Stress field orientation controls fault leakage at a natural CO₂ reservoir

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10   **Abstract.**

Travertine deposits at the St. Johns Dome natural CO₂ reservoir in Arizona, USA, document a long (>400 ka) history of surface leakage of CO₂ from a subsurface reservoir. Travertine deposits are concentrated along surface traces of faults implying that there has been a structural control on the migration pathway of CO₂ rich fluids. Here, for the first time, we combine slip tendency and fracture stability to analyse the geomechanical stability of the reservoir-bounding Coyote Wash Fault for three

15   different stress fields and predict areas with high leakage risks. We find that these areas coincide with the travertine deposits on the surface indicating that high permeability pathways as a result of critically stressed fracture networks exist in both fault damage zones and around a fault tip. We conclude that these structural features control leakage. Importantly, we find that even without in-situ stress field data the known leakage points can be predicted using geomechanical analyses, despite the complex tectonic setting. Thus, even though acquiring high quality stress field data for secure subsurface CO₂ or energy storage is a

20   must, a first order assessment of leakage risks during site selection can be made with limited stress field knowledge.

1 Introduction

The successful subsurface storage of fluids in sedimentary basins is key for technologies such as Carbon Capture and Storage (CCS), cited as a cost-effective tool for climate change mitigation, or for energy storage, required to balance the intermittency of future energy systems relying on renewable sources(Alcalde et al., 2018; Matos et al., 2019; Scott et al., 2013). The integrity

25   of such engineered subsurface storage sites is controlled by a range of geological, geochemical, and geotechnical factors. One major concern is that impermeable caprock seals may be bypassed by faults and naturally occurring, or induced, fracture networks which can form preferential fluid pathways and may lead to the rapid migration of the stored fluid (e.g. CO₂, H₂, methane) to shallow aquifers or the atmosphere (IPCC, 2005; Shipton et al., 2004; Song and Zhang, 2012). Indeed, selection criteria for subsurface storage sites commonly cite the need of minimal faulting and/or low permeability faults in the vicinity

30   of a storage site (Chadwick et al., 2008; IEA GHG, 2009; Miocic et al., 2016). However, within sedimentary basins, which are

key areas for geological storage of fluids, faults will occur naturally close to or within a storage complex and thus predictability of whether a fault will act as barrier to fluid flow or not is key for an accurate risk assessment.

Whether a fault zone is sealing or non-sealing is dependent on the structure and composition of the fault zone and the mechanics of the faulting (Faulkner et al., 2010). In a widely used simple conceptual model for fault zones strain is localized in the fault

35   core which is surrounded by a damage zone of fractures; fault zones can have a single high-strain core (Chester and Logan, 1986) or contain several cores (Choi et al., 2016; Faulkner et al., 2003). The damage zone and the fault core have contrasting mechanical and hydraulic properties with the fault core being often rich in phyllosilicates which typically have low permeability while open fractures in the damage zone can have a substantially higher permeability than the host rock (Caine et al., 1996; Cappa, 2009; Faulkner and Rutter, 2001; Guglielmi et al., 2008). Lateral fluid migration across the fault zone is

40   thus controlled by the permeability of the fault gouge within the fault core(s), which is dependent on the host rock composition and the shear strain. Contrastingly, vertical fluid migration is governed by the connectivity and permeability of the fractures in the damage zone. A significant amount of research has been focused on understanding the mechanisms and parameters that control the composition, and continuity of fault gouges as well as their permeability for different fluids as they have the potential to form effective seals (Davatzes and Aydin, 2005; Lehner and Pilaar, 1997; Lindsay et al., 1993; Miocic et al.,

45   2019b; Vrolijk et al., 2016; Yielding et al., 1997). The damage zone permeability is controlled by the permeability of the host rock and the presence and geometric composition of macro-scale fracture networks and deformation and compaction bands which decrease in frequency with increasing distance from the fault core (Mitchell and Faulkner, 2009; Shipton et al., 2002). Outcrop studies have shown flow channelling and emphasise the strong spatial and temporal heterogeneity of fault zones (Bond et al., 2017; Burnside et al., 2013; Dockrill and Shipton, 2010; Eichhubl et al., 2009; Schulz and Evans, 1998; Soden et al.,

50   2014). If fracture networks or faults are close to failure due to tectontically induced changes in the stress conditions or changes in pore pressure vertical fluid flow is enhanced (Barton et al., 1995; Wiprut and Zoback, 2000). Geomechanical parameters such as slip tendency (Morris et al., 1996) or fracture stability can be used to assess the potential of vertical fluid flow.

The need for improved understanding of fracture networks and the potential of fracture reactivation and/or hydromechanically fracturing of caprock due to the injection of CO_2 has been highlighted by experiences at existing industrial CO_2 storage projects.

55   At the Sleipner storage site, fractures in thin caprock layers appear to control the size and extend of the CO_2 plume (Cavanagh et al., 2015). The storage site of In Salah, Algeria, where between 2004 and 2011 around 4 million tons of CO_2 were injected into an anticlinal structure at ~1,800 m depth, has been the focus of many studies on fracture reactivation and hydraulic fracturing of caprocks as observations at the end of the injection period suggested that pressure had migrated vertically into the caprock (Bond et al., 2013; Michael et al., 2010; Rutqvist et al., 2010; Stork et al., 2015). The existing data indicate that

60   injection pressures were too high for the low permeability reservoir rock and hydraulically fractured the reservoir and the lower caprock, potentially also reactivating pre-existing fracture networks related to small scale faults (White et al., 2014).

To study how vertical fluid flow along fault zones may be related to geomechanical parameters we examine the naturally occurring CO_2 reservoir of the St. Johns Dome, located at the border of Arizona and New Mexico. At this site migration of

fluids from the subsurface reservoir to the surface is directly linked to faults which extend through the caprock (Gilfillan et al.,
65   2011; Moore et al., 2005), with leakage having occurred for at least 420 ka and is still ongoing (Miocic et al., 2019a; Priewisch
et al., 2014).

