Statistics of the Seismic Sequence and Rupture Directivity of the M5.5 Earthquake in Orkney, South Africa

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Abstract. Despite our general knowledge of earthquake processes, it is still not fully understood how earthquake ruptures nucleate and propagate and why they stop. Also, the controlling factors of the frequency and of the size of earthquakes are subject of ongoing research. We aim to address these questions with a comprehensive study of seismicity in deep South African gold mines. We find here the unique situation that the seismicity consists of both induced earthquakes and aftershocks triggered

5   by the M5.5 Orkney earthquake which occurred in August 2014. We separate the cataloged seismicity and group the events into three classes: the aftershock sequence, seismicity induced by fluids, and seismicity induced by mining activities. We examine statistical properties of earthquakes in each of the three classes. We conclude that the magnitude statistics of both aftershocks and induced earthquakes are influenced by the finite size and geometry of the rock volume of stress perturbation resulting in an absence of larger magnitude events. The magnitude frequency distributions obey the Lower Bound model of magnitude

10   probability. The statistics of dynamic stress drop of aftershocks and induced earthquakes satisfy log-normal distributions but the value range is different, it means that aftershocks are generally characterized by higher stress drops. Another key aspect in our study is the imaging of the propagating rupture of the M5.5 earthquake. We apply the back projection imaging approach using seismological data from two different local networks and retrieve similar results. We conclude that the rupture of the M5.5 earthquake propagated predominantly unilaterally, nearly from North to South over a distance of about 6km. The hypocenters

15   of the aftershock sequence are situated unilaterally in respect to the hypocenter of the main shock and are aligned to the South which confirms the obtained rupture propagation image and the directivity of the main shock.

1 Introduction

On 5 August 2014, an earthquake with local magnitude $M5.5$ occurred beneath the Moab Khotsong gold mine in Orkney which is located 150 km southwest of Johannesburg in South Africa. The Orkney earthquake was the second largest recorded

20   seismic event in South Africa and the largest within an active mining area (Manzunzu et al., 2017). South Africa is generally regarded as a stable continental region (SCR) with a rare number of earthquakes exceeding magnitude 4.5. The majority of the seismicity observed in active mine districts is induced by mining operations and the largest earthquakes have been recorded with magnitudes around $M4$ (Richardson and Jordan, 2002). Hence the occurrence of the Orkney earthquake in this intraplate continental context is atypical leaving room for speculations about its nature (Manzunzu et al., 2017). The earthquake was

25 atypical also because it was a sinistral strike slip event on an unknown geological structure below the mine whereas the seis-
micity in the mining horizons are typically normal faulting events.

Earthquakes induced by activities and operations in deep-level mines are a topic of research since decades starting with the
pioneering works of Spottiswoode and McGarr (1975) and Gay and Ortlepp (1979). In the 1990, seismic monitoring and quan-
titative seismic hazard assessment became routine work in South African gold mines (Mendecki, 1996; Mendecki et al., 1999;

30 Van Aswegen et al., 2005). Substantial progress in the understanding of mining-induced earthquakes was achieved within the
last decade mainly by the JAGUARS (Japanese German Acoustic Emission Research in South Africa) and the SATREPS (Sci-
ence and Technology Research Partnership for Sustainable Development) projects (see e.g. Nakatani et al., 2008; Yabe et al.,
2009; Moriya et al., 2015; Naoi et al., 2011, 2015; Ziegler et al., 2015). Highly sensitive acoustic emission sensors were addi-
tionally deployed in the gold mines. This enabled analysis of scaling laws of events occurring in deep mines and the verification

35 that scaling relations involving event size, rupture dimension and earthquake stress drop retain their validity also in case of ex-
tremely small, high-frequency acoustic emission (AE) events (Boettcher et al., 2009; Kwiatek et al., 2010, 2011; Naoi et al.,
2014). It was also concluded that the combined analysis of different types of earthquakes in mining environment including
aftershocks leads to unusual shape and curvature of the magnitude frequency distribution (Boettcher et al., 2009). Naoi et al.
(2014) argued that the size distributions of seismic events occurring in the vicinity of mining fronts follow a Gutenberg-Richter

40 law but with minor variations of the b value with time.The monitoring of AE also allowed to study the nucleation process of a
$M2.2$ earthquake in a deep mine in South Africa documented by Yabe et al. (2015). The authors concluded that the earthquake
ruptured a preexisting plane of weakness and the AE activity during six month prior to the main shock already delineated the
rupture plane in advance.

In August 2016, the International Continental Drilling Programme (ICDP) has approved drilling into M2.0-M5.5 seismogenic

45 zones from deep gold mines in South Africa (DSeis). The drilling project aims to probe quasi-static and dynamic ruptures,
to recover cores fully, to measure stress, and to log boreholes and cores (Voosen, 2017). Several holes have been successfully
drilled and the last drilling was completed in June 2018. The targets also included the seismogenic zone of the Orkney earth-
quake (Figure 1).

Here, we present our results of a comprehensive study of the seismicity observed in the Orkney area. It includes not only the

50 aftershock sequence of the M5.5 earthquake but also induced earthquakes due to mining operations. Essentially, we study the
magnitude statistics of the different classes of seismicity in order to identify characteristic patterns that might be supportive in
hazard assessment and mitigation. Additionally, we establish a rough connection between the area of aftershock activity and
the area of stress perturbation of the M5.5 earthquake using the back projection imaging technique for tracking the propagating
rupture of this earthquake.

