Answer to the reviewer on:

Report #1

Submitted on 14 May 2022

Anonymous Referee #1

Referee #1: This paper shows the analysis of seismicity recorded at three arrays in Fogo and Brava, Cape Verde. The results are interesting, and apparently well supported by the data. Below you can find some suggestions to improve the manuscript.

With the data shown, the classification of the seismic events into proper earthquakes and hybrid events is very arguable. It is taken as an a priori assumption, surely based on previous works. Of course the concept of a hybrid event is flexible and you can find in the bibliography a variety of events named as "hybrids". But in this case, the waveform and spectrogram of the hybrid event shown (figure 8, which I assume is your best example) are not different from what you would expect for a shallow earthquake, where S phases are often hard to identify. Moreover, I do not see a clear, sustained, low-frequency coda, but a coda that is similar to the earthquake coda shown for example in figure S6. Now, not all earthquakes without clear S phases are hybrid events; the name suggests a particular source mechanism related to fluid-driven resonances. Perhaps you could start talking just about two groups of seismic events, based on the waveform characteristics, and in the discussion elucidate whether the second group can be hybrids.

Answer: We thank the reviewer for appreciating our work and for the valuable comments and suggestions.

The earthquake of Figure S6 shows the peak in energy in the 10-20Hz band. Below 10 Hz, there is little energy, mainly above 5 Hz. This is clearly not the case for the hybrid event. Figure 8 shows a comparison of a shallow earthquake on Fogo with a hybrid event. The spectra of the two events are significantly different. One important difference is the smooth transition from higher to lower frequencies. The earthquake shows the main energy in the 10 to 25Hz band and there is no clear transition to lower energies observable. As already stated by the referee, this observation is also based on previous studies, which found the same type of event and identified it as hybrid event.

The traces of the hybrid events at stations several km away (e.g. at array CG) do not show any S-phase as well. This is different when compared to earthquakes occurring on Fogo. Especially on Brava, we observe many earthquakes very close to the array, which are likely shallow. These events display S-phases as well at the array stations.

However, in the discussion section, we intensively describe possible causes for the observed hybrid events on Fogo. There are studies, which propose a different source mechanism for hybrid events, independent of fluid motions (Harrington and Brodsky, 2007). For example, brittle failure in unconsolidated material might cause these events and lead to their characteristic signature. Even for long-period events, there may be origins without the fluid-driven source mechanism (Bean et al., 2014). This discussion can be found in our manuscript in lines 336-347.

Nevertheless, as the origin of the events remains unclear, we follow the referee's suggestions and now use this terminology more carefully (see lines 50-53, 70-72, 244-247, 273).

R1: Another disputable point is that the beamforming calculations do not include any information about the quality of the beam. The beam energy is normalized to the maximum value in the apparent slowness grid selected, and therefore we cannot see if the waveform fit is good or not. For example, it is obvious that the beamforming shown in figure 3 represents a good fit, with all waveforms aligned and looking similar. On the other hand, figures S2.1 and S2.2 are not a good fit. Therefore the peaks in figures S2.1 and S2.2 are less meaningful than the peak in figure 3. But since figures are normalized to the maximum, this difference is obliterated. I suggest to plot a magnitude that shows how adequate is the beam, how similar the waveforms are, and how well the delays are adjusted by the apparent slowness vector. For example: you could define the beam energy as the energy of the beam (as in the equation in line 125) divided by the average energy of the individual traces. In this way, we should see a high peak only when the beam is ok as in figure 3, and smaller peaks in the other cases (and perhaps this may account for the multiple peaks, which may not be significant).

A.: Yes, we have considered this carefully. Indeed, there are events, where the traces are not perfectly aligned. This leads then typically to sidelobes (as in Fig. S2.2), which then again are included in the analysis (lines 156-159). Typically, a waveform alignment of low quality leads to a large standard deviation. Additionally, due to site effects the arrays have different frequency characteristics. In general, we observed that the higher the frequency of the traces in general, the lower the quality of the alignment. If a weighting is implemented according to the individual energy values, this would lead to a down weighting of arrays, which show higher frequencies. We, therefore, decided to judge the beam quality based on the standard deviation of the individual backazimuths.

R1: Related to the beamforming analysis, you should address the effect of the duration of the stacking window. You say that "the variation of the stacking window are accounted for as described in Section 3.2" (lines 170-171). And "The selection of the stacking window has a larger contribution to possible errors and is thus included in the analysis" (lines 177-178). But this is not true, because your procedure includes just the variations of the starting time of the stacking window within some pre-defined range, and you never actually address the effect of the window length. The duration of the stacking window is a crucial parameter that may affect drastically the performance of the method, and you should explain how you select the optimum value.

A.: The selection of the optimum value of the window length has been carefully addressed during development and application of the single-array time-domain approach. In these first evaluations of the method, we found that the optimum stacking window length is in the range of one or two period lengths. One advantage of the method we use is that the stacking window can be kept very narrow around the phase of interest. The stacking window length is thus kept as small as possible to produce a reliable result. The analyst judges whether a result is reliable. This is done by evaluating the resulting stack of the traces, their proper alignment and if the sum trace displays a clear signal onset. Furthermore, by the evaluation of the resulting energy plot, its uncertainties, sidelobes, and by the reproducibility of the result during several trials.

We have also specified these considerations in the revised manuscript (lines 119-122, 189-195).

R1: You should also clarify some comments about the stacking window: "a stacking window with the length of one or two periods of the signal" (line 112); "some stacking windows are as small as 0.6 s" (line 136). These are not true, in the figures we can see that stacking windows are smaller, 0.1-0.15 s in figures 3 and S2.1 (around 1 period), <0.1 s in figure S2.2 (barely half period). Such short time windows will yield higher values of the beam energy, but this is not always an improvement because it is more likely to obtain secondary peaks as in figure S2.2.

A.: Thank you very much for pointing this out, indeed there is an error in line 136, the correct value is 0.06s. We have corrected this in the revised manuscript (line 145).

The signal is chosen in reference to the central station. In Figure S2.2 the stacking window of trace no. 1 includes one period. It is surely hard to see in the figure, but there is already a second maximum contained in the stacking window. It is easier to see e.g. at trace no. 6. However, it is correct, that array CG shows higher frequencies than the other arrays, which increases the uncertainties of the array analysis of this particular array.

R1: I do not understand how you transfer the standard deviation of the back-azimuth into the energy plot. In table S13 the values of back-azimuth are large, often above 90 deg. And if the std is so large, the values themselves must be even more variable. How can this distribution transform into the relatively narrow beams shown in the figures? And then all that confusing process with the percentages, described in figure S3. Why not just extracting directly from the energy plot the maximum value of energy corresponding to each back-azimuth direction?

A.: This must be a misunderstanding. Of course the beam becomes wider for larger standard deviations and vice versa.

The standard deviation of the backazimuth can indeed be rather high for some events. To include this information, the width of the beam during the multi-array analysis is determined by the standard deviation of the backazimuth. A larger standard deviation leads to a broader beam. If three broad beams (in a worst case scenario) intersect each other the area of the possible event location becomes very large.

We reduce the beam width successively in steps of 1% of the standard deviation. Thus, the beams narrow down, approaching the maximum peak in energy and the area of intersection becomes smaller, such that the epicenter of the event and its uncertainties can be better defined.

The event shown in Fig. 3 and 4 occurred on 22 July 2017, 23:35. For this event e.g. array AF shows a standard deviation of backazimuth of 72.8°, which is rather large. This results in a broader beam at array AF during the multi-array analysis. This is shown in Fig. 4, where the array AF has the broadest beams of the three arrays. We have included a corresponding information in the revised manuscript (lines 162-164).

We do not extract the energy values from the energy contour plots, because this can lead to some down weighting of individual arrays as discussed above.

R1: In the discussion, the network locations are still taken as ground truth. This is evidenced in several comments along the text (e.g. lines 344-345), and in the comparisons of figure 5, where you talk about deviations in the range of 7-9 degrees. But the network locations also have uncertainties, as shown in figure S5. You should consider these uncertainties for the comparison of the results. For example, in the case of BR most earthquakes are close to the array and the uncertainty in the network-estimated back-azimuth is large. I would not be so sure that the real back-azimuth is the one obtained from the network locations. Perhaps the proper way to go is to add the uncertainty limits of the estimates shown in figure 5, so we can see the range of azimuth and slowness corresponding to each estimate, and compare the performances of the array and the network locations.

A.: Accounting for individual errors would make Fig. 5 unreadable. Here, we are interested in a general overview to show systematic deviations between network and array localizations. Individual errors therefore do not increase the information content of the Figure.

R1: About the small apparent slowness of hybrids, suggesting a deep source: your method is biased to find high apparent slowness values, since you limit the analysis to apparent slowness smaller than 0.3 s/km (line 115). If the events are shallow they might have larger apparent slowness, that fall outside of your analysis range. If that is the case, your current solutions would be just secondary peaks. You could try extending the analysis to larger apparent slowness, for example using a grid from -1 to 1 s/km or even larger, just to make sure that your finding of large apparent velocities for the hybrids is robust.

A.: This is a valid point. We carefully tested higher slowness values when establishing the method. Of course these slowness values were tested for the different types of events, which we observed. A slowness range of -1 to 1 s/km does not lead to secondary peaks and the observed high apparent velocity remains at the observed value. However, we added a corresponding figure to the supplements and included a corresponding information in the manuscript (lines 282-284).

R1: - figure 1, why not showing all array configurations?

A.: The arrays are always designed in the same way. To increase the visibility of Figure 1, we decided to show only one array configuration.

R1: - 70-71: the sentence should be at the end of the paragraph, since systematic deviations can be observed only after you compare them with the network locations.

A.: We rephrased the sentence in the revised manuscript (lines 71-72).

R1: - 86: why would you expect this range of frequency of 5-10 Hz? The real spectra clearly extend to higher frequencies, as we can see in the figures.

A.: This is a value based on previous observations and on a pilot study we conducted earlier (Leva et al., 2020). It is a mean frequency range of the first phases of the events detected. For hybrid events the dominant frequency of the first signal onset is typically smaller than that of earthquakes. Additionally, we also observe harmonic tremors in a frequency range of 1 to 5 Hz. When the arrays were designed, this event type was also considered. Therefore, the mean frequency range is between 5 to 10 Hz.

We have added a corresponding explanation to the revised manuscript (lines 84-86).

R1: - 145: definition of error as X=(1-std/360)*100. This is counter-intuitive, a large std implies a small error, and viceversa? (e.g. std=360 deg => X=0%).

A.: Thanks for the comment. In the revised manuscript we corrected the formula and clarified the text (lines 153-156).

R1: - 157: geographical coordinates?

A.: Thank you for pointing this out, we corrected this in the revised manuscript (line 168).

R1: - 166: length of the stacking window.

A.: Yes, this is correct and mentioned in the text.

R1: - The velocity model of Vales et al. 2014 is used for some applications, however it is modified ad hoc for no apparent reason, for example:
184-185: "For this standard localization technique, we apply the velocity model from Vales et al. (2014)". The model is used as is for the network locations.
216: "For this we apply a two-layer velocity model with a mean crustal and a mean mantle velocity, derived from Vales et al. (2014)". The model is simplified to two layers for the calculation of distances of earthquakes with S-P times, why not use it as before?
301-302: "The velocity model is adapted from Vales et al. (2014) with velocity increasing at steps of 0.1 km". The model is modified at steps of 0.1 km to obtain the depth of the hybrids, how is it modified? Smoothing the velocity?

A.: The simplified velocity model used for the single-array analysis is based on a two-layer (crust, mantle) model. The method used here for the distance estimation is a rather simple approach and there is no necessity for a detailed velocity model.