2 Geological setting

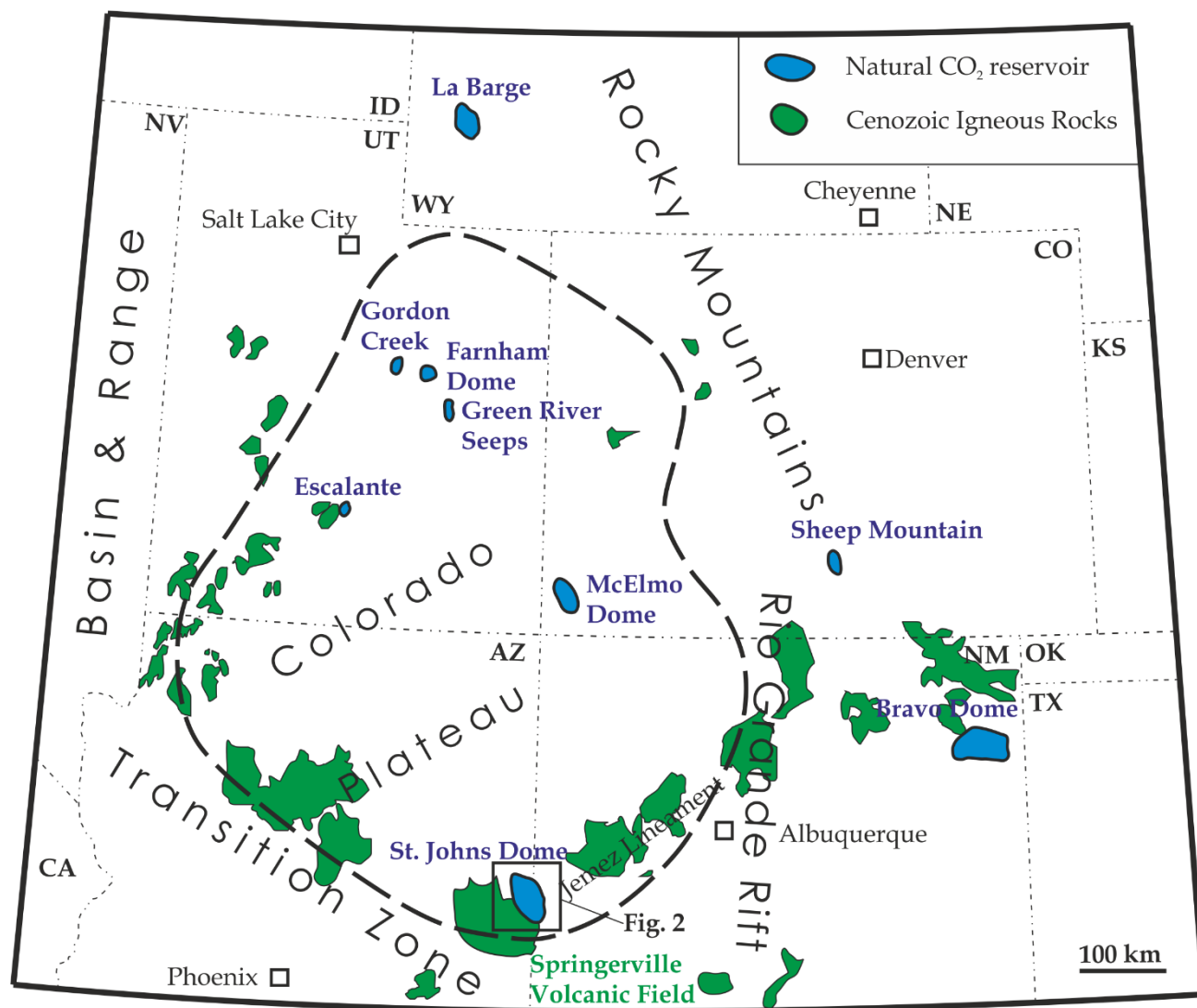
The St. Johns Dome (or Springerville-St. Johns Dome) natural CO_2 reservoir has more than 4.7 x10^{10}m³ of recoverable CO_2
and is located on the south-eastern edge of the Little Colorado River Basin on the Colorado Plateau (Fig. 1) near to the
70   Transition Zone to the Basin and Range and Rio Grande Rift tectonic provinces (Bashir et al., 2011; Rauzi, 1999). It is one of
sixteen known naturally occurring CO_2 reservoirs on the Colorado Plateau and one of the few known naturally occurring CO_2
reservoirs world-wide where fluids are leaking to the surface (Gilfillan et al., 2008, 2009; Miocic et al., 2016). The CO_2
reservoir lies within a broad, northwest-trending asymmetrical anticline that is dissected by the steeply dipping northwest-
southeast trending Coyote Wash fault (Fig. 2, Moore et al., 2005; Rauzi, 1999). This major fault seems also to form the western
75   boundary to the productive portion of the formally commercially exploited gas field. Normal displacement across the fault
ranges from less than 30 m (Salado Springs) to more than 200 m at the apex of the Cedar Mesa Anticline, 25 km SE of Salado
Springs. Much of the fault displacement is thought to be of Paleogene age when Laramide compressional tectonics led to
monoclinal folding of the Phanerozoic strata and the reactivation of older basement structures such as the Coyote Wash Fault
on the Colorado Plateau (Marshak et al., 2000). The Permian reservoir rocks (siltstones, sandstones and limestones)
80   discordantly overlay Precambrian granites (Fig. 3), are relatively shallow at 400-700 m depth and the CO_2 is present as gas
(Gilfillan et al., 2011). Anhydrite and mudstone beds within the Permian rocks divide the reservoir into several producing
zones while Triassic and Cretaceous calcareous shales and mudstones act as seals (Fig. 3). The Permian strata includes, from
youngest to oldest, the San Andres Limestone Glorieta Sandstone and the Supai Formation which consist of the Corduroy
Member, Fort Apache Member, Big A Butte Member and Amos Wash Member. A detailed geological description of the
85   Permian Rocks can be found in Rauzi (1999). The current gas-water-contact is at 1425 m above sea level and the reservoir is
not filled to spill. The surface rocks are mainly Triassic to Quaternary sediments, volcanic rocks and travertine deposits (Fig.
2). To the north-west the CO_2 reservoir is boarded by the Holbrook basin (Harris, 2002; Rauzi, 2000) and it is closely associated
with the Plio-Pleistocene Springerville volcanic field which lies just to the south and south-west of the CO_2 reservoir (Crumpler
et al., 1994; Sirrine, 1956). The volcanic field consists of more than 400 individual vents and related flows. Basaltic volcanism
90   began around 9 Ma and the youngest flows that are about 0.3 Ma old can be found 8 km northwest of Springerville (Condit et
al., 1993; Condit and Connor, 1996). As the CO_2 within the reservoir is of magmatic origin (Gilfillan et al., 2008, 2009, 2011),
charging of the reservoir is thought to originate from degassing of magma underneath the volcanic field, with CO_2 migrating
along fractures and faults through the basement into the reservoir (Miocic et al., 2019a).

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Figure 1: Map showing the location of natural CO₂ reservoirs on the Colorado Plateau and adjacent areas. The St. Johns Dome reservoir is located on the southern edge of the plateau, next to the Springerville Volcanic field. After Moore et al. (2005) and Bashir et al. (2011). Black rectangle indicates outline of figure 2.
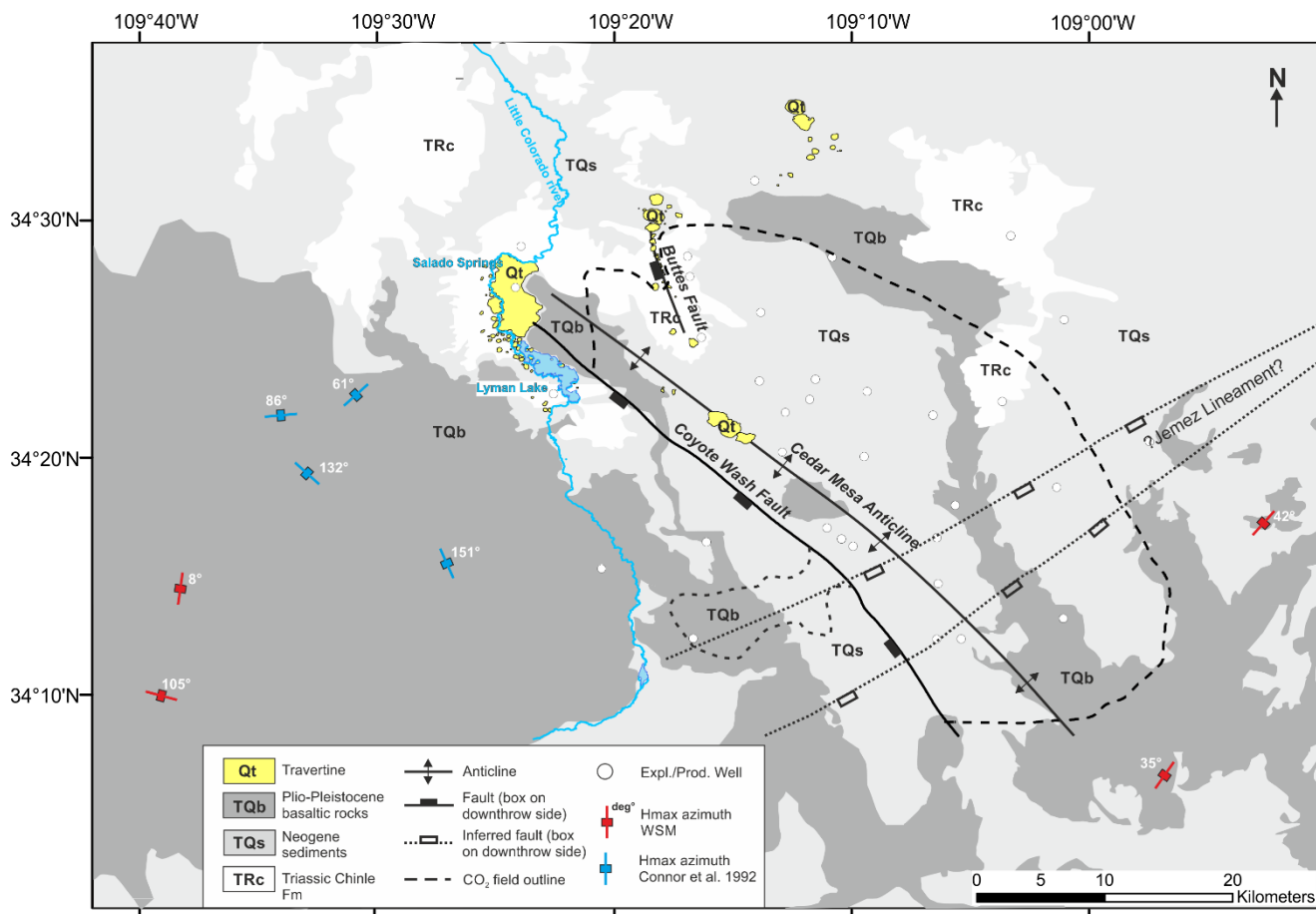
Figure 2: Geological map of the St. Johns Dome natural CO₂ reservoir showing the present-day extent of the CO₂ reservoir, the location of the travertine deposits, orientation of the studied faults, and the location of exploration and production wells used to build the subsurface model. Stress field markers indicate the azimuth of maximum horizontal stress after the world stress map (Heidbach et al., 2016) in red and Connor et al. 1992 in blue.
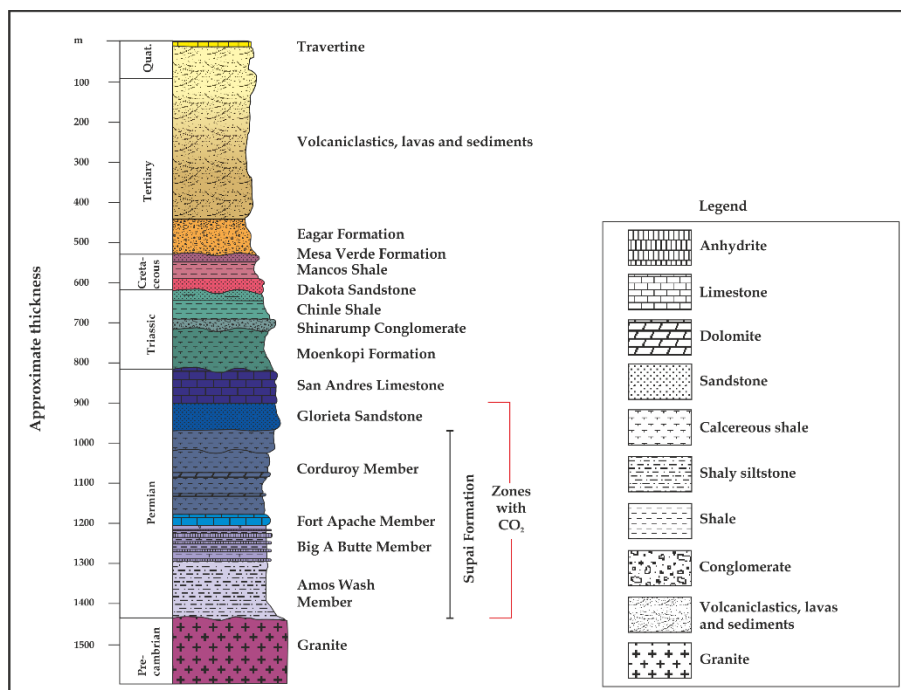
Figure 3: General stratigraphic column of the St. Johns-Springerville area. Note that Cretaceous and younger deposits are often thinner than described in this figure. CO₂ occurs in the Permian strata. After Rauzi (1999) and Embid (2009).

