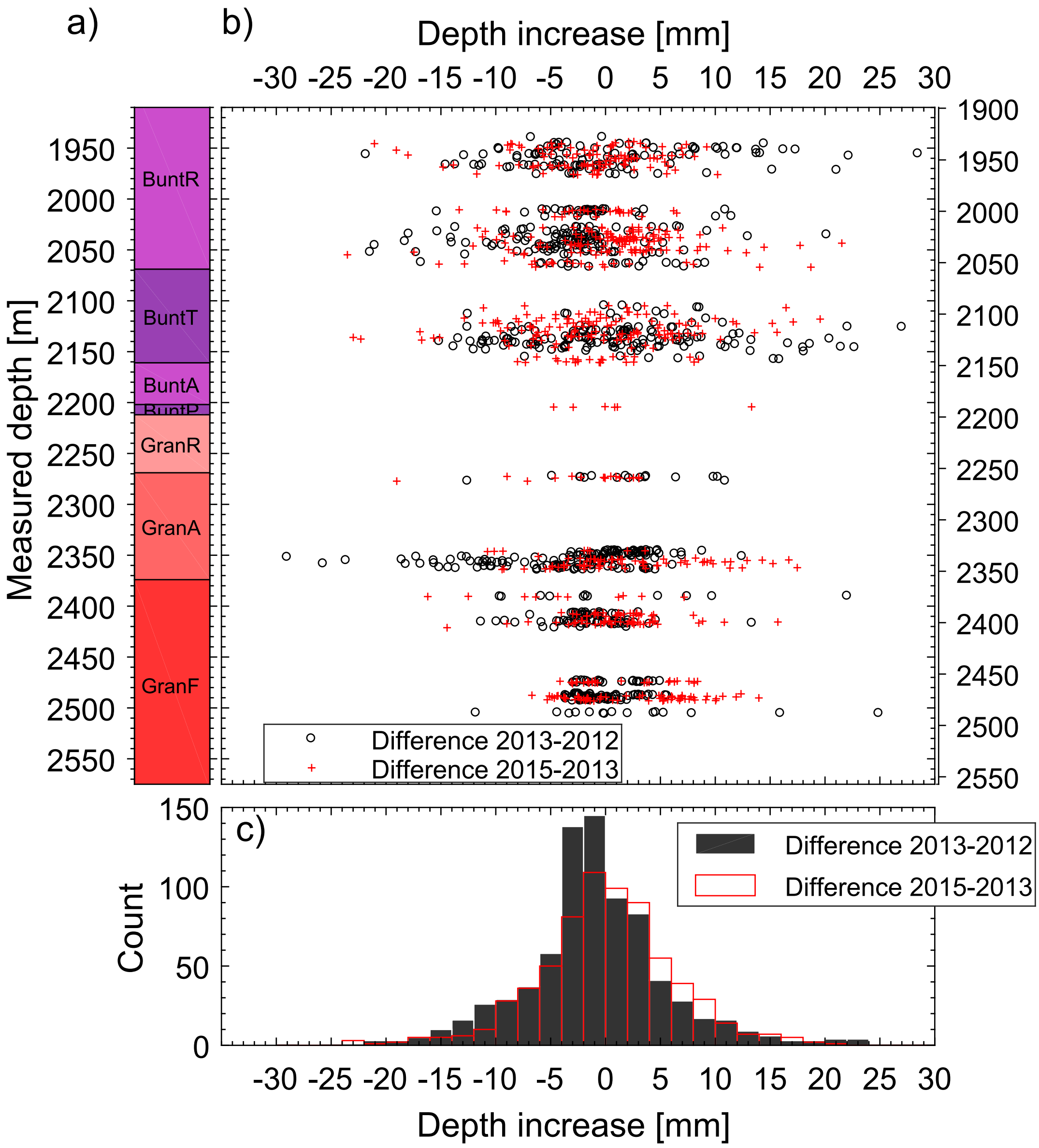
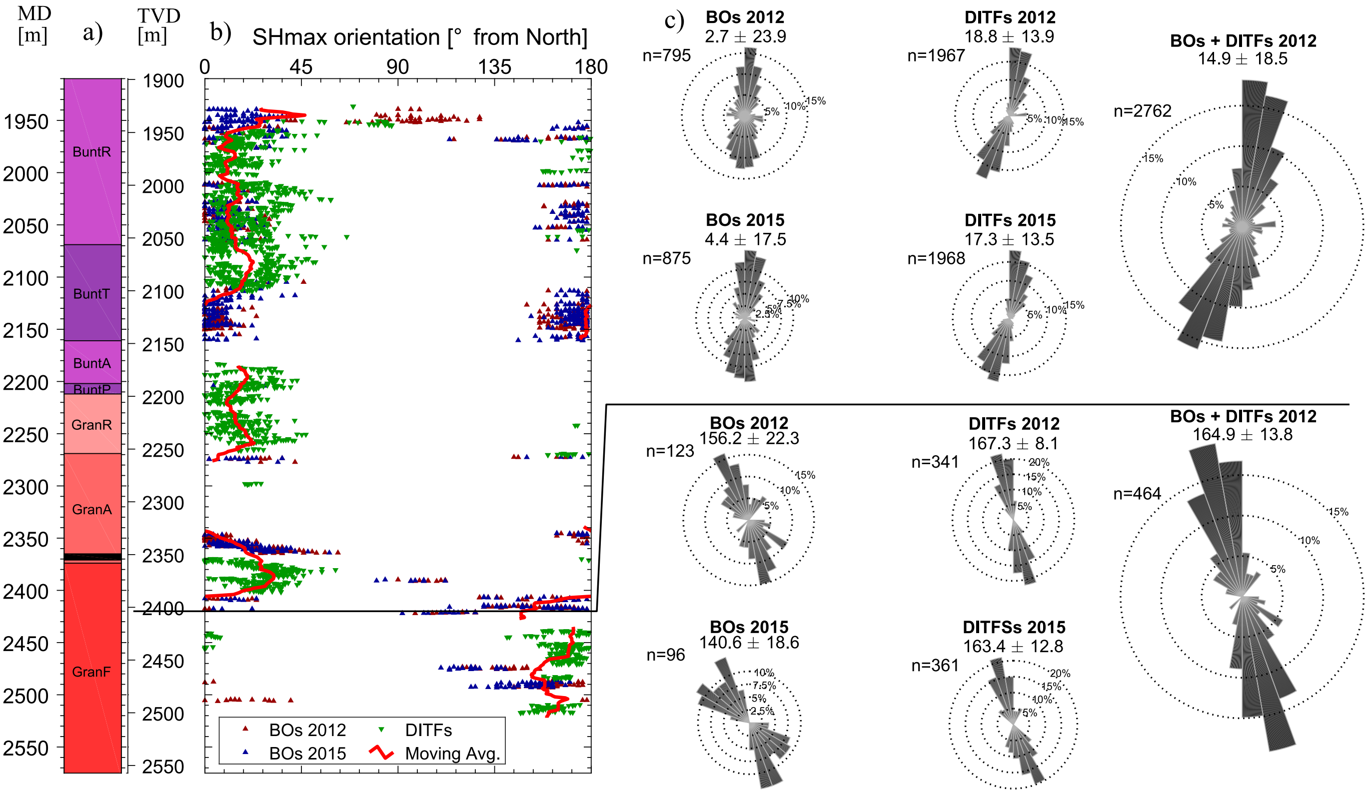