During the standard localization method, no modification of the velocity model is necessary. However, this is not the case when the depth of the hybrid events is determined with the approach described in the manuscript. Here it turned out that the steps of the velocity model of Vales et al. (2014) are too large and it was not possible to derive a reliable estimate of the depth when applying Snell's law. Therefore, we have chosen rather small steps to intersect the velocity layers determined by Vales et al. (2014).

We have added a corresponding explanation to the revised manuscript (lines 324-325).

R1: - 199: if station elevations are to be discussed, you should provide them!

A.: We included the station locations and elevations in the supplementary material.

R1: - figure S1: you show the array response at 7.5 Hz, however the events have energy at higher frequencies. Therefore the energy peaks are generally narrower than the array response. Therefore, it would be better to see the response at 10 or 15 Hz?

A.: In the case of a time domain array analysis the array response for a single frequency is not as significant as in the frequency-domain approach. In the time domain a relatively broad frequency band is implicitly incorporated. This leads to a suppression of the sidelobes, which may appear in the array response function for single frequencies.

Nevertheless, we have added a new figure displaying the array response function for 5, 7.5 and 15Hz to the supplement. Additionally, we included a corresponding explanation in the revised manuscript (lines 90-95).

R1: - 336: site effects?

A.: We corrected this in the revised manuscript (line 359).

Multi-array analysis of volcano-seismic signals at Fogo and Brava, Cape Verde

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Abstract. Seismic arrays provide tools for the localization of events without clear phases or events outside of the network, where the station coverage prohibits classical localization techniques. Beamforming allows the determination of the direction (backazimuth) and the horizontal (apparent) velocity of an incoming wavefront. Here we combine multiple arrays to retrieve event epicenters from the area of intersecting beams without the need to specify a velocity model. The analysis is performed

10   in the time domain, which allows to select a relatively narrow time window around the phase of interest while preserving frequency bandwidth. This technique is applied to earthquakes and hybrid events in the region of Fogo and Brava, two islands of the southern chain of the Cape Verde archipelago. The results show that the earthquakes mainly originate near Brava whereas the hybrid events are located on Fogo. By multiple-event beam-stacking we are able to further constrain the epicentral locations of the hybrid events in the northwestern part of the collapse scar of Fogo. In previous studies, these events were

15   attributed to shallow hydrothermal processes. However, we obtain relatively high apparent velocities at the arrays, pointing to either deeper sources or to complex ray paths. For a better understanding of possible errors of the multi-array analysis, we also compare slowness values obtained from the array analysis with those derived from earthquake locations from classical (local network) localizations. In general, the results agree well. Nevertheless, there occur some systematic deviations of the array-derived backazimuth and slowness values that can be quantified for certain event locations.

20   **1 Introduction**

Many typical volcano-seismic signals, such as long-period events or tremors, lack clear and impulsive phases. To retrieve information about the characteristics of these events, including their hypocenters, multiple small-aperture seismic antennas have been utilized in past studies at different volcanoes. For example, Almendros et al. (2001a,b) were able to resolve a detailed 3D image of the source region of long-period events at Kilauea, Hawaii, using three arrays. The same arrays were used to

25   discriminate between different wave field components of Kilauea volcano, such as background tremor or surface waves (Almendros et al., 2002). The source of explosion quakes at Stromboli volcano, Italy, could be located using two seismic antennas (La Rocca et al., 2004). Also, Etna in Italy has been the subject to multi-array studies. For example, Saccorotti et al. (2004) deployed two arrays in 1999 to locate sources of tremor during a decreasing eruptive activity. The tremor of the 2004-2005 eruption has been the subject of the double seismic antenna study of Di Lieto et al. (2007). Almendros et al. (2007)

30   provided a model of the possible causes of seismicity during the seismic crisis of Teide volcano, Tenerife, in 2004 using three

arrays. The sources and mechanism of vulcanian explosions of Ubinas volcano, Peru, were analyzed with two seismic antennas by Inza et al. (2014). In 2014 the VolcArray study has been performed at Piton de la Fournaise, La Réunion, with three seismic arrays, each consisting of 49 stations (Brenguier et al., 2016). By applying array techniques and ambient noise cross correlations, multipath body waves could be separated and direct and reflected surface waves were extracted (Nakata et al.,

35 2016). The data from the same arrays have also been used by Mao et al. (2019) who monitor relative changes of the velocity in the shallow crust and by Takano et al. (2020) who are able to resolve velocity changes below the detection limit of geodetic measurements from ballistic waves. These examples represent only a small selection of multi-array studies at volcanoes, however, they are indicative of a wide range of possible applications.

In this study we use multiple seismic arrays to investigate the seismic activity of Fogo and Brava. The two islands are located

40 in the southwest of Cape Verde (see inset Fig. 1), about 700 km west of Senegal in the Atlantic Ocean. Their volcanic origin is attributed to a mantle plume beneath the islands (Courtney and White, 1986). Fogo volcano shows frequent eruptions with intervals of about 20 years, where the last took place from November 2014 to February 2015 (González et al., 2015). This is in contrast to the other volcanoes of the Cape Verde islands, which did not experience eruptions since the settlement in 15th century. Nevertheless, there is evidence for volcanic activity beneath and around the western islands of both (northern and

45 southern) chains of the Cape Verde. The activity occurs either beneath the islands or offshore in fields of submarine volcanic cones, including the Cadamosto Seamount southwest of Brava (Faria and Fonseca, 2014; Vales et al., 2014; Leva et al., 2020). It also involves the high seismic activity beneath and around Brava. This seismicity is characterized by a shift in location over time and frequent variations in the intensity of the seismic activity (Leva et al., 2020).

Despite the frequent volcanic eruptions, Fogo shows a rather low rate of seismicity compared to its neighbour Brava. In Fogo,

50 we mainly ~~find~~detect seismic events with a transition from high to low frequencies and without clear S-phases. In a previous study by Faria and Fonseca (2014), ~~T~~this type of event has been described as hybrid event, as it combin~~ing the~~es features of a volcano~~-~~ tectonic event in the signal onset and of a long-period event with respect to the coda (see, e.g., McNutt, 2000; Wassermann, 2012). ~~For the localization of the epicenter of these events we perform a time-domain multi-array analysis.~~

There exists a variety of different methods to locate typical volcanic seismic events, depending on the characteristics of the

55 signals. Emerging events, for example, have been analysed using probabilistic approaches to determine the maximum likelihood of the source location (e.g. Saccorotti et al., 1998; Métaxian et al., 2002; La Rocca et al., 2004). Some of these methods involve the application of a velocity model (e.g. Almendros et al., 2001a,b). In the present study however, we focus on events which show a clear signal onset. It allows us to clearly identify the same phase at all stations and arrays. For events with a pronounced first arrival classical beamforming can be applied and works well (Rost and Thomas, 2002; Schweitzer et

60 al., 2012). This method can be described as a delay-and-sum procedure in the time domain. The time-domain multi-array analysis has the advantage of being independent of velocity models. The velocity structure is often very complex in volcanic environments and there is, so far, no detailed 3D-velocity model available for Fogo volcano or Brava island. The time-domain array analysis allows for the incorporation of a narrow time window while including a broad frequency band (Leva et al., 2020). As a result, the central peak of energy of the corresponding (broad-band) array transfer function, tends to become more

65 narrow with sidelobes suppressed (Singh and Rümpker, 2020). Traces are shifted and stacked in the time domain to increase the SNR and to retrieve information about the incoming wavefront (i.e. the backazimuth and the magnitude of the horizontal slowness, which corresponds to the inverse of the apparent velocity (Rost and Thomas, 2002)). Including multiple arrays allows the localization of the event in the area of the intersected beams. In our study, we operated three arrays, two on Fogo, one on Brava, and seven short-period single stations from January 2017 to January 2018. We focus on volcano-tectonic

70 earthquakes originating in the study area around Brava and Fogo, and on certain volcano-seismic events on Fogo, which we, in accordance with Faria and Fonseca (2014), interpret as hybrid events occurring on Fogo. However, due to ray bending, it is possible to observe systematic deviations in backazimuth and slowness values can be observed. To investigate these deviations at the three arrays of our study, we compare multi-array localizations with locations derived from standard (network-based) localization techniques (see, e.g., Krüger and Weber, 1992). These standard techniques are based on the picking of P- and S-

75 phases. For this comparison, earthquakes occurring within the network are chosen, i.e. earthquakes beneath or close to Brava or Fogo.

2 Network

From 18 January 2017 to 12 January 2018 we operated a total of 37 seismic stations on Fogo and Brava (see Fig. 1). Our network comprised three arrays each consisting of 10 stations. Two arrays were deployed on Fogo close to the villages of

80 Achada Furna (AF) and Curral Grande (CG), the third one on Brava (BR). Another seven stations were operated as single short-period stations to complement the network – two on Brava, five on Fogo. All seismometers used are three-component instruments.

The design of the arrays is based on the array transfer function (in terms of frequency and the slowness components). The frequencies are chosen between 5 and 10 Hz, corresponding to mean dominant frequencies of the local events. These events

85 include volcano-tectonic earthquakes as well as other types of volcano-seismic signals such as hybrid events and harmonic tremors, the latter characterized by a frequency range between 1 to 5 Hz. Each array is circular and consists of a central station with two concentric rings with diameters of 700 m and 350 m, respectively. Four of the ten stations at each array are equipped with broad-band stations, the other six stations with 4.5 Hz short-period sensors (see lower right inset map in Fig. 1). As we expect events with mean frequencies between 5 and 10 Hz, the array is optimized for mean frequencies of 7.5 Hz. The array

90 transfer function for 7.5 Hz is shown in the supplementary material (Fig. S1c). It shows a single sharp maximum of energy and only minor secondary peaks. The circular shape of the array leads to a circular, symmetric peak in energy, which allows the detection of incoming wave fronts from any direction. For comparison, the transfer functions for 5 and 15 Hz are shown in Fig. S1 as well. These conventional array responses are calculated for single frequencies. However, when performing a time-domain array analysis, a broader frequency band is incorporated implicitly. The stacking of (integration over) the array

95 response of a wider frequency band leads to a much narrower peak. Additionally, sidelobes are strongly suppressed.

3 Method and data analysis

Criteria for applicability of the classical localization of local earthquakes are clear phases of the signal and a network distributed around the origin of the signal. If these criteria are not met, array techniques can help to locate the seismic event. By beamforming, the backazimuth and the magnitude of horizontal slowness are determined. For this purpose, the coherent part of the signal is shifted in time and summed up (Rost and Thomas, 2002). This method is based on the assumption, that the wavefront approaching the array approximates a plane wave, which is a valid assumption if the distance between array and source is considerably larger than about ten times the wavelength of the signal (Schweitzer et al., 2012).

Performing an array analysis for local events using only one array necessitates an epicentral distance estimation. In a previous study, we determined the epicentral distance based on the S–P travel-time difference. We also assumed a simplified two-layer velocity model and a fixed event depth (for details see Leva et al., 2020). However, this approach may cause significant uncertainties in the localization due the choice of the velocity model. In the present study, to overcome this limitation, we perform a multi-array analysis. A beam, pointing towards the epicenter, is determined at each array. Transferring these beams onto a map, allows to determine an area of overlap. This area provides the epicentral location of an event. A main advantage of this method is its independence of a velocity model. Details of this method are described in the following sections.

