Thermal history of the East European Platform margin in Poland based on apatite and zircon low temperature thermochronology

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Abstract. The Phanerozoic tectono-thermal evolution of the SW slope of the East European Platform (EEP) in Poland is reconstructed by means of thermal maturity, low temperature thermochronometry and thermal modelling. We provide a set of new thermochronometric data and integrate stratigraphic and thermal maturity information to constrain the burial and thermal history of sediments. Apatite fission track analysis (AFT) and zircon (U-Th)/He (ZHe) thermochronology have been carried out on samples of sandstones, bentonites, diabase and crystalline basement rocks collected from 17 boreholes located in central and NE Poland. They penetrated sedimentary cover of the EEP subdivided from the north to south into the Baltic, Podlasie and Lublin Basins. The average ZHe ages from Proterozoic basement rocks as well as Ordovician to Silurian bentonites and Cambrian to lower Carboniferous sandstones range from 848 ± 81 Ma to 255 ± 22 Ma with a single early Permian age of 288 Ma, corresponding to cooling after a thermal event. The remaining ZHe ages represent partial reset or source ages. The AFT ages of samples are dispersed in the range of 235.8 ± 17.3 Ma (Middle Triassic) to 42.1 ± 11.1 Ma (Paleogene) providing a record of Mesozoic and Cenozoic cooling. The highest frequency of the AFT ages is in the Jurassic and Early Cretaceous prior to Alpine basin inversion. Thermal maturity results are consistent with the SW-ward increase of the Palaeozoic and Mesozoic sediments thickness. An important break in a thermal maturity profile exists across the base Permian-Mesozoic unconformity. Thermal modelling showed that significant heating of Ediacaran to Carboniferous sedimentary successions occurred before the Permian with maximum paleotemperatures in the earliest and latest Carboniferous for Baltic-Podlasie and Lublin Basins, respectively. The results obtained suggest an important role of early Carboniferous uplift and exhumation at the SW margin of the EEP. The SW slope of the latter was afterward overridden in the Lublin Basin by the Variscan orogenic wedge. Its tectonic loading interrupted Carboniferous uplift and caused resumption of sedimentation in the late Viséan. Consequently, a thermal history of the Lublin Basin is different from that in the Podlasie and Baltic Basins, but similar to other sections of the Variscan foreland, characterised by maximum burial at the end of Carboniferous. The Mesozoic thermal history was characterised by gradual cooling from peak temperatures at the transition from Triassic to Jurassic due to decreasing heat flow. Burial caused maximum paleotemperatures in the SW part of the study area, where the EEP was covered by an extensive sedimentary pile. However, farther NE, due to low temperatures caused by shallow burial, the impact of fluids can be detected by vitrinite reflectance, illite/smectite and thermochronological data. Our new results emphasise the importance of using multiple low-

temperature thermochronometers and thermal modelling in connection with thermal maturity analysis to elucidate the near-surface evolution of platform margins.

1 Introduction

The boundary between the old and thick East European Platform (EEP) and the younger and thinner Palaeozoic Platform of Western Europe is entirely concealed beneath Palaeozoic and Mesozoic basins filled with several kilometres thick sedimentary successions (Figs. 1-4; Guterch et al., 1986, 1999; Pharaoh, 1999; Grad et al., 2002; Mazur et al., 2015). These basins form an extensive platform cover resting upon the SW slope of the East European Craton, comprising Paleoproterozoic to Mesoproterozoic crystalline basement (Krzemińska et al., 2017). The tectonic and thermal evolution of the area was punctuated by several successive phases of extension and shortening related to the Caledonian and Variscan orogenies, early Permian continental rifting and Late Cretaceous-Paleocene basin inversion (Krzywiec, 2002, 2009; Mazur et al., 2005). The timing and scale of these processes have not been yet properly constrained owing to the scarcity of thermochronological studies performed in the area and incomplete sedimentary record on the SW margin of the EEP that is due to the intervening periods of exhumation (Fig. 2). Although quantitative estimates of exhumation have been previously performed based on thermal maturity of organic matter, they provided divergent results due to the absence of thermochronological data (e.g., Botor et al., 2002; Poprawa et al., 2010). However, in the Scandinavian part of the EEP, the discordant zircon U-Pb ages, onlap of Palaeozoic sediments, geomorphological analyses and low-temperature thermochronology indicate that the Precambrian basement was buried beneath extensive Palaeozoic and Mesozoic sedimentary successions (Lidmar-Bergström, 1993, 1996; Hansen, 1995; Larson and Tullborg, 1998; Larson et al., 1999; Hendricks and Redfield, 2005; Söderlund et al., 2005; Green and Duddy, 2006; Larson et al., 2006; Hendriks et al., 2007; Japsen et al., 2016, 2018; Guenthner et al., 2017).

Thermal maturity, low temperature thermochronometry and thermal modelling are used in this study to reconstruct the Phanerozoic tectono-thermal evolution of the south-western margin of the East European Platform in Poland. We provide a set of new thermochronometric data and integrate stratigraphic and thermal maturity information to constrain the burial and thermal history of sediments. Using apatite fission-track (AFT) and zircon (U-Th)/He (ZHe) analyses, we reconstructed the burial and exhumation paths for different sedimentary successions along the SW margin of the EEP (Fig. 1). The low-temperature thermochronological methods applied are sensitive in the temperature range of c. 60-120°C (AFT) and c. 130-220 °C (ZHe; Reiners, 2005; Guenthner et al., 2013; Guenthner, 2020), allowing to investigate thermal history of rocks at shallow crustal levels (e.g., Farley, 2002; Armstrong, 2005; Reiners, 2005; Green and Duddy, 2012; Ketcham et al., 2007a; Guenthner et al., 2013; Wauschkuhn et al., 2015; Guenthner, 2020; Milesi et al., 2020). Therefore, these methods provide new constraints on the main subsidence and exhumation episodes in the Phanerozoic history of the region.
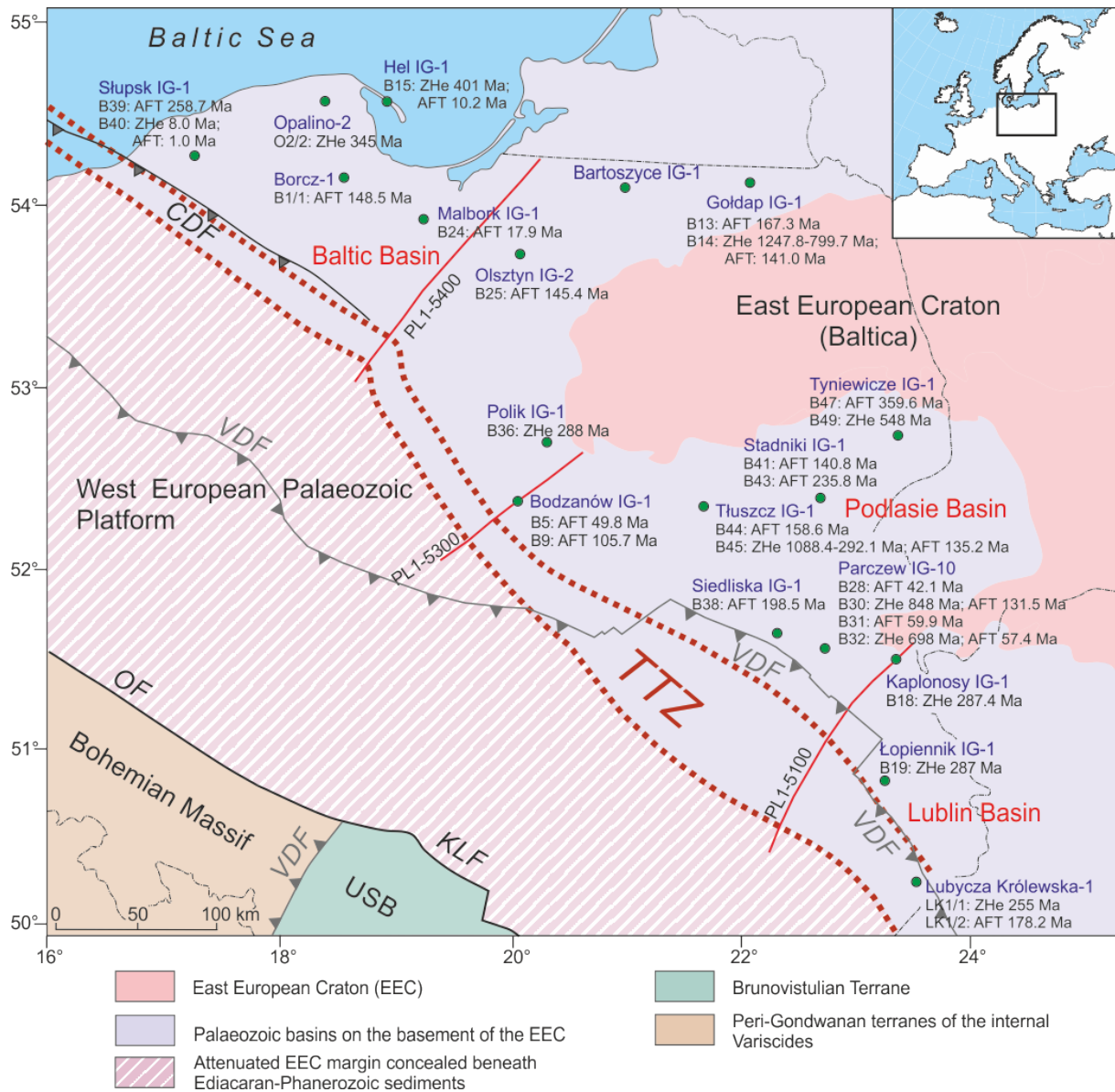
Figure 1. Simplified tectonic map of central and north-eastern Poland (modified from Mazur et al., 2018b) showing the location of borehole samples and the AFT and ZHe ages obtained. Location of PolandSPAN^{TM} seismic profiles (PL1-5100, 5300, 5400) used in Figure 2 is indicated. CDF - Caledonian Deformation Front; KLF - Kraków-Lubliniec Fault; OF - Odra Fault; VDF - Variscan Deformation Front; TTZ - Teisseyre-Tornquist Zone; USB - Upper Silesia Block.