55 Recently, a more comprehensive earthquake catalog compared to the initial catalog was compiled and provided by Anglogold
Ashanti (Pty) Ltd. The Anglogold Ashanti in-mine routine seismic monitoring systems have been in operation since 1990s, with
several tens of underground stations at 2-3 km depth, in the deepest levels of Moab Khotsong, Great Noligwa and Kopanang
gold mines. The new catalog covers the time period from August 2014 to September 2017. It contains approximately 70,000
earthquakes which can be subdivided into three classes. These are (1) the aftershock sequence with about 15,000 earthquakes,

60 (2) mining induced seismicity with about 7,000 earthquakes, and (3) fluid induced seismicity with nearly 48,000 earthquakes. Figure 2 shows the hypocenter locations of the three classes of earthquakes.

2 Magnitude statistics of the M5.5 aftershock sequence

First, we study the statistics of magnitudes of the aftershock sequence which we educe from the whole catalog. The sequence contains approximately 15,000 events and their hypocenter locations are shown in Figure 3. The area of aftershock activity

65 extends over 5km in NNW-SSE direction, 0.5km perpendicular to the strike direction and 1.5km in vertical direction dipping to the South. The M5.5 main shock is situated at the northern edge of the aftershock cloud. The locations of aftershocks with respect to the main shock location raises the speculation that the rupture of the main shock propagated unilaterally to the South. The magnitude frequency distribution of the aftershock sequence illustrated in Figure 4(a). We estimate the b-value of the Gutenberg-Richter law (Gutenberg and Richter, 1944) using the maximum likelihood method introduced by Utsu (1966)

70 and later modified by Bender (1983). The corresponding graph atop the observed magnitude distribution reveales that the number of aftershocks with a magnitude $M > 1.5$ deviates from the strict power law of the Gutenberg-Richter relation. Such a phenomenon is often observed for the magnitude frequency distribution of fluid induced seismicity where the seismically active rock volume is restricted to the domain of pore pressure perturbation (Shapiro et al., 2011, 2013; Shapiro, 2015). A finite perturbed rock volume influences the magnitude statistics of fluid induced earthquakes. Here, we find similarly for the

75 M5.5 aftershock sequence. The occurrence of aftershocks is restricted in space to the domain of stress perturbation caused by the preceding main shock. This leads to a finite size and to a nearly fixed geometry of the seismically active rock volume which influences correspondingly the magnitude statistics of the aftershock sequence. Therefore, we apply the so-called Lower Bound model of the magnitude probability (Shapiro et al., 2013), which captures the effect of a finite perturbed rock volume on the magnitude statistics, to the Orkney aftershock sequence. We conclude that the lower bound model represents a better

80 approximation to the observed magnitude frequency distribution as shown in Figure 4(b)). Furthermore, with the application of the Lower Bound model we retrieve a geometry-corrected estimate of the b-value of the magnitude frequencies of aftershocks which we report in Table 1. Note, that the difference in b-value is marginal only but it influences seismic hazard assessment. The Lower Bound model yields also an estimate of the maximum possible magnitude of an earthquake occurring within the spatially confined rock volume of a stress perturbation. In the present case of the aftershock sequence of the Orkney earthquake,

85 the maximum possible magnitude of an aftershock is estimated to be $M_{max} = 2.4$. This agrees with the observation, however, it contradicts Båth law (Båth and Duda, 1964; Båth, 1981) which empirically states that the magnitude of the largest aftershock is around one order lower than the nagnitude of the preceding main shock. Additionally, the lower bound model provides an alternative way to derive the stress drop $\Delta\sigma$ (Shapiro et al., 2013). This parameter is an averaged representation for the aftershock sequence and it is also reported in Table 1.

90 Now, we compare the statistical characteristics of magnitudes of the aftershock sequence to other seismicity observed in the gold mines in the Orkney area. These are two different types of earthquakes caused by human operations, either induced by fluid or induced by mining, respectively. Figure 5 depicts the corresponding magnitude frequency distributions as well as the

Gutenberg-Richter law using maximum likelihood b-value estimates. The magnitude frequency distributions are influenced in such way that larger magnitudes are under-represented in comparison to the power-law of the Gutenberg-Richter relation. It

95   reflects again the influence of a finite perturbed rock volume on the magnitude statistics. We find that this influence is more pronounced for mining induced seismicity. We apply the Lower Bound model of magnnitude probabilities to the magnitude frequency distributions of induced earthquakes (Figure 6) which result in geometry-corrected estimates of the Gutenberg-Richter b-value and the maximum possible magnitude, and in estimates of averaged stress drop $\Delta\sigma$ (Table 1). Fluid induced and mining induced events differ clearly regarding the statistics of magnitudes. We obtain with the Lower Bound model

100   the highest b-value and lowest stress drop estimates for fluid induced seismicity. In contrast, mining induced seismicity is characterized by the lowest b-value resulting therefore in a higher estimate of the maximum possible magnitude.