3.1 Beamforming

The array analysis is performed in the time domain. The time-domain analysis is equivalent to the incorporation of a wide frequency band, while the stacking window is kept narrow around the relevant phase, e.g. the first arrival of the incoming signal (Singh and Rümpker, 2020; Leva et al., 2020). Events are detected by a trigger algorithm from the traces and chosen by the analyst for the array analysis. Traces are first band-pass filtered within the dominant frequencies of the signal (Fig. 2). The cut-off frequencies are chosen in view of the waveform spectrogram. Following, an analysis window is chosen by the analyst around the first onset of the signal. For the local events we analyze in this study this is typically in the range of one to two seconds. This is shown in Fig. 3a for an example earthquake at array AF. Later, traces will be shifted within this window in reference to the trace of the central array station. A stacking window (in red in Fig. 3a) with the length of one or two periods of the signal around the signal onset marks the phase, for which the beamforming is performed. This stacking window is chosen as narrow as possible around the first onset of the signal. Several tests showed that a length of one to two periods leads to the best results of the stacked traces and the resulting energy. However, the length is chosen individually for each array and each event. All windows are chosen in reference to the central array station. The trace of the central station is kept fixed during the time-shift of the remaining traces. This time-shift is performed by a grid search with slowness values from -0.3 to 0.3 s/km and a grid size of 124×124. For each grid node traces are shifted accordingly and summed up. The resulting energy is defined by

$$E(t) = \int\limits_{t_1}^{t_2} \left[\frac{1}{M} \sum_{i=1}^{M} x_i(t + \boldsymbol{r_i} \cdot \boldsymbol{s}) \right]^2 dt \,,$$

following Harjes and Henger (1973), where the waveform at station i is given by x_i and the number of stations by M. The vector $\boldsymbol{r}_i$ contains the coordinates of the array stations in reference to the central station and the slowness vector is given by its horizontal components $\boldsymbol{s} = (s_x, s_y)$. The resulting contour plot of the energy is shown in Fig. 3b. The slowness components s_x and s_y corresponding to the maximum energy are further used to determine slowness and backazimuth of the event. Slowness, apparent velocity, and backazimuth are estimated using the expressions $s = \sqrt{s_x^2 + s_y^2}$, $v_{app} = 1/s$, and $BAZ = (180°/\pi)\arctan(s_x/s_y)$, respectively. The traces, which are shifted according to the determined slowness of the event, and the corresponding sum trace are displayed in Fig. 3c. The analogue procedure for array BR is shown in Fig. S2.1a-c and for array CG in Fig. S2.2a-c.

The horizontal slowness (or ray parameter) is related to the angle of incidence by $\sin(i) = s \cdot v_c$, with the velocity v_c of the upper layer beneath the array. Thus, the lower the slowness, the steeper the wavefront arrives at the array. For near-vertical angles of incidence, the slowness s becomes close to zero and the apparent velocity approaches infinity.

3.2 Multi-array analysis

After determining the energy grid of each array, the beams are intersected in the next step to obtain the earthquake epicenter from the multi-array analysis. For the error determination, the standard deviation of the maximum energy is determined. This is estimated in dependence of the chosen stacking window by randomly varying the start and end times of the stacking window 100 times by values between -0.2 s and 0.2 s. The values of ±0.2 s for the variation of the start and end times of the stacking window were chosen after performing tests with values between ±0.1 s, ±0.2 s and ±0.5 s. For the variation of ±0.1 s the standard deviation becomes very small. There is nearly no deviation from the original result, which means the resulting error is very likely underestimated and not reliable. Regarding the fact that some stacking windows are as small as 0.06s the variation of ±0.5 s proves to be too large and often leads to stacking windows far away from the signal phase of interest. Backazimuth and slowness values, thus, exhibit deviations that are far too large. Examples of the stack of the 100 energy estimations are shown in Fig. 3d for array AF, Fig. S2.1d for array BR and Fig. S2.2d for array CG. In some cases (like in Fig. S2.1d for array BR and in Fig. S2.2d for array CG) the main beam broadens, pointing to a higher sensitivity of the event at the specific array to the choice of the stacking window. If the stack of the 100 energy estimations is comparable to the original energy stack ~~(like in Fig. 3d for array AF)~~, the choice of the stacking window has nearly no impact on the determined beam.

The standard deviation of the slowness value is used as the error of the slowness at each array. In the next step the beam is determined. First, the standard deviation of the backazimuth is ~~determined~~ estimated as described above and converted into an error X in percent ($X\,[\%] = $ ~~100%~~ $(std[°]/360°) \cdot 100\%$). This error is then transferred to the contour plot of energy, where the energy values within this error range are determined. This yields minimum and maximum backazimuth values which frame ~~these energy values around~~ the maximum peak in energy. The beam width is defined by these values, which implies that it accounts for possible uncertainties and may be asymmetric with respect to the maximum energy value. Additionally, in this way small sidelobes are included in the multi-array analysis, if their energy values are larger than the error (in percent)

determined before. This can be seen in Fig. 4, where a small beam at array CG points to the south-southwest. The two values
within which the beam is plotted are referred to as the outer range of the beam from here on. Due to possible errors and
deviations of the main beam, we do not expect that all three main beams intersect in the same point where "main beam" refers
to the beam with the maximum energy. The backazimuth range defined above is used to plot the related beam. A large standard
deviation leads to a broad beam (see Fig. 4, where array AF shows a standard deviation in backazimuth of ±72.8°). To ensure
that a localization remains possible even for beams with relatively large standard deviations, However, a depiction of the beam
energy with higher resolution would be desirable. To achieve this, we further intersect the broad beam in steps of 1% of the
error estimated from the standard deviation. Thus, we assign values from 1 (broadest beam) to 100 to each of the steps. This
is further shown in the supplementary Fig. S3.

Now these beams are transferred to a map spanning the geological geographical coordinates of the research area, with the array
location as origin of the corresponding beam. The maximum value, which can theoretically be reached when intersecting the
three beams, is used to normalize their values. After intersecting the beams, the area with the highest probability of the event
location is determined. The last step is to choose a narrower section of the map that includes the arrays and the most likely
epicenter determined in the previous step (Fig. 4).

We choose a confidence interval of 90% of the maximum value of the intersected area as error for the multi-array analysis.

3.3 Error considerations

Different factors have an influence on the uncertainty of the result of the multi-array analysis. They can be divided into two
categories: uncertainties related to the parameters of the array analysis and due effects of the ray paths. Parameters in
connection with the analysis are the frequency range of the data and the length of the stacking window. Effects along the ray
path from the source to the array, such as heterogeneities, can result in a systematic deviation of backazimuth and horizontal
slowness at the array.

To test for the influence of the chosen stacking window (start and end time) and the frequencies, multiple repetitions of the
analysis at one array are computed with a random variation of these parameters. Results concerning the variation of the stacking
window are accounted for as described in Sect. 3.2. The same analysis has been performed for varying cut-off frequencies. For
earthquakes the lower frequency is randomly varied between 2 and 8 Hz, the upper frequency between 15 and 30 Hz. For
hybrid events the variation was between 1 and 4 Hz and 10 to 20 Hz for the lower and upper frequencies, respectively. The
analysis is done 100 times and the resulting standard deviation is again used to display the energy beam in the multi-array
analysis. The results for varying cut-off frequencies show a minor influence on the backazimuth, as demonstrated in the
supplementary Fig. S4. We conclude that, for a given stacking window, the variation of the frequency band can be neglected
in the error determination. The selection of the stacking window has a larger contribution to possible errors and is thus included
in the analysis, the results of which are displayed in Fig. 4. However, the choice of the stacking windows is performed
individually for each array and event. One advantage of the time-domain array analysis is the possibility to focus on the very
first phase of the signal. Therefore, it is preferable to select the stacking window as narrow as possible, as long as it

encompasses the first arrival. The analyst chooses the value of the window length according to the outcome of the stacked traces and the energy plot. If the traces are not aligned properly, the sum trace displays no clear signal onset or the energy plot shows large uncertainties, as expressed, for example, by large or several sidelobes. If the result is not reproducible after several trials, it is discarded from further analysis.

Velocity heterogeneities beneath the arrays or along the ray paths can possibly lead to a systematic bias in slowness and backazimuth determination (Rost and Thomas, 2002). This deviation of horizontal slowness and backazimuth at the arrays can be determined by comparing backazimuth and slowness values with those derived from a different localization technique (e.g. Krüger and Weber, 1992; Schweitzer et al., 2012). With respect to local events, we decided to locate earthquakes with a classical analysis (using SEISAN (Havskov and Ottemoller, 1999), based on the HYPOCENTER code of Lienert et al., 1986) by including all single stations of our network and one station of each array. For this standard localization technique, we apply the velocity model from Vales et al. (2014). To ensure the reliability of the classical localization, only earthquakes within or very close to the network were used. This comprises only earthquakes beneath Brava, Fogo and those located between the islands. Additionally, we only used results for which the rms values and errors of the classical analysis are small (rms < 0.25 s, errors < 5 km in longitude, latitude and depth). In total, a number of 13 events fulfilled all criteria and could be used for the comparison. Figure S5 contains a map showing the locations of the classically located earthquakes including error bars. The corresponding reference backazimuth and magnitude of horizontal slowness of these events (determined using the velocity model of Vales et al. (2014)) are compared to the respective values of the array analysis. The resulting vectors, pointing with a blue line from the backazimuth and slowness value of the array analysis (red points) towards the respective values from the classical analysis, are displayed in Fig. 5. In the range of 240°-270°, array AF systematically yields backazimuths pointing too far to the south by about 7° and array CG shows backazimuth values too far to the north with a mean deviation of about 9°. The values at the array on Brava show a variety of deviations due to the many directions of incoming waves. It appears, that backazimuth values in the range of 270°-360° point too far to the north. However, for the comparison with the classical localization, there are only four events within this range, prohibiting a reliable statement on systematic deviations.

The station elevation differences of the array stations can have an impact on the result of the array analysis. Therefore, we carefully tested possible influences under the assumption of the different station elevations according to Schweitzer et al. (2012). It turned out that the station elevation differences are small enough to be neglected.

For a successful localization with multiple arrays certain requirements need to be fulfilled. For example, the stacking windows at each array should contain the same phase of the signal (Almendros et al., 2002). To ensure this, we perform the multi-array analysis on the first onset of the signal. Additionally, the occurrence of strong sidelobes in the energy stack must be avoided as the occurrence of secondary peaks results in two or even more beams at one array. This may lead to event mislocations. Furthermore, the occurrence of strong sidelobes generally indicates higher uncertainties in results. Regarding the intersection of the beams additional considerations must be taken into account. If beams trend almost parallel, the epicenter will be located far away with a large uncertainty in distance (see Fig. 6a,b). Furthermore, if two beams point from one array to another, the whole area between the arrays will be a potential source region, leading to large errors in the localization (see Fig. 6c,d). In

these two cases the third array is of particular importance, as it will strongly narrow down the area of the likely source. If the third array does not provide any additional information in such cases the localization of the corresponding event must be discarded due to the high level of uncertainty. Also, considering that the backazimuth and horizontal slowness show small but systematic deviations, it is not unlikely to find a result, where the three beams do not overlap in the same area. To be able to assess the reliability of the location obtained during the analysis, information about the epicentral distance are added to the map of the intersecting beams. This can be used especially for the analysis of earthquakes: Here, S–P travel-time differences are determined for each array. From these traveltime differences the epicentral distance of the event to each array is estimated. For this we apply a two-layer velocity model with a mean crustal and a mean mantle velocity, derived from Vales et al. (2014), and a fixed event depth of 5 km (see Leva et al., 2020). The fixed event depth has been defined after estimating a mean event depth from previous studies of the region around Brava (Faria and Fonseca, 2014). The resulting epicentral distances are indicated by circles which are plotted on the map with the intersecting beams (see Fig. 7). However, this information about the epicentral distance is not included in the localization, as we want to retrieve the source epicenter without applying a velocity model. It only serves as a reference for the analyst to evaluate whether or not the estimated source location is reasonable. Due to the lack of S-phases this estimate is not used during the analysis of hybrid events. However, here the array locations with respect to the event locations is very favourable, as the beams intersect almost perpendicular. This prevents the occurrence of parallel trending beams and beams pointing towards each other.