a)
b)
Depth increase [mm]
-30 -25 -20 -15 -10 -5 0 5 10 15 20 25 30
Measured depth [m]
1950
2000
2050
2100
2150
2200
2250
2300
2350
2400
2450
2500
2550
BuntR
BuntT
BuntA
GranR
GranA
GranF
TVD [m]
1900
Difference 2013-2012
Difference 2015-2013
c)
Count
150
100
50
0
Difference 2013-2012
Difference 2015-2013
-30 -25 -20 -15 -10 -5 0 5 10 15 20 25 30
Depth increase [mm]

Figure 11: Evolution in orientation of the maximum principal stress ~~with measured depth~~in function of Measured Depth (MD) and Vertical Depth (TVD) in GRT-1, in 2012 and 2015. a) Simplified lithologies along GRT-1 borehole (see Fig. 8 for the legend). b) Orientation of *SH* from the azimuth of maximum radial extension of the breakouts (BOs) from the dataset of 2012 (in blue) and of 2015 (in red) acquired in GRT-1. In green, orientation of *SH* from the azimuth of Drilling Induced Tensile Fractures (DITFs). The red line is a moving average of the orientation data. c) From the datasets displayed in panel b), orientation in rose diagrams.



MD [m]
a)
TVD [m]
b)
SHmax orientation [° from North]
0
45
90
135
180
BuntR
BuntT
BuntA
BuntP
GranR
GranA
GranF
BOs 2012
BOs 2015
DITFs
Moving Avg.
c)
BOs 2012
2.7 ± 23.9
n=795
DITFs 2012
18.8 ± 13.9
n=1967
BOs + DITFs 2012
14.9 ± 18.5
n=2762
BOs 2015
4.4 ± 17.5
n=875
DITFs 2015
17.3 ± 13.5
n=1968
BOs 2012
156.2 ± 22.3
n=123
DITFs 2012
167.3 ± 8.1
n=341
BOs + DITFs 2012
164.9 ± 13.8
n=464
BOs 2015
140.6 ± 18.6
n=96
DITFSs 2015
163.4 ± 12.8
n=361

Figure 12: Stabilized wellhead pressure [MPa] as a function of flow rate [L.s^{-1}], measured during the hydraulic stimulation of the GRT-1 well in 2013 (after Baujard et al., 2017).