We also compare the statistics of stress drops for the three different classes of seismicity (Figure 7). We find it remarkable that stress drops are similar distributed, precisely log-normal distributions, independent of the event type. However, the value ranges are different, it means the stress drop of aftershocks is approximately two orders of magnitude higher than the stress

105   drop of induced earthquakes. We observe the lowest stress drop values for the seismicity that is induced by fluid. Finally, we analyze how the cataloged earthquake stress drops relate to the averaged stress drop estimate resulting from the Lower Bound model. The median values of static and dynamic stress drops are summarized in Table 1 and we find that the Lower Bound derived stress drops are reasonable values and are in good agreement with the median of dynamic stress drops. This finding holds for all three classes of seismicity.

110   ## 3   Imaging the M5.5 rupture

There exist few methods to track the propagation and the directivity of earthquakes ruptures including the back projection imaging technique. We apply the Microseismic Rupture Propagation Imaging (MRPI) technique (Folesky et al., 2015) to the Orkney earthquake. The method is based on the back-projection technique (Ishii et al., 2005; Krüger and Ohrnberger, 2005) but was modified to cope with the special conditions imposed by local seismic networks. For tracking the rupture propagation, we

115   back-project the recorded wavefield to a grid of possible source locations. Mathematically, we stack the absolute values of P-wave amplitudes of displacement seismograms recorded at a selected number of stations as functions of time for every potential source grid point corrected by the predicted theoretical source-receiver traveltime (Ishii et al., 2007). Theoretical traveltimes are obtained from the finite difference eikonal solver of Podvin and Lecomte (1991) which is implemented in the NonLinLoc software (Lomax, 2011). NonLinLoc is also used to calculate hypocenter location, origin time, and traveltime residuals. The

120   residuals are used to shift the P-wave amplitude traces to account for velocity model uncertainties. Our results of the inversion for hypocenter location and origin time differ from results previously known from other sources (CGS, Manzunzu et al. (2017), and Moyer et al. (2017)). However, the computed traveltimes for the relocated earthquake hypocenter are now more consistent with the observation and a consistent traveltime-grid is needed for the application of MRPI.

At first, we use the main shock registrations at stations installed in the deep mines. The networks are operated by Anglogold

125   Ashanti (Pty) Ltd (AGA) using a seismic monitoring system commercially available from the Institute of Mine Seismology

(Pty) Ltd (IMS) which is specialized in seismic monitoring of industrial underground activities (Mendecki, 1996). High-quality waveform data with good signal-to-noise ratios were recorded at $4.5Hz$ 3-component geophones with a sampling rate of $6000Hz$. Figure 8 shows the spatial distribution of stations of the three in-mine networks and the cut-out P-waveforms which we use in the MRPI method. To invert for the hypocenter location and to compute the traveltime-grid, we apply a homogeneous

130   1-layer velocity model with $v_P = 5.89km/s$ and $v_S = 3.69km/s$ for P-wave and S-wave,respectively, that was provided by AGA and IMS.

The results of the MRPI method for wavefields from the in-mine stations are illustrated in Figure 9, for one image we use absolute values of P-wave amplitudes and for the other image the envelopes. Both show the maximum of stacked back-projected energy in space and time providing a track of the rupture propagation. The rupture approximately propagated unilaterally from

135   N to SSW in the beginning (bluish colors) and then continues in S direction (greenish and yellowish colors). At the end of the rupture track (reddish colors), we observe a lot of scatter which correlates with an amplitude decrease of the energy stack value and it indicates the end of the rupture process.

Comparing the rupture track with the spatial distribution of aftershocks, we find that the elongation of the aftershock cloud in N-S direction fits well the rupture length. But, the aftershock cloud is shifted southwards by one to two kilometers which

140   we identify by comparing Figures 9 and 10(a). Also, the strike directions of both the cloud and the rupture track are slightly different. One reason for this finding could be the spatial distribution of stations of the in-mine network which is suboptimal because of the limited azimuthal coverage with respect to the earthquake location (Figure 8(a)). However, Manzunzu et al. (2017) determined focal mechanisms for the M5.5 earthquake and a selected number of aftershocks (Figure 10(b)). They found that the main shock occurred on a steeply dipping strike-slip fault with a strike direction of 182.6° N. The strike direction of

145   the rupture track coincides in a first approximation with the reported focal mechanism.

Additionally, we also apply the back projection imaging using wavefields recorded at the surface network. Figure 11 shows the spatial distribution of stations of this network which is operated by the Council for Geoscience South Africa (CGS) and cut-out recordings of P-waveforms that we use for the MRPI. The azimuthal coverage of the station distribution is slightly enhanced in comparison to the in-mine network but still not perfect. The network is equiped mainly with 3-components broadband

150   seismometers, data are recorded with $200Hz$ sample rate. Since we consider here the surface network, we apply a more advanced but still rather simple velocity model which is listed in Table 2. It consists of two layers with homogeneous velocities and it was build based on input from AGA and Moyer et al. (2017).

The resulting energy stack images are illustrated in Figures 12. Interestingly, by ignoring the scatter points at the beginning and at the end, we observe that the rupture nucleates around our determined earthquake hypocenter and propagates first in

155   Northwest direction for about 2 km where it turns and then propagates to the South. Such an observed complexity of rupture behavior indicating an initial and a main rupture phase was also reported by Ellsworth (*pers. communication*). He carried out a slip inversion for the M5.5 earthquake on a vertical plane and he found that the initial rupture origin is located to the South of the main rupture origin which propagated mainly southward.