4 Results

The majority of the recorded events are local volcano-tectonic earthquakes mainly occurring in the area of Brava. However, we also observe a different type of ~~hybrid~~ events which are recorded by the stations on Fogo. These events are characterized by a smooth transition from higher (15–40 Hz) to lower (1–10 Hz) frequencies and a lack of S-phases. As these are characteristics of hybrid events (e.g. McNutt, 2000; Wassermann, 2012), we follow previous studies on the seismicity of Fogo (Faria and Fonseca, 2014) and use the same terminology. Figure 8 shows traces and spectrograms of the~~se~~ two types of events. In the following we will focus on events, which were initially detected by a trigger algorithm and selected for further analysis by visual inspection.

4.1 Volcano-tectonic earthquakes

The volcano-tectonic earthquakes on average occur 8 times a day (see Fig. 9a). The rate of seismicity frequently increases, leading to phases with elevated seismic activity. 2709 earthquakes were recorded from 18 January 2017 to 12 January 2018, 112 of which could be located using multi-array techniques. The earthquakes mainly occurred around Brava (Fig. 10). The reason for the discrepancy in the number of detected and located earthquakes are manifold. Many smaller earthquakes are recorded with our stations only on Brava, thus precluding the multi-array analysis, as for this at least two arrays must detect the event. As described in Sect. 3.3, there may be cases, where the result of the array analysis must be discarded and cannot be used for further analysis. Thus, the multi-array analysis can only be performed for events with stable results for the

backazimuth determination. If the energy grid shows e.g. strong sidelobes or the choice of slightly different stacking windows for the same event leads to strongly different results, the result of this array for this particular event is discarded. Additionally, at least two arrays must show reliable and stable results, which further reduces the number of located events. The recordings of the stations on Fogo show a rather high frequency content with the main frequencies between 10 to 30 Hz (Fig. S6a). On Brava the dominant frequencies of the same event are lower and range between 2 and 20 Hz. The corresponding spectrum is shown in the supplementary material (Fig. S6b).

The mean apparent velocities at the arrays on Fogo are in the range of 7.1 km/s for events originating close to Brava. For such a distance between event and array, the ray is first propagating downwards from the source. In a medium with lateral homogeneous velocities, the apparent velocity of this ray measured at the array is equivalent to the velocity at the ray turning point. Apparent velocities <8 km/s thus point to a ray turning point within the crust (velocity model taken from Vales et al., 2014), indicating crustal depths of the earthquakes. Note that the array on Brava shows higher apparent velocities for the same earthquakes with a mean of 10.8 km/s. However, array BR is located closely to the sources, which results in a steeper angle of incidence (and smaller slowness) compared to the arrays located on Fogo.

The supplementary material contains a map with error bars of the analyzed earthquakes (Fig. S7).

4.2 Hybrid events

As described above, Tthe hybrid events observed on Fogo (see Fig. 9b) are characterized by high frequencies (15–40 Hz) at the beginning of the signal, followed by low frequencies (1–10 Hz) and a lack of clear S-phases. The signals mainly last about 20 to 30 seconds, some last up to 1 minute, and usually reach station CV10 first, where they also show the largest amplitudes. Figure 8b shows an example event recorded at a broad-band station of the array AF. Vertical traces of such an event are displayed in the supplementary material (Fig. S8). The spectrograms of all components are shown in Fig. S9 and reveal the low frequency coda, where more energy occurs in the 1-10 Hz band than before the event onset. As the hybrid events were only recorded by the stations on Fogo, they were located using the arrays AF and CG. We observe 125 hybrid events, 12 of which could be located. Figure 11a shows the resulting epicenters, in or close to the collapse scar of Fogo, Chã das Caldeiras. The events exhibit rather high apparent velocities, in average 7.8 km/s at array AF and 8.4 km/s at array CG. The mean errors of these velocities are 2.9 km/s and 2.8 km/s at array AF and CG, respectively. These high apparent velocities are not biased by the choice of the slowness range during the array analysis. We also tested a slowness range between -1 to 1s/km, which does not change the outcome (see Fig. S10.1 and S10.2). To determine the source location of the hybrid events, we superimpose the beams of all localizations of the hybrid events. This is shown in Fig. 11b, where the area with the highest likelihood for hybrid-event occurrence is marked by a white line (corresponding to 80% of the maximum of the relative sum of energies). We find this area in the northwestern part of Chã das Caldeiras.

5 Discussion

Most earthquakes occur around and beneath Brava and the seismic activity shows several periods with increased seismicity
(Fig. 9a). This is a common observation for the seismicity around Brava (Faria and Fonseca, 2014; Vales et al., 2014; Leva et
al., 2020). The earthquakes originate in the crust as derived from the apparent velocities measured at the arrays. Performing
the time-domain array analysis allows for the determination of the epicenter of small local earthquakes ($M_L < 0.5$), although
the P-wave arrival is not clearly visible at all stations. However, their combination during the beamforming results in a clear
P-phase onset of the sum trace. The application of the time-domain array analysis is favourable in such a case, as a wide
frequency band can be chosen to optimize the SNR. It is worth noting, that the frequency content of the earthquake recordings
on Brava generally exhibit lower dominant frequencies (supplementary material Fig. S6b) than the recordings of the same
events on Fogo (Fig. S6a). This is surprising, as higher frequencies are typically more attenuated. On the other hand,
observation of high-frequency tremor around Fogo has been described by Heleno et al. (2006). These authors report on the
conservation of high frequencies in a tremor signal even at larger distances (about 15 km) from the source. In September we
observe some earthquakes beneath Fogo (see Fig. 10) which occur within the shallow crust according to the apparent velocities
measured at the arrays and the S–P travel-time differences. These events are located close to the area, where deep subcrustal
earthquakes have been observed in August 2016 (see Fig. S11~~0~~; Leva et al., 2019). Nevertheless, due to their large difference
in depth and the long amount of time between these two occurrences, we cannot establish a link between them (as due to the
transport of magma from depth into the crust).

Apart from the earthquakes in September, Fogo mainly shows volcanic seismic signals, which are best described as hybrid
events (in total 125 in 2017). Their origin is located in the northwestern part of the collapse scar of Fogo and on top of the
Bodeira wall, which surrounds large parts of the collapse scar Chã das Caldeiras. It has been discussed in previous studies (e.g.
McNutt, 2000; Wassermann, 2012) that these events are caused by a combination of source mechanisms relevant for volcano-
tectonic earthquakes and long-period events. One such hypothesis is a volcano-tectonic earthquake, which triggers the
oscillation of a fluid-filled cavity (McNutt, 2000). At Fogo, hybrid events have been detected before (Faria and Fonseca, 2014).
They were attributed to hydrothermal processes at shallow depths (several hundred meters), due to the interaction of rainwater
and hot rock. This hypothesis is based on the seasonal variation of the number of hybrid events and a water table found at 370
m depth in the Chã das Caldeiras. We observe a variation in the number of events over the year of observation and compared
it with the amount of precipitation per month in 2017. The corresponding figure is shown in the supplementary material (Fig.
S12~~1~~). We find an increase of hybrid events from February to March and from September to November. The precipitation
shows a small peak in March, which might correspond to the peak of hybrid events. However, the strongest peak of
precipitation occurs in August. This does not directly correlate with the maximum peak in the number of hybrid events, which
occurs in November. From this, we conclude that a causal relationship between precipitation rates and the occurrence of hybrid
events cannot be established.

320 High apparent velocities of the hybrid events indicate steep angles of incidence, possibly pointing to a deep seated source. With the multi-array analysis applied in this study, it is not possible to estimate the depth of the events, as we do not include a velocity model. However, for an estimate of the source depth of the hybrid events, it is possible to derive the ray path under consideration of a simple velocity model and the angle of incidence. The velocity model is adapted from Vales et al. (2014) with velocity increasing at steps of 0.1 km. For this approach the modification of the velocity model was necessary to allow

325 for an incremental ray bending according to Snell's law. The ray path is traced back from the angle of incidence at the array until the epicentral distance is reached. This simple model yields event depths of 5 to 14 km. The uncertainties are relatively large. For some events the depth estimates at the two arrays can lead to different results which points to inconsistencies likely related to the insufficient (1D) velocity model. Additionally we considered other velocity models, which might be better suited regarding the expected complex velocity structure. Adapting a velocity model for Etna (Almendros et al., 2000) yields event

330 depths between 10 and 20 km. The use of the velocity model for the caldera of Tenerife (Lodge et al., 2012) yields results between 3.5 and 15 km. This shows the significant impact of the velocity model on the estimation of the angle of incidence at the array and the computed ray path. The event depths estimated from the slowness values observed at the arrays and the different velocity models would be significantly deeper than the depths reported in previous studies. There can be several reasons for such an observation. It is possible, that the source of the events has shifted to greater depths after the eruption of

335 Fogo in 2014. This might also explain, why there is no direct correlation between the precipitation data of 2017 and the number of hybrid events. Another possibility is that the wave field is affected by path effects caused by the complex structure of the volcanic edifice (Kedar et al., 1996). These authors suggest that a single pulse can trigger seismic waves, which then interact with heterogeneities in the elastic, loosely consolidated surrounding layers of the volcanic edifice, leading to complex harmonic seismic signals at the receiver. Such an effect is hard to discriminate from an oscillating resonator. Finally, Harrington and

340 Brodsky (2007) provide the explanation that hybrid events are not necessarily caused by fluid motion, but by brittle failure. Low rupture velocities and strong path effects result in the long low-frequency coda. Similar effects of low rupture velocities in unconsolidated volcanic material have also been suggested to cause the signature of long-period events, rather than fluid-driven source mechanisms (Bean et al., 2014). On Fogo, brittle failure at shallow depths could be caused by gravity loadings in the collapse scar after the latest eruption. In view of these previous studies and of our observations, i.e. the clear signal onset,

345 the lack of S-phases and the smooth transition from high to low frequencies without the appearance of definite dominant frequencies, we suggest that scattering effects along the ray path may explain the distinct appearance of the hybrid events on Fogo.

A complex ray path might also affect the slowness measured at the arrays. Almendros et al. (2001a) evaluate the influence of a complex 3D velocity structure of Kilauea, Hawaii, on the apparent velocity recorded at a seismic array. The results point to

350 a reduction of the slowness values in comparison to a homogenous velocity model. It is likely that the complex velocity structure of Fogo has an impact on the ray path and thus leads to slowness variations. This bias could possibly result in smaller slowness values and, thus, explain the high apparent velocities we measure. However, the assumed uncertainties of the apparent velocities are rather large and should cover this bias. In addition to these considerations, we observe strong differences in the

amplitudes at the stations. The amplitudes of hybrid events at station CV10 in the collapse scar are nearly twice as large as the amplitudes of the other stations on Fogo, not located this close to the source. The second station CV14 in the collapse scar was only operational during the last three months of the study. However, for the few events detected in this period, the amplitudes at CV14 are in the range of those at CV10, but the signal arrives slightly later than at station CV10. If the events would actually occur in depths of 5 to 14 km, we would not expect such a large difference in the amplitude ratios. For earthquakes occurring on Fogo or Brava, the amplitudes at station CV10 are in the range of the amplitudes at the other stations. A bias due to site effects at this station is thus unlikely. We thus conclude, that despite the high apparent velocities, the hybrid events should actually originate from shallower depths, as already suggested by previous authors (Faria and Fonseca, 2014). Nevertheless, a hydrothermal origin may not be necessary to explain their occurrence and their real cause remains unclear. The use of a high resolution 3D velocity model or a dedicated dense network of stations placed near the observed epicenters could contribute to a better understanding of these events, as it would allow for a more precise depth estimate.