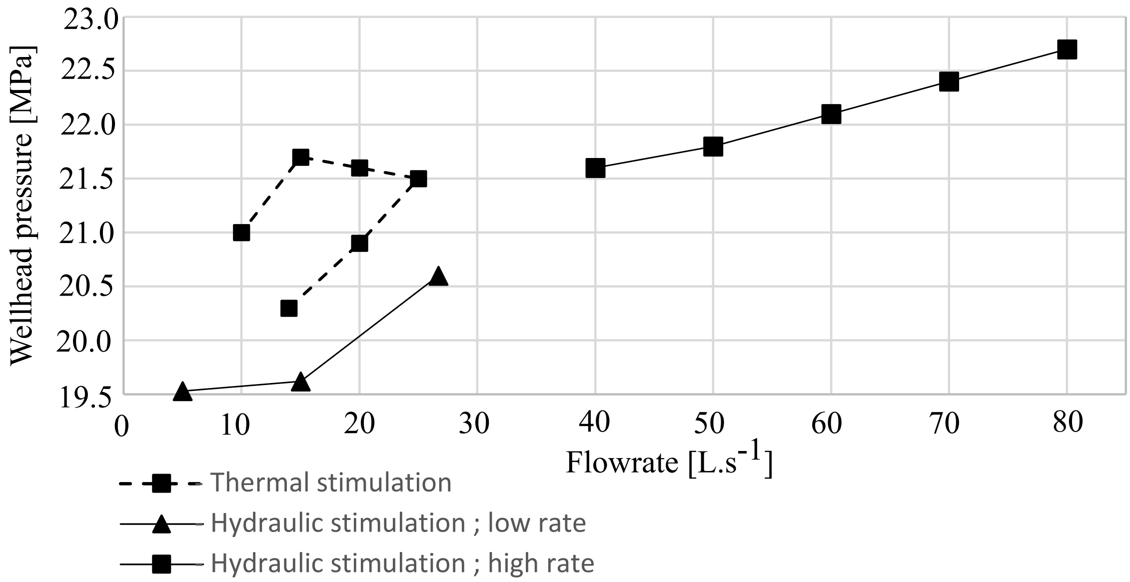
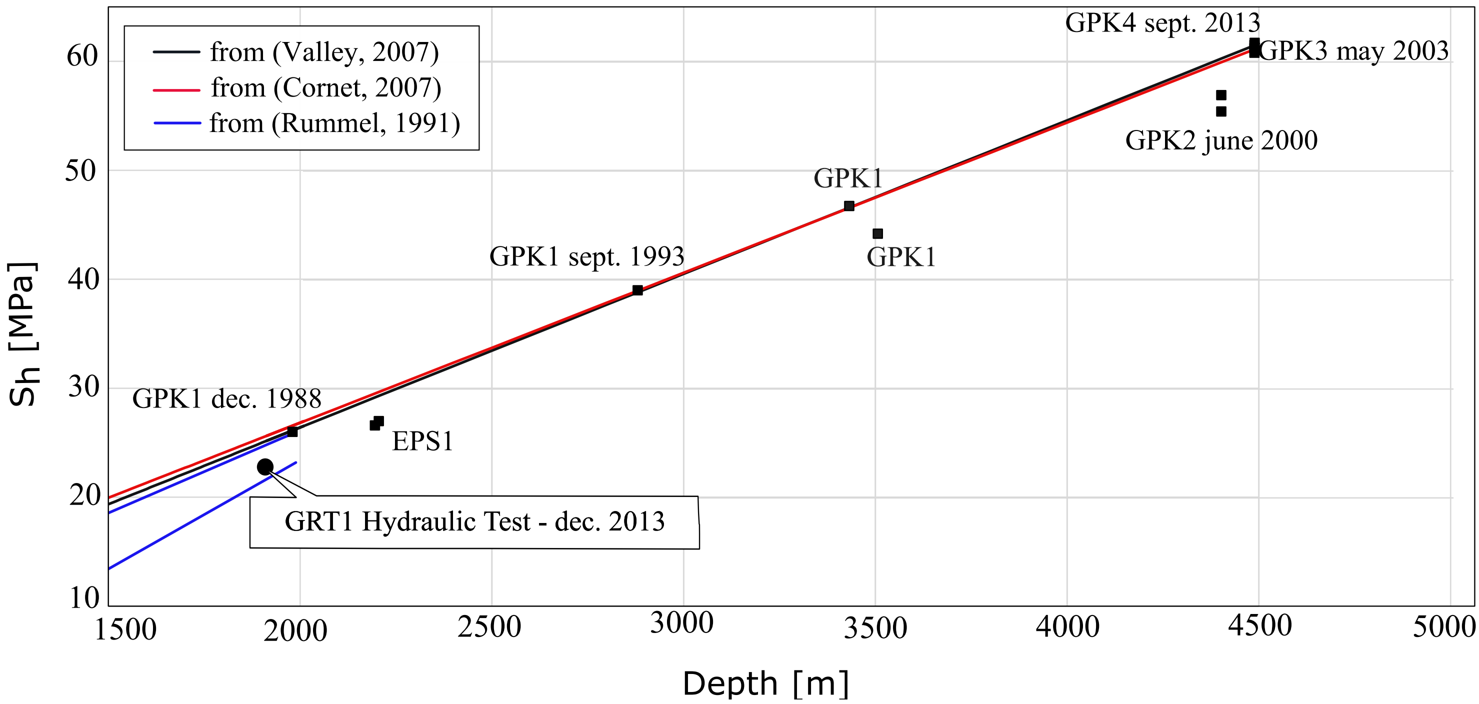

Wellhead pressure [MPa]
23.0
22.5
22.0
21.5
21.0
20.5
20.0
19.5
0
10
20
30
40
50
60
70
80
Flowrate [L.s-1]
Thermal stimulation
Hydraulic stimulation ; low rate
Hydraulic stimulation ; high rate