2.1 Expression, and timing, of fluid flow

The travertine deposits at the St. Johns Dome are an expression of CO_2-charged fluids from the subsurface to the surface as travertine formation occurs when CO_2-rich fluids outgas CO_2 as they migrate upwards to shallower depths and lower pressure, resulting in $CaCO_3$ supersaturation and carbonate precipitation. The travertine deposits cover a surface area of more than 30 km^2, spread out over more than 300 km² (Fig. 2), making them one of the greatest concentrations of travertine deposits in the U.S.. Spatially, the travertine deposits are particularly concentrated in a 10 km long zone between Salado Springs and Lyman Lake (Fig. 2, Gilfillan et al., 2011; Moore et al., 2005). This area, where present day travertine formation occurs (Priewisch et al., 2014), is bound by the Coyote Wash Fault and the distribution of the travertine deposits and active springs suggests that the local groundwater hydrology has been influenced by the Coyote Wash Fault. Analyses of surface springs, groundwater wells and CO_2 wells with respect to the CO_2 composition, water composition and noble gas concentrations have shown that samples along the Coyote Wash Fault trace and are influenced by waters from depth that have been enriched in mantle derived [3]He and Ca (Gilfillan et al., 2011; Moore et al., 2005). Several modelling approaches emphasise the importance of the Coyote Wash Fault for CO_2 and He migration from the Supai Formation to the surface (Allis et al., 2004; Keating et al., 2014) as in all models migration of gas to the surface occurs only if the fault forms a permeable section through the cap rocks. Soil-flux measurements indicate that there is no diffuse CO_2 leakage through the cap rocks, highlighting the role of faults for vertical fluid flow (Allis et al., 2005). In addition to the occurrences along the NW tip of the Coyote Wash Fault, travertine mounds

are found to follow the trace of the Buttes Fault over a distance of more than 7 km (Fig. 2). Travertine mounds are also found North of the present-day extend of the CO_2 reservoir, with no clear link to structural elements. It is notable that there are no

125 indications for fluid migration in the southern half of the reservoir.

U-series dating of the travertine mounds shows that leakage of CO_2 from the reservoir to the surface has occurred for at least 420 ka (Miocic et al., 2019a; Priewisch et al., 2014). Several of the samples analysed by Miocic et al. (2019a) fall outside the dating limitations of the U-Th method (~500 ka), which indicates that leakage may have occurred over much longer time-scales which is not surprising given the age of the Springerville volcanic field from where the magmatic CO_2 is almost certainly

130 sourced. Individual seeps along the Buttes Fault have lifespans of up to 200 ka and the massive travertine platform between Salado Springs and Lyman lake has at least a similar lifetime. Volume calculations indicate that the subsurface reservoir is constantly recharged as several times the volume of CO_2 stored in the current reservoir has leaked in the past. However, due to the long timeframe of recorded leakage, only a very low percentage (0.1 to 0.001 %) of the reservoir volume is leaking annually and thus the site could still be seen as a suitable carbon storage site from a climate mitigation point of view (Miocic

135 et al., 2019a).

These observations illustrate that fluid migration at the St. Johns Dome occurs along fault zones and once migration pathways have been established they are spatially stable for long periods (>100 ka). This is in contrast to other fault controlled fluid migration pathways on the Colorado Plateau for which it is suggested that these stay open only episodically for few thousands of years after rapid fault opening/movement and subsequently heal (Frery et al., 2015). Similar cyclic reopening and healing

140 of fractures governing fault zone permeability has been recorded by travertine deposition at other active fault zones (Brogi et al., 2010). Spatially and temporally stable migration pathways are concerning for subsurface storage sites and thus the processes controlling vertical fault zone permeability at the St. Johns Dome are analysed in the following.

3 Methods

In order to investigate the mechanisms governing the vertical fluid flow at the St. Johns Dome, a geomechanical analysis of

145 the Coyote Wash Fault was conducted using slip tendency and fracture stability approaches. Slip tendency (Ts) is a method that allows for a fast assessment of the tendency of a surface to undergo slip under the present-state effective stress field. It is the ratio of resolved shear stress to resolved normal stress on a surface (Morris et al., 1996):

$$Ts = \frac{\tau}{\sigma_n}$$

where τ is the shear stress and σ_n the effective normal stress acting on the fault plane. Slip is likely to occur on a surface if $Ts \geq$

150 μ_s with μ_s being the coefficient of static friction which is generally assumed to be 0.6 (Byerlee, 1978; Moeck et al., 2009; Sibson, 2003). Fracture stability (Fs) is the increase in pore pressure that is required to reduce the effective stresses such that the fault plane is forced into failure. In contrast to slip tendency, Fs takes rock properties such as cohesion and tensile strength into account and thus fault rock composition needs to be known.

A 3D geological model of the St. Johns Dome was built based on published geological maps, well data from 37 exploration
and production wells (well logs, horizon markers) as well as previously published reservoir horizon maps. Between wells a
constant stratigraphic thickness was assumed. The current gas-water-contact is at 1494 m above sea level and is assumed to
be horizontal. Geomechanical analysis of the model was conducted with industry standard software (TrapTester®). For Fs
analysis two different fault rocks were taken into account: claysmear (cohesion C = 0.5 MPa, coefficient of internal friction
μ=0.45) and phyllosilicate (C=0,5, μ=0.6). For stress field no in-situ stress measurement was available, however in addition
to world stress map data (Heidbach et al., 2016) the nearby Springerville volcanic field can be used to derive the orientation
of the horizontal stresses as presented in the following.

3.1 Stress field at the St. Johns Dome

The location of the St. Johns Dome reservoir at the margin of the Colorado Plateau and just north of the Basin and Range
province has significant impact on the stress field in the study area. It is clear that the regional stress field is highly variable as
shown by the available stress field data in the vicinity of the St. Johns Dome (50 km radius) from the world stress map
combined with a regional study on volcanic vent orientation in the Springerville volcanic field (Table 1) (Connor et al., 1992;
Heidbach et al., 2016). To the south and south-east of the CO_2 field, maximum horizontal stress (SH_{max}) is oriented NE-SW
while west of the reservoir SH_{max} orientation is highly variable, ranging from NW-SE to E-W (Fig. 2). While these datapoints
are associated with an uncertainty of at least ±15°, the orientation of the stress field for the St. Johns Dome faults is difficult
to constrain. A normal faulting regime (vertical stress (S_v) > maximum horizontal stress (SH_{max}) > minimum horizontal stress
(Sh_{min})) is assumed as indicated by the world stress map and works by Kreemer et al. (2010) and Wong and Humphrey (1989)
for this area of the Colorado Plateau. Integration of density logs (wells 10-29-31 and 11-16-30) gives a magnitude of S_v of 23
MPa/km. Minimal horizontal stress in normal faulting regimes is typically about 65-85% of the vertical stress (Hillis, 2003),
which gives a magnitude of Sh_{min} in the range of 15 to 19.5 MPa with the magnitude of SH_{max} set between S_v and Sh_{min}.

As the reported stress field measurements seem to form three clusters (Tab 1.), three different stress fields were defined (Tab.
2): Stress field A is similar to the stress fields indicated by measurements 7 and 8, with SH_{max} having an azimuth of 140°, stress
field B is similar to the stress fields indicated by measurements 2 to 4 with an SH_{max} azimuth of 50°, and stress field C is
oriented similar to the stress field measurements 5 and 6. The single north-south stress field measurement (ID=1) was not
considered further. For the ~NW-SE trending Coyote Wash Fault these stress fields also represent the most (A), least (B) and
an intermediate (C) likely cases for fault reactivation. Geomechanical analysis was conducted under all three defined stress
fields.

Table 1: Table listing published stress field indicators around the St. Johns Dome. WSM=World stress map. They form three clusters: IDs 2-4, 5&6, 7&8.

