RESPONSE TO EDITOR

Comment 1
In the revised version, you mention in lines 30-31 of p7 "we conclude that foredeep parallel extension has occurred in the foredeep of the Pyrenean belt since the Paleocene and until the end of convergence" Do you consider here that Sigma 3 is negative as proposed in Figure 1b and introduction ? Extension is an unclear deformation term not synonym to tension or extensional stresses (i.e. negative stresses). Clarify this in the text please.
Response
We consider sigma 3 negative. The new text is: "Thus, we conclude that foredeep parallel tension has established in the foredeep of the Pyrenean belt since the Paleocene and until the end of convergence"

Comment 2
On this negative stresses as shown in Figure 1b, although we can agree on your interpretation, the paper suffers considering the significant contributions from experimental tests which have been compared to natural joints from the past decade. You mention extensional stresses (negative) but what about splitting without negative stresses (and even with a slightly compressive sigma 3) such as demonstrated in dry axi-symmetric, oedometric, plane strain and poly-axial experiments by Chemenda et al. (JGR,2011) and Jorand et al (Tectonophysics 2012) ? These studies shows joints formed under dry contraction without negative sigma 3, which are not so far than uniaxial splitting fractures observed in triaxial cells (e.g. Holzhausen and Johnson, 1979), but here clearly without the triaxial boundary effect mentionned by Fakhimi and Hemami (2015).
Response
We have no doubt that it is possible to replicate the morphology of a joint at the specimen size in an experimental apparatus using a Sigma3 > 0. However, we have some concerns about the possibility of upscaling such an experimental result at the basin scale and for tens of meters long systematic joints. Also, the occurrence of orthogonal cross-joints is not compatible with compressive sigma 3. We have added this text: "This indicates that E-W striking joints are cross-joints formed perpendicular to, and about synchronously with, the N-S striking joint set and that N-S joints formed in response to a negative (tensile) minimum stress (e.g. Bai and Gross, 1999; Bai et al., 2002)"

Comment 3
A common species of joints show very low displacement gradients compared to other fractures (veins, faults) (Pollard and Aydin, 1998; Schultz et al., 2008), which also support the general fact that joint sets do not require significant amount of negative stresses perpendicular to them. Have you measured the mean opening of the observed fractures ? Can this help to discuss this point ?
Response
We have not collected joint aperture data

Comment 4
I recommend you to better support the hypothesis mentionned in lines 31-32 p2 and 1-3 p3, which only relies on one reference, while others works previously described stress permutation during LPS. For example, stress permutation in foreland basin has been proposed from field observations and stress path calulations by Soliva et al. (2013), and reused with nearly the same concept in Fossen's book 2015 version. Addition of such references is just a fair strengthening of the hypothesis on which the work relies
Response
We have added this text " This is evidenced by the occurrence of bedding-perpendicular pressure solution-vein pairs (e.g Railsback and Andrews, 1995; Evans and Elmore, 2006; Quintà and Tavani, 2012; Weil and Yonkee, 2012) and/or conjugate strike-slip faults at a high angle to bedding

(e.g. Marshak et al., 1982, Hancock, 1985, Erslev, 2001; Lacombe et al., 2006; Amrouch et al., 2010, Weil and Yonkee, 2012) occurring in foreland areas and in the adjacent fold and thrust belts worldwide, although in many cases structures associated with this strike-slip regime do not develop during layer parallel shortening (Soliva et al., 2013). “

Transverse jointing in foreland fold-and-thrust belts: a remote sensing analysis in the eastern Pyrenees

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Abstract

Joint systems in the eastern portion of the Ebro Basin of the eastern Pyrenees enjoy near continuous exposure from the frontal portion of the belt up to the external portion of its associated foredeep. Utilizing orthophoto mosaics of these world class exposures, we have manually digitized over 30000 joints within a 16x50 km study area. The mapped traces exhibit orientations that are dominantly perpendicular to the trend of the belt (transverse) and, subordinately, parallel to the belt (longitudinal). In particular, joints systematically orient perpendicular to the trend of the belt both in the frontal folds and in the inner and central portion of the foredeep basin. Longitudinal joints occur rarely, with a disordered spatial distribution, exhibiting null difference in abundance between the belt and the foredeep. Joint orientations in the external portion of the foredeep become less clustered, with adjacent areas dominated by either transverse or oblique joints. Our data indicates that joints in the studied area formed in the foredeep in response to a foredeep-parallel stretching, which becomes progressively less intense within the external portion of the foredeep. There, the minimum stress direction becomes more variable, evidencing poor contribution of the forebulge-perpendicular stretching on stress organization.