Being independent of any velocity model and able to locate the epicenters of events without clear onset of phases or offshore, outside of the network, are strong advantages of the utilization of multiple seismic arrays. However, there are certain limitations of the multi-array analysis. The backazimuth and slowness determined with the arrays on Fogo and Brava show a systematic deviation, which has been estimated by a comparison with classically located events. The number of reference events (in total 13) is too small for a correction of backazimuth and slowness values during the analysis. However, some relevant conclusions can still be drawn for the utilization of the multi-array technique. At the arrays AF and CG on Fogo wavefronts arrive from a range of backazimuths of 240° to 270° (see supplementary Fig. S13.1). Within this range the backazimuth values show a mean deviation to the south of 7° at array AF and a mean deviation to the north of 9° at array CG. For the array BR on Brava observed backazimuth values cover a wide range (see supplementary Fig. S13.1) and slowness values can be small for events close to the array. Figure 5 shows larger deviations of backazimuth and slowness for events with horizontal slowness values below 0.1 s/km. The question arises, whether the results of array BR should generally be discarded when they show horizontal slowness values below 0.1 s/km. However, the beams related to the arrays on Fogo can easily trend almost parallel, leading to an over-estimated epicenter distance (based on comparison with the S–P travel-time difference, see Fig. 7). Therefore, the beam of the array on Brava is essential, as it usually locates the event closer to the expected location. This is shown in Fig. 6a,b. Generally, the errors in the events location, which result from the uncertainties of the backazimuth determination at array BR are by far smaller than the errors when using only the arrays on Fogo. The distance estimated from the S–P travel-time difference serves as verification of the epicentral distance determined by the multi-array analysis (see Fig. 7). This is especially helpful when only two arrays are available for a localization. Thus, a multi-array analysis using only two arrays is still possible, but might lead to a certain amount of earthquakes that cannot be located due to the deviation of backazimuth values. For the hybrid events on Fogo, the determination of the deviation vectors is not possible due to the lack of reference localizations. The distribution of backazimuth values of the hybrid events is displayed in the supplementary material (Fig. S13.2). The backazimuth values clearly indicate a location close to or in the collapse scar of Fogo. Nevertheless, a possible deviation should not lead to large errors in the localization, because of the location of the arrays with respect to the source region.

6 Conclusion

From January 2017 to January 2018, we operated three arrays on Fogo and Brava to apply a time-domain multi-array analysis
390   for seismic events occurring in this region. This application allows the epicentral event localization without assuming a velocity
model. This is a significant advantage in volcanic environments, where the velocity structure is difficult to constrain.
Additionally, we are able to determine the epicenter of offshore earthquakes outside of the network and hybrid events without
clear S-phases. Although the application of the time-domain multi-array analysis has many benefits, it is necessary to evaluate
possible errors of the localization, which may result from systematic deviations of backazimuth and slowness values
395   determined at the arrays. These deviations can be caused by heterogeneities along the ray path. To determine the deviations of
backazimuth and slowness values, we compare them to those derived from a classical earthquake analysis. It turns out, that
the number of reference events is too small for a reliable correction. We therefore allow for relatively large location
uncertainties to cover the possible deviations.

A large number of volcano-tectonic earthquakes are located beneath and around Brava. As reported previously (Faria and
400   Fonseca, 2014; Vales et al., 2014; Leva et al., 2020), we observe several periods of elevated seismic activity and a frequent
shift of locations around the island. Additionally, a few earthquakes occur beneath Fogo in the shallow part of the crust. Some
of them occur in the shallow crust in approximately the same epicentral area as deep subcrustal earthquakes of 2016 (Leva et
al., 2019). However, a conclusion concerning a possible link between these two occurrences could not be made due to the
rareness of such earthquakes. However, the majority of seismic events beneath Fogo are hybrid events. As shown by a joint
405   analysis of the events, their epicenters are close to the northwestern part of the Chã das Caldeiras and beneath the Bodeira
wall. These events show larger apparent velocities than the volcano-tectonic earthquakes recorded with the arrays on Fogo.
Most likely, these high values result from the influence of the topography and the complex velocity structure of the volcanic
edifice, leading to a possible bias in the slowness determination. Additionally, the station CV10 located in the Chã das
Caldeiras shows significantly larger amplitudes than the remaining stations on Fogo. We believe that the origin of the hybrid
410   events is not as deep as the high apparent velocities would suggest. However, the origin remains unclear due to the lack of
information about the depth. The application of a precise 3D velocity model or a dedicated local network could shed further
light on the depth and thus on the possible source mechanism of these events.

In addition to the volcano-tectonic earthquakes and the hybrid events, we detected isolated instances of volcanic tremor, which
we have not yet analyzed in detail. This will be subject of forthcoming studies.

415   **Data availability**

The data are available for download at GEOFON (https://geofon.gfz-potsdam.de, GEOFON, 2021). Please refer to Wölbern
and Rümpker (2020, https://doi.org/10.14470/4W7562667842).

Author contributions

The study and the setup of the seismic arrays were initiated and conceived by IW and GR. IW was also responsible for project

420 administration. CL analyzed the data and prepared the figures. CL wrote the manuscript as part of her PhD under supervision of GR. The manuscript was revised by GR and IW. All authors took part in the field work.

Competing interests

The authors declare that they have no conflict of interest.