4   Discussion and Conclusions

160   The combination of three sources of earthquake data − waveform recordings from a surface network, from a deep in-mine network, and a comprehensive catalog with about 70,000 earthquakes − allowed for detailed investigations of seismogenic processes in a spatially confined, seismically active volume. We separated the cataloged seismicity and grouped the events into three different classes: (1) the aftershock sequence of the Orkney M5.5 earthquake ($\approx$ 15,000 EQ), (2) seismicity induced by fluids ($\approx$ 48,000 EQ), and (3) seismicity induced by mining activity ($\approx$ 7,000 EQ). We studied statistical properties of earth-

165   quakes in each of the three classes and compared jointly the results. We focused on the magnitude statistics and on the statistics of dynamic stress drop. Another key aspect in our study was the imaging of the propagating rupture of the M5.5 earthquake. We applied successfully the back-projection method using seismological recordings from two different, independently operating networks and retrieved comparable results.

We concluded that the rupture of the M5.5 earthquake propagated predominantly unilaterally, nearly from North to South over

170   a distance of about $6\,km$. In general, the coherency of waveforms is high for seismic networks located at teleseismic distances to the earthquake source which facilitates application of the back-projection imaging. In our case, however, the waveforms were recorded at spatially suboptimally distributed stations of local monitoring system. Hence the waveforms can be quite variable due to earthquake radiation pattern, higher frequencies and local heterogeneity of seismic velocities and attenuation structure. To avoid this complexity, there exists a different and simpler geometrical method, the P-wave polarization stacking

175   (PPA), for imaging the propagating rupture front. The method was introduced by Bayer et al. (2012) and we recently applied it successfully to seismicity on a regional scale in N-Chile (Folesky et al., 2018a) and on a local scale in Italy (Folesky et al., 2018b). Wehner et al. (2018) tested the feasibility for the situation in Orkney and they concluded that the P-wave polarization stacking approach can be considered another method for tracking the rupture front of the $M5.5$ earthquake. The resulting final image of the rupture track revealed a southward propagation over a distance of about $4\,km$ (Wehner, 2018). The images

180   of back-projected seismic energy using the surface recordings indicated a complex rupture process of the Orkney earthquake. Two nucleation slip patches, an initial and a main rupture origin, appeared to locate on the rupture surface. They were seperated horizontally by around $2\,km$ in which the initial origin is south of the main origin. An assimiable observation supporting our findings was reported by Moyer et al. (2017). They identified an immediate foreshock precursory to the main shock which was located in about $1.4\,km$ horizontal distance south of the main shock. Integrating both observations supports the conclusion that

185   the interaction of two patches led to the M5.5 Orkney earthquake. Such a phenomenon is well-studied for subduction zone earthquakes. It was discussed and concluded that interactions and synchronization between multiple patches along the subduction interface led to the $M_w9.0$ Tohoku-oki earthquake in March 2011 (see Aochi and Ide, 2011; Hori and Miyazaki, 2011; Kiser and Ishii, 2012, e.g.). A similar complex triggering model was reported by Kiser and Ishii (2011) also for the $M_w8.8$ Maule earthquake which occurred in 2010.

190   The hypocenters of the aftershock sequence of the M5.5 earthquake are situated unilaterally in respect to the hypocenter of the main shock and are aligned to the South. In general, this observation confirms the obtained rupture propagation image and the directivity of the main shock. Furthermore, the extent of the aftershock cloud in N-S direction is in agreement with the length
of the imaged rupture track even though the two are offset slightly. We found that the magnitude statistics of both aftershocks and induced earthquakes are affected by the finite size and geometry of the rock volume of stress perturbation which inhibits

195 the occurrence of larger magnitude events. This influence is most pronounced for the magnitudes of aftershocks whereas it is least for fluid induced seismicity. The observed magnitude frequency distributions of aftershocks and induced earthquakes were fitted with the Lower Bound model of magnitude probabilities resulting in geometry-corrected b-values of the Gutenberg-Richter relation. The computed maximum possible magnitude of an earthquake occurring in given finite perturbed rock volume is in the case of aftershocks and mining induced seismicity significantly lower than an expected maximum magnitude based on

200 a classical non-truncated power-law Gutenberg-Richter law. The statistics of stress drop of aftershocks as well as of induced earthquakes comply with a log-normal distribution but the value range is different, it means aftershocks are generally characterized by higher stress drops and seismicity induced by fluid shows the lowest stress drops. The averaged stress drops, estimated with the Lower Bound model, retain the ranking of the three different classes of seismicity. Interestingly, the Lower Bound stress drops are closer to the median of dynamic stress drops than to the median of static stress drops. Overall, our findings and

205 conclusions have basically the potential to contribute to an improved assessment and a possible mitigation of seismic hazard in active mining districts. In particular, we think that the Lower Bound model combined either with the Omori-Utsu model for aftershock productivity or with a Poissonian model for induced earthquakes would be a first step towards more realistic hazard estimations of mining operations.

Code and data availability. Software codes in order to process and analyze the seismological data were written in MATLAB©®and are

210 available upon request from the corresponding author. The seismological data from the CGS network are downloadable from the CGS webpage and are also available from the supplementary material of Moyer et al. (2017). The in-mine networks data used in this research (catalog as well as stations recordings) are not publicly available. The data were provided by AngloGold Ashanti Ltd. and Harmony Gold Ltd. based on a collaboration within the framw of the DSeis drilling project.