Introduction

Fractures can be effective pathways for fluid flow (e.g. Laubach et al., 2019), thus impacting production of hydrocarbons (Barr et al., 2007; Engelder et al., 2009; Questiaux et al., 2010) and geothermal water (Haffen et al., 2013; Vidal et al., 2017), the pathways and fates of contaminants released from deep geological radioactive waste repositories (Berkowitz et al., 1988; Iding and Ringrose, 2010) and the sustainable management of groundwater (Masciopinto and Palmiotta, 2013). Associated with crustal tension, joints are ubiquitous open-mode fractures occurring in a range of tectonic settings, including collisional belts. In collisional settings, layer bending and stretching during the growth of thrust-related anticlines has conventionally been invoked as the principal causative process for the development of joints oriented approximately parallel (e.g. Ramsay, 1967; Murray, 1968) and perpendicular (e.g. Dietrich, 1989; Lemiszki et al., 1994) to the trend of the belt and of the thrust related anticlines. However, the frequently observed obliquity between joints (and other meso-scale structures) and the trend of the hosting anticlines (e.g. Tavani et al., 2019; Beaudoin et al., 2020), along with the documented occurrence of joints in exposed forelands (e.g. Dunne and North, 1990; Zhao and Jacobi, 1997; Billi and Salvini, 2003; Whitaker and Engelder, 2006), has more recently led to the conclusion that, in many cases, joints and other kinds of extensional fractures exposed in thrust and fold belts have developed prior to folding and thrusting, within the foreland region (e.g. Doglioni, 1995; Zhao and Jacobi, 1997; Tavarnelli and Peacock, 1999; Lash and Engelder, 2007; Branellec et al., 2015; Basa et al., 2019; Giuffrida et al., 2019; Martinelli et al., 2019; Carrillo et al., 2020), where joints are layer-perpendicular and commonly oriented parallel (longitudinal) and perpendicular (transverse) to the belt-foredeep-forebulge trend (Tavani et al., 2015).

A partially unresolved question in foreland deformation relates to the development of transverse joints, which requires a tensile minimum stress oriented parallel to the foredeep. Even in arched systems, the forebulge, the foredeep, and the belt tend to be nearly parallel to each other locally (e.g. the Hellenic arc, the Apennines-Calabrian arc, the Betic-Rif arc). The shortening direction in the inner portion of the foredeep (subjected to layer parallel shortening) and the stretching direction in the forebulge (where bulge-perpendicular stretching induced by lithospheric outer arc extension operates) are nearly parallel in a belt-perpendicular transect (Fig. 1). In addition, in the innermost portion of the foredeep, where layer parallel shortening operates, the $\sigma 2$ is typically vertical and the $\sigma 3$ is positive, horizontal, and parallel to the trend of the belt (Tavani et al., 2015). This is evidenced by the occurrence of bedding-perpendicular pressure solution-vein pairs (e.g Railsback and Andrews, 1995; Evans and Elmore, 2006; Quintà and Tavani, 2012; Weil and Yonkee, 2012) and/or conjugate strike-slip faults at a high angle to bedding (e.g. Marshak et al., 1982, Hancock, 1985, Erslev, 2001; Lacombe et al., 2006; Amrouch et al., 2010, Weil and Yonkee, 2012) occurring in foreland areas and in the adjacent fold and thrust belts worldwide, although in many cases structures associated with this strike-slip regime do not develop during layer parallel shortening (e.g. Soliva et al., 2013). Given the above, there should be an area in between the inner portion of the foredeep and the peripheral bulge where σ1 becomes vertical and where σ3 is still horizontal and parallel to the belt (Fig. 1a). This scenario could explain the development of layer-perpendicular transverse extensional structures, with transverse extensional faults or veins expected to develop where σ3 is positive. In this scenario, transverse joints occurring in this zone of localized tension could only develop as cross-joints (e.g. Gross, 1993) of the longitudinal set formed in the forebulge. This simplified model does not explain the documented occurrence of transverse joints in areas where longitudinal joints do not occur or represent the cross-joint set (e.g. Zhao and Jacobi, 1997; Quintà and Tavani, 2012). This framework does not admit a simple stress permutation in the foredeep and requires a negative (tensile) σ3 connected to a foredeep-parallel stretching component (Fig. 1b). Lithospheric bending of the foredeep, both along the horizontal plane (e.g. Doglioni, 1995) and along the vertical plane parallel to the trench (Quintà and Tavani, 2012), has been invoked as a process able to produce foredeep-parallel stretching.