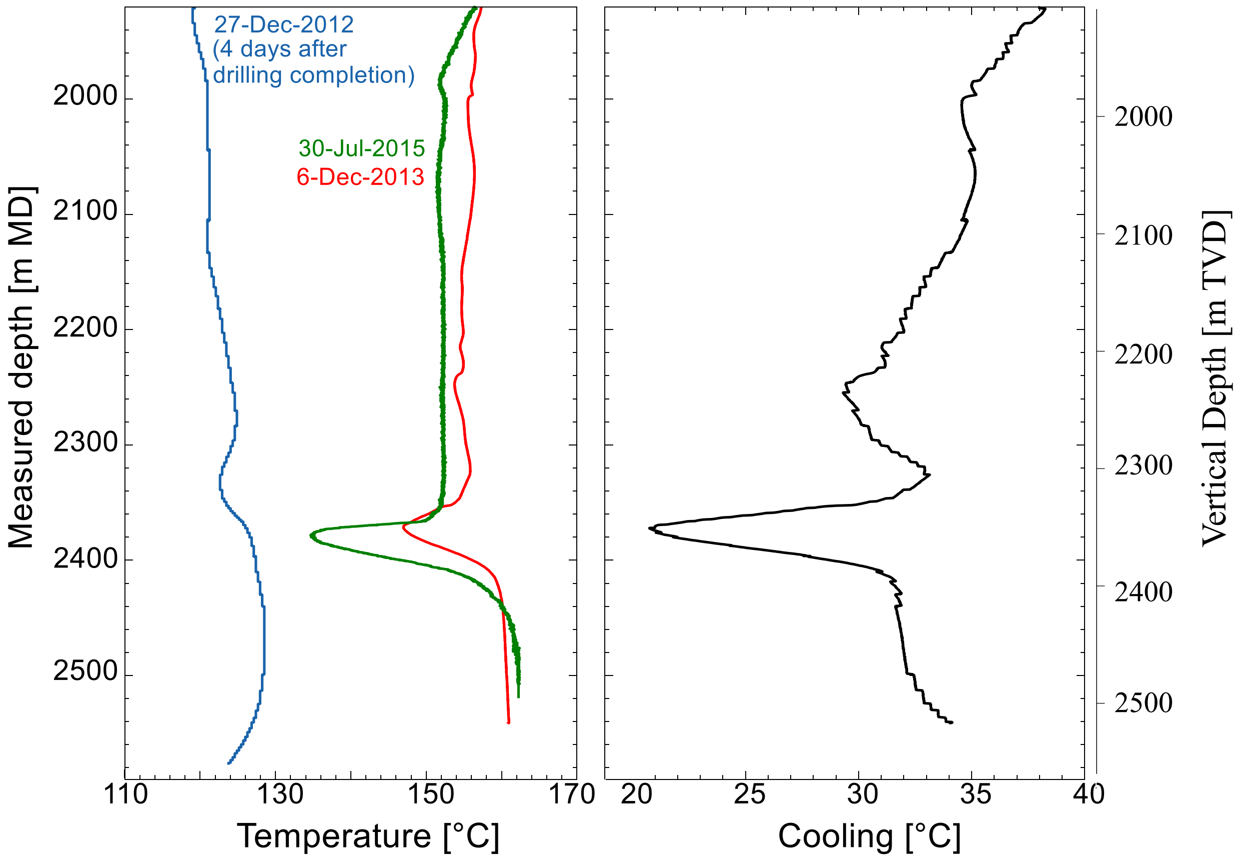
Figure 13: Minimal horizontal stress *Sh* [MPa] ~~with~~as a function of vertical depth (TVD) measured at the Soultz-sous-Forêts site from ~~three studies obtained from~~the analysis of high-volume injections in the GPK-1, GPK-2, GPK-3 and EPS-1 wells. The lower bound for the minimal horizontal stress *Sh* obtained from the analysis of the ~~hydraulic~~ wellhead pressure measured during the stimulation of the well GRT-1 in Rittershoffen is represented for comparison ~~(~~as a black circle~~)~~.

~~**Figure 13: Stabilized wellhead pressure [MPa] for each flowrate step measured during the stimulation operations conducted in GRT-1 in 2013, targeting a depth of 1913 m (TVD), as a function of flow rate [L.s^{-1}] (after Baujard et al., 2017).**~~