ID	Lat.	Long.	SHmax azimuth(°)	Shmin azimuth (°)	Error	Faulting regime	Source
1	34.230	-109.630	8	98	± 25	NF	WSM, 2016
2	34.070	-108.930	35	125	± 25	NF	WSM, 2016
3	34.250	-108.870	42	132	± 25	NF	WSM, 2016
4	34.248	-109.527	61	151	± 15	NF	Connor et al., 1992
5	34.313	-109.579	86	176	± 15	NF	Connor et al., 1992
6	34.140	-109.680	105	195	± 15	NF	WSM, 2016
7	34.269	-109.557	132	222	± 15	NF	Connor et al., 1992
8	34.202	-109.449	151	241	± 15	NF	Connor et al., 1992

Table 2: Stress fields used for the geomechanical modelling.

Stress field	S_v	SHmax		Shmin	
	MPa/km	MPa/km	Azimuth (°)	MPa/km	Azimuth (°)
A	23	20	140	16	50
B	23	22	50	16	140
C	23	20	100	16	10

4 Geomechanical controls on vertical fluid migration

The results of the geomechanical analysis of the Coyote Wash Fault highlight that the orientation of the stress field has a major impact on both the slip tendency (Fig. 4) and fracture stability (Fig. 5). Slip tendency indicates that for stress field A most parts of the fault are close to slip (Ts>0.5), for stress field C the fault is only intermediately stressed (0.3<Ts<0.5), and for stress field B the fault is far away from failure (Ts<0.2). Similarly, only slight increases of pore pressure are needed to force the fault into failure under stress fields A and C and a claysmear fault rock (0.95 MPa and 1.33 MPa, respectively). The pore pressure increase needed to force the fault into failure in case of stress field B is much higher at 6.21 MPa. Note that slip tendency in both stress fields A and C is higher at the NW tip of the fault than in the SE section of the fault (Fig. 4), indicating that failure is more likely to occur in the NW. This is also true for the spatial distribution of fracture stability which for stress fields A and B (most and least likely to fail) and a claysmear fault rock is illustrated in figure 6.

The results of the geomechanical analysis show that the bounding fault of the St. Johns Dome CO_2 reservoir is intermediately to critically stressed for two of the three modelled stress fields (A and C). For the same stress fields a weak fault rock within the Coyote Wash Fault zone results in fracture stabilities which range from less than 1 MPa to 1.33 MPa. The most critically stressed areas are located at the NW tip of the fault while the SE part of the fault is relatively stable for all studied stress fields.

The Fs values of 0.95 MPa and 1.33 MPa translate to an additional CO_2 column of ~110 m and ~160 m, respectively. Currently
205   the reservoir is not filled to spill and the 3D geological model indicates that the reservoir interval at the NW part of the reservoir
could hold an additional ~300 m of CO_2 column. Thus, additional filling of the reservoir with a third to half more CO_2 by
volume could lead to fault failure and vertical fluid migration along the fault. Evidence that the reservoir has held more CO_2
in the past is provided by older travertine deposits located outside the present day extent of the subsurface reservoir (Fig. 2,
Miocic et al., 2019a) and the fact that higher paleo-reservoir pressures have been implied by a geochemical study (Moore et
210   al., 2005). These higher reservoir pressures were likely enough to bring the NW part of the fault close to failure and we suggest
that the permeability of fracture networks within the critically stressed fault damage zone was therefore increased (Barton et
al., 1995; Ito and Zoback, 2000; Min et al., 2004). In order to sustain the long periods of leakage recorded by the spatially
stable travertine deposition the fault must be critically stressed for similarly long periods. Indeed, volume calculations on how
much CO_2 must have leaked to the surface based on the travertine deposits show that one to two orders of magnitude more
215   CO_2 was lost from the reservoir than it can hold (Miocic et al., 2019a). It is suggested that the continuous influx of magmatic
CO_2 degassing from beneath the Springerville Volcanic Field into the reservoir caused the fault to be close to being critically
stressed – a reasoning also encouraged by this study.
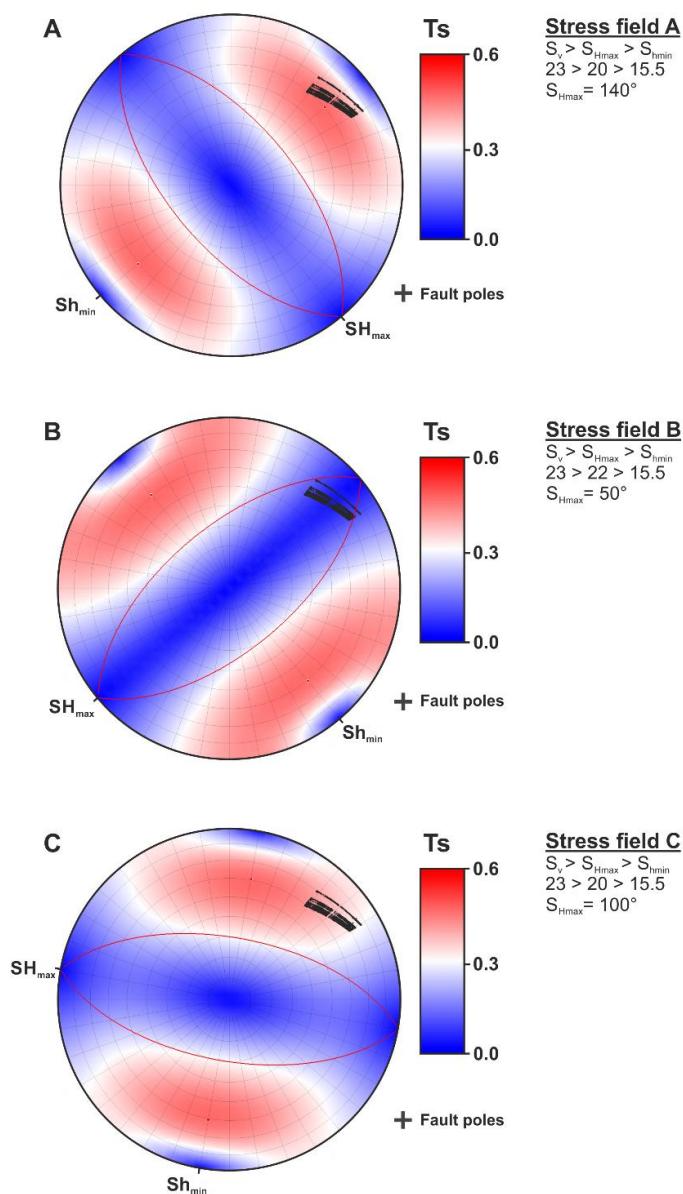
Figure 4: Stereoplots illustrating the slip tendency for the Coyote Wash Fault (black crosses) for the three different stress fields. The
220 **fault is at least partly close to failure for stress fields A and C while stress field B results in very low slip tendencies. Particularly the**
NW fault tip (indicated on the stereoplots) where travertine deposits are found at the surface is close to failure.

The geomechanical analysis also demonstrates that a change of the fault orientation within the stress field should not be

underestimated and can lead to failure along one part of a fault while large parts of the fault are geomechanically stable. The

strike direction of the Coyote Wash Fault changes from ESE-NWN in the southern part of the fault to NW-SE in the northern

225 section and this change in strike is enough to render the northern section critically stressed (in two of the stress fields modelled;

A & C) - with leakage pathways being the result (Fig. 7). Higher paleo gas columns within the reservoir likely contributed the

forcing the fault into failure at the northern fault tip. However, some sections of the fault in the SE also have relatively low fracture stability values (Fig. 6A) which translate to only 10s of meters more supported gas column than the NW section. Yet, there are no indications for past or present leakage in the SE part of the St. Johns Dome. We argue that the stress field

230 orientation in the SE is different from the stress field orientation in the NW area of the St. Johns Dome and that as a result the fault is far from failure. This is supported by stress field measurements in the vicinity of the southern edge of the reservoir (Fig. 2, Tab. 1) which imply a NE-SW SH_{max} orientation, similarly to stress field B in this study for which the Coyote Wash Fault is far from critically stressed.