Continuous exposures across the entire foreland region of the eastern Pyrenees allows investigation of the primary mechanism responsible for transverse joint development described above. Tens to hundreds of meters long joints affect the sedimentary sequence of the Ebro foredeep basin (Fig. 2a), and are found tilted within the frontal structures of the Pyrenean belt (Fig. 2b). These pre-folding joints are exceptionally exposed and mappable from orthophotos (Fig. 2c-f), from which they can be traced almost continuously from the external portion of the foredeep until the thrust belt. We have remotely mapped 30059 joints traces from the aforementioned orthophoto dataset and obtained their azimuthal distributions across the study area. Subsequently, this extended lineament database has been used to constrain the causative mechanism behind transverse jointing in the Ebro foredeep basin.

Geological framework

The study area is situated in the eastern Ebro foreland basin within an area connecting the eastern Pyrenees with the Catalan Coastal Ranges (Fig. 3a). The Pyrenees is an EW-striking orogenic system that formed as the Iberian and European plates collided from Late Cretaceous to Miocene times (Roest and Srivastava, 1991; Rosenbaum et al., 2002; Muñoz, 1992, 2002), and constitutes an asymmetric, doubly vergent orogenic wedge above the northward subduction of the Iberian lithosphere beneath the European plate (Chevrot et al., 2018). As a result, the Ebro basin formed as a flexural foreland developed on the downgoing Iberian plate at the southern margin of the chain (Beaumont et al., 2000). In the study area, to the south of the Ebro Basin (Fig. 3a), the Catalan Coastal Ranges developed as a Paleogene intraplate left-lateral transpressional system (López-Blanco, 2002; Santanach et al., 2011). The low-displacement character of thrusts and the absence of an associated foredeep both evidence the limited importance of this range. A series of NE-SW and NW-SE trending extensional faults strike parallel and perpendicular to the Catalan Coastal Range respectively, which are a result of the late Oligocene-Miocene opening of the NW Mediterranean basin (Vegas, 1992; Sàbat et al., 1995; Granado et al., 2016a).

The study area where joint traces have been digitized is delimited to the north by the frontal Pyrenean thrusts and by the Eocene Bellmunt anticline (Figs. 3a,b). This anticline comprises the Paleocene to upper Eocene foredeep infill. Immediately to the south of the anticline (i.e.< 5 km), this multilayer becomes sub-horizontal and thins southward where the Paleozoic to Mesozoic foredeep floor gently rises up and becomes exposed (Fig. 3a). There, this pre-orogenic succession is slightly tilted to the north by uplift in the footwall of NW-SE striking ~~L~~late Oligocene to Neogene extensional faults. Further to the southwest, this succession is affected by the Paleogene contractional structures of the Catalan Coastal Ranges (Fig. 3b).