2 Geological setting

The SW slope of the EEP is concealed beneath a thick cover of Ediacaran and Phanerozoic sediments overlying the Paleoproterozoic crystalline basement (Figs. 2, 3). The thickness of sediments gradually increases south-westwards with the most rapid depth-to-basement increase within the Teisseyre-Tornquist Zone (TTZ; Mikołajczak et al., 2019). The latter corresponds to a transition between the thick crust of the EEP and the thinner crust underlying the West European Palaeozoic Platform (Fig. 1) (Grad et al., 2002; Guterch et al., 2010; Mazur et al., 2015, 2018a). The TTZ represents a c. 50 km wide zone of the Moho uplift by c. 6 km and a coincident basement downward slope in the range of 9-11 km (Mazur et al., 2015; Mikołajczak et al., 2019). The depth to crystalline basement reaches 8-10 km in the NW and central segments of the TTZ and 14-16 km in its SE part (Mikołajczak et al., 2019). The basement is even deeper beneath the Mid-Polish Trough, an elongated Permian-Mesozoic depocentre, that adjoins the TTZ from the SW (Fig. 4; Dadlez et al., 1995). The trough was subsequently inverted into the Mid-Polish Swell during a Late Cretaceous to earliest Paleogene compressional event (Dadlez et al., 1995; Krzywiec, 2002).

The sedimentary cover deposited on the SW slope of the EEP comprises Ediacaran-early Palaeozoic, Devonian-Carboniferous, and Permian-Mesozoic successions. The lower Palaeozoic sequence has the largest thickness and lateral extent (Figs. 2, 3). Its lower portion was developed on a passive continental margin of Baltica, whereas the upper part represents the rapidly subsided Silurian foredeep of the Caledonian orogen (Poprawa et al., 1999; Mazur et al., 2018b; Poprawa, 2019). At the transition from the Silurian to Devonian, the north-western section of this fordeep was included in the distal part of the Caledonian orogenic wedge comprising a NW-SE-oriented belt of deformed Upper Ordovician and Silurian sedimentary rocks onshore and offshore NW Poland (e.g., Dadlez et al., 1994; Mazur et al., 2016). The fold belt is entirely concealed beneath a 1-4 km thick cover of largely undeformed younger sediments and documented based on borehole data. Toward the east, the Caledonian Deformation Front (CDF) delineates a sharp contact between the Caledonian fold belt and the undeformed sediments of the early Palaeozoic basin. This basin is almost completely buried beneath younger sediments with Permian-Mesozoic, Carboniferous and Devonian strata overlying the top of Silurian in NW, central and SE Poland, respectively (Figs. 2, 3). The top of Silurian represents an unconformity in NW and central Poland. In SE Poland, the unconformity is younger, and located at the top of Lochkovian, being formed due to a widespread late Caledonian compressional event (Narkiewicz, 2007; Krzywiec et al., 2017a; Poprawa et al., 2018; Mazur et al., 2018b). The initially uniform early Palaeozoic basin was subsequently separated by differential exhumation in three distinct parts: the Baltic, Podlasie and Lublin Basins, straddling the SW slope of the EEP from the NW to SE (Fig. 1).

The post-Caledonian succession overlying the EEP in SE Poland comprises Lower Devonian thick terrestrial clastics and Middle Devonian relatively thin (200 m) and laterally variable alluvial to open-shelf deposits (Narkiewicz, 2007). Localised subsidence during the Late Devonian led to formation of a 2.5 km thick sedimentary succession filling the NW-SE oriented Lublin Trough (Narkiewicz, 2007). This extensional event was terminated by significant uplift during the "Bretonnian" tectonic phase at the transition from the Devonian to Carboniferous that caused widespread exhumation in the range of 1000-2000 m (Poprawa, 2006; Krzywiec et al., 2017a). The lack of Devonian strata over the remaining part of the SW slope of the EEP in Poland is usually attributed to post-Variscan exhumation in the latest Carboniferous and early Permian (Żelichowski, 1987;
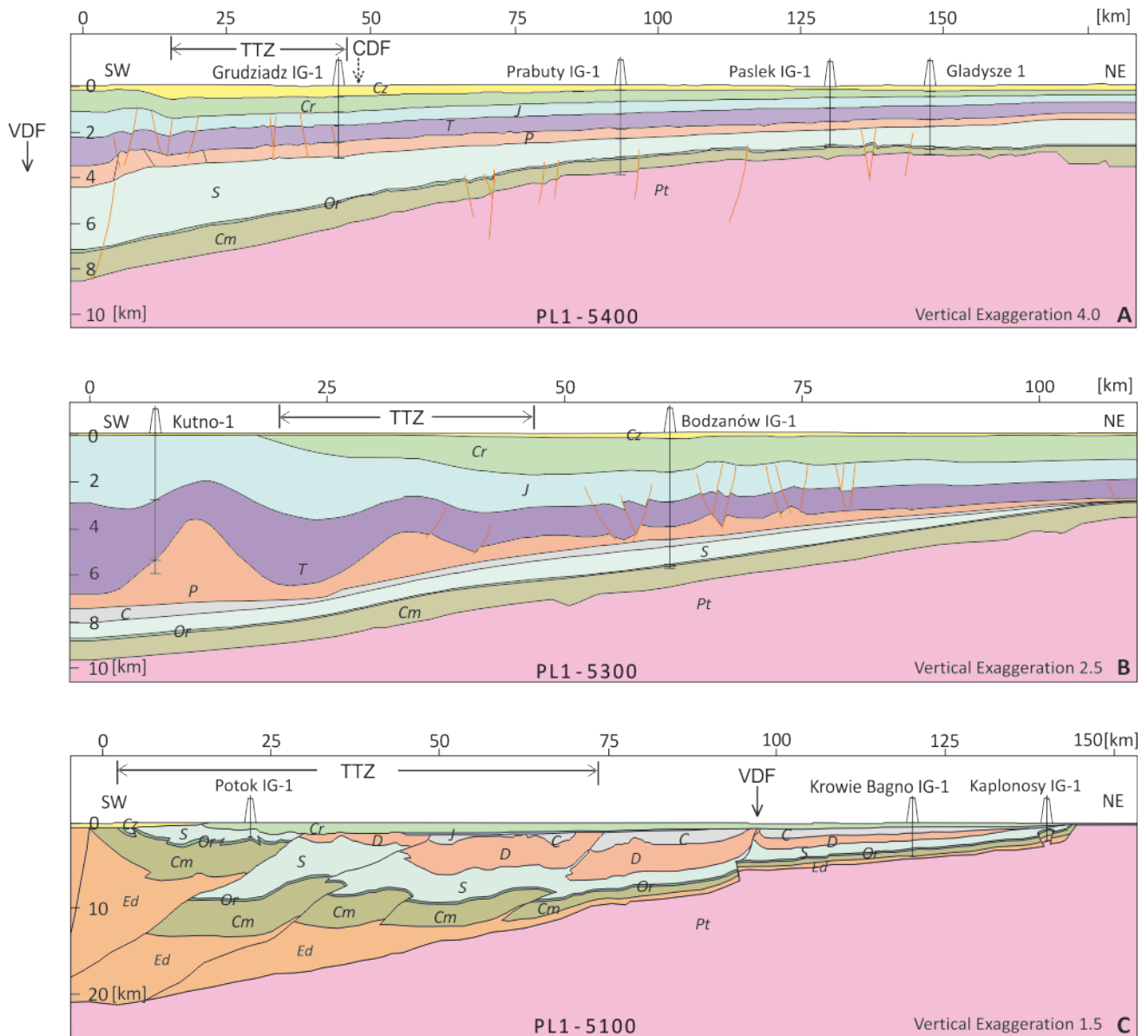
Figure 2. Vertically exaggerated geological sections across the south-western slope of the East European Platform, showing the structure of sedimentary cover and top of the crystalline basement. Cross-sections are built upon the PolandSPAN^{TM} seismic profiles PL1-5400 (A), PL1-5300 (B) and PL1-5100 (C) for the Baltic, Podlasie and Lublin Basins, respectively. Modified from Mazur et al. (2018a); (A) and Mazur et al. (2015); (B) and (C). Abbreviations of chronostratigraphic units: Pt - Paleoproterozoic; Ed - Ediacaran; Cm - Cambrian; Or - Ordovician; S - Silurian; D - Devonian; C - Carboniferous; P - Permian; T - Triassic; J - Jurassic; Cr - Cretaceous; Cz - Cenozoic.