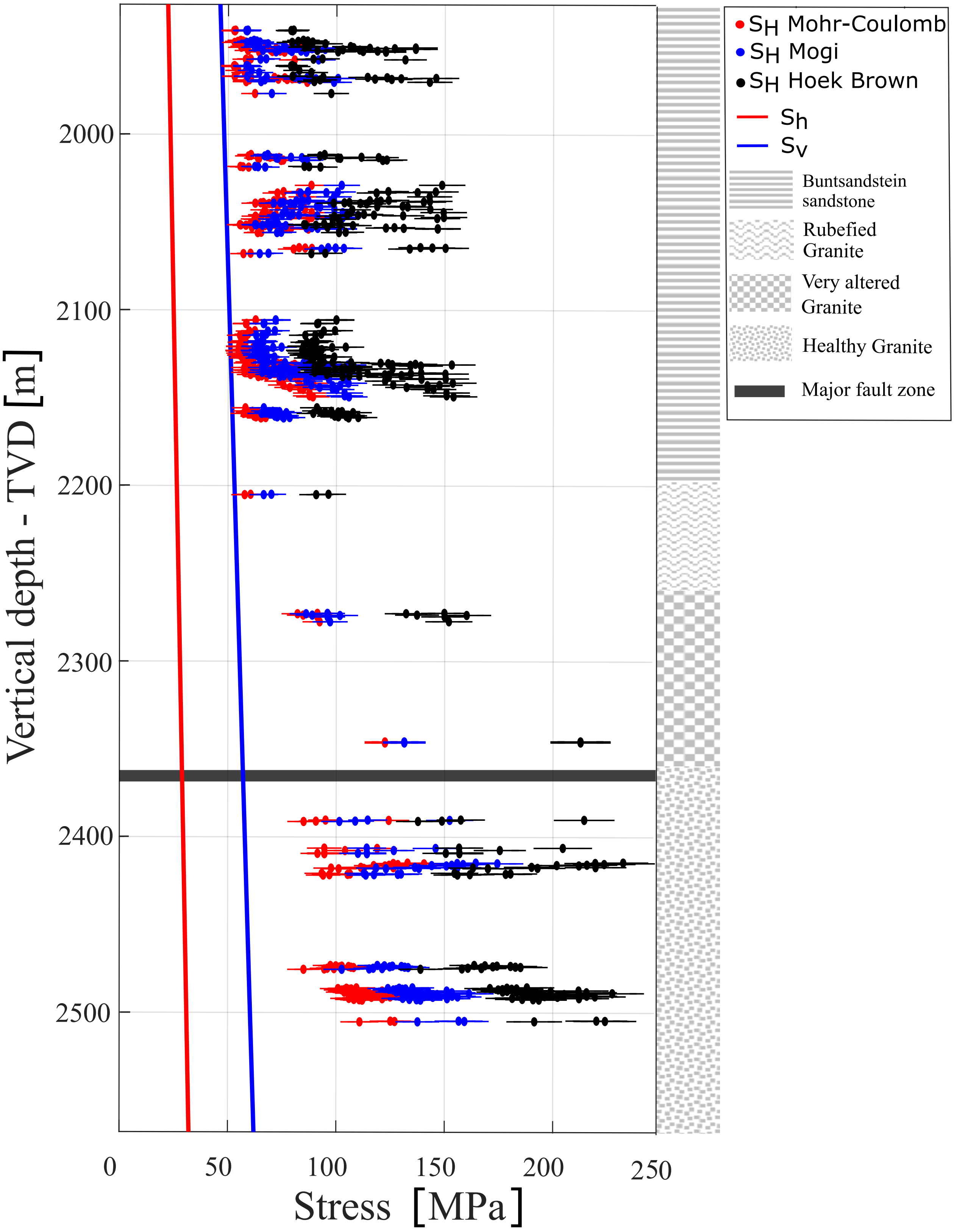
from (Valley, 2007)
from (Cornet, 2007)
from (Rummel, 1991)
GPK4 sept. 2013
GPK3 may 2003
GPK2 june 2000
GPK1
GPK1 sept. 1993
GPK1
GPK1 dec. 1988
EPS1
GRT1 Hydraulic Test - dec. 2013
Sh [MPa]
Depth [m]
60
50
40
30
20
10
1500
2000
2500
3000
3500
4000
4500
5000

Figure 14: Left panel: variation of temperature [°C] ~~with~~in function of Measured Depth (MD) or Vertical Depth (TVD), estimated from the temperature log acquired in 2015 in GRT-1 (green curve), plotted along with the temperature log acquired in 2013 (red curve). The temperature log acquired four days after drilling completion (blue curve) enables to estimate the temperature at the borehole wall during drilling. Right panel: estimation of the difference in temperature between the wellbore temperature and the borehole wall temperature after completion *Δt* used in the evaluation of the thermal stress components.



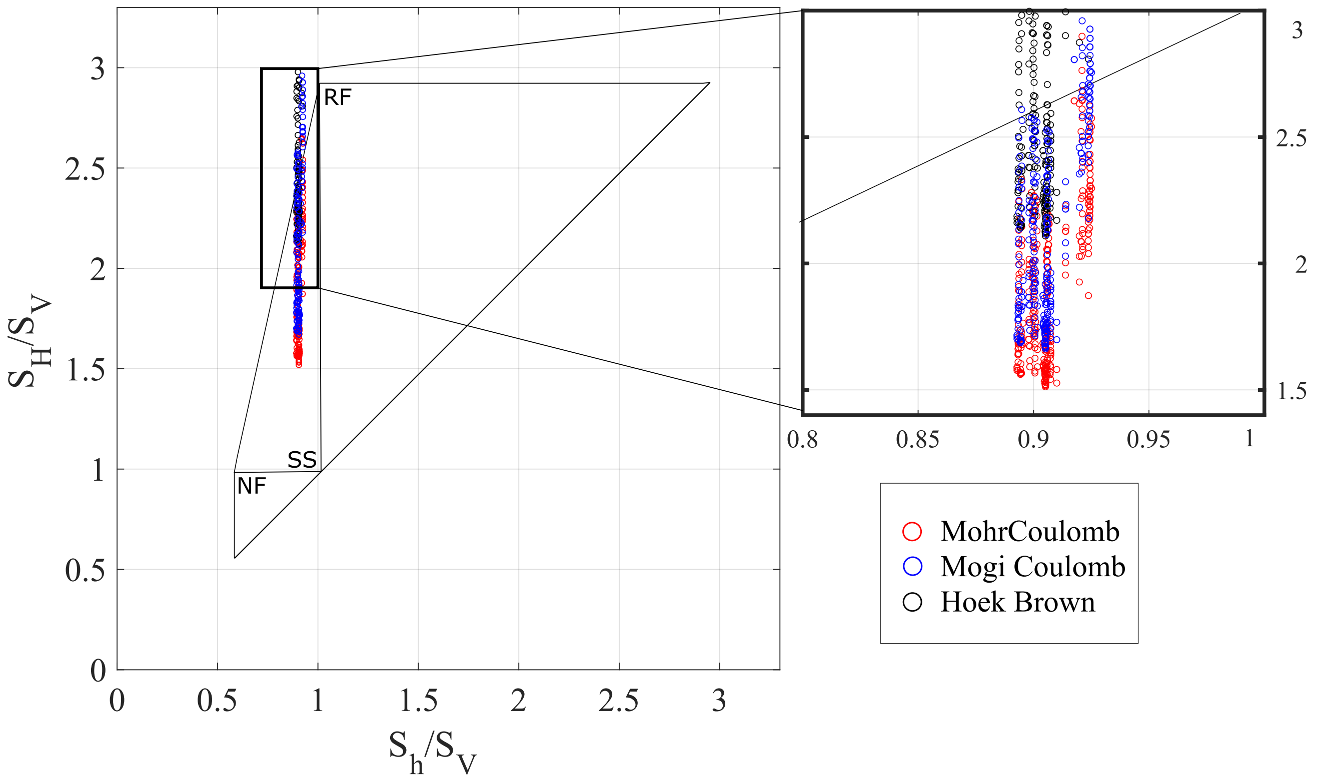
27-Dec-2012
(4 days after
drilling completion)
30-Jul-2015
6-Dec-2013
Measured depth [m MD]
2000
2100
2200
2300
2400
2500
110
130
150
170
Temperature [°C]
20
25
30
35
40
Cooling [°C]
Vertical Depth [m TVD]