Author contributions. CD carried out the data analysis and wrote the manuscript. JF provided the imaging tools used in this study, and JF

215 and JK assisted substantially in the interpretation of the obtained results. SS and HO are PI of the ICDP funded project DSeis − *Drilling into Seismogenic Zones of M2.0-M5.5 Earthquakes in Deep South African Gold Mines.*

Competing interests. The authors declare that they have no conflict of interest.

Acknowledgements. The research presented in this paper was funded by the German Research foundation, DFG, project numbers SH 55/16-1 and KU 2484/4-1. We thank all institutions that are involved with operating the AGA and CGS seismic monitoring networks and providing

220 the earthquake data and additional information. We thank AngloGold Ashanti Ltd. and Harmony Gold Ltd. for the permission to publish this study.

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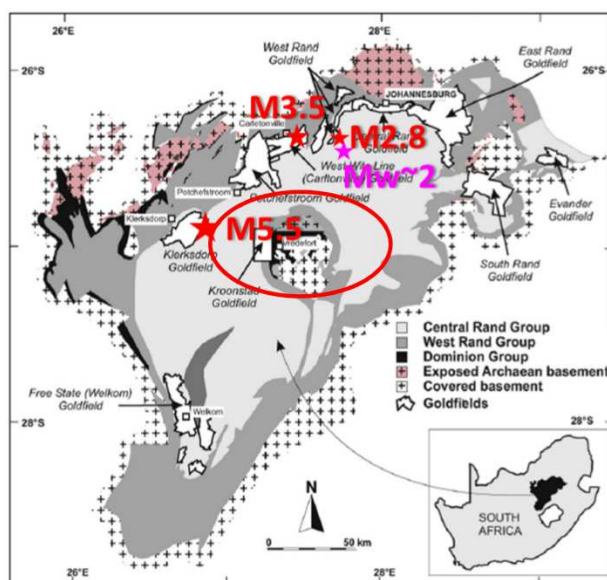
Figure 1. Simplified geological map of the Witwatersrand basin, also showing the locations of gold fields and the locations of the DSeis drilling targets marked in red.
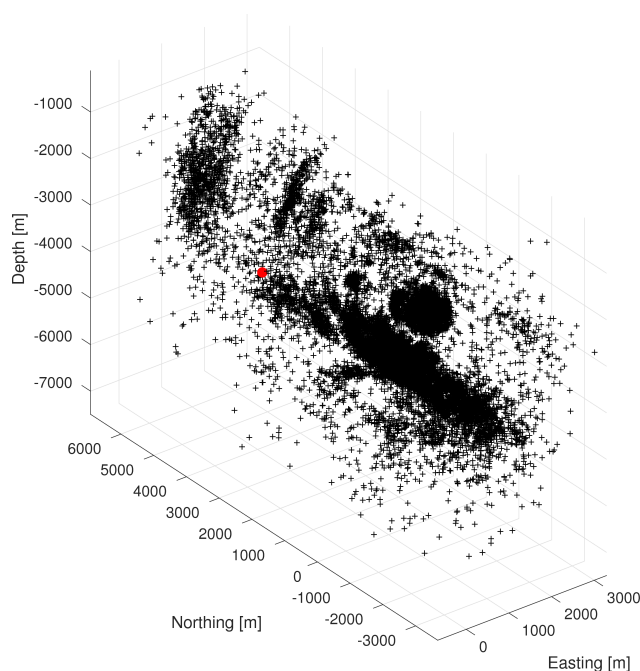
Figure 2. Hypocenter locations of about 70,000 earthquakes in the Orkney area in the DSeis catalog which is derived from the in-mine network. Coordinates are given in meters relative to the location of the Moab Khotsong mine shaft collar (0;0;1315). Red dot marks the hypocenter of the M5.5 earthquake.
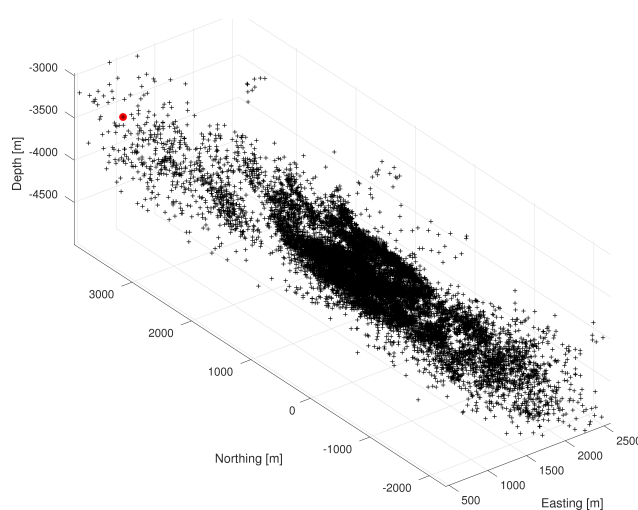