Narkiewicz, 2007). Based on facies patterns, it was accordingly postulated that Devonian and early Carboniferous successions had originally extended over considerable stretches of the EEP slope (Żelichowski, 1987). However, the presence of upper
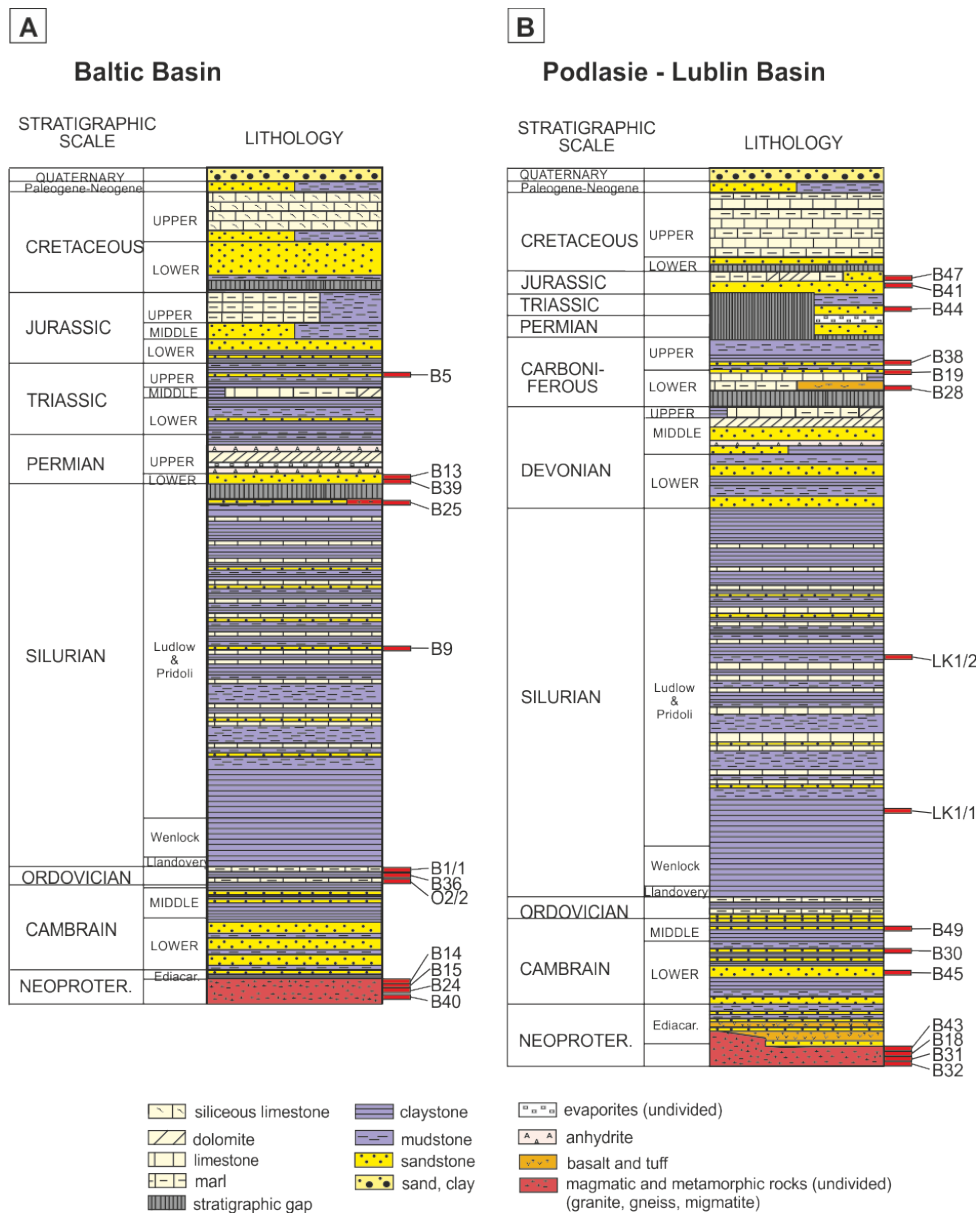
Figure 3. Simplified chronostratigraphic logs for the (A) Baltic Basin and (B) Podlasie-Lublin Basins the location of samples indicated (based on various sources including Żelichowski, 1987; Modliński et al., 2010; Poprawa, 2010)

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90 Carboniferous sediments directly overlying Silurian strata in central Poland points out to a role of early Carboniferous uplift and exhumation across the entire SW slope of the EEP. The top of Carboniferous represents a major unconformity throughout the

study area. Its origin is usually attributed to final stages of Variscan shortening (Mazur et al., 2010) although the unconformity is also developed over the areas of the EEP that are virtually unaffected by Variscan deformation.

During the Permian and Mesozoic, the study area was located in the marginal part of the Polish Basin that onlapped the SW slope of the EEP (Kutek and Głazek, 1972; Pożaryski and Brochwicz-Lewiński, 1978; Dadlez et al., 1995; Kutek, 2001). Up to a few kilometres of sediments were deposited at that time, and the thickness is rapidly decreasing towards the interior of the EEP (Świdrowska et al., 2008). Bulk of the Mesozoic infill in a distal, eastern part of the Polish Basin located beyond the TTZ consists of the Oxfordian-Kimmeridgian dominantly carbonate succession covered by the Albian transgressive clastics and thick Cenomanian-Maastrichtian pelagic carbonates (Fig. 3; Krzywiec, 2002, 2009). In contrast to the Mid-Polish Trough, where significant Late Cretaceous basin inversion took place (Botor et al., 2018; Łuszczak et al., 2020), the part of the Polish Basin extending over the TTZ and adjacent area of the EEP experienced only mild to little inversion (Krzywiec, 2009; Krzywiec et al., 2017a,b). This is probably related to the rather shallow depth to the rigid Precambrian basement. Cenozoic strata, forming the uppermost part of the sedimentary section, are composed of varied, poorly consolidated deposits attaining a maximum thickness of a few hundreds of meters.