Vertical migration of fluids through fracture networks or corridors can be classified by their location in (1) the fault damage
235 zone, (2) at the fault tip, and (3) at the crest of a fold (e.g. Ogata et al., 2014). As evidenced by travertine deposits vertical fluid migration at the St. Johns Dome occurred at all three types of fracture networks (Fig. 2), but considerably larger volumes of fluid migrated through fracture networks linked to faults. This indicates that, at least at this site, faults are a higher risk factor for leakage than other migration pathways such as fracture networks in anticline structures or capillary leakage though a caprock. Based on travertine volumes the largest volumes of leakage occurred at the NW tip of the Coyote Wash Fault, in the
240 area between Lyman Lake and Salado Springs (Figs. 2, 7, Miocic et al., 2019a). This indicates high permeability fracture networks within the damage zone close to the fault tip as predicted by numerical models (Backers and Moeck, 2015; Zhang et al., 2008). The lack of similar leakage pathways at the SE tip of the fault can be attributed to the different stress field orientation. For geological storage the occurrence of large-scale leakage at fault tips is also concerning as displacement at fault tips usually is low and as such fault tips are not seismically resolvable and may remain undetected. Thus seismically resolved faults should
245 be extended beyond the normally picked extend to include the fault tips. Similarly, faults with low displacement such as the Buttes fault, for which significant fault related leakage has been recorded, are not detectable on seismic data. This highlights the need for a good structural understanding of any geological storage site to ensure that fault tips and small faults are identified and incorporated into the geological model.
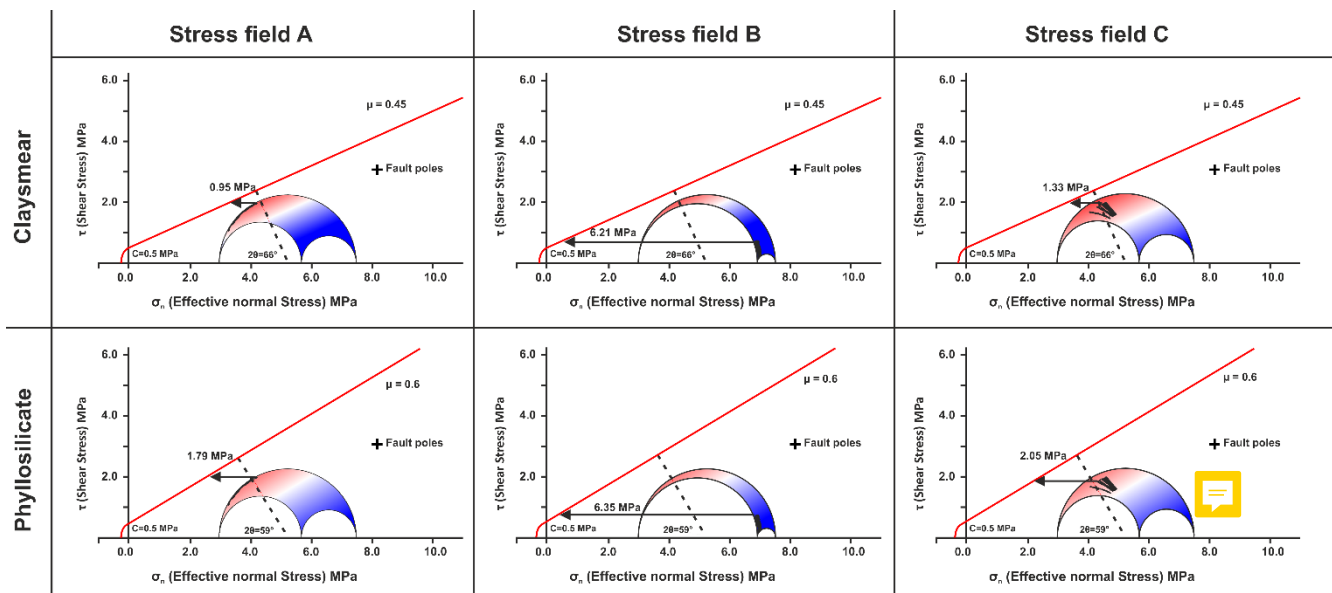
Figure 5: Mohr diagrams illustrating the fracture stability for the Coyote Wash Fault within the three stress fields for two types of fault rocks. Black arrows indicate the increase in pore-pressure needed to force the fault into failure.
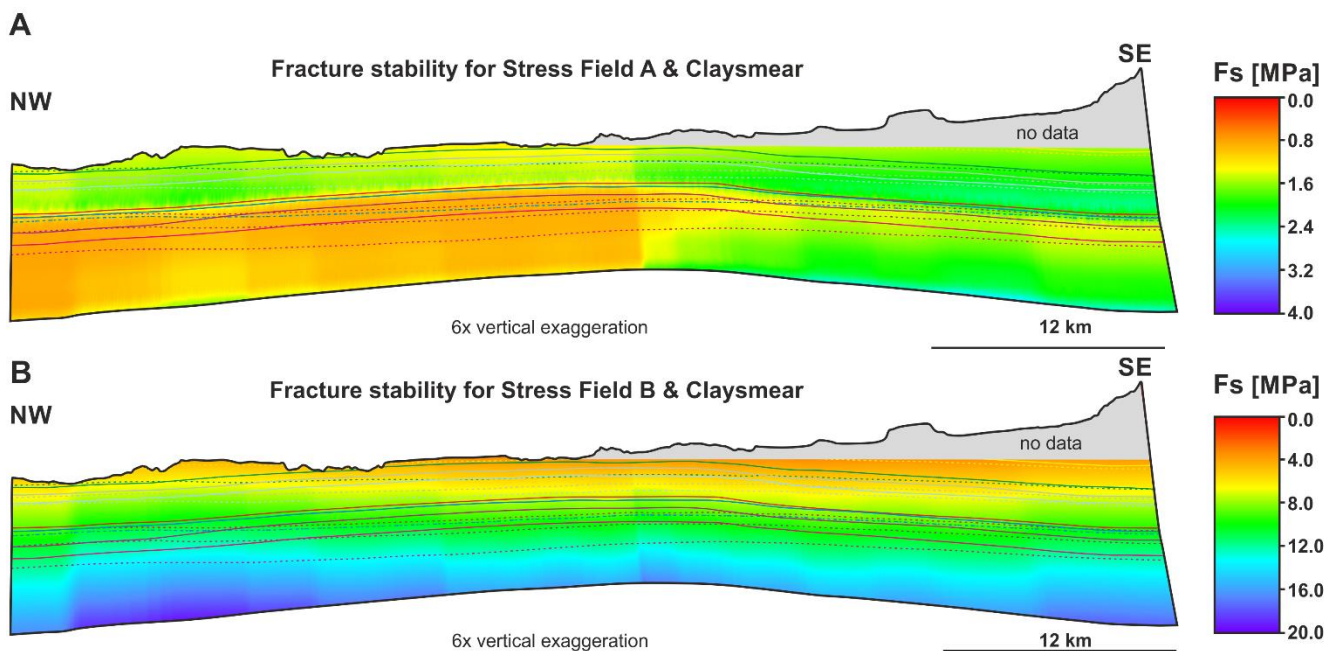
Figure 6: Fracture stability plotted onto the Coyote Wash Fault surface. (A) for stress field A and a claysmear fault rock. (B) For stress field B and a claysmear fault rock. Note the differences in the colour scale. Fs is more critical in the NW part of the fault for stress field A while Fs for stress field B is far from critical. Dashed lines represent hangingwall, straight line footwall stratigraphy.
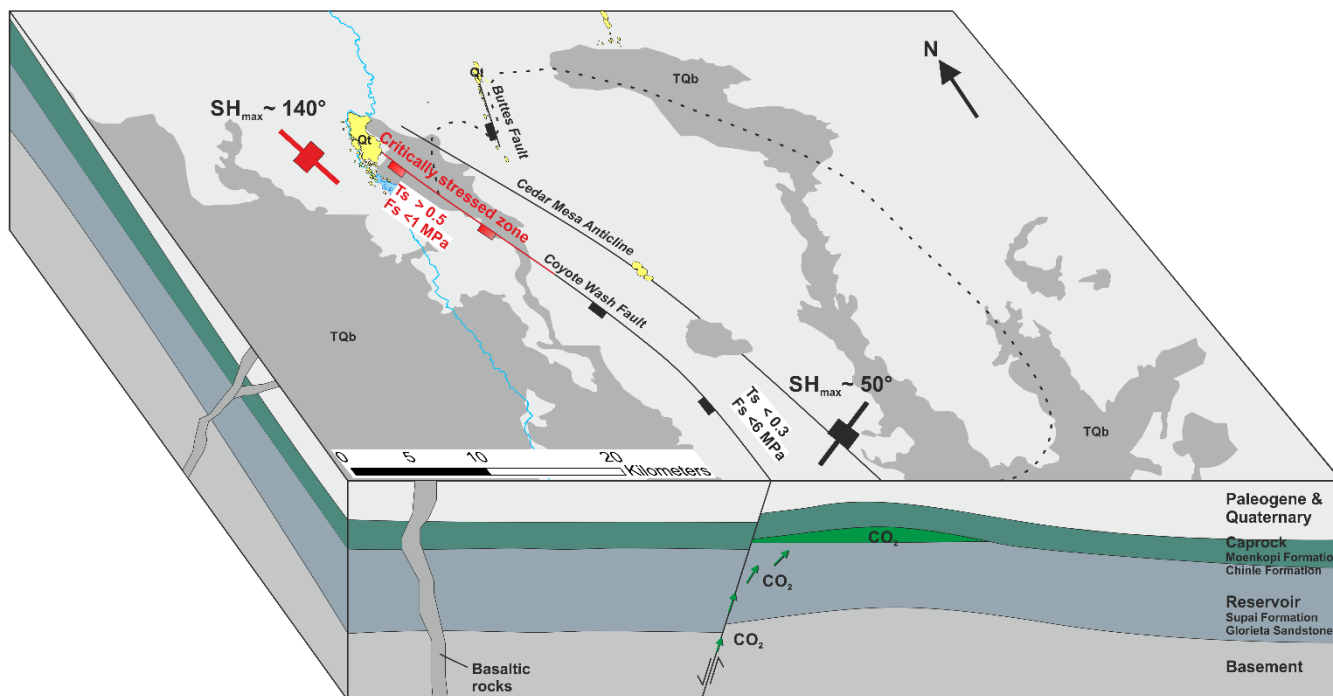
Figure 7: Block diagram illustrating the geological setting of the St. Johns Dome area. The fault-bound CO$_2$ reservoir is constantly
filled with CO$_2$ from depth and leakage of CO$_2$ from the reservoir to the surface occurs at the NW tip of the Coyote Wash Fault.
The leakage is geomechanically controlled as the stress field orientation changes along the strike of the fault as well as the strike
direction of the fault. At the NW tip of the fault slip stability values of >0.5 and fracture stability of less than 1 MPa indicate a
critically stressed fault. See figure 2 for a complete legend, ~10x vertical exaggeration.