Figure 3. Hypocenter locations of aftershocks of the M5.5 earthquake. Coordinates are given in meters relative to the location of the Moab Khotsong mine shaft collar (0;0;1315). Red dot marks the hypocenter of the M5.5 earthquake.
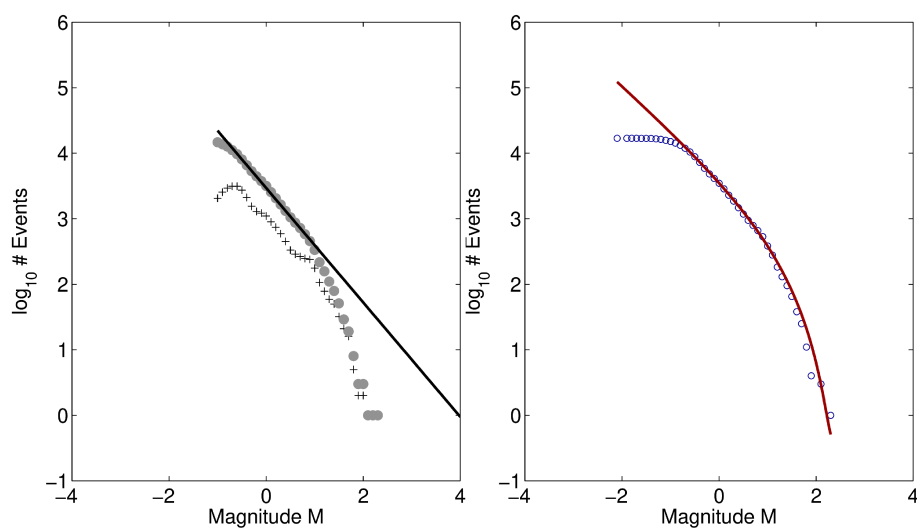
Figure 4. (Left) Magnitude frequency distribution of the aftershock sequence. The b value of the Gutenberg-Richter relation (solid line) was computed using the formula of Bender (1983). (Right) Application of the Lower Bound model (solid line) of magnitude probabilities in finite volume (Shapiro et al., 2013) to the observed magnitude frequencies yields a geometry-corrected b value.
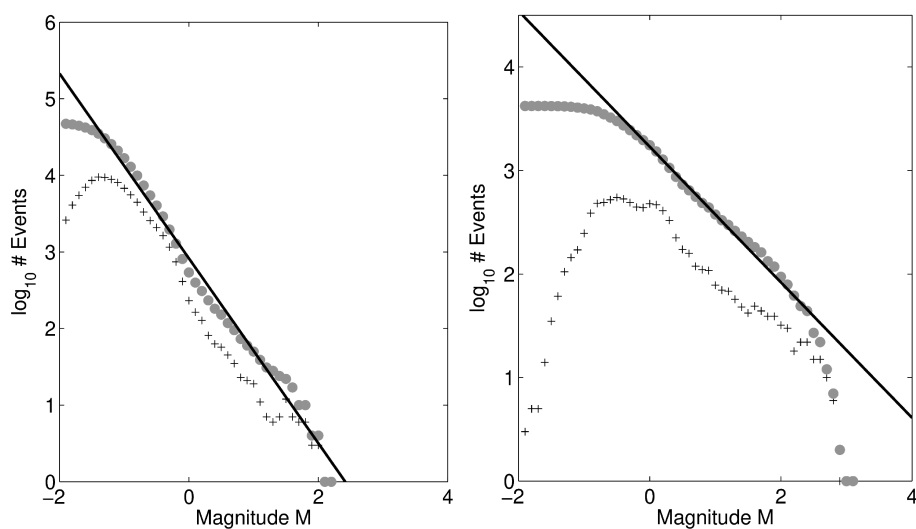
Figure 5. Magnitude frequency distributions of earthquakes induced by human operations in the Orkney area, (left) seismicity induced by fluid, and (right) mining induced seismicity. The b value of the Gutenberg-Richter relation (solid line) was computed using the formula of Bender (1983).
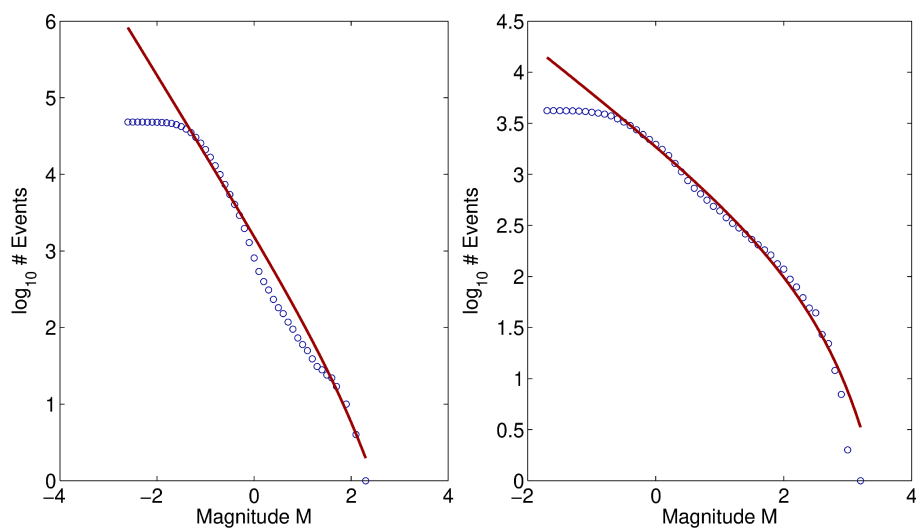
Figure 6. (Application of the Lower Bound model (solid line) of magnitude probabilities in finite volume (Shapiro et al., 2013) to the observed magnitude frequencies yields a geometry-corrected b value.
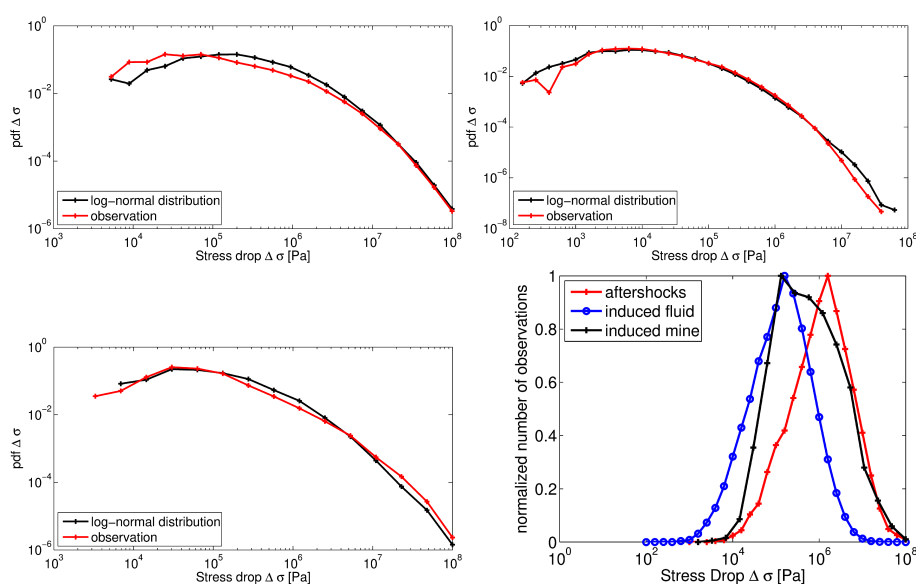
Figure 7. Probability density function of dynamic stress drop of aftershocks (top left), seismicity induced by fluid (top right), and seismicity induced by mining (bottom left), red lines correspond to the observations, black lines are best-fitting log-normal distributions. (Bottom right) Comparison of normalized histograms of dynamic stress drop of earthquakes for the three different classes of seismicity in the Orkney area.