In the field, joints are constantly bedding-perpendicular, regardless of the bedding dip (Fig. 2a,b), and they are characterized by the occurrence of either a single set (Fig. 2c) or by a ladder pattern (Fig. 2d,e). In the latter case, the few E-W striking joints are almost everywhere perpendicular to the N-S striking set and abut on it (Fig. 2e). This indicates that E-W striking joints are cross-joints formed perpendicular to, and about synchronously with, the N-S striking joint set and that N-S joints formed in response to a negative (tensile) minimum stress (e.g. Bai and Gross, 1999; Bai et al., 2002) .

Data and Methods

Joints have been digitized within the 16.4x49.2 km area displayed on Figure 3, on 25 and 50 cm/px orthophotos provided by the Spanish Instituto Geográfico Nacional via the PNOA (Plan Nacional de Ortofotografía Aérea) project (https://pnoa.ign.es/). The open source geographical information system QGIS 3.4 has been used to manually digitize 30059 joint traces. A collection of these traces seen on orthophotos is provided in Figure 2c-f. Joints have been digitized in Bartonian and, subordinately, in Lutetian and Priabonian sedimentary rocks (Fig. 3c). The NE and SE portions of the study area are highly vegetated (Fig. 3d,e) and only a few joint traces have been mapped there. Quaternary sediments unaffected by joints crop out in the central portion of the study area (Fig. 3d).

Digitized traces have lengths ranging from 2 to 100 m, with an average of ~20 m (Fig. 3f). The frequency distribution of trace trend shows that the vast majority of the mapped joints are approximately N-S striking (Fig. 3f). A second subordinate set corresponds to E-W striking joints, which in the field occur as cross-joints (Fig. 2d). The frequency distribution of trace trend is not symmetrical around these orthogonal sets, due to the presence of a third less abundant set composed of NW-SE oriented joints (Fig. 2e,f). The NW-SE striking joints are mostly seen in the southern portion of the study area, in exposures where the N-S striking set is not occurring. As mentioned above, exposures are frequently characterized by a single set (Fig. 2c,d). At many locations, the dominant set is accompanied by associated cross-joints (Fig. 2d). Very rarely, two mutually oblique sets occur in the same exposure (Fig. 2f).

In order to evaluate the variability of joint traces, the 16.4x49.2 km study area has been divided into meshes of equilateral triangular elements with edge lengths of 1025m (Mesh 1) and 1640m (Mesh 2). At each node, mean value and variance of trace trends has been computed using a circular moving window with a radius of 1200m (Mesh 1) and 1900m (Mesh 2). The radius of the circular moving window is set to these values for two reasons: 1) it is two orders of magnitude longer than the average length of joints; 2) it is large enough to ensure that most of the nodes have data number >20, as only nodes with data number > 20 have been analyzed. Since trace trends are circular data with an angle (α) over a period of 180°, in agreement with Mardia (1975) we used equations 1 to 4 to derive at each node the circular mean value (Mv_π), the circular variance (V_π) and the resultant length (R_π; $V_\pi = 1-R_\pi$), the latter spanning from 0 (unclustered distribution) to 1 (perfectly clustered distribution). In the presence of cross-orthogonal joint sets, it is also useful using a period of π/2, thus modifying Mardia's equations and introducing the $Mv_{\pi/2}$ and $R_{\pi/2}$ parameters, which are computed using Equations 5 to 8.

$$C_{\pi}=\frac{\sum_{i=1}^{n}\cos\left(2\alpha_{i}\right)}{n}\ (1)\ ;\ S_{\pi}=\frac{\sum_{i=1}^{n}\sin\left(2\alpha_{i}\right)}{n}\ (2)\ ;\ V_{\pi}=1-R_{\pi}=1-\sqrt{C_{\pi}^{2}+S_{\pi}^{2}}\ (3)\ ;\ Mv_{\pi}=\frac{Arc\tan\left({}^{S_{\pi}}/{}_{C_{\pi}}\right)}{2}\ (4)$$

$$C_{\pi/2}=\frac{\sum_{i=1}^{n}\cos\left(4\alpha_{i}\right)}{n}\ (5)\ ;\ S_{\pi/2}=\frac{\sum_{i=1}^{n}\sin\left(4\alpha_{i}\right)}{n}\ (6)\ ;V_{\pi/2}=1-R_{\pi/2}=1-\sqrt{C_{\pi/2}^{2}+S_{\pi/2}^{2}}\ (7)\ ;\ Mv_{\pi/2}=\frac{Arc\tan\left({}^{S_{\pi/2}}/{}_{C_{\pi/2}}\right)}{4}\ (8)$$