While the geomechanical analysis highlights the role of critically stressed faults for fluid migration at the St. Johns Dome, it

is missing in-situ stress field data from within the CO$_2$ reservoir. Such data is crucial for a detailed and reliable study of fracture

and fault stability (e.g. Becker et al., 2019), however there are cases where such in-situ data is missing and a geomechanical

analysis may be needed (Henk, 2005). In particular during the site selection and appraisal of subsurface storage sites a

preliminary geomechanical analysis based on existing stress field data can identify potentially critically stressed faults. The

lack of in-situ data can be compensated by studying several potential stress fields (as in this study) and including uncertainties

into the geomechanical analysis. For the latter, uncertainties in the stress field orientation and magnitude and in the fault

orientation should be included. Statistical approaches such as Bayesian or Markov Chain Monte Carlo modelling can be useful

to identify uncertainty thresholds and to determine the precision by which the geomechanical parameters need to be known in

order to have reliable fault and fracture stability predictions (Bao et al., 2013; Chiaramonte et al., 2008; McFarland et al.,

2012).

For the geomechanical prediction of permeable fracture networks and thus leakage pathways at the St. Johns Dome the stress

field orientation is integral. The location of the natural CO$_2$ reservoir at the edge of the Colorado Plateau is the likely reason

for the stress field orientation change, with clear changes in crustal composition and strength in the vicinity of the St. Johns

Dome (Hendricks and Plescia, 1991; Qashqai et al., 2016). The study area is located at the intersection of the NE-NNE trending

Jemez Lineament, a tectonically active zone that is characterized by Paleogene-Quaternary extension and volcanism (Fig. 1),
280 and the ESE tending Arizona Transition Zone (Aldrich and Laughlin, 1984; Kreemer et al., 2010). Additionally, the presence
of salt deposits in the Holbrook Basin north of the study area may also impact the local stress field (Neal and Colpitts, 1997;
Rauzi, 2000). The complex tectonic setting at the St. Johns Dome and the associated uncertainties for geomechanical modelling
further highlight the need for good site selection criteria for engineered fluid storage sites to ensure that only reservoirs with
well understood structural framework are chosen.

285 **5 Implications for geological storage applications**

Geomechanical modelling suggests that vertical fluid migration from the reservoir to the surface at the St. Johns Dome natural
CO_2 reservoir is controlled by fracture networks in the damage zone and tip of close to critically stressed faults. We propose
that constant filling of the reservoir with CO_2 from mantle sources increased the pore pressure within the reservoir and further
reduced the stability of close to critically stressed faults, leading to the leakage of large volumes of CO_2 over the time-span of
290 several 100 kas. While the leakage rates at the St. Johns Dome are low enough to render the faulted site an adequate CO_2 store
for climate mitigation, similar leakage rates could impede geological storage of methane or hydrogen as any surface leakage
of these flammable gases could be hazardous, in particularly at onshore storage sites.
For fault-bound subsurface storage sites for CO_2 or energy the history of geomechanically controlled leakage at the St. Johns
Dome clearly illustrates the need for a good understanding of regional and local stress fields and faults. In particular the stress
295 state of faults and fault related fracture networks prior to fluid injection needs to be well understood in order to reduce the risk
of vertical fluid migration through fractured caprock. ~~For subsurface storage sites it is recommended~~ to select areas where
there are no significant regional stress field changes as these complicate geomechanical predictions. Indeed, in-situ stress data
from wells are key for any advanced leakage risk prognosis. As the pore pressure increase within the reservoir depends on the
buoyancy of the injected fluid, site selection and appraisal needs to be fluid specific. To further understand the leakage
300 mechanisms at the St. Johns Dome geomechanical modelling of the Buttes Fault, combined with an uncertainty assessment, is
recommended. More detailed dating of the travertine deposits could reveal at which structural setting (fault tip vs fault damage
zone) failure occurred first and provide insights into the time dynamics of leakage.

Author contribution

JM and SG designed the research project which was carried out by JM with help from GJ and input from SG. JM prepared the
305 manuscript with contributions from all co-authors.

Competing interests

The authors declare that they have no conflict of interest.

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References

Alcalde, J., Flude, S., Wilkinson, M., Johnson, G., Edlmann, K., Bond, C. E., Scott, V., Gilfillan, S. M. V., Ogaya, X. and
Haszeldine, R. S.: Estimating geological CO 2 storage security to deliver on climate mitigation, Nature Communications, 9(1),
315   2201, doi:10.1038/s41467-018-04423-1, 2018.

Aldrich, M. J. and Laughlin, A. W.: A model for the tectonic development of the Southeastern Colorado Plateau Boundary,
Journal of Geophysical Research: Solid Earth, 89(B12), 10207–10218, doi:10.1029/JB089iB12p10207, 1984.

Allis, R., Bergfeld, D., Moore, J., McClure, K., Morgan, C., Chidsey, T., Heath, J. and McPherson, B.: Implications of results
from CO2 flux surveys over known CO2 systems for long-term monitoring, in Fourth Annual Conference on Carbon Capture
320   and Sequestration., 2005.

Allis, R. G., Moore, J. and White, S. P.: Reactive Multiphase behavior of CO2 in Saline Aquifers beneath the Colorado Plateau,
Quaterly Technical Report, University of Utah, Salt Lake City., 2004.

Backers, T. and Moeck, I.: Fault tips as favorable drilling targets for geothermal prospecting–a fracture mechanical perspective,
International Society for Rock Mechanics and Rock Engineering., 2015.

325   Bao, J., Xu, Z., Lin, G. and Fang, Y.: Evaluating the impact of aquifer layer properties on geomechanical response during CO2
geological sequestration, Computers & Geosciences, 54, 28–37, doi:10.1016/j.cageo.2013.01.015, 2013.

Barton, C. A., Zoback, M. D. and Moos, D.: Fluid flow along potentially active faults in crystalline rock, Geology, 23(8), 683–
686, doi:10.1130/0091-7613(1995)023<0683:FFAPAF>2.3.CO;2, 1995.

Bashir, L., Gao, S. S., Liu, K. H. and Mickus, K.: Crustal structure and evolution beneath the Colorado Plateau and the southern
330   Basin and Range Province: Results from receiver function and gravity studies, Geochem. Geophys. Geosyst., 12(6), Q06008,
doi:10.1029/2011GC003563, 2011.

Becker, I., Müller, B., Bastian, K., Jelinek, W. and Hilgers, C.: Present-day stress control on fluid migration pathways: Case
study of the Zechstein fractured carbonates, NW-Germany - ScienceDirect, Marine and Petroleum Geology, 103, 320–330,
doi:https://doi.org/10.1016/j.marpetgeo.2019.03.002, 2019.

335   Bond, C. E., Wightman, R. and Ringrose, P. S.: The influence of fracture anisotropy on CO2 flow, Geophys. Res. Lett., 40(7),
1284–1289, doi:10.1002/grl.50313, 2013.

Bond, C. E., Kremer, Y., Johnson, G., Hicks, N., Lister, R., Jones, D. G., Haszeldine, R. S., Saunders, I., Gilfillan, S. M. V.,
Shipton, Z. K. and Pearce, J.: The physical characteristics of a CO2 seeping fault: The implications of fracture permeability
for carbon capture and storage integrity, Int. J. Greenh. Gas Con., 61, 49–60, doi:10.1016/j.ijggc.2017.01.015, 2017.

340   Brogi, A., Capezzuoli, E., Aqué, R., Branca, M. and Voltaggio, M.: Studying travertines for neotectonics investigations: Middle–Late Pleistocene syn-tectonic travertine deposition at Serre di Rapolano (Northern Apennines, Italy), Int J Earth Sci (Geol Rundsch), 99(6), 1383–1398, doi:10.1007/s00531-009-0456-y, 2010.

Burnside, N. M., Shipton, Z. K., Dockrill, B. and Ellam, R. M.: Man-made versus natural CO2 leakage: A 400 k.y. history of an analogue for engineered geological storage of CO2, Geology, 41(4), 471–474, doi:10.1130/G33738.1, 2013.

345   Byerlee, J.: Friction of rocks, PAGEOPH, 116(4–5), 615–626, doi:10.1007/BF00876528, 1978.

Caine, J. S., Evans, J. P. and Forster, C. B.: Fault zone architecture and permeability structure, Geology, 24(11), 1025–1028, 1996.

Cappa, F.: Modelling fluid transfer and slip in a fault zone when integrating heterogeneous hydromechanical characteristics in its internal structure, Geophys. J. Int., 178(3), 1357–1362, doi:10.1111/j.1365-246X.2009.04291.x, 2009.