3   Previous thermal history studies

The Ediacaran to Mesozoic strata on the SW slope of the EEP were investigated by reflectance of organic particles (including vitrinite), Rock-Eval analysis, Conodont Colour Alteration Index (CAI) and Thermal Alteration Index (Botor et al., 2019a,b, and references therein). These studies revealed a systematic zonation in the degree of diagenesis of Ediacaran to lower Palaeozoic rocks that is oriented NE-SW, parallel to the TTZ (Fig. 4; Kanev et al., 1994; Nehring-Lefeld et al., 1997; Swadowska and Sikorska, 1998; Grotek, 1999, 2006, 2016; Skręt and Fabiańska, 2009; Więcław et al., 2010, 2012). In all stratigraphic units, thermal maturity increases toward SW correspondingly to the increase of burial depth. Accordingly, mean random vitrinite reflectance (VR) for the base of Phanerozoic sequences in the Baltic, Podlasie and Lublin Basins attains 5.0, 1.3 and 3.4% VR, respectively (Grotek, 1999, 2005, 2006, 2016). Conodont CAI for Ordovician-Devonian rocks in the interior of the EEP attains 1-1.5 (Drygant, 1993; Nehring-Lefeld et al., 1997; Poprawa, 2010), the value that is indicative of paleotemperatures between 50 and 90 °C, according to the CAI calibration by Epstein et al. (1977). Toward the SW, the Ordovician CAI values gradually increase and reach level 5 (i.e., over 300 °C) within the TTZ (Drygant, 1993; Nehring-Lefeld et al., 1997). In the Baltic Basin, the analogous zonation was documented by the illite-smectite ratio in shales (Środoń et al., 2009). In the Ukrainian part of the EEP (Podole area, see further discussion in section 6.4), Silurian bentonite samples were examined by clay mineralogy, illite K-Ar and AFT dating (Środoń et al., 2013). Apatite samples yielded mid-Late Cretaceous FT ages (107-63 Ma), significantly younger than the stratigraphic ages (423-417 Ma; Środoń et al., 2013). Illite K-Ar data constrained the age of maximum paleotemperatures at 390-310 Ma (Middle Devonian-late Carboniferous), comparable to the range of ages obtained for bentonites from the Baltic Basin (382-294 Ma). Środoń et al. (2009, 2013) interpreted these data as indicative of deep burial under a Devonian-Carboniferous sedimentary cover. However, new illite K-Ar ages show an early Carboniferous age (350-325 Ma)
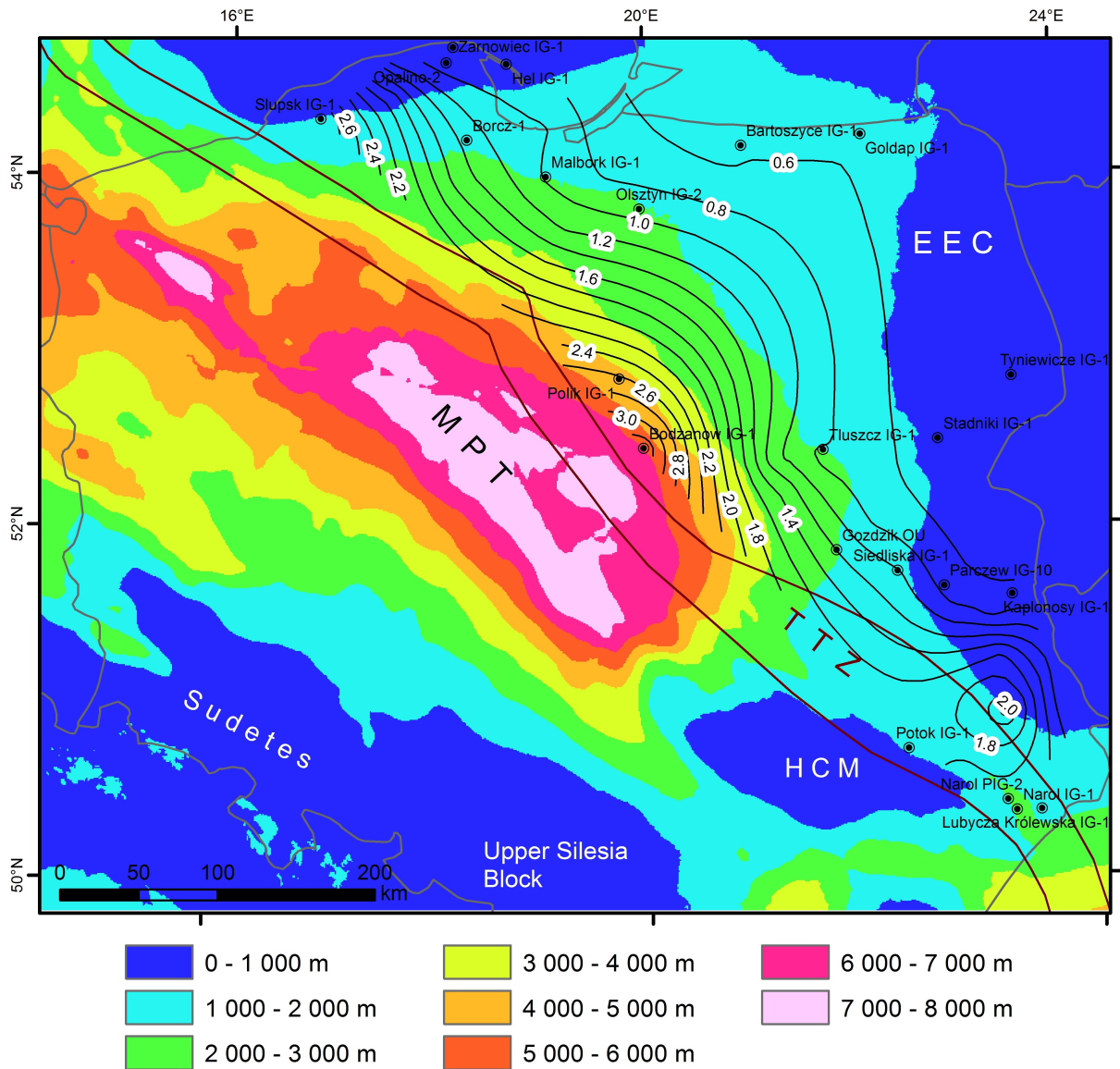
Figure 4. Permian-Cenozoic sediments isopach map for Poland adapted from Lamarche and Scheck-Wenderoth (2005); Maystrenko and Scheck-Wenderoth (2013); Mazur et al. (2021). Isolines show mean vitrinite-equivalent reflectance model for the base of Silurian in NE Poland modified from Papiernik et al. (2019). EEC - East European Craton; HCM - Holy Cross Mountains; MPT - Mid-Polish Trough; TTZ - Teisseyre-Tornquist Zone.

of maximum temperatures for most of the area, except for the SE part of the Lublin Basin, where K-Ar data revealed younger,
125   late Carboniferous to Permian ages for a thermal climax (Kowalska et al., 2019).

Permian and Mesozoic rocks in the Baltic Basin show much lower thermal maturity compared to the lower Palaeozoic sediments, usually below 0.6 % VR (Grotek, 1999, 2006; Poprawa et al., 2010). Farther SE, thermal maturity is variable in the Carboniferous and Devonian-Carboniferous strata of the Podlasie and Lublin Basins, respectively (Poprawa and Pacześnia, 2002; Grotek, 2005; Botor et al., 2002, 2019a,b). However, the Permian-Mesozoic sediments consistently show low thermal maturity below 0.5% VR in these basins (Grotek, 2005, 2006, 2016).

Maturity modelling, based on downhole VR profiles (Fig. 5) showed that the thermal evolution in the Lublin Basin was related to burial in Devonian-Carboniferous times with the most prominent temperature increase in the late Carboniferous (Botor et al., 2002; Karnkowski, 2003; Kosakowski et al., 2013; Botor, 2018; Botor et al., 2019a,b). A fluid flow influence on thermal maturity in this area was suggested by Poprawa and Żywiecki (2005) based on combined analysis of fluid inclusions and modelling of organic maturation. In the Baltic Basin, maturity modelling has shown that the increase of burial and temperature was continuous from the late Silurian to early mid-Carboniferous. Toward the west, more important was late Silurian burial, whereas in the east Devonian and early Carboniferous heating prevailed (Poprawa et al., 2010; Botor et al., 2019a). In Baltic Basin, an alternative possibility was proposed by Poprawa et al. (2010), who suggested that peak temperatures were related to burial and an associated advectional/convectional thermal event in the Late Cretaceous.

4 Samples and methods

Samples of sandstones, bentonites, diabase and crystalline basement rocks were collected from 17 boreholes located along the SW slope of the EEP (Fig. 1) and their geological details are provided in Table 1. Single samples were available from most of boreholes, apart from the Gołdap IG-1 and Tłuszcz IG-1, each of which was represented by two samples. Stratigraphic ages of the samples range from the Proterozoic to Middle Jurassic and their depth varies between 416 m and 5110 m below surface (Tab. 1). Corrected present-day temperatures in Table 1 are based on analysis of geophysical data available (Górecki et al., 2006a,b). Apatite and zircon crystals were separated from core samples using conventional crushing, sieving, magnetic and heavy liquids separation techniques. However, most of the processed 70 samples did not contain any apatite or zircon crystals suitable for thermochronology. Some core samples were too small to yield enough crystals, particularly apatites. Consequently, 12 samples were suitable for ZHe and 21 for AFT analyses.

4.1 Zircon (U-Th)/He analyses

Zircon crystals were handpicked following the recommendation of (Farley, 2002). Selected crystals were characterized by euhedral shape with two pyramidal terminations, and a width of >65 μm. The crystals were then photographed, measured for physical dimensions and loaded in Pt microtubes. The helium was extracted at 1000 °C under high vacuum using a diode laser and measured by isotope dilution using a Hidden 3F triple-filter mass spectrometer at Göttingen University. A "re-extract" was run after each sample to verify complete degassing of the crystals. Following degassing, samples were retrieved from the gas extraction line and spiked with calibrated ^{230}Th and ^{233}U solutions. Zircon crystals were dissolved in teflon bombs using a mixture of double distilled 48% HF and 65% HNO$_3$ at 220 °C for 5 days. Spiked solutions were analysed as 0.4 ml solutions
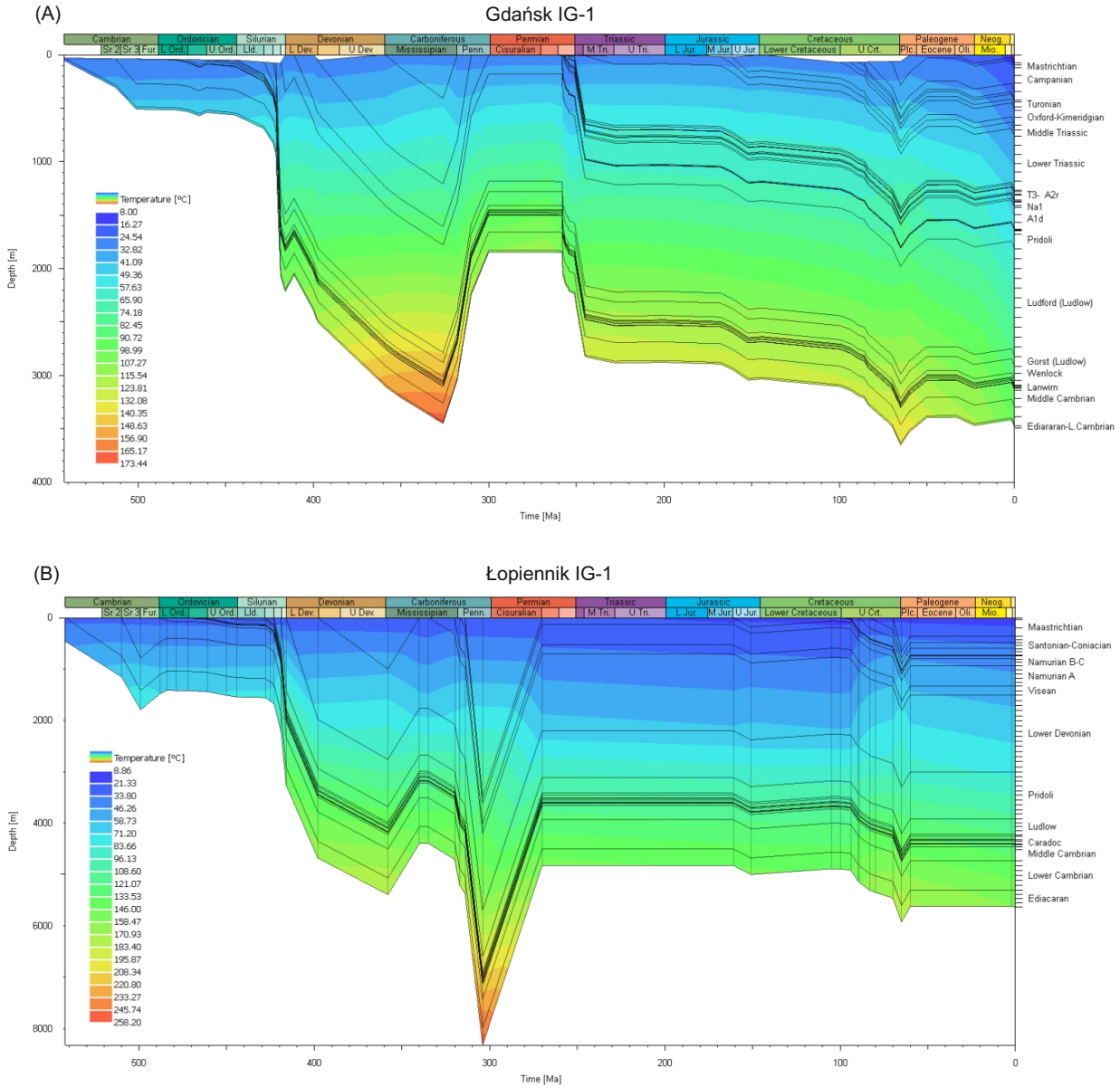
Figure 5. Representative burial history plots for (A) Baltic Basin, and (B) Lublin Basin after Botor et al. (2019a). In the Baltic and Lublin Basins, a maximum temperature in the lower part of the Ediacaran-Palaeozoic sedimentary succession was reached in the early and latest Carboniferous, respectively.