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Figures
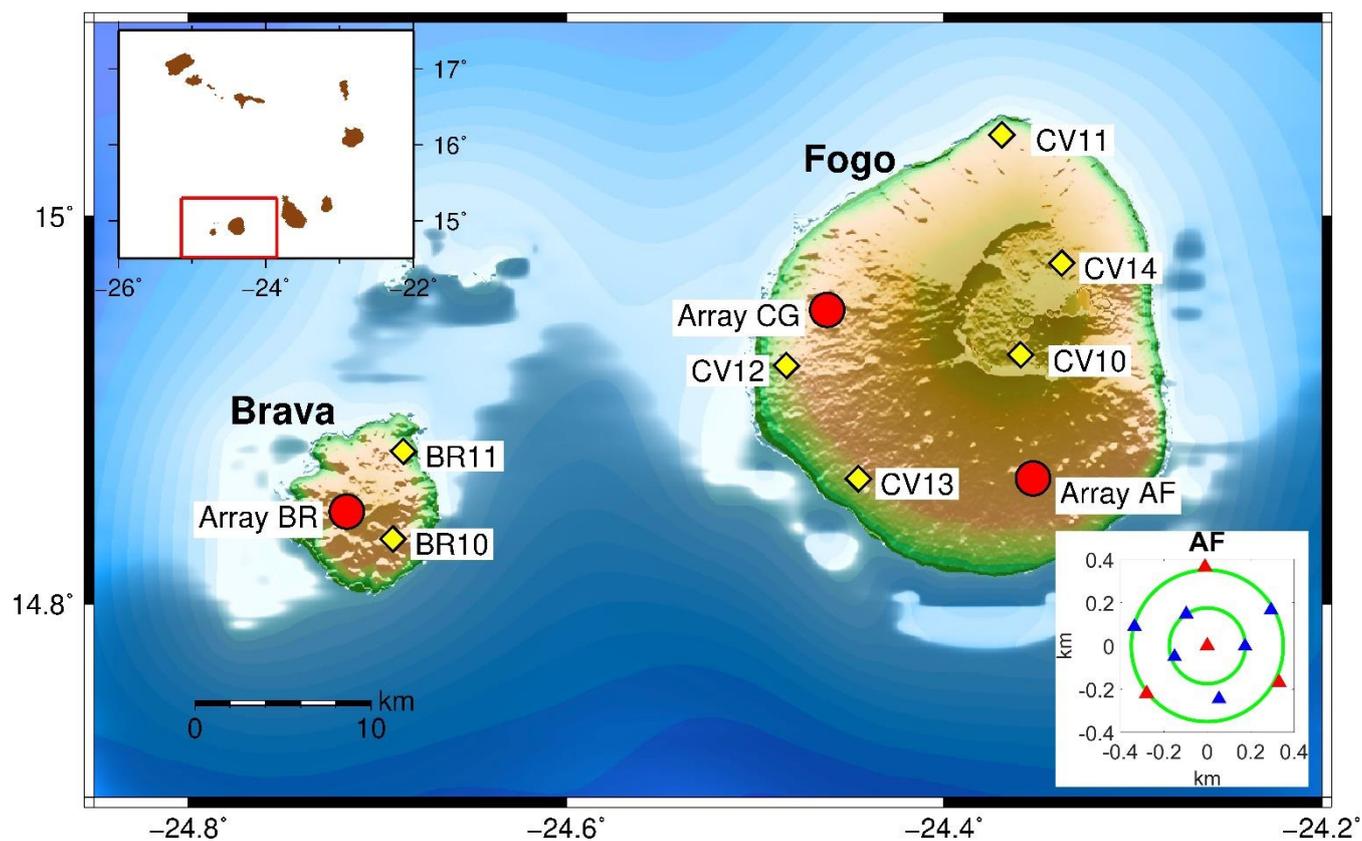
Figure 1: Station configuration on Fogo and Brava from January 2017 to January 2018. Red circles: array locations; yellow diamonds: short-period single stations. Left inset: Cape Verde, current section around Fogo and Brava marked in red. Right inset: setup of the array AF, red: broad-band stations, blue: short-period stations. The arrays BR and CG are designed in the same way. Topography and bathymetry data are from Ryan et al. (2009).
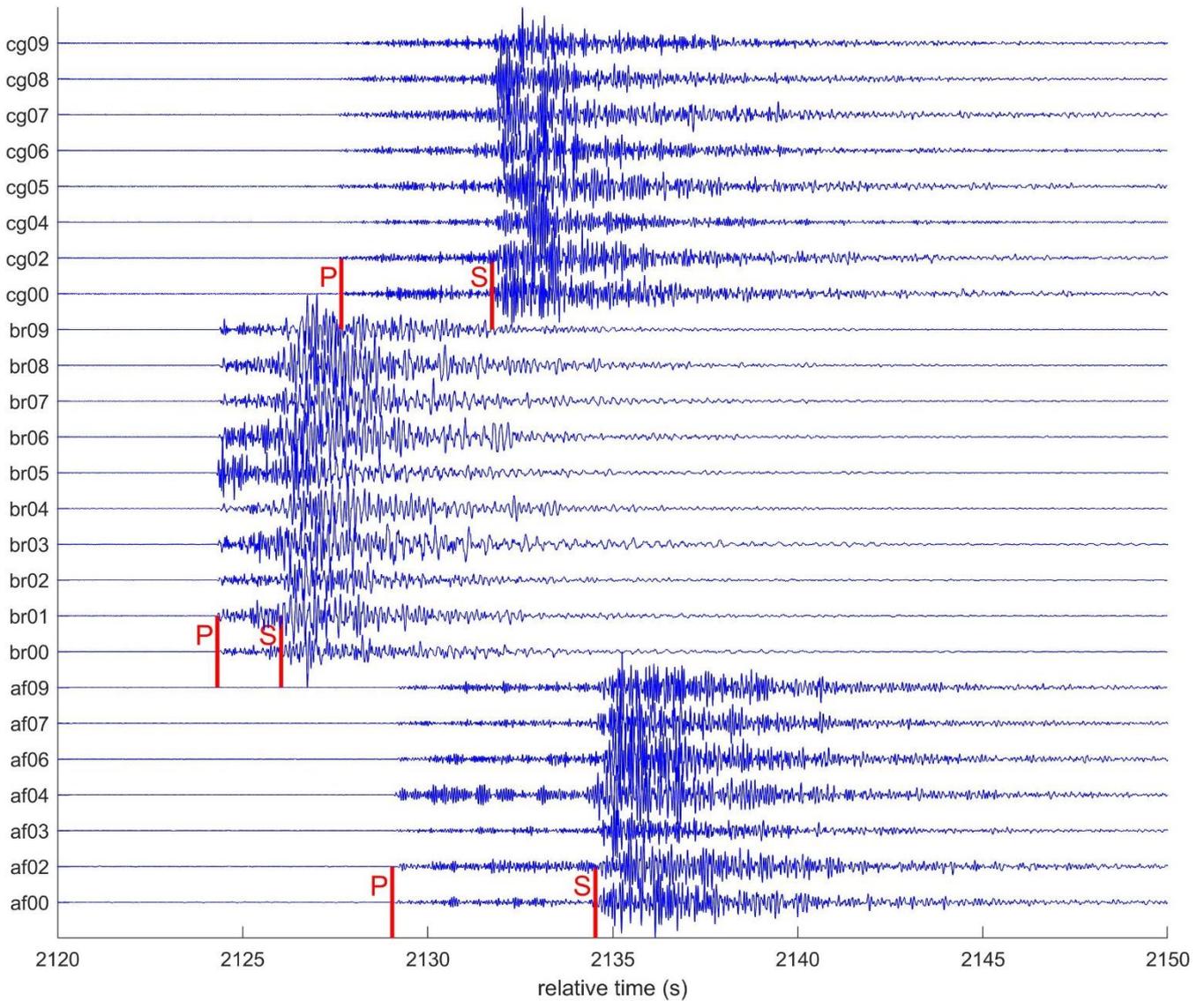
Figure 2: Z-components of the seismogram of an earthquake on 22 July 2017 (23:35 UTC) before the array analysis is performed.
Traces are filtered individually according to the spectrogram of each array. Filters applied here are: 2–20 Hz at array AF, 2–24 Hz
at array BR and 2–21 Hz at array CG. Red lines mark the P- and S-phases at the central array stations AF00, BR00 and CG00,
respectively.
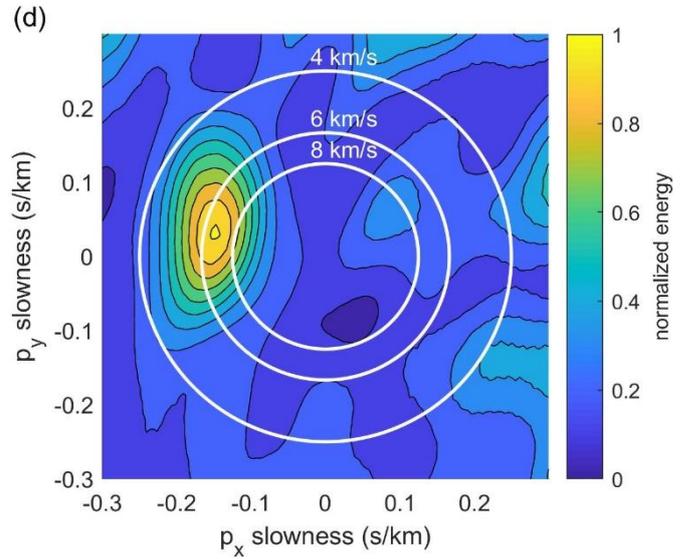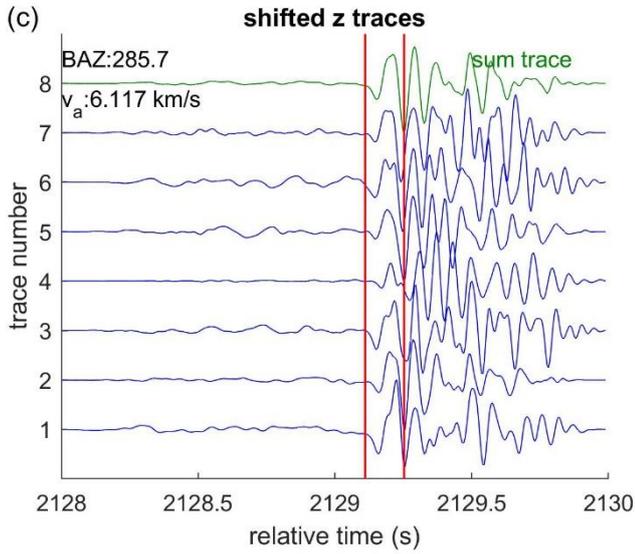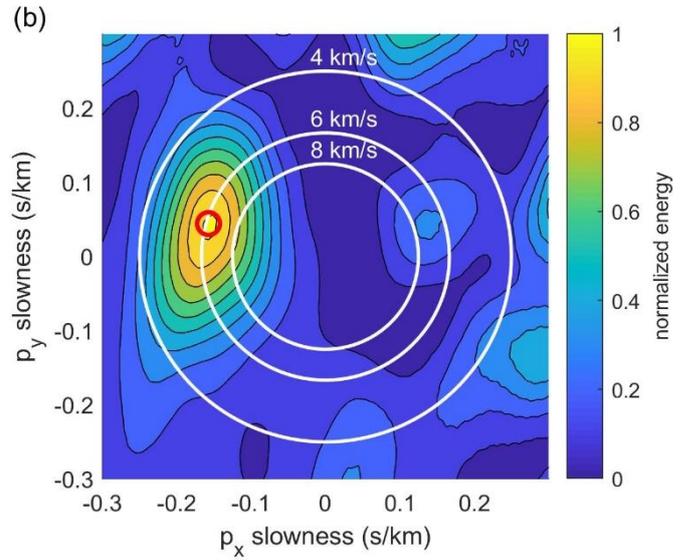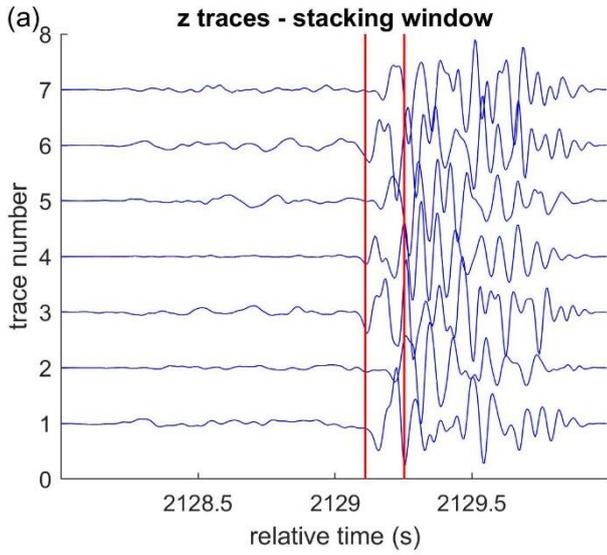
Figure 3: Time-domain array analysis of an earthquake on 22 July 2017 (23:35 UTC) at the array AF. (a) Analysis window of 2 s length with the stacking window marked in red. Traces are displayed before shifting and stacking and are filtered between 2 and 20 Hz. (b) Resulting time-domain energy stack. Red circle: maximum beam energy. (c) Time-shifted traces. The upper green trace represents the sum trace. (d) To retrieve the standard deviation of the backazimuth, the stacking window is varied 100 times by values between -0.2 and 0.2 s. The standard deviation is estimated from the 100 resulting backazimuth values. Shown here is the stack of the 100 energy plots.
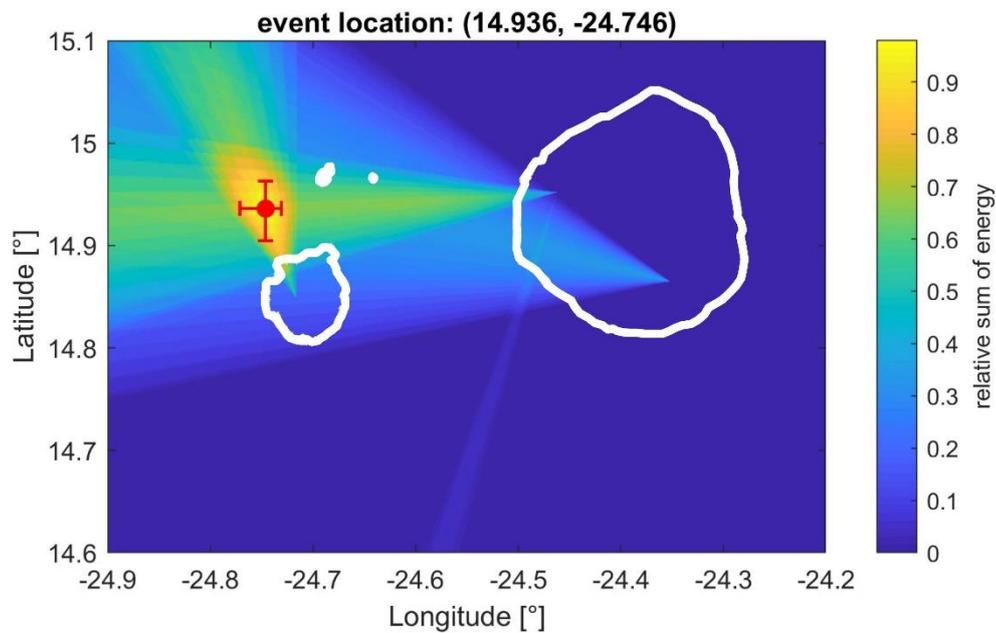
Figure 4: Intersection of the beams projected on a map section, including the coordinates of the arrays and the location of the determined epicenter. The intersected beams correspond to the example earthquake on 22 July 2017 (23:35 UTC). The small beam pointing south-southwest from array CG results from a sidelobe with energy values in the range of the error corresponding to the standard deviation of the backazimuth. Red circle: event location with error bars. Topography and bathymetry data are from Ryan et al. (2009).