350   Cavanagh, A. J., Haszeldine, R. S. and Nazarian, B.: The Sleipner CO2 storage site: using a basin model to understand reservoir simulations of plume dynamics, First Break, 33(6), 61–68, 2015.

Chadwick, A., Arts, R., Bernstone, C., May, F., Thibeau, S. and Zweigel, P.: Best practice for the storage of CO2 in saline aquifers - observations and guidelines from the SACS and CO2STORE projects, British Geological Survey, Nottingham, UK., 2008.

355   Chester, F. M. and Logan, J. M.: Implications for mechanical properties of brittle faults from observations of the Punchbowl fault zone, California, PAGEOPH, 124(1–2), 79–106, doi:10.1007/BF00875720, 1986.

Chiaramonte, L., Zoback, MarkD., Friedmann, J. and Stamp, V.: Seal integrity and feasibility of CO2 sequestration in the Teapot Dome EOR pilot: geomechanical site characterization, Environ Geol, 54(8), 1667–1675, doi:10.1007/s00254-007-0948-7, 2008.

360   Choi, J.-H., Edwards, P., Ko, K. and Kim, Y.-S.: Definition and classification of fault damage zones: A review and a new methodological approach, Earth-Science Reviews, 152, 70–87, doi:10.1016/j.earscirev.2015.11.006, 2016.

Condit, C. D. and Connor, C. B.: Recurrence rates of volcanism in basaltic volcanic fields: An example from the Springerville volcanic field, Arizona, Geological Society of America Bulletin, 108(10), 1225–1241, doi:10.1130/0016-7606(1996)1082.3.CO;2, 1996.

365   Condit, C. D., Crumpler, L. S. and Aubele, J. C.: Lithologic, age group, magnetopolarity, and geochemical maps of the Springerville Volcanic Field, East-Central Arizona, U.S. Dept. of the Interior, U.S. Geological Survey., 1993.

Connor, C. B., Condit, C. D., Crumpler, L. S. and Aubele, J. C.: Evidence of regional structural controls on vent distribution: Springerville Volcanic Field, Arizona, J. Geophys. Res., 97(B9), 12349–12359, doi:10.1029/92JB00929, 1992.

Crumpler, L. S., Aubele, J. C. and Condit, C. D.: Volcanics and neotectoniccharacteristics of the Springerville volcanic field, 370   Arizona, in New Mexico Geological Society Guidebook, 45th Field Conference, edited by R. M. Chamberlin, B. S. Kues, S. M. Cather, J. M. Barker, and W. C. McIntosh, pp. 147–164., 1994.

Davatzes, N. C. and Aydin, A.: Distribution and nature of fault architecture in a layered sandstone and shale sequence: An example from the Moab fault, Utah, AAPG Memoir, (85), 153–180, 2005.

Dockrill, B. and Shipton, Z. K.: Structural controls on leakage from a natural CO2 geologic storage site: Central Utah, U.S.A,
375   Journal of Structural Geology, 32(11), 1768–1782, doi:10.1016/j.jsg.2010.01.007, 2010.

Eichhubl, P., Davatz, N. C. and Becker, S. P.: Structural and diagenetic control of fluid migration and cementation along the
Moab fault, Utah, AAPG Bulletin, 93(5), 653–681, doi:10.1306/02180908080, 2009.

Embid, E. H.: U-series dating, geochemistry, and geomorphic studies of travertines and springs of the Springerville area, east-
central Arizona, and tectonic implications, MSc thesis, The University of New Mexico, Albuquerque., 2009.

380   Faulkner, D. R. and Rutter, E. H.: Can the maintenance of overpressured fluids in large strike-slip fault zones explain their
apparent weakness?, Geology, 29(6), 503–506, doi:10.1130/0091-7613(2001)029<0503:CTMOOF>2.0.CO;2, 2001.

Faulkner, D. R., Lewis, A. C. and Rutter, E. H.: On the internal structure and mechanics of large strike-slip fault zones: field
observations of the Carboneras fault in southeastern Spain, Tectonophysics, 367(3–4), 235–251, doi:10.1016/S0040-
1951(03)00134-3, 2003.

385   Faulkner, D. R., Jackson, C. A. L., Lunn, R. J., Schlische, R. W., Shipton, Z. K., Wibberley, C. A. J. and Withjack, M. O.: A
review of recent developments concerning the structure, mechanics and fluid flow properties of fault zones, J. Struct. Geol.,
32(11), 1557–1575, doi:10.1016/j.jsg.2010.06.009, 2010.

Frery, E., Gratier, J.-P., Ellouz-Zimmerman, N., Loiselet, C., Braun, J., Deschamps, P., Blamart, D., Hamelin, B. and Swennen,
R.: Evolution of fault permeability during episodic fluid circulation: Evidence for the effects of fluid–rock interactions from
390   travertine studies (Utah–USA), Tectonophysics, doi:10.1016/j.tecto.2015.03.018, 2015.

Gilfillan, S. M. V., Ballentine, C. J., Holland, G., Blagburn, D., Lollar, B. S., Stevens, S., Schoell, M. and Cassidy, M.: The
noble gas geochemistry of natural CO2 gas reservoirs from the Colorado Plateau and Rocky Mountain provinces, USA,
Geochimica et Cosmochimica Acta, 72(4), 1174–1198, doi:10.1016/j.gca.2007.10.009, 2008.

Gilfillan, S. M. V., Lollar, B. S., Holland, G., Blagburn, D., Stevens, S., Schoell, M., Cassidy, M., Ding, Z., Zhou, Z.,
395   Lacrampe-Couloume, G. and Ballentine, C. J.: Solubility trapping in formation water as dominant CO2 sink in natural gas
fields, Nature, 458(7238), 614–618, doi:10.1038/nature07852, 2009.

Gilfillan, S. M. V., Wilkinson, M., Haszeldine, R. S., Shipton, Z. K., Nelson, S. T. and Poreda, R. J.: He and Ne as tracers of
natural CO2 migration up a fault from a deep reservoir, International Journal of Greenhouse Gas Control, 5(6), 1507–1516,
doi:10.1016/j.ijggc.2011.08.008, 2011.

400   Guglielmi, Y., Cappa, F. and Amitrano, D.: High-definition analysis of fluid-induced seismicity related to the mesoscale
hydromechanical properties of a fault zone, Geophys. Res. Lett., 35(6), L06306, doi:10.1029/2007GL033087, 2008.

Harris, R. C.: A review and bibliography of karst features of the Colorado Plateau, Arizona, Open-File Report, Arizona
Geological Survey., 2002.

Heidbach, O., Rajabi, M., Reiter, K., Ziegler, M. and WSM Team: World Stress Map Database Release 2016. V.1.1., GFZ
405   Data Service, doi:http://doi.org/10.5880/WSM.2016.001, 2016.

Hendricks, J. D. and Plescia, J. B.: A review of the regional geophysics of the Arizona Transition Zone, Journal of Geophysical
Research: Solid Earth, 96(B7), 12351–12373, doi:10.1029/90JB01781, 1991.

Henk, A.: Pre-drilling prediction of the tectonic stress field with geomechanical models, First Break, 23(11), 53–57, 2005.

Hillis, R. R.: Pore pressure/stress coupling and its implications for rock failure, Geological Society, London, Special
410 Publications, 216(1), 359–368, doi:10.1144/GSL.SP.2003.216.01.23, 2003.

IEA GHG: CCS Site Characterisation Criteria, IEA Greenhouse Gas R&D Programme., 2009.

IPCC: IPCC Special report on Carbon Dioxide Capture and Storage, Cambridge University Press, New York, USA Cambridge,
UK., 2005.

Ito, T. and Zoback, M. D.: Fracture permeability and in situ stress to 7 km depth in the KTB scientific drillhole, Geophysical
415 Research Letters, 27(7), 1045–1048, doi:10.1029/1999GL011068, 2000.

Keating, E., Newell, D., Dempsey, D. and Pawar, R.: Insights into interconnections between the shallow and deep systems
from a natural CO2 reservoir near Springerville, Arizona, International Journal of Greenhouse Gas Control, 25, 162–172,
doi:10.1016/j.ijggc.2014.03.009, 2014.

Kreemer, C., Blewitt, G. and Bennett, R. A.: Present-day motion and deformation of the Colorado Plateau, Geophys. Res.
420 Lett., 37(10), L10311, doi:10.1029/2010GL043374, 2010.

Lehner, F. K. and Pilaar, W. F.: The emplacement of clay smears in synsedimentary normal faults: inferences from field
observations near Frechen, Germany, in Norwegian Petroleum Society Special Publications, vol. Volume 7, edited by P.
Møller-Pedersen   and   A.G.   Koestler,   pp.   39–50,   Elsevier.   [online]   Available   from:
http://www.sciencedirect.com/science/article/pii/S0928893797800057, 1997.

425 Lindsay, N. G., Murphy, F. C., Walsh, J. J. and Watterson, J.: Outcrop Studies of Shale Smears on Fault Surface, in The
Geological Modelling of Hydrocarbon Reservoirs and Outcrop Analogues, pp. 113–123, Blackwell Publishing Ltd. [online]
Available from: http://dx.doi.org/10.1002/9781444303957.ch6, 1993.