by isotope dilution on a Perkin Elmer Elan DRC II ICP-MS with an APEX micro-flow nebulizer. Procedural U and Th blanks by this method are usually very stable in a measurement session and below 1.5 pg. Sm, Pt, Zr and Ca were determined by
160 external calibration. The oxide formation rate and the PtAr - U interference was always monitored, but the effects of isobaric argides were negligible relatively to the signal of actinides. The ejection correction factors (F_t) were determined for the single crystals by a modified algorithm of Farley et al. (1996) using an in-house spread sheet.

4.2 Apatite fission-track analyses

Apatite grains were mounted in epoxy resin on glass slides and polished to expose internal grain surfaces. Spontaneous tracks
165 were revealed by 5N HNO$_3$ at 21 °C for 20 seconds. Neutron fluence was monitored using CN5 uranium dosimeter glass. Thin flakes of low-U muscovite were used as external detectors. Samples together with age standards (Fish Canyon, Durango, and Mount Dromedary apatite) were irradiated with a thermal neutron nominal flux of 9×10^{15}n/cm^2 at the Oregon State University TRIGA reactor (USA). After irradiation, the muscovite detectors were etched in 40 HF for 45 minutes at 20 °C to reveal the induced tracks. AFT analysis was carried out at the Institute of Geological Sciences, Polish Academy of Sciences
170 in Kraków (Poland), using the external detector method and the ζ age calibration method in order to determine the fission track ages (Hurford and Green, 1983; Hurford, 1990). Track counting and length measurement of horizontal confined fission tracks and the etch pit diameter (Dpar) values were carried out by means of optical microscopy at 1250× magnification using a Nikon Eclipse E-600, equipped with a motorized stage, digitizing tablet, and drawing tube controlled by program FTStage 4.04 (Dumitru, 1993). All quoted AFT ages are central ages with 1σ error (Galbraith and Laslett, 1993). The degree to which
175 individual AFT grain ages in a sample belong to a single population was assessed by a χ^2 statistic (Galbraith, 1981). Values of P(χ^2) below 5% indicate a statistically significant spread in single grain ages and the presence of more than one population (Galbraith, 1981, 1990; Galbraith and Laslett, 1993). Crystals chosen for confined track measurements had a well-polished surface, parallel to the c-axis. For each sample, as many confined track lengths as possible were measured (Gleadow et al., 1986). Data analyses and age calculations were based on a Zeta value ζCN5 of 348.18 ± 6.52 for CN5 accomplished by using
180 Trackkey 4.2 software (Dunkl, 2002). We used confined tracks non-corrected to the c-axis as normalizing fossil track lengths to a personalized zero-length is a questionable procedure for eliminating discrepancies between measured and predicted lengths (e.g., Green and Duddy, 2012).

4.3 Thermal modelling

The modelling of thermal history was performed using HeFTy software (Ketcham, 2005). The program requires input data
185 such as measured AFT age, track length distribution, kinetic parameter as Dpar for apatite and apparent ZHe age, radioactive elements concentration and diameter of the dated crystals to define "acceptable" time-temperature paths that pass statistical criteria and conform to a possible set of user-defined geological constraints (Ketcham, 2005). Thermal histories were modelled using the fission track annealing model (Ketcham et al., 2007b). Randomly generated thermal histories predict the AFT age and length parameters and compare them to the measured data. Modelling was run until the software calculated at least 100
190 acceptable fits. An acceptable fit corresponds to thermal histories representing the t-T paths that give a goodness of fit (GOF)

Table 1. Locality, stratigraphy, lithology, temperature, and thermal maturity of samples.

Sample	Well	Depth (m)	Stratigraphy	Lithology	VR (%)	VR$_{eqv}$ (%)	Max. pTemp. (°C)	Temp. (°C)
B1/1	Borcz-1	3707,0	Ordovician Upper	bentonite	1,15	1,00	135	65
B13	Gołdap IG-1	1090,0	Permian (Saxon)	sandstone	0,39		60	35
B14	Gołdap IG-1	1645,0	Proterozoic	granitoid	0,82	0,76	113	45
B15	Hel IG-1	3489,0	Proterozoic	gneiss	1,30	1,11	144	100
B18	Kaplonosy IG-1	1940,0	Proterozoic	gneiss	0,95	0,85	123	55
B19	Łopiennik IG-1	1509,8	Carboniferous lower	sandstone	1,24		153	50
B24	Malbork IG-1	3675,0	Proterozoic	granitoid	1,64	1,36	160	95
B25	Olsztyn IG-2	2176,0	Silurian upper	microsyenite	0,55	0,56	89	50
B28	Parczew IG-10	1047,0	Carboniferous lower	diabase	0,60		94	30
B30	Parczew IG-10	2129,0	Cambrian lower	sandstone	1,03	0,91	128	50
B31	Parczew IG-10	2304,0	Proterozoic	granitoid	1,46	1,23	152	75
B32	Parczew IG-10	2307,0	Proterozoic	migmatite	1,46	1,23	152	75
B36	Polik IG-1	4495,0	Ordovician Upper	bentonite	3,55	2,75	217	80
B38	Siedliska IG-1	1300,0	Carboniferous (Bashkirian)	sandstone	0,58		92	35
B39	Słupsk IG-1	1146,3	Permian (Rotliegend)	sandstone	0,54		86	40
B40	Słupsk IG-1	5110,4	Proterozoic	gneiss	4,92	3,75	242	145
B41	Stadniki IG-1	505,5	Jurassic Middle	sandstone	n.d.	n.d.	n.d.	22
B43	Stadniki IG-1	1560,0	Proterozoic	gneiss	n.d.	n.d.	120**	45
B44	Tłuszcz IG-1	1577,0	Triassic Lower	sandstone	0,56		89	36
B45	Tłuszcz IG-1	2482,0	Cambrian lower	sandstone	0,86	0,79	116	55
B47	Tyniewicze IG-1	416,0	Jurassic Middle	sandstone	n.d.	n.d.	n.d.	25
B49	Tyniewicze IG-1	523,0	Cambrian middle	sandstone	0,48	0,51	81	30
B5	Bodzanów IG-1	3448,0	Triassic Upper	sandstone	0,82		119	75
B9	Bodzanów IG-1	4969,0	Silurian upper	sandstone	3,80	2,93	222	100
LK1/1	Lubycza Krolewska-1	2790,8	Silurian upper	bentonite	2,28	1,82	184	72
LK1/2	Lubycza Krolewska-1	2455,8	Silurian upper	bentonite	2,05	1,66	176	64
O2/2	Opalino-2	2911,0	Ordovician Upper	bentonite	1,02	0,90	127	60

VR - in the adjacent depth level in which VR was measured in shale (Nehring-Lefeld et al., 1997; Grotek, 1999, 2005, 2006, 2016). *Values for the magmatic/metamorphic rocks are from adjacent above sedimentary rock. Vreqv - the relationship between graptolite reflectance and equivalent vitrinite reflectance can be expressed by the equation: VReqv = 0.73 R(graptolite + vitrinite-like)low + 0.16 in the Ediacaran-Silurian sediments (Petersen et al., 2013). Max. pTemp. - maximum paleotemperature calculated applying (Barker and Pawlewicz, 2005) equation where Max. pTemp.=(LN(VR)+1,68)/0,0124. 120** °C in Stadniki IG-1 from illite/smectite data of Lower Cambrian (Kowalska et al., 2019). Temp. - present-day average temperature of sample in the given borehole, n.d. - no data.

value greater than 0.05 for both the age and the length distribution (Ketcham, 2005). The minimum statistic value above 0.05 means that all statistics pass the 95% confidence test. When the minimum is above 0.5, the statistical precision limit, the model is termed good. For a comprehensive overview of fission-track methods and their modelling techniques, see more details in Donelick (2005); Ketcham (2005) and Ketcham et al. (2007b). For the helium diffusion kinetics in zircon, we applied the model by Guenthner et al. (2013). Chronostratigraphic subdivisions and absolute ages refer to Gradstein et al. (2012) throughout the text.