Figure 15: *in-situ* stress state components *Sh*, *SV* and *SH* [MPa]. Maximum horizontal stresses *SH* are inverted with three distinctive failure criteria for the images ~~of~~acquired in 2013 ~~of~~in GRT-1~~ well~~. Error bars are calculated considering the error on the measurement of the breakout width, on the estimates of the elastic parameters and on the *Sh* and *SV* trends~~ with depth~~. The ~~background pattern represents~~right column illustrates the four major lithological units retained in the model and the horizontal band locates the major fault zone crossed by GRT-1.



SH Mohr-Coulomb
SH Mogi
SH Hoek Brown
Sh
Sv
Buntsandstein sandstone
Rubefied Granite
Very altered Granite
Healthy Granite
Major fault zone
2000
2100
2200
2300
2400
2500
Vertical depth - TVD [m]
0
50
100
150
200
250
Stress [MPa]

Figure 16: Normalized stress polygon defining stress states (*SH/SV*, *Sh/SV*) at a depth of 2500m in GRT-1, according to a Coulomb law with a coefficient of friction *μ*=1. The borders of the polygon correspond to an active fault situation. According to Anderson's faulting theory, *RF* – reverse faulting – *SS* – strike slip regime – and *NF* – normal faulting – refer to the Anderson's faulting regimes. It is plotted along with the stresses (*SH/SV* – *Sh/SV*) calculated from the image of the GRT-1 of 2013, for three different failure criteria (~~colored~~ circles in color).



RF
SS
NF
S_H/S_V
S_h/S_V
MohrCoulomb
Mogi Coulomb
Hoek Brown

Figure 17: *in-situ* stress components *Sh, SV* and *SH* [MPa] in the deviated well GRT-2. *SH* stresses are inverted using a Mohr Coulomb failure criterion and represented as a function of the vertical depth (TVD) for the images acquired in 2014 and 2015. Error bars are calculated considering the errors on the measurements of the breakout widths, on the elastic parameters and on the *Sh* and *Sv* trends. The ~~background pattern represents~~right column illustrates the ~~major~~ lithological unit ~~crossed by~~retained in the ~~well~~model.