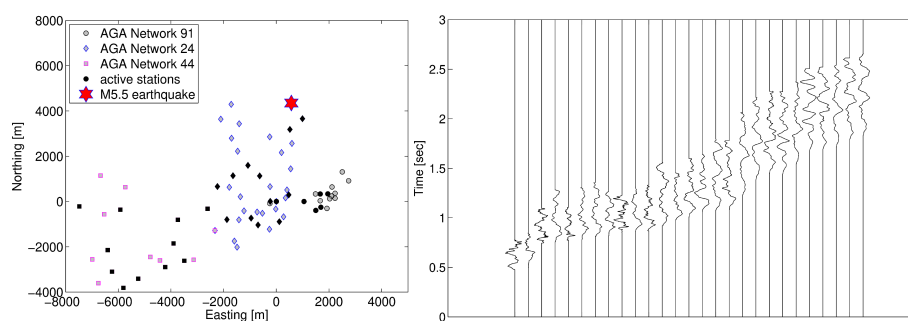
Figure 8. (Left) Map of seismic monitoring stations of the in-mine networks operated by AGA. Coordinates are given in meters relative to the location of the Moab Khotsong mine shaft collar (Longitude $26.802°$, Latitude $-26.867°$). Stations marked in black are used for imaging. (Right) P-wave amplitudes on the vertical component of displacement seismograms from the AGA stations which recorded the main shock. Time $0.0\,sec$ is the inverted origin time of the main shock (10:22:35.10 UTC).
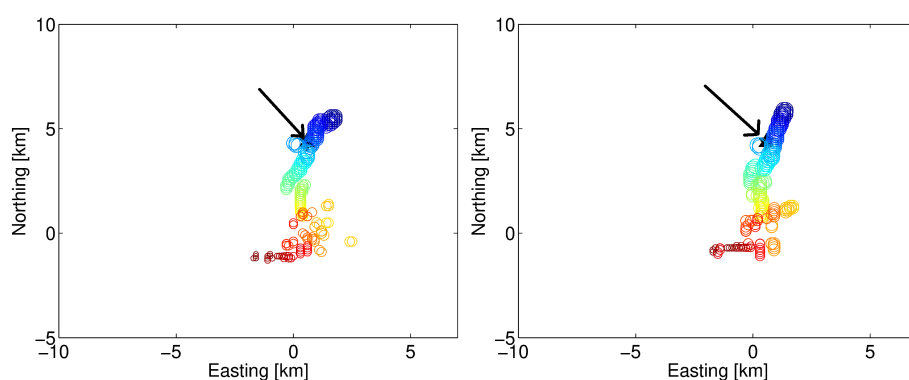
Figure 9. Back projection imaging of the M5.5 rupture using data from the in-mine network, (left) absolute values of P-wave amplitudes and (right) envelopes of P-wave. Coordinates are given in meters relative to the location of the Moab Khotsong mine shaft collar (0;0). The black star (indicated by black arrow) marks the location of the M5.5 earthquake (0.57;4.35). Color (blue - red) correspond to the time scale from 0-1.5 sec.
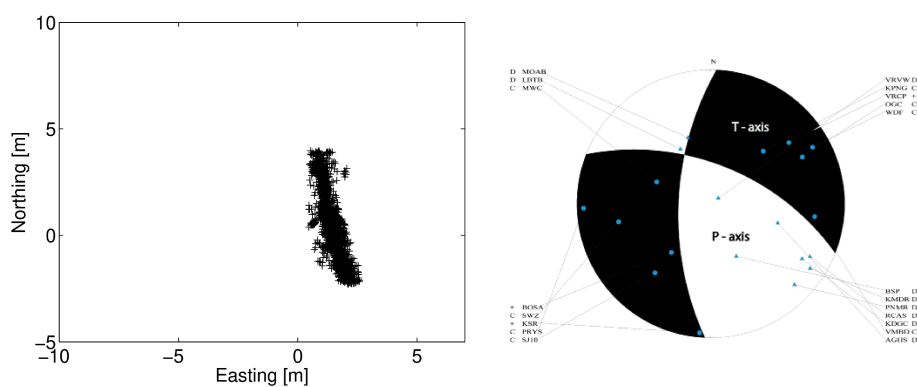
Figure 10. (Left) Map view of locations of the cloud of aftershocks for comparing with the MRPI results illustrated in Figure 9. (Right) Focal mechanism solution of one of the aftershocks which is representative for the M5.5 Orkney earthquake (see also Table 1 in Manzunzu et al., 2017).
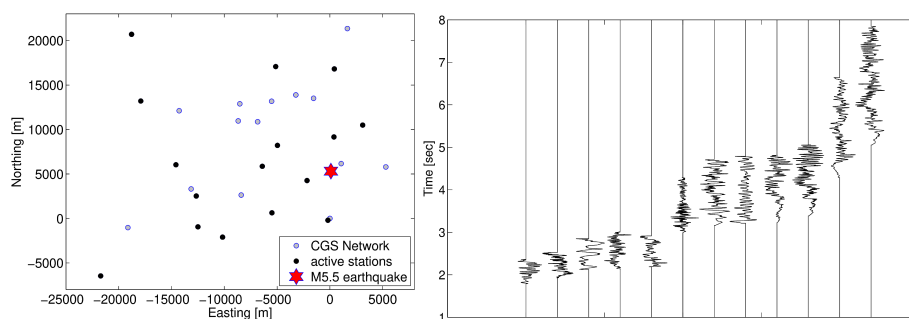