5 Results

5.1 Zircon (U-Th)/He ages

The stratigraphic ages of the samples are from the Proterozoic to Triassic (Tab. 1). Single grain ZHe ages are widely dispersed being in some cases older than a stratigraphic age (Tab. 2). These grains represent zircons that avoided resetting or are only

partially reset. Many others single grain ages, although younger than a stratigraphic age, are scattered broader than their individual uncertainties (Tab. 2). This effect must be also related to partial reset of studied crystals or secondary disturbance of the isotopic system. No mean ages are calculated for samples, where the dispersion of single grain ages exceeds estimated uncertainties (Tab. 2). Two versions of a mean age are calculated for sample B19: (1) excluding one single grain age that is older than a stratigraphic age, and (2) additionally excluding the youngest single grain age that is far beyond the uncertainties of the remaining grains (Tab. 2). The average ZHe ages range from 848 ± 81 Ma to 255 ± 22 Ma, except for sample B40 in which a measured ZHe is 8 Ma (Tab. 2, Fig. 1). However, this young ZHe age comes from a borehole where the present-day temperature is c. 145°C (Tab. 1), thus practically from the partial retention thermal conditions (ZPRZ, c. 130-220°C; Reiners, 2005).

Certain samples do not show a significant Phanerozoic reset (Tab. 2, Fig. 1) and e.g., the B14 Proterozoic granitoid sample has 3 single-grain ZHe ages which scatter from 1248 to 800 Ma. The B32 sample of Proterozoic migmatite has 3 single grain ZHe ages which scatter from 804 to 596 Ma and a mean ZHe age of 698 Ma. The B30 sample of lower Cambrian sandstone also has 3 single grain ZHe ages which are in the range of 905 to 791 Ma and a mean ZHe age of 848 Ma. Two Cambrian samples suggest partial reset of ZHe ages in the Phanerozoic due to some single grain ages younger than a depositional age. In the B49 sample of mid-Cambrian sandstone, a mean ZHe age is 548 Ma with single grain ages of 654, 509 and 481 Ma. The B45 sample of lower Cambrian sandstone has 4 single-grain ZHe ages which scatter from 1088 to 292 Ma. Two zircon have an age younger than a stratigraphic age of the sample. These results are consistent with the low thermal maturity of lower Palaeozoic strata in the NE part of the EEP and negligible or very low re-heating during Phanerozoic (Grotek, 1999, 2006, 2016; Skręt and Fabiańska, 2009; Pletsch et al., 2010).

All other ZHe mean ages are significantly younger than the stratigraphic age and thus are considered partially or totally reset (Tab. 2, Fig. 1). These samples occur along the SW slope of the EEP relatively close to the TTZ. In the Baltic Basin, sample B-15 (Hel IG-1 borehole) of Proterozoic gneiss reveals 3 single grain ages of 460, 378 and 365 Ma that suggest partial reset and yield a mean ZHe age of 401 Ma. A second sample from the Baltic Basin, the O2/2 sample of Upper Ordovician bentonite (Opalino-2) has an early Carboniferous mean ZHe age of 345 Ma. In the Polik-Bodzanów area (west of Podlasie Basin), the B36 Ordovician bentonite sample (Polik IG-1) has a tightly clustered early Permian mean ZHe age of 288 Ma. In the Lublin Basin, the B19 sample (lower Carboniferous sandstone, Łopiennik IG-1) has ZHe single grain ages from 562 to 187 Ma suggesting partial reset. In the B18 (Kaplonosy IG-1) sample only one of 4 zircon crystals are euhedral with no inclusions, showing a ZHe age of 287 Ma (early Permian). Sample LK1/1 (Silurian bentonite, Lubycza Królewska-1) has a late Permian mean ZHe age of 255 Ma. A wide scatter of Palaeozoic ZHe ages between Devonian and Permian results from the mixture of partially and fully reset zircon grains. Therefore, the ZHe mean ages probably do not reflect the timing of a particular tectonic event. The early Permian ages (B36, LK1/1 and possibly B19) represent the delayed effect of cooling after Carboniferous uplift and erosion. An alternative explanation would be cooling after the earliest Permian phase of continental rifting and possibly associated increased heat flow. Sample O2/2 might be with reservation interpreted as directly reflecting the time of cooling after early Carboniferous exhumation. The results obtained do not bear any record of Mesozoic tectonic events that may have caused thermal effects below the resolution of the ZHe method.

Table 2. Zircon (U-Th)/He data.