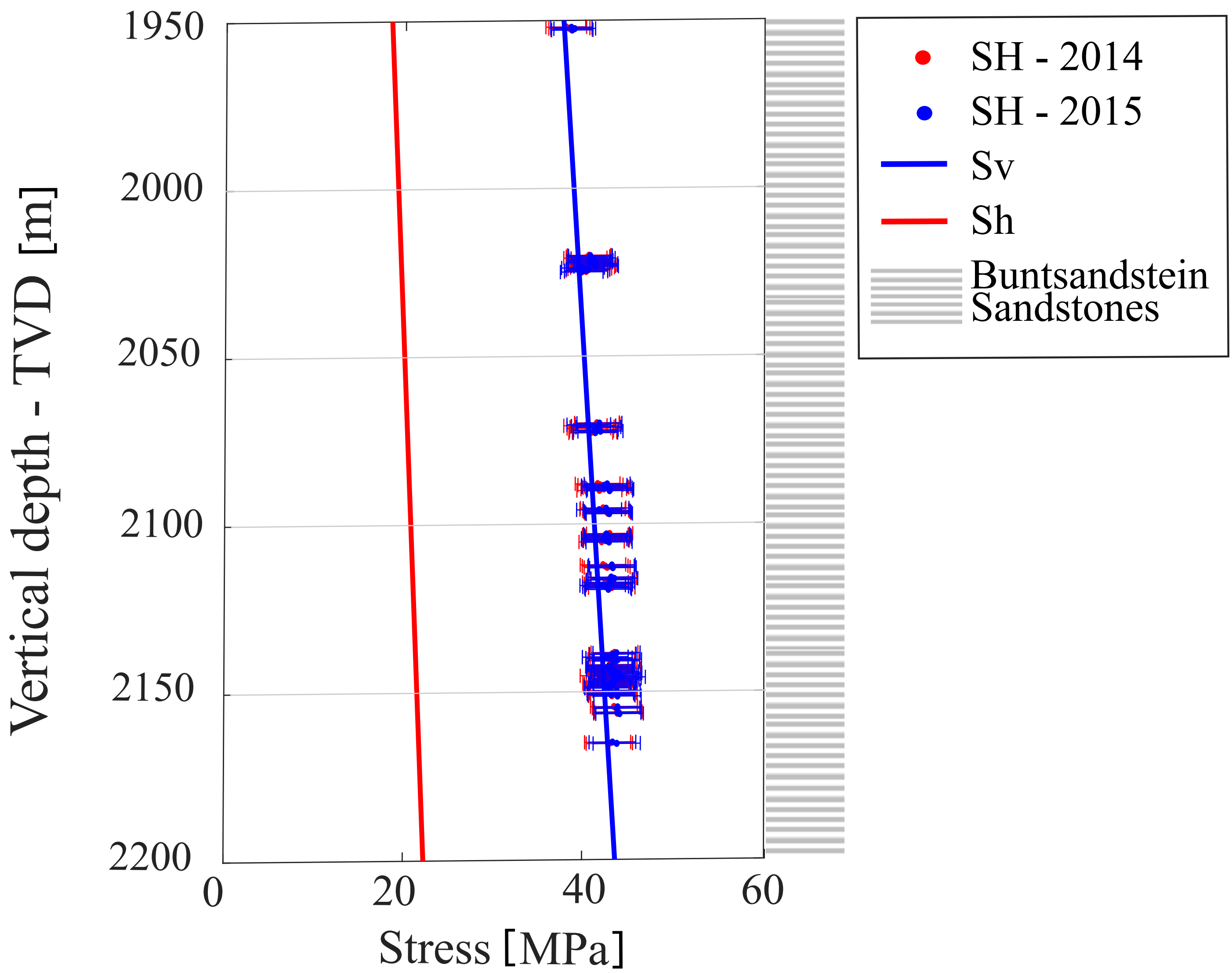

1950
2000
2050
2100
2150
2200
Vertical depth - TVD [m]
0
20
40
60
Stress [MPa]
SH - 2014
SH - 2015
Sv
Sh
Buntsandstein
Sandstones

Figure 18: Panel a. shows the *in-situ* stress components *Sh*, *SV* and *SH* [MPa] in the deviated wells GRT-1 and GRT-2. *SH* stresses [MPa] inverted with a Mohr-Coulomb criterion are obtained from the analysis of the images acquired in 2012 – 2013 and 2015 (~~plain colored~~respectively black, blue and red circles) in GRT-1 and in 2014 and 2015 (~~empty colored circles~~respectively black and red crosses) in GRT-2, as a function of vertical depth~~.~~ (TVD). The ~~background shows the~~right column illustrates the four major lithological units retained in the model. Panel b. ~~shows~~is a histogram with 1 MPa bins representing the difference ~~in~~between the *SH* stresses ~~between~~ measured in GRT-1 in 2015 and in 2012~~ in GRT-1 in a histogram with 1 MPa bins.~~ .