By using these four parameters together, instead for example of the classical k-means clustering analysis, it is possible to derive important considerations on the distribution of polymodal distributions in which two mutually orthogonal sets do occur, as illustrated in Figure 4. We compare the Mv and the R parameters computed using the π and π/2 periods. In the first example two parallel traces are analyzed, resulting in $Mv_{\pi/2} = Mv_{\pi}$, and $R_{\pi/2} = R_{\pi} = 1$ (i.e. circular variance = 0). When data dispersion is slightly increased (Example 2), $Mv_{\pi/2}$ is still equal to Mv_{π}, whereas $R_{\pi/2}$ decreases faster than R_{π}. Further increase of data dispersion (Example 3), in an asymmetric distribution (i.e. non-orthogonal sets), causes additional decrease of $R_{\pi/2}$ with respect to R_{π}, and slight divergence between $Mv_{\pi/2}$ and Mv_{π}. In the presence of a cross-orthogonal subset, the statistical usefulness of $R_{\pi/2}$ becomes evident, as illustrated in Example 4. In this case, R_{π} rapidly approaches zero, suggesting high dispersion (i.e. unrepresentative Mv_{π}), whereas $R_{\pi/2}$ is essentially unaffected with respect to Example 2, indicating low dispersion and a representative $Mv_{\pi/2}$. However, the use of the π/2 period only returns results in the 0 to π/2 range, so that NW-SE trending traces result in a NE-SW trending mean value ($Mv_{\pi/2}$), as shown in the Example 5. In summary, Mv_{π} is useful to derive the mean direction, whereas $R_{\pi/2}$ and R_{π} should be used in conjunction to discriminate between populations in which oblique sets occur ($R_{\pi/2} < R_{\pi}$) from those in which two perpendicular sets occur ($R_{\pi/2} > R_{\pi}$).

Results

Figure 5 displays attribute maps generated from Mesh 1 and Mesh 2 using the trace orientation parameters described above. Both Mesh 1 and Mesh 2 have $R_{\pi} > 0.5$ across almost the entire study area, with the only exception being in its NW corner. For $R_{\pi/2}$, in addition to the NW corner, the central portion of the study area has $R_{\pi/2} < 0.5$. Noteworthy differences between R_π and $R_{\pi/2}$ occur: (1) in the NW corner ($R_\pi << R_{\pi/2}$) and (2) in the central area ($R_\pi >> R_{\pi/2}$). The first area corresponds to a vegetated folded and faulted zone (Fig 3b-e). Consequently, we consider the dataset poorly reliable in this location. In the central portion of the study area, the difference between R_π and $R_{\pi/2}$ (which increases as data dispersion increases) is less pronounced in Mesh 1 than Mesh 2. Thus, data dispersion increases with increasing the search window size (1200m for Mesh 1 and 1900m for Mesh 2), evidencing that joint orientation is changing in this area. In the rest of the analyzed foreland, $R_{\pi/2}$ has values similar to R_π, indicating approximately unimodal data distribution within this region, and poor spatial organization of the longitudinal cross-joints.

Distribution of Mv_π relates to the prevalence of NS-striking joints in the northern and central portion of the study area. Towards the south, patches characterized by both N-S and NW-SE-striking joints occur. High values of R_π and $R_{\pi/2}$ are characteristic of such subareas, which as previously mentioned, is indicative of unimodal joint trace distributions.