Sample	Stratigraphic age [Ma]	He vol. [ncc]	1s [%]	U mass [ng]	1s [%]	conc. [ppm]	Th mass [ng]	1s [%]	conc. [ppm]	Th/U ratio	Sm mass [ng]	1s [%]	conc. [ppm]	Ejection correct. (Ft)	*Uncorr.* He-age [Ma]	Ft-Corr. He-age [Ma]	2s [Ma]	Stratig. Age [Ma]	ZHe unweighted aver. ± 1 s.e. [Ma]	[Ma]	Zircon crystal quality	eU [ppm]
B14 z1		150,35	1,3	1,426	1,8	155	1,588	2,4	173	1,11	0,140	8,0	15	0,82	655,5	799,7	53,4				eu, i	196
B14 z2	Proterozoic	156,46	1,3	0,991	1,8	142	1,306	2,4	187	1,32	0,042	10,4	6	0,74	927,0	1247,8	107,9				eu, i	186
B14 z3		89,52	1,3	0,616	1,8	70	0,748	2,4	85	1,21	0,013	18,5	1	0,72	873,1	1216,7	113,6	>540	n.d. for aver.		eu	90
B15 z1		51,63	1,3	0,868	1,8	142	2,764	2,4	454	3,18	0,096	9,2	16	0,76	275,7	365,2	30,3				eu clean	249
B15 z2	Proterozoic	88,52	1,3	0,993	1,8	168	5,004	2,4	845	5,04	0,081	9,6	14	0,72	330,1	460,4	43,3				eu, i	366
B15 z3		55,02	1,3	0,847	1,8	130	3,382	2,4	521	3,99	0,044	11,4	7	0,72	271,8	378,4	35,3	>540	401	30	eu, i	253
B18 z2		67,46	1,6	2,379	1,8	323	1,637	2,4	222	0,69	0,094	9,2	13	0,69	198,7	287,4	29,5				eu	375
B18 z3	Proterozoic	17,18	1,6	1,857	1,8	294	1,839	2,4	291	0,99	0,160	8,0	25	0,69	61,8	89,1	9,1				eu frac	363
B-18 z4		35,742	1,7	1,501	1,8	676,9	0,683	2,4	308	0,45	0,073	9,3	33	0,71	175,4	248,1	24,6				eu, si,	749
B-18 z5		26,174	1,6	1,634	1,8	711,8	0,418	2,4	182	0,26	0,034	18,2	15	0,74	123,8	167,1	15,1	>540	n.d. for aver.		eu, si,	755
B19 z1		30,27	1,6	0,916	1,8	115	0,564	2,4	71	0,62	0,040	10,2	5	0,80	234,2	291,3	21,6				eu, si	132
B19 z2	Carboniferous lower	161,46	1,6	2,470	1,8	257	1,082	2,4	113	0,44	0,096	8,6	10	0,84	470,9	562,0	37,4				elips	284
B19 z3		94,18	1,6	2,196	1,8	241	1,074	2,4	118	0,49	0,028	13,3	3	0,84	310,1	368,1	24,1				eu, i	268
B-19 z4		27,657	1,6	0,733	1,8	118,8	0,910	2,4	147	1,24	0,021	21,6	3	0,79	237,0	301,0	23,3		*320	24	eu, si, fl	153
B-19 z5		12,232	1,7	0,472	1,9	33,7	0,695	2,4	50	1,47	0,117	9,2	8	0,84	157,0	187,1	12,3	340	**287	43	big, eu, si, f	45
B30 z2	Cambrian lower	132,33	1,6	1,132	1,8	95	1,220	2,4	103	1,08	0,088	8,3	7	0,80	727,4	905,2	66,3				eu, clean	119
B30 z3		108,93	1,6	1,119	1,8	186	1,144	2,4	190	1,02	0,060	10,1	10	0,78	617,4	791,3	62,6	530	848	81	eu, clean	230
B32 z1		110,94	1,6	1,260	1,8	203	0,599	2,4	96	0,48	0,033	11,6	5	0,77	621,3	803,8	65,7				eu, i, frac	225
B32 z2	Proterozoic	79,62	1,6	1,250	1,8	280	0,457	2,4	102	0,37	0,017	13,1	4	0,78	466,2	596,1	47,5				eu, i, frac	304
B32 z3		68,85	1,6	0,941	1,8	228	0,379	2,4	92	0,40	0,028	12,2	7	0,76	528,2	694,9	59,2	>540	698	60	eu, i, frac	250
B36 z1		11,35	1,6	0,473	1,8	79	0,238	2,4	40	0,50	0,013	15,4	2	0,63	175,0	277,7	33,3				eu, si	88
B36 z2	Ordovician Upper	15,54	1,6	0,548	1,8	82	0,443	2,4	66	0,81	0,022	13,4	3	0,65	193,9	299,9	34,5				eu, i, irr	98
B36 z3		5,89	1,6	0,235	1,9	48	0,122	2,4	25	0,52	0,013	14,6	3	0,63	182,1	287,6	34,4	450	288	6	eu	54
B40 z1		0,84	1,5	1,405	1,8	420	0,524	2,4	157	0,37	0,024	8,5	7	0,72	4,6	6,4	0,6				eu, few i	457
B40 z2	Proterozoic	0,32	1,7	1,065	1,8	376	0,206	2,4	73	0,19	0,008	12,6	3	0,70	2,4	3,4	0,4				eu, few i	393
B40 z3		1,38	1,4	1,205	1,8	323	0,203	2,4	54	0,17	0,007	14,5	2	0,68	9,1	13,3	1,4	>540	8	3	eu, few i	336
B45 z1		29,89	1,6	0,523	1,8	137	0,413	2,4	108	0,79	0,020	14,5	5	0,74	386,3	524,0	47,4				eu, i, frac	162
B45 z2	Cambrian lower	28,20	1,6	0,962	1,8	122	0,262	2,4	33	0,27	0,011	18,8	1	0,77	223,7	292,1	24,5				eu, i	129
B45 z3		113,58	1,6	1,078	1,8	145	0,284	2,4	38	0,26	0,058	9,8	8	0,70	766,9	1088,4	108,7				eu, i	155
B-45 z4		18,921	1,6	0,279	2,0	141,9	0,163	2,4	83	0,59	0,002	69,8	1	0,70	474,7	678,7	69,0				eu, si	161
B-45 z5		26,207	1,6	0,447	1,9	205,6	0,448	2,4	206	1,00	0,073	10,7	34	0,70	380,5	544,5	54,9	530	n.d. for aver.		eu, si,	254
B49 z1		24,02	1,6	0,486	1,8	127	0,204	2,4	53	0,42	0,041	10,1	11	0,75	361,1	481,4	42,3				eu, si	140
B49 z2	Cambrian middle	35,98	1,6	0,548	1,8	148	0,270	2,4	73	0,49	0,039	11,4	10	0,72	467,5	653,7	63,2				eu, i	165
B49 z3		33,19	1,6	0,526	1,8	87	0,612	2,4	101	1,16	0,111	9,2	18	0,78	396,9	509,1	40,4	520	548	53	elips	111
LK1/1 z1		15,66	1,4	0,679	1,8	195	0,428	2,4	123	0,63	0,023	8,5	7	0,75	163,9	219,1	19,0				eu, i	224
LK1/1 z2	Silurian upper	20,46	1,4	0,687	1,8	133	0,391	2,4	76	0,57	0,028	7,9	5	0,73	213,5	293,6	27,0				eu, i	151
LK1/1 z3		26,77	1,3	0,962	1,8	300	0,788	2,4	245	0,82	0,040	7,2	12	0,75	190,0	253,2	21,5	425	255	22	eu, i	358
02/2 z1		38,09	1,3	1,073	1,8	133	0,452	2,4	56	0,42	0,026	12,1	3	0,73	261,4	356,2	32,0				eu	146
02/2 z2	Ordovician Upper	14,21	1,6	0,466	1,8	190	0,293	2,4	119	0,63	0,019	16,2	8	0,72	216,0	298,3	28,1				eu	218
02/2 z3		42,86	1,6	1,154	1,8	319	0,346	2,4	96	0,30	0,026	14,4	7	0,73	280,2	381,7	35,1	450	345	25	eu	342

Amount of helium is given in nano-cubic-cm in standard temperature and pressure. Amounts of radioactive elements are given in nanograms. Ejection correct. (Ft) - correction factor for alpha-ejection according to Farley et al. (1996) and Hourigan et al. (2005). Uncertainties of helium and the radioactive element contents are given as 1 sigma, in relative error %. Uncertainty of the single grain age is given as 2 sigma in Ma and it includes both the analytical uncertainty and the estimated uncertainty of the Ft. Uncertainty of the sample average ages are in 1 standard error, as (SD)/(n)1/2; where SD = standard deviation of the age replicates and n = number of age determinations; eU has been calculated with the formula $eU = U + 0.235 \cdot Th$; eu - euhedral, i - inclusions, si - small inclusions, frac - fracture, elips - eliptic shape of grain , fi fluid inclusions. *320 ZHe age without B19z2, and B19z5 (562 Ma and 187.1 Ma). **287 ZHe age without B19z2 (562 Ma).

5.2 Apatite fission track data

The results of AFT analyses, performed on 21 samples, are presented in Table 3, Figure 1 and Appendix (Fig. A1). A stratigraphic age of samples ranges from Proterozoic to Jurassic (Tab. 1). The quality of AFT data is varied, but in most samples as a maximum 20 apatite grains were studied (Tab. 3). The average uranium content is from 3 to 70 ppm. All the AFT ages represent unimodal age populations as shown by high P (χ^2), younger than their stratigraphic age except for sample B47 (Tab. 3, Fig. 1). Central AFT ages range from 359.6 ± 27.2 to 1.0 ± 0.7 Ma. Except the samples characterised by high present-day temperature, related to their depth in boreholes (Tab. 1) and having impact on apatite tracks annealing (B15, B24, B40), the central AFT ages range from 359.6 ± 27.2 to 42.1 ± 11.1. The AFT age of sample B47 (359.6 ± 27.2) is inherited from a source of detritus since it is much older than a stratigraphic age of the sample (Middle Jurassic; Tab. 1). Therefore, a scatter of ages interpreted as partially or fully reset is between 258.7 ± 32.1 (latest Permian) and 42.1 ± 11.1 (Paleogene; Tab. 3). Some caution must be attached to 4 samples yielding Paleogene ages. Three of them (B5, B31 and B32) are from boreholes characterised by a present-day temperature of 75 °C, i.e., within the apatite partial annealing zone. For sample B28, having an AFT age of 42.1 ± 11.1 Ma, no track length measurements are available whereas its central age is only based on 7 grains with some scatter.

Horizontal confined track lengths have been measured in 16 samples. Due to the relatively low uranium content only eight samples yielded the track numbers >50 (Tab. 3). A mean track length (MTL) ranges from 10.2 ± 0.4 to 13.4 ± 0.5 μm (Tab. 3) and distributions are from wide to relatively narrow (Fig. A1). Most samples show a negative skewness between -0.1 and -1.6 with tails toward shorter track lengths, apart from 3 samples (B25, B39, and B43). The standard deviation values of track length range from 1.3 to 2.0 μm except for sample B1/1 (SD =0.8; Tab. 3).

A total of 1150 Dpar values were determined for all apatite grains used in this study (Tab. 3). The mean Dpar values of the samples are in the range of 1.8 ± 0.3 to 2.7 ± 0.4 μm (Tab. 3) and exhibit skewness in the range between 1.16 and -0.45 except for sample B44 (-2.86).

5.3 Thermal maturity

Thermal maturity data put key constraints on the interpretation of thermochronological results and ensuing thermal modelling. Besides a countrywide compilation of thermal maturity data (Fig. 4), we discuss in detail profiles of key boreholes that penetrate the sedimentary cover of the East European Platform. We also refer to thermochronological results obtained on samples that were collected from these boreholes. Corrected present-day temperatures in Figures 6 – 12 are based on analysis of geophysical data available (Górecki et al., 2006a,b). Although, a number of AFT and ZHe samples per well with is usually low (in many wells just a single sample), but integration of thermochronological and thermal maturity data provides meaningful results.

Thermal evolution recognised in some boreholes reveals features that are usually considered an effect of fluid flow overprinted on burial diagenesis. The possible influence of advective heat transfer is inferred based on a set of characteristic features that were documented from other sedimentary basins (Ziagos and Blackwell, 1986; Middleton et al., 1994; Lampe et al., 2001; Green and Duddy, 2012): (1) palaeotemperatures being much higher than predicted from the burial history under conditions of

Table 3. Apatite fission track data.