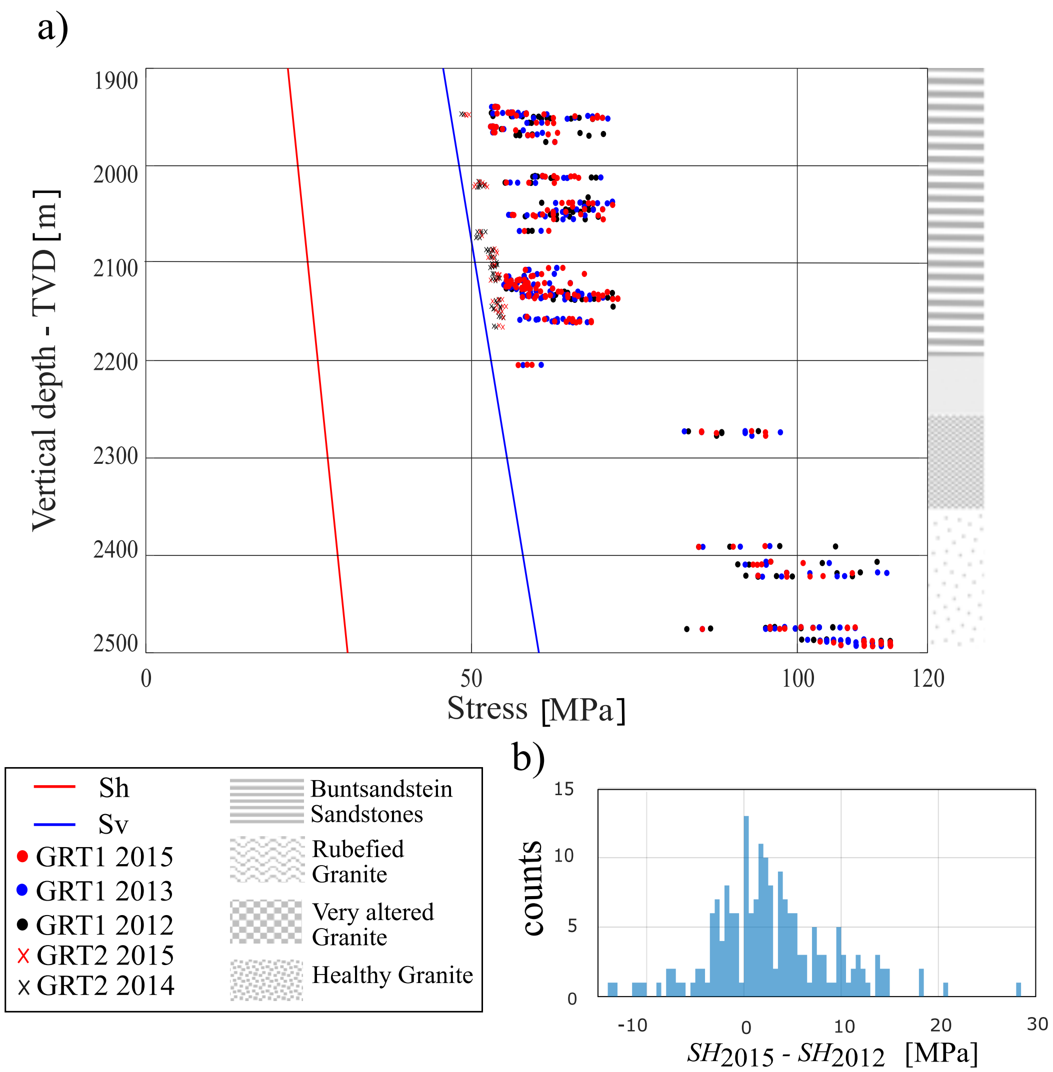

a)
Vertical depth - TVD [m]
1900
2000
2100
2200
2300
2400
2500
0
50
100
120
Stress [MPa]
Sh
Sv
GRT1 2015
GRT1 2013
GRT1 2012
GRT2 2015
GRT2 2014
Buntsandstein Sandstones
Rubefied Granite
Very altered Granite
Healthy Granite
b)
counts
15
10
5
0
-10
0
10
20
30
SH2015 - SH2012 [MPa]

Table 1: Data acquired in GRT-1 and GRT-2 and specificities of UBI acquisition programs.

Well	Acquisition Date	Stimulation	Logging depth range [m~~;~~ - MD] [m - TVD]	Transducer diameter [inch]
~~GRT1~~ GRT-1	30-Dec-2012	4 days after drilling completion	1913~~.00~~ - 2568~~.00~~ 1902 - 2550	4.97
	9-Dec-2013	1 year after drilling completion 5 months after THC stimulation	1912~~.00~~ - 2531~~.16~~ 1901 - 2513	2.92
	30-Jul-2015	2.5 years after drilling completion 2 years after THC stimulation~~.~~	~~1910.96 - 2499.9~~1911 - 2500 1900 - 2483	4.97
~~GRT2~~ GRT-2	23-Jul-2014	Four days after drilling completion	2118~~.00~~ - 2531~~.22~~ 1869 - 2196	4.97
	29-Jul-2015	1 year after drilling completion	2111~~.00~~ - 2869~~.23~~ 1863 - 2464	4.97

Table 2: Elastic (Poisson ratio) and strength parameters (used in the Mohr-Coulomb, Mogi-Coulomb and Hoek Brown failure criteria) for the four geological units retained in the model, for both GRT-1 and GRT-2 wells, as a function of measured depth (MD~~).~~) and vertical depth (TVD). Elastic and strength parameters for granites are based on a data compilation of tests conducted on samples from Soultz-sous-Forêts. For the Buntsandstein sandstones, we use usual strength parameters based on Hoek & Brown (1997).

Depth ~~(~~GRT-1 [m - MD~~)~~ ~~[m]~~ ~~GRT1~~[m - TVD]	Depth ~~(~~GRT-2 [m - MD~~)~~ ~~[m]~~ ~~GRT2~~[m - TVD]	Geology: Stratigraphy	Geology: Lithology	Elastic and strength Parameters: E [GPa]	ν [-]	Cohesion *C* [MPa]	Internal Friction ~~(θ)~~	UCS [MPa]	~~(a, b)~~ Mogi Coulomb (a, b)	Hoek Brown *mi*
1799-2212 1789-2197	2022-2479 1792-2155	Buntsandstein	Sandstones (argilic)	22 ±2	0.22	24 ±5	35°	92±14	(18 ±3, 0.54)	19
2212-2269 2197-2254	2479-2629 2155-2274	Granitic Basement	Ruberfied Granite	54 ±2	0.26	23 ±5	40°	100 ±15	(13 ±3, 0.68)	20
2269-2374 2254-2358	2629-2881 2274-2473		~~Hydrothermally~~Hydro- altered Granite	40 ±2	0.26	29 ±5	40°	125 ±17	(17 ±3.5, 0.68)	23
2374-2580 2357-2561	2881-3196 2473-2723		Low altered Granite	54 ±2	0.26	32 ±5	45°	155 ±20	(21 ±3.5, 0.68)	27

Table 3: Mean density retained for each lithological layer and vertical depth (TVD) in each well.

Description	Depth in ~~GRT1~~GRT-1 [m]	Depth in ~~GRT2~~GRT-2 [m]	Volumetric mass [kg.m^{-3}]
Tertiary	0	0	2350
	1172	1166.5	
Jurassic	1172	1166.5	2440
	1447	1431.5	
Keuper	1447	1431.5	2700
	1653	1637	
Muschelkalk	1653	1637	2750
	1798	1793.5	
Top Buntsandstein	1798	1793.5	2610
	1855	1850	
Mean Buntsandstein	1855	1850	2520
	2147	2109	
Bottom Buntsandstein	2147	2109	2540
	2198	2167	
Granitic basement	2198	2167	2570
	2568	2707.5	