Marshak, S., Karlstrom, K. and Timmons, J. M.: Inversion of Proterozoic extensional faults: An explanation for the pattern of
Laramide and Ancestral Rockies intracratonic deformation, United States, Geology, 28(8), 735–738, doi:10.1130/0091-
430 7613(2000)28<735:IOPEFA>2.0.CO;2, 2000.

Matos, C. R., Carneiro, J. F. and Silva, P. P.: Overview of Large-Scale Underground Energy Storage Technologies for
Integration of Renewable Energies and Criteria for Reservoir Identification, Journal of Energy Storage, 21, 241–258,
doi:10.1016/j.est.2018.11.023, 2019.

McFarland, J. M., Morris, A. P. and Ferrill, D. A.: Stress inversion using slip tendency, Computers & Geosciences, 41, 40–46,
435 doi:10.1016/j.cageo.2011.08.004, 2012.

Michael, K., Golab, A., Shulakova, V., Ennis-King, J., Allinson, G., Sharma, S. and Aiken, T.: Geological storage of CO2 in
saline aquifers—A review of the experience from existing storage operations, International Journal of Greenhouse Gas Control,
4(4), 659–667, doi:10.1016/j.ijggc.2009.12.011, 2010.

Min, K.-B., Rutqvist, J., Tsang, C.-F. and Jing, L.: Stress-dependent permeability of fractured rock masses: a numerical study,
440 International Journal of Rock Mechanics and Mining Sciences, 41(7), 1191–1210, doi:10.1016/j.ijrmms.2004.05.005, 2004.

Miocic, J. M., Gilfillan, S. M. V., Roberts, J. J., Edlmann, K., McDermott, C. I. and Haszeldine, R. S.: Controls on CO2 storage
security in natural reservoirs and implications for CO2 storage site selection, Int. J. Greenh. Gas Con., 51, 118–125,
doi:10.1016/j.ijggc.2016.05.019, 2016.

Miocic, J. M., Gilfillan, S. M. V., Frank, N., Schroeder-Ritzrau, A., Burnside, N. M. and Haszeldine, R. S.: 420,000 year
445 assessment of fault leakage rates shows geological carbon storage is secure, Sci. Rep., 9(1), 769, doi:10.1038/s41598-018-36974-0, 2019a.

Miocic, J. M., Johnson, G. and Bond, C. E.: Uncertainty in fault seal parameters: implications for CO_2 column height retention and storage capacity in geological CO_2 storage projects, Solid Earth, 10(3), 951–967, doi:https://doi.org/10.5194/se-10-951-2019, 2019b.

450 Mitchell, T. M. and Faulkner, D. R.: The nature and origin of off-fault damage surrounding strike-slip fault zones with a wide range of displacements: A field study from the Atacama fault system, northern Chile, Journal of Structural Geology, 31(8), 802–816, doi:10.1016/j.jsg.2009.05.002, 2009.

Moeck, I., Kwiatek, G. and Zimmermann, G.: Slip tendency analysis, fault reactivation potential and induced seismicity in a deep geothermal reservoir, Journal of Structural Geology, 31(10), 1174–1182, doi:10.1016/j.jsg.2009.06.012, 2009.

455 Moore, J., Adams, M., Allis, R., Lutz, S. and Rauzi, S.: Mineralogical and geochemical consequences of the long-term presence of CO2 in natural reservoirs: An example from the Springerville–St. Johns Field, Arizona, and New Mexico, U.S.A., Chemical Geology, 217(3–4), 365–385, doi:10.1016/j.chemgeo.2004.12.019, 2005.

Morris, A., Ferrill, D. A. and Henderson, D. B.: Slip-tendency analysis and fault reactivation, Geology, 24(3), 275–278, 1996.

Neal, J. T. and Colpitts, R. M.: Richard Lake, an evaporite-karst depression in the Holbrook Basin, Arizona, Carbonates
460 Evaporites, 12(1), 91–97, doi:10.1007/BF03175807, 1997.

Priewisch, A., Crossey, L. J., Karlstrom, K. E., Polyak, V. J., Asmerom, Y., Nereson, A. and Ricketts, J. W.: U-series geochronology of large-volume Quaternary travertine deposits of the southeastern Colorado Plateau: Evaluating episodicity and tectonic and paleohydrologic controls, Geosphere, 10(2), 401–423, doi:10.1130/GES00946.1, 2014.

Qashqai, M. T., Afonso, J. C. and Yang, Y.: The crustal structure of the Arizona Transition Zone and southern Colorado
465 Plateau from multiobservable probabilistic inversion, Geochemistry, Geophysics, Geosystems, 17(11), 4308–4332, doi:10.1002/2016GC006463, 2016.

Rauzi, S.: Carbon Dioxide in the St. Johns-Springervile Area, Apache County, Arizona, Open File Report, Arizona Geological Survey., 1999.

Rauzi, S. L.: Permian Salt in the Holbrook Basin, Arizona, Open-File Report, Arizona Geological Survey., 2000.

470 Rutqvist, J., Vasco, D. W. and Myer, L.: Coupled reservoir-geomechanical analysis of CO2 injection and ground deformations at In Salah, Algeria, International Journal of Greenhouse Gas Control, 4(2), 225–230, doi:10.1016/j.ijggc.2009.10.017, 2010.

Schulz, S. E. and Evans, J. P.: Spatial variability in microscopic deformation and composition of the Punchbowl fault, southern California: implications for mechanisms, fluid–rock interaction, and fault morphology, Tectonophysics, 295(1–2), 223–244, doi:10.1016/S0040-1951(98)00122-X, 1998.

475 Scott, V., Gilfillan, S., Markusson, N., Chalmers, H. and Haszeldine, R. S.: Last chance for carbon capture and storage, Nature Clim. Change, 3(2), 105–111, 2013.

Shipton, Z. K., Evans, J. P., Robeson, K. R., Forster, C. B. and Snelgrove, S.: Structural Heterogeneity and Permeability in Faulted Eolian Sandstone: Implications for Subsurface Modeling of Faults, AAPG Bulletin, 86(5), 863–883, doi:10.1306/61EEDBC0-173E-11D7-8645000102C1865D, 2002.

480 Shipton, Z. K., Evans, J. P., Kirschner, D., Kolesar, P. T., Williams, A. P. and Heath, J.: Analysis of CO2 leakage through 'low-permeability' faults from natural reservoirs in the Colorado Plateau, east-central Utah, Geological Society, London, Special Publications, 233(1), 43–58, doi:10.1144/GSL.SP.2004.233.01.05, 2004.

Sibson, R. H.: Brittle-failure controls on maximum sustainable overpressure in different tectonic regimes, AAPG Bulletin, 87(6), 901–908, 2003.

485 Sirrine, G. K.: Geology of the Springerville-St. Johns area, Apache County, Arizona, PhD thesis, University of Texas., 1956.

Soden, A. M., Shipton, Z. K., Lunn, R. J., Pytharouli, S. I., Kirkpatrick, J. D., Do Nascimento, A. F. and Bezerra, F. H. R.: Brittle structures focused on subtle crustal heterogeneities: implications for flow in fractured rocks, Journal of the Geological Society, 171(4), 509–524, doi:10.1144/jgs2013-051, 2014.

Song, J. and Zhang, D.: Comprehensive Review of Caprock-Sealing Mechanisms for Geologic Carbon Sequestration, Environ.
490 Sci. Technol., 47(1), 9–22, doi:10.1021/es301610p, 2012.

Stork, A. L., Verdon, J. P. and Kendall, J. M.: The microseismic response at the In Salah Carbon Capture and Storage (CCS) site, International Journal of Greenhouse Gas Control, 32, 159–171, doi:10.1016/j.ijggc.2014.11.014, 2015.

Vrolijk, P. J., Urai, J. L. and Kettermann, M.: Clay smear: Review of mechanisms and applications, J. Struct. Geol., 86, 95–152, doi:10.1016/j.jsg.2015.09.006, 2016.

495 White, J. A., Chiaramonte, L., Ezzedine, S., Foxall, W., Hao, Y., Ramirez, A. and McNab, W.: Geomechanical behavior of the reservoir and caprock system at the In Salah CO2 storage project, Proceedings of the National Academy of Sciences, 111(24), 8747–8752, 2014.

Wiprut, D. and Zoback, M. D.: Fault reactivation and fluid flow along a previously dormant normal fault in the northern North Sea, Geology, 28(7), 595–598, doi:10.1130/0091-7613(2000)28<595:FRAFFA>2.0.CO;2, 2000.

500 Wong, I., G. and Humphrey, J.: Contemporary seismicity, faulting, and the state of stress in the Colorado Plateau, Geological Society of America Bulletin, 101(9), 1127–1146, doi:10.1130/0016-7606(1989)101<1127:CSFATS>2.3.CO;2, 1989.

Yielding, G., Freeman, B. and Needham, D. T.: Quantitative fault seal prediction, AAPG Bulletin, 81(6), 897–917, 1997.

Zhang, Y., Schaubs, P. M., Zhao, C., Ord, A., Hobbs, B. E. and Barnicoat, A. C.: Fault-related dilation, permeability enhancement, fluid flow and mineral precipitation patterns: numerical models, Geological Society, London, Special
505 Publications, 299(1), 239–255, doi:10.1144/SP299.15, 2008.