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(a) **AF**

(b) **AF**

(c) **BR**

(d) **CG**

(e) **CG**

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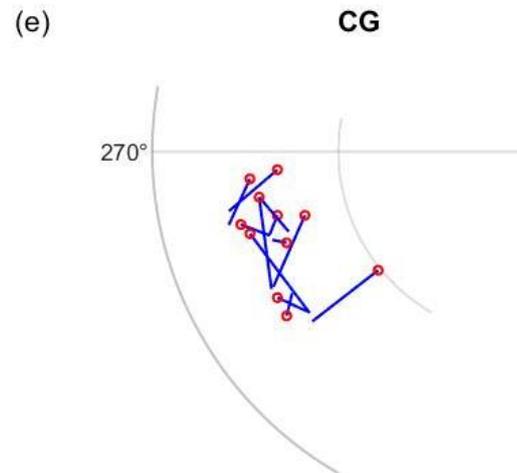
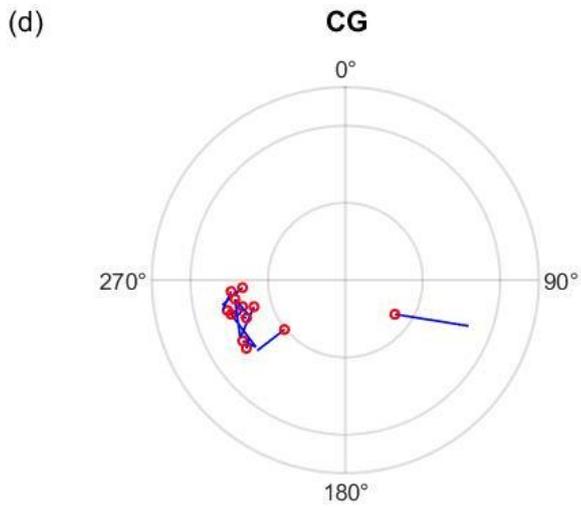
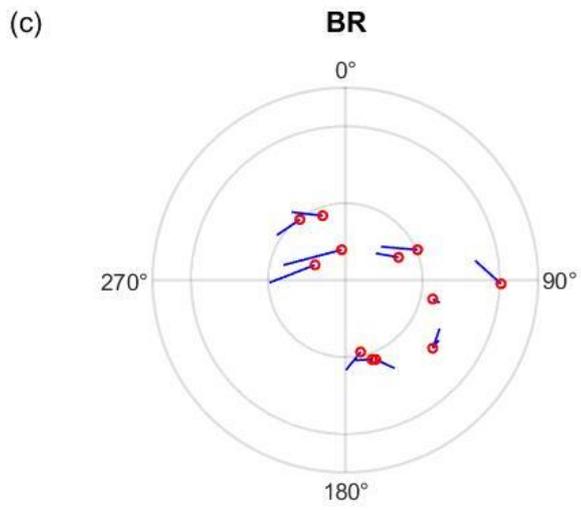
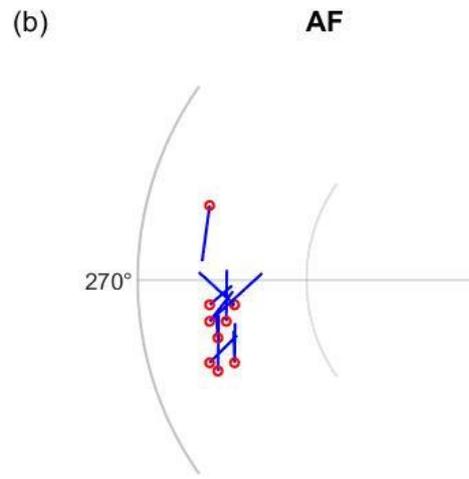
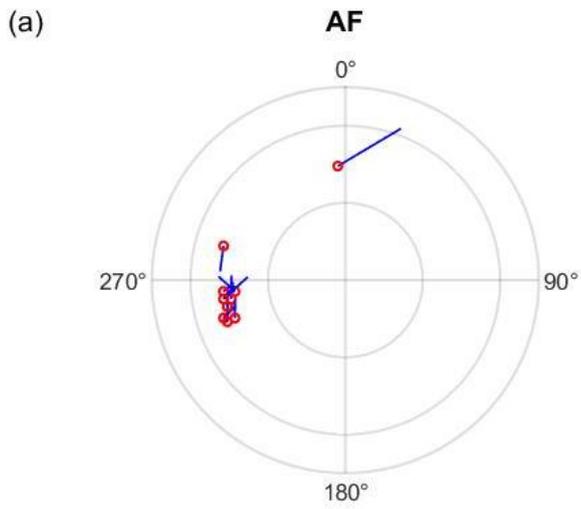


Figure 5: Deviations of backazimuth (BAZ) and horizontal slowness. Red points: Backazimuth and horizontal slowness values of the array analysis. Blue lines point towards the corresponding reference values of the standard localization. (a) Deviations of BAZ and horizontal slownesses, determined at array AF. Different radii correspond to slowness values of 0.1 s/km, 0.2 s/km and 0.25 s/km, respectively. (b) Deviations of BAZ in the range of 235° to 305° and horizontal slownesses at array AF. The radii correspond to slowness values of 0.1 s/km and 0.2 s/km, respectively. (c) Same as in (a) for array BR. (d) Same as in (a) for array CG. (e) Deviations of BAZ in the range of 210° to 280° and horizontal slownesses at array CG. The radii correspond to slowness values of 0.1 s/km and 0.2 s/km, respectively.

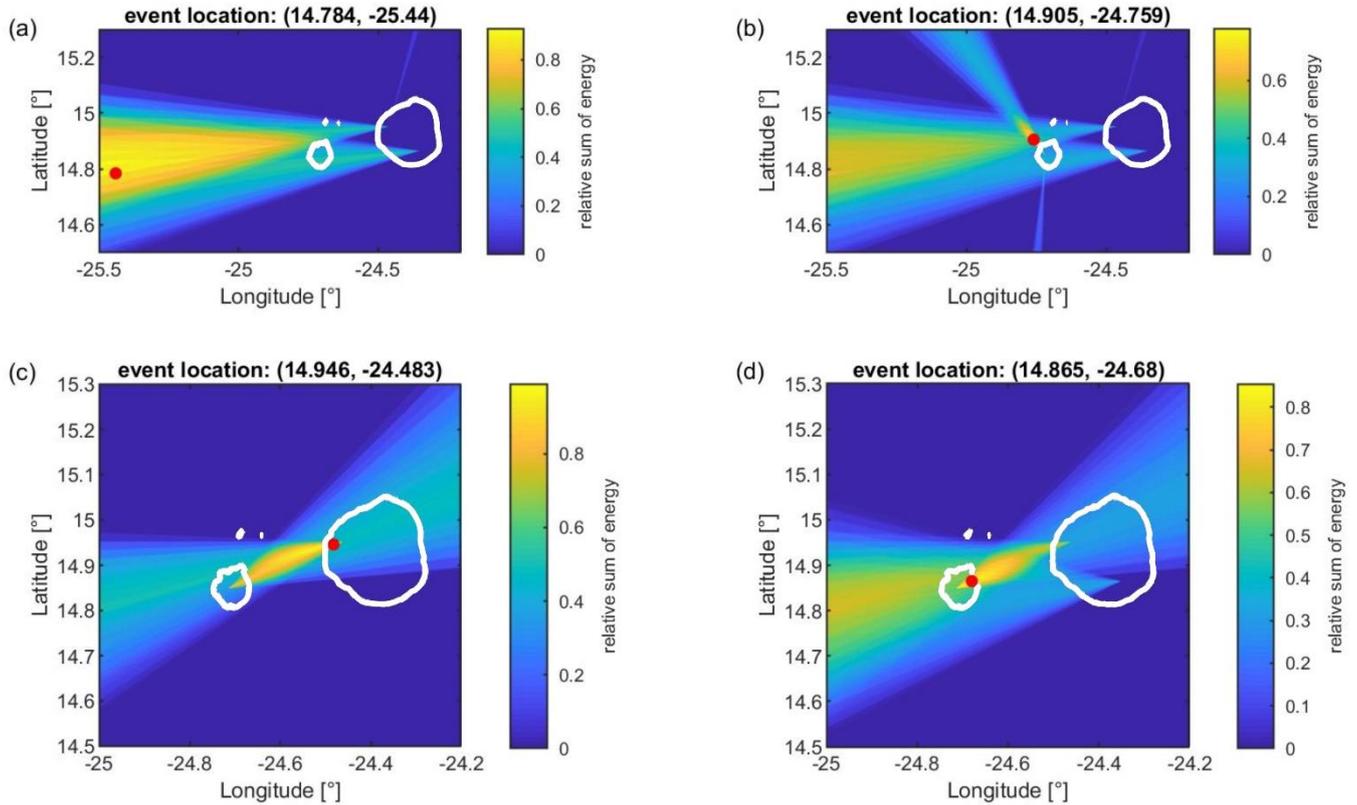



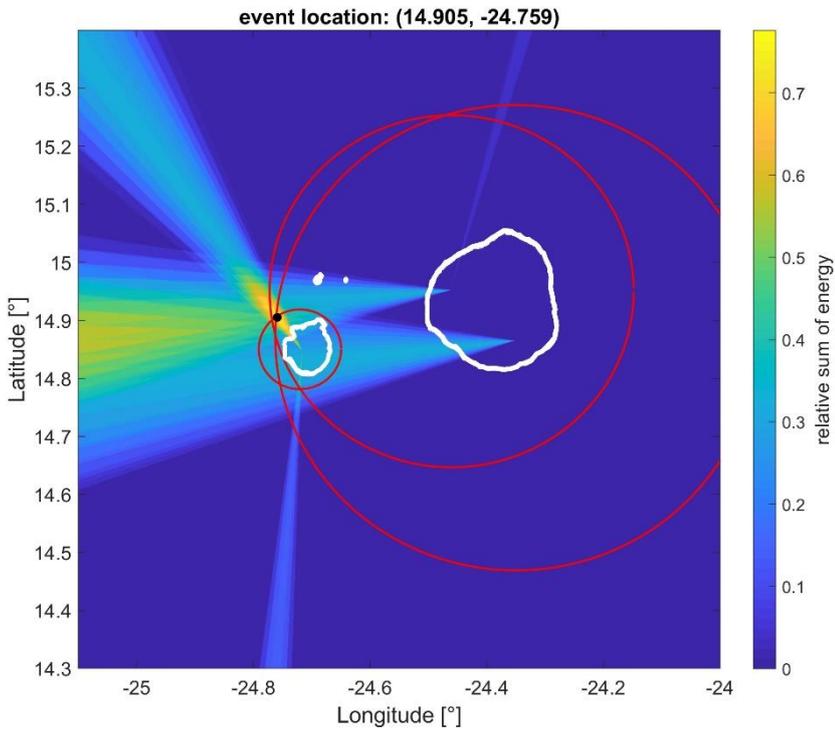
Figure 6: Examples of problematic localizations due to unfavourable source-receiver configurations. (a), (b) Intersection of the beams, (a) without the beam of array BR and (b) with the beam of array BR included. In the case of parallel trending beams (a) the localization of the event is distorted and the beam of the third array is needed (b).
(c), (d) Intersection of the beams, (c) without the beam of array AF and (d) with the beam of array AF included. In the case of beams pointing from one array to another, (c), the region of elevated energy spans a large area between the two arrays. In this case the beam of the third array is needed for a proper localization (d).

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Figure 7: Verification of the event localization using additional travel-time information. Black circle: location of the event derived from the intersecting beams; red circles: epicentral distances of the event estimated from S–P travel-time differences observed at the three arrays. The circles give an estimate of the expected distance of the event to the array providing a tool to better judge the reliability of the event location. Note, that this representation only serves as a support for the analyst. The final event location is only
575 **based on the multi-array analysis.**