Discussion

Remotely sensed and mapped joint traces in the eastern portion of the Pyrenees-Ebro system show systematic distributions of azimuthal orientations. In the frontal portion of the belt and within the foredeep, joints are mostly transverse (i.e. N-S-striking), with limited occurrence of E-W-striking longitudinal cross-joints. Approaching the southern border of the foredeep, joints exhibit both N-S and NW-SE orientations where the pre-Cenozoic floor of the foredeep is exposed. Joints affect the Lutetian to Priabonian foredeep infill, and are found tilted together with strata within the Bellmunt anticline (Fig. 2b), which is Priabonian in age (Burbank et al., 1992). Given the above, jointing in the study area must have taken place during or before the Priabonian. The occurrence of the N-S striking joints within the Ebro foredeep is documented also to the west of the study area (e.g. Turner and Hancock, 1990), where joint emplacement affects up to the Miocene (Arlegui and Simon, 2001). Transverse joints, striking approximately NNE-SSW occur also to the SW in Bartonian to Priabonian strata cropping out at the boundary between the Ebro basin and the Catalan Coastal Ranges (Alsaker et al., 1996). These data indicate that transverse joints systematically developed in the foredeep basin of the E-W oriented Pyrenean belt. Locally, the process of transverse jointing occurred until the early Miocene (i.e. until the end of mountain building). Also, prethrusting transverse extensional faults of upper Paleocene to lower Eocene age occur a few tens of km to the NE of the study area (Carrillo et al., 2020), being presently incorporated into the Pyrenean belt. Thus, we conclude that foredeep-parallel ~~extension~~tension has established~~occurred~~ in the foredeep of the Pyrenean belt since the Paleocene and until the end of convergence. Transverse joints documented in this work clearly represent foredeep-related structures, which can develop by (i) N-S oriented layer-parallel shortening (LPS) or (ii) E-W oriented along foredeep stretching (Tavani et al., 2015). LPS is to be excluded, as the state of stress in this case would include a positive minimum stress (Fig. 1). In agreement, LPS-related extensional structures can form only due to fluid pressure contribution and they include mm- to cm-long fractures filled with calcite (which is removed from pressure solution seams, Fig. 1; Tavani et al., 2015 and references therein). The type (joints with no calcite infill) and size (tens of m-long) of transverse extensional structures described here are incompatible with the layer-parallel shortening mechanism.

The relatively constant orientation of joints along the strike of the foredeep, the occurrence of appreciable dispersion at the outer border of the foredeep, and the remarkably poor abundance of longitudinal joints allow us to derive two major conclusions:

(1) The almost linear trend of the eastern Pyrenees facilitates the exclusion of planar arching (e.g. Doglioni, 1995; Zhao and Jacobi, 1997) as the causative process for generating foredeep-parallel stretching (i.e. required to establish the negative σ3 responsible for transverse jointing). Arching along the vertical plane parallel to the trench (Quintà and Tavani, 2012) represents instead a viable mechanism for generating along-foredeep stretching. This is analogous, albeit at a larger scale to the process of release faulting described by Destro (1995). The basic concept behind this mechanism is the following: when a straight line joining two fixed points - the tips of a fault in the case of Destro (1995) or the edges of the foredeep in the case of Quintà and Tavani (2012) – becomes an arch, there is stretching (Fig. 1b), which causes extensional stress parallel to the direction of elongation. In essence, this mechanism is expected to operate in any doubly plunging foredeep, particularly at its lateral edges, such as in the study area (Fig. 3a).

(2) Extension in the peripheral bulge, which is documented from many active and fossil foredeep basins (e.g. Bradley and Kidd, 1991; Ranero et al., 2003; Tavani et al., 2015; Granado et al., 2016b), including the lower Eocene foredeep basin presently incorporated into the Pyrenees (e.g. Martinez et al., 1989; Pujadas et al., 1989) appears to be weakly influential

at the southern border of the study area (i.e. the upper Eocene peripheral bulge). Indeed, the observed longitudinal joints are characteristically subordered, forming locally as cross-joints to the identified transverse set within the study area (Fig. 2d). The transverse set becomes less organized at the southern margin of the foredeep, where patches of N-S and NW-SE dominated domains do occur. This evidences the absence of a major forebulge-perpendicular extension capable of systematically reorienting σ3 at the external foredeep edge.