Figure 11. (Left) Map of seismic monitoring stations of the surface network operated by CGS. Coordinates are given in meters relative to the location of the Moab Khotsong mine shaft collar (Longitude $26.802°$, Latitude $-26.867°$). Stations marked in black are used for imaging. (Right) P-wave amplitude on the vertical component of displacement seismograms from the CGS stations which recorded the main shock. Time $0.0\,sec$ is the inverted origin time of the main shock (10:22:34.75 UTC).
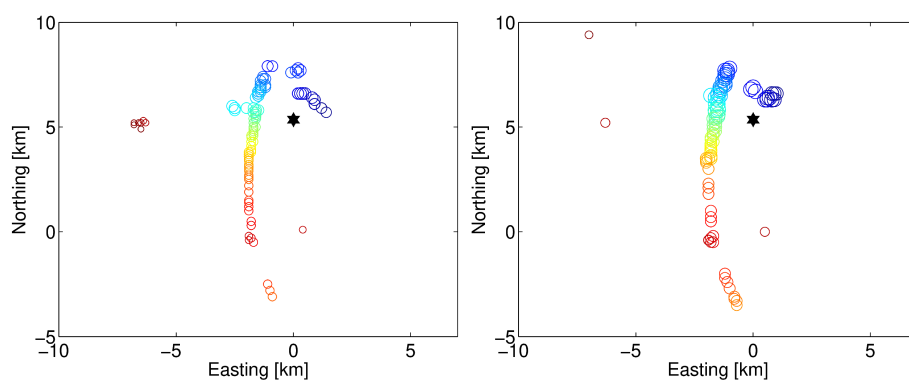
Figure 12. Back projection imaging of the M5.5 rupture using data from the surface network, (left) absolute values of P-wave amplitudes and in (right) envelopes of P-wave. Coordinates are given in meters relative to the location of the Moab Khotsong mine shaft collar (0;0). The black star marks the location of the M5.5 earthquake (0.1;5.3). Color (blue - red) correspond to the time scale from 0-1.5 sec.

Table 1. Summary of retrieved parameters of the magnitude statistics of seismicity in the Orkney area.

	aftershock sequence	fluid induced	mining induced
completeness magnitude M_c	-0.6	-0.5	-0.4
maximum likelihood b-value	0.84	0.96	0.61
Lower Bound b-value	0.73	1.08	0.52
Lower Bound M_{max}	2.4	3.3	3.8
Lower Bound stress drop $\Delta\sigma$	$700\,kPa$	$140\,kPa$	$500\,kPa$
median of static stress drops	$420\,kPa$	$40\,kPa$	$80\,kPa$
median of dynamic stress drops	$950\,kPa$	$120\,kPa$	$340\,kPa$

Table 2. Velocity model used in NonLinLoc to locate the M5.5 earthquake and to compute the traveltime-grid for MRPI using the CGS waveform data (a.s.l. − above sea level, b.s.l. − below sea level).

layer	top of layer	P-wave velocity v_P	S-wave velocity v_S
1	200 m a.s.l.	3.79 km/s	2.39 km/s
2	800 m b.s.l.	5.96 km/s	3.61 km/s