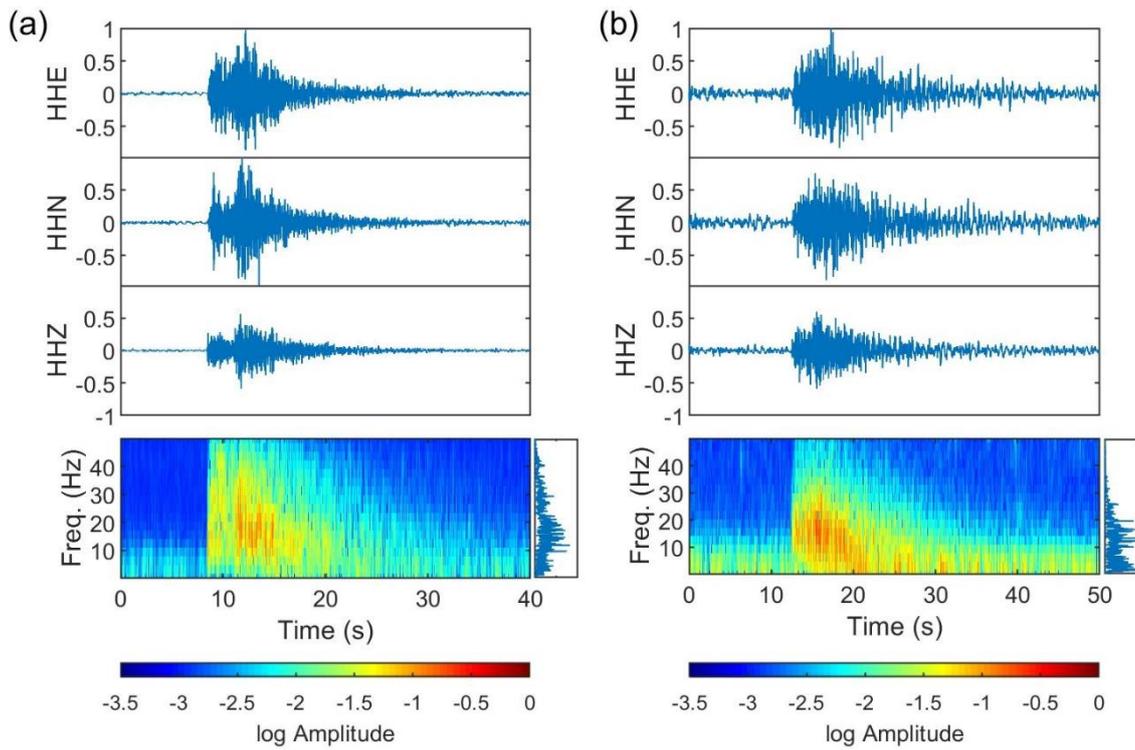



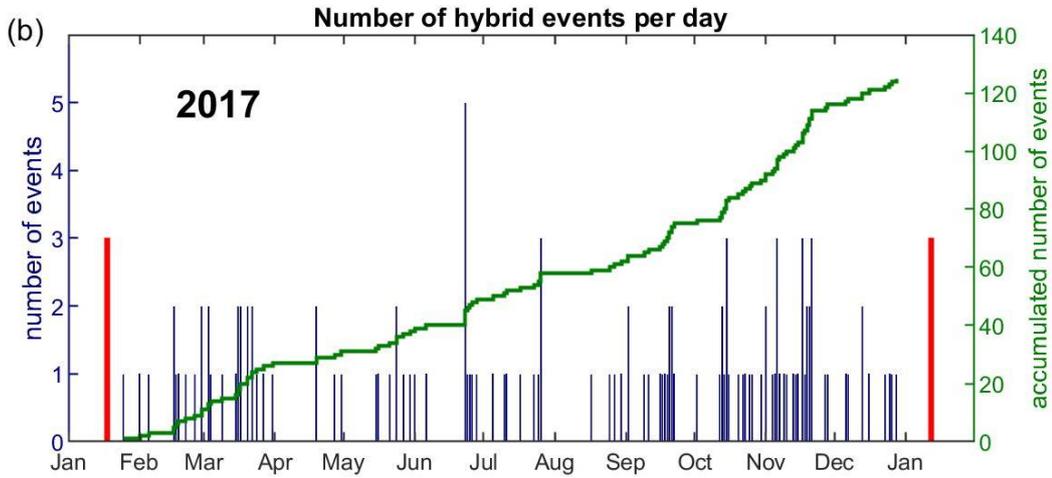
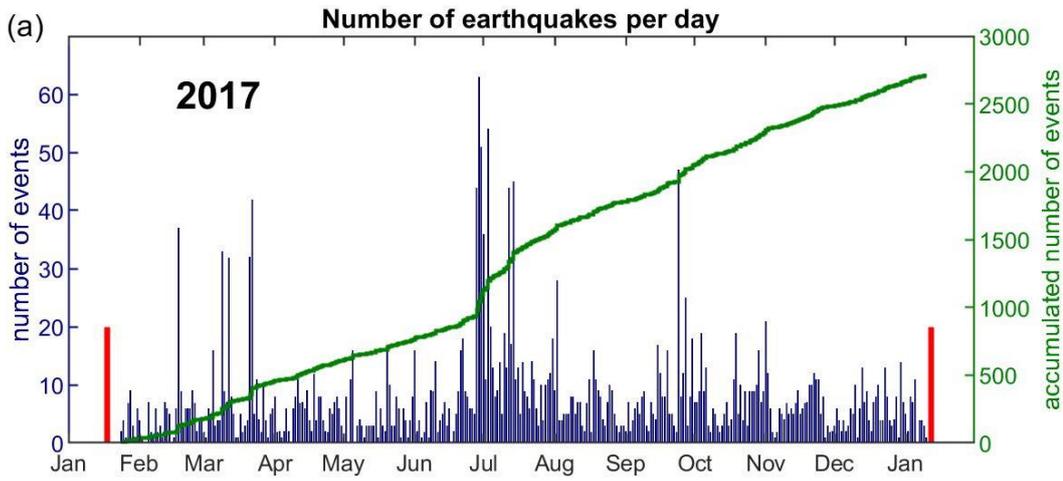
Figure 8: Comparison of a volcano-tectonic earthquake and a hybrid event. (a) Top: Traces of an earthquake beneath Fogo on 18
Nov 2017 (04:19 UTC; event location: 15.022 -24.349) recorded at the broad-band station AF02 of array AF on Fogo. Bottom:
spectrogram of the vertical component and overall frequency content (panel on the right). (b) Traces of a hybrid event recorded on
17 Aug 2017 (02:54 UTC; event location: 14.969 -24.400) at the same station. Traces are filtered between 1 and 50 Hz to remove
ocean-generated noise. Compared to the earthquake, the hybrid event shows no clear S-phase and more energy in the 1-10 Hz band
(of the coda).

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Figure 9: (a) Number of earthquakes per day. Green line: accumulated number of earthquakes. The recordings range from 18 January 2017 to 12 January 2018. (b) Number of hybrid events per day. Green line: accumulated number of hybrid events during the same time period.
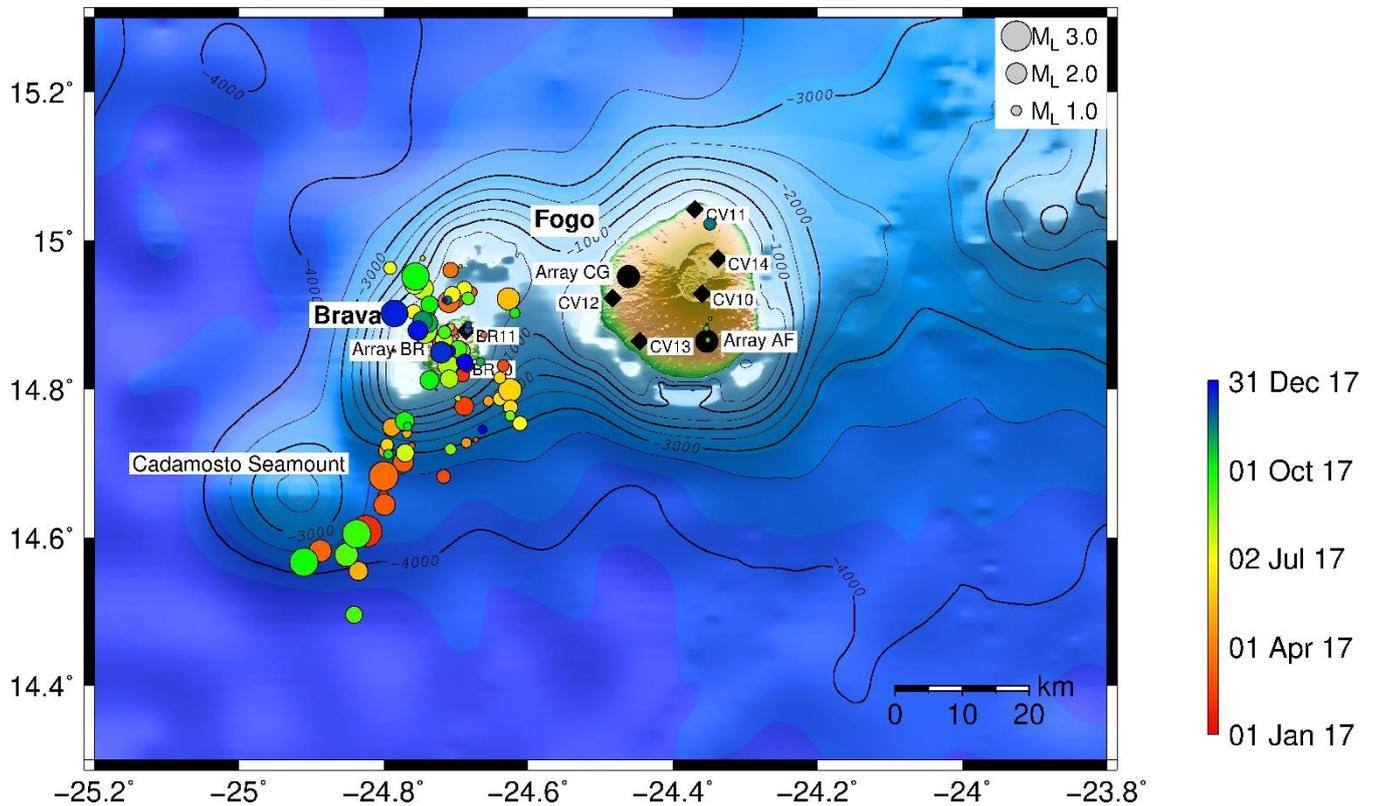
Figure 10: Earthquake locations from 18 Jan 2017 to 12 Jan 2018. Black circles: array locations; black diamonds: short-period single stations. Topography and bathymetry data are from Ryan et al. (2009).
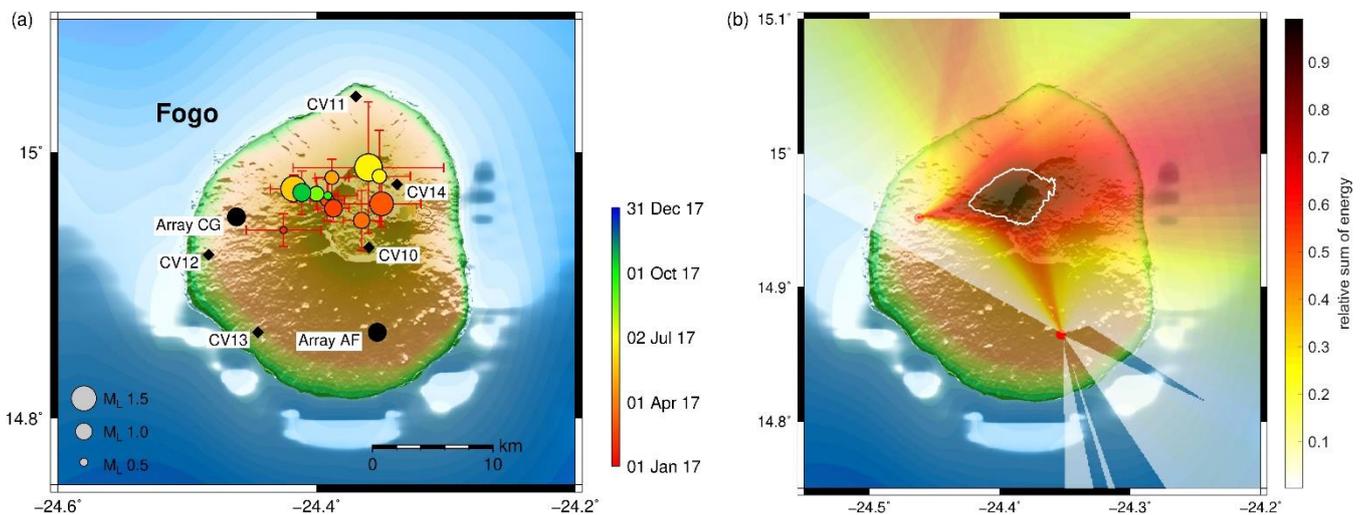
Figure 11: (a) Locations of hybrid events detected between 18 Jan 2017 and 12 Jan 2018 and located with the arrays on Fogo. Black circles: array locations; black diamonds: short-period single stations. (b) Superimposed beams of all the hybrid localizations. The

white line corresponds to 80% of the maximum of the relative sum of energy and indicates a region of high likelihood for the occurrence of hybrid events. Topography and bathymetry data are from Ryan et al. (2009).