Conclusions

Analysis of remotely sensed and mapped joints in the eastern Pyrenees and in the adjacent Ebro foreland basin indicates that the emplacement of the dominant joint set within the area, which strikes perpendicular to the trend of the foredeep occurred prior to folding and developed in response to along-strike stretching caused by the plunging shape of the foredeep. Joints developed in response to flexuring of the lithosphere at the peripheral bulge do not occur in the area, suggesting that this mechanism has limited relevance to the observed joint system. This is confirmed by the variability of joint orientations observed at the foredeep external edge, negating the occurrence of a major forebulge-perpendicular extension able to systematically orient the stress field at the foredeep edge.

Data availability

Digitized traces in shapefile format and bedding and joint data in csv format are in the supplementary materials

Author contributions

ST, PG, AC, TS, JMC, and JAM contributed equally to the elaboration of the manuscript.

Competing interests

The authors declare that they have no conflict of interest.

Acknowledgements

We thank Roger Soliva, Eric Salomon, and an anonymous referee for useful suggestions. The Institut de Recerca Geomodels and the Geodinàmica i Anàlisi de Conques research group (2014SGR467SGR) acknowledges financial support from the Agència de Gestió d'Ajuts Universitaris i de Recerca (AGAUR) and the Secretaria d'Universitats i Recerca del Departament d'Economia i Coneixement de la Generalitat de Catalunya. Financial support form projects CGL2017-87631-P and PGC2018-093903-B-C22 from Ministerio de Ciencia, Innovación y Universidades/Agencia Estatal de Investigación/Fondo Europeo de Desarrollo Regional, Unión Europea is also acknowledged.

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Captions

Figure 1

Scheme showing the architecture of a foreland fold-and-thrust belt and adjacent foredeep basin, with syn-orogenic fracture patterns in the different structural domains of the orogenic system. (A) The foredeep state of stress is governed by the permutation between the state of stress in the layer-parallel shortening and peripheral bulge domains. (B) The foredeep state of stress is controlled by the along-strike stretching of the foredeep.

Figure 2

Examples of pre-folding joints within the studied area. (A) N-S striking joint with plumose structures in the foredeep sediments (42°02'39.7"N; 2°13'54.9"E). (B) Tilted N-S and E-W striking joints in the southern limb of the Bellmunt anticline, with the red arrow indicating an E-W striking joint abutting on a N-S striking joint (42°05'39"N; 2°17'41.5"E). The density contours of poles to bedding and joints (in their present day orientation and after unfolding) refer to data collected in the Bellmunt anticline area. (C to F) Examples of joints seen on orthophotos. (C) Transverse joints. (D) N-S striking transverse joints with subordinate E-W striking cross-orthogonal joints. (E) NW-SE striking joints. (F) Rare example of multiple oblique sets occurring at the same exposure. Orthophotos are available from the Spanish Instituto Geográfico Nacional (https://pnoa.ign.es/).

Figure 3

(A) Simplified geological map of the western Pyrenees and Catalan Coastal Ranges (based on the Geological Map of Catalunya scale 1:250’000; https://www.icgc.cat/en/Downloads), with N-S geological cross-section (modified from Parés et al., 1999). (B) Detailed geological map of the study area, with digitized joints (C). (D) Orthophoto (https://pnoa.ign.es/) of the study area, with digitized joints (E). (F) Frequency distribution of joint traces trend and length.

Figure 4

Examples of joint patterns and resultant Mv and R parameters calculated for the π and π/2 periods. For the five examples, we show the map view of the joints, the azimuthal frequency, and the sin-cos coordinates of the resultant values of Mv and R. Note that the distance from the center is proportional to R.

Figure 5

Results of circular statistics analysis for both Meshes 1 and 2. Length of traces of Mv_{π} is proportional to R_{π}. Color code refers to R_{π} and $R_{\pi/2}$, whereas the orientation of traces is the Mv. See text for details.