Sample	Nc	Dosimeter		Spontaneous		Induced		P (χ²)	U	Central age	MTL ± 1 s.e.	SD	No*	Dpar mean
		ρd	Nd	ρs	Ns	ρi	Ni	[%]	[ppm]	[Ma] ± 1σ	(µm)	(µm)		(µm)
B1/1	3	1.039	3120	1,813	49	2,183	59	98.5	25.6	148.5 ± 29.0	11.01 ± 0.66	0.9	2	2.3 (12)
B5	20	1.015	3049	0,399	112	1,409	396	91.5	19.0	49.8 ± 5.5	11.39 ± 0.64	1.8	8	2.2 (78)
B9	5	1.031	3095	0,469	38	0,789	64	98.7	8.7	105.7 ± 21.8	N.D.	N.D.	N.D.	N.D.
B13	20	1.025	3079	1,648	569	1,735	599	99.4	24.3	167.3 ± 10.7	13.06 ± 0.15	1.3	72	2.6 (83)
B14	20	1.035	3109	0,654	227	0,827	287	100.0	10.1	141.0 ± 13.0	11.98 ± 0.21	1.5	49	2.6 (80)
B15	20	1.024	3075	0,304	156	5,328	2736	80.0	69.8	10.2 ± 0.9	11.28 ± 0.29	1.9	42	2.7 (79)
B24	20	1.032	3100	0,431	144	4,321	1443	100.0	50.4	17.9 ± 1.6	10.80 ± 0.19	1.6	65	2.6 (80)
B25	20	1.047	3144	0,266	142	0,329	176	100.0	3.9	145.4 ± 16.8	13.42 ± 0.51	1.5	9	2.7 (83)
B28	7	1.039	3119	0,067	18	0,285	77	74.7	3.4	42.1 ± 11.1	N.D.	N.D.	N.D.	N.D.
B30	19	1.048	3153	0,447	177	0,613	243	100.0	8.2	131.5 ± 13.4	12.05 ± 0.37	1.4	14	2.4 (76)
B31	20	1.018	3058	0,697	318	2,053	937	91.9	25.3	59.9 ± 4.2	10.94 ± 0.25	2.0	67	2.6 (80)
B32	20	1.022	3070	0,929	586	2,869	1809	91.5	34.9	57.4 ± 3.1	11.51 ± 0.19	1.5	61	2.6 (81)
B38	20	1.027	3083	1,485	504	1,317	447	100.0	17.5	198.5 ± 13.9	11.64 ± 0.2	1.7	70	2.3 (80)
B39	14	1.035	3106	1,108	167	0,756	114	98.9	10.1	258.7 ± 32.1	10.23 ± 0.40	1.8	23	1.8 (54)
B40	20	1.019	3062	0,007	2	1,336	361	99.7	17.7	1.0 ± 0.7	N.D.	N.D.	N.D.	N.D.
B41	2	1.033	3101	1,265	19	1,598	24	97.7	19.0	140.8 ± 43.4	N.D.	N.D.	N.D.	N.D.
B43	20	1.031	3096	0,733	495	0,548	370	99.8	6.8	235.8 ±17.3	12.34 ± 1.01	1.4	73	2.4 (80)
B44	20	1.029	3092	1,507	449	1,682	501	85.7	21.2	158.6 ± 11.1	12.69 ± 0.17	1.3	59	2.6 (80)
B45	11	1.041	3127	1,113	141	1,476	187	100.0	18.5	135.2 ± 15.5	11.15 ± 0.40	1.9	21	2.5 (44)
B47	20	1.017	3093	1,847	610	0,884	292	95.2	11.2	359.6 ± 27.2	13.13 ± 0.20	1.5	56	2.3 (80)
LK1/2	3	1.038	3116	0,212	14	0,212	14	100.0	2.6	178.2 ± 67.5	N.D.	N.D.	N.D.	N.D.

Nc - number of measured crystals, ρs - density of spontaneous tracks (x 10^6 tracks for cm^{-2}); Ns - number of counted spontaneous tracks; ρi - density of induced tracks in external detector (mica) (x 10^6 tracks for cm^{-2}); Ni - number of counted induced tracks; ρd - density of induced tracks in external detector which cover dosimeter (glass CN5)(x 10^6 tracks for cm^{-}2); Nd - numbers of counted tracks. P (χ^2) [%] - probability homogeneity apatite population were showed by the test agreement χ^2 value (Galbraith, 1981). AFT age± 1 σ error is a central age of sample (Galbraith and Laslett, 1993) counted using calibration method zeta (Hurford 1983, 1990) and dosimeter (glass) CN5. Data analyses and age calculations were based on a Zeta value is 348.18 ± 6.52 (A. Anczkiewicz), and accomplished by using Trackkey 4.2 software (Dunkl, 2002). Crystals chosen for confined track measurements had a well-polished surface, parallel to the c-axis. U - amount of uranium was calculated by software Trackkey 4.2. MTL (µ m± SE) - mean confined length. No - number of measured confined tracks. SD - standard deviation. Dpar mean - mean etch pit diameter, in brackets are number of etch pit diameters measured.

vertical conductive heat transfer, (2) palaeotemperature profiles that fluctuate markedly, suggesting fluid-driven heat transfer along certain horizons and (3) discrepancies between palaeotemperatures derived from fluid inclusions, illite/smectite studies and kinetically dependent thermal maturity indicators such as vitrinite reflectance and apatite fission track analysis. In this study, such characteristics is documented and discussed in the following boreholes: Gołdap IG-1, Bartoszyce IG-1, Tłuszcz IG-1, and Siedliska IG-1. In some other wells as Opalino-2 and Lubycza Królewska-1 fluid flow influence on thermal evolution is also possible. Particularly, fluid flow heat transfer can explain high exhumation estimates inferred in some wells.

5.3.1 Opalino-2

Sample O2/2 of Ordovician bentonite from the Opalino-2 borehole gave single ZHe ages of 381.7, 356.2 and 298.3, Ma and an average ZHe age of 345 Ma (Fig. 6a, Tabs. 1–2). The equivalent-VR value for the O2/2 sample is 0.9% that can be re-calculated into a maximum paleotemperature of c. 127°C (Tab. 1). This estimate is close to the lower sensitivity range of the

zircon (U-Th)/He method and may explain only partial reset of the ZHe ages. The illite K-Ar age from the bentonite horizon is 336-329 Ma (Kowalska et al., 2019) that might be close to the timing of maximum paleotemperature. The VR profile is too short and scattered to be suitable for calculation of VR gradient. Therefore, paleogeothermal gradient of 20°C/km was assumed for the Opalino-2 well that is equivalent to the paleogradient estimated for the Żarnowiec IG-1 well. In several adjacent wells, the discontinuity of VR profile between Silurian and Permian (base Permian unconformity) suggests the pre-Permian development of thermal maturity (Fig. 6b; Botor et al., 2019a). In the nearby Żarnowiec IG-1 borehole (Fig. 4), a paleogeothermal gradient during maximum burial (20°C/km) seems to be similar to a present-day value (18°C/km) as both are almost parallel (Fig. 6) after omitting effects of pressure retardation. A paleogeotermal gradient of 20°C/km applied to the Silurian-early Carboniferous succession in the Opalino-2 borehole results in c. 5 km of exhumation.

5.3.2 Gołdap IG-1

Two samples, Proterozoic granite (B14) and Permian (Rotliegend) sandstone (B13), were taken from the Gołdap IG-1 borehole (Fig. 7a, Tabs. 1, 3). The older sample (B14) shows a younger AFT age (141 Ma) and lower MTL (12.0 μm with SD 1.5) compared to the younger early Permian sample (B13) - an AFT age of 167 Ma and MTL of 13.1 μm with SD 1.3. The VR value in the Permian rocks (c. 0.4% VR) give a maximum paleotemperature of 60 °C (Fig. 7), which is too low to cause a full reset of the AFT thermochronometer. Also, illite/smectite data suggest a similar paleotemperature below 70 °C (Kowalska et al., 2019). Furthermore, a temperature of 60 °C is consistent with limited sedimentary burial of sample B13 (e.g., Botor et al., 2019a). No episodes of significant Mesozoic exhumation are documented by the Gołdap IG-1 well, whose profile is full of sedimentary gaps but did not reveal any erosional unconformity. A significant discrepancy between VR$_{eq-derived}$ paleotemperature (c. 75 °C) and illite/smectite-derived paleotemperature (160 °C) exists in the nearby Bartoszyce IG-1 well (Fig. 4). A similar discrepancy between the maximum palaeotemperatures evaluated from illite-smectite and biomarkers was detected for the EEP by Derkowski et al. (2021). These authors interpreted a diagenetic pattern in the Ediacaran sediments of the EEP as the result of short-lasting pulses of potassium-bearing hot fluids, effectively promoting illitization in porous rocks without altering the organic matter (Derkowski et al., 2021). Correspondingly, we suggest that the reset of the AFT age in sample B13 was achieved in temperatures close to the upper limit of the apatite partial annealing zone in the presence of fluids. This interpretation is supported by paleotemperature profile of the nearby Bartoszyce IG-1 well, where a bell-shape paleotemperature profile suggests transient fluid flow in Triassic rocks (Fig. 7d). The fluids hypothesis does not explain resetting of the AFT age in sample B14, representing the crystalline basement. However, this sample, due to deeper burial, was subjected to a temperature of c. 75 °C in the Mesozoic. In the B14 sample, ZHe single grain ages are dispersed in the range of 1216-799 Ma suggesting no re-heating in the Phanerozoic and temperatures below a helium partial retention zone (below 130 °C). An unconformity exists between the middle Cambrian sandstones and Ordovician (Arenig) mudstones that is characterised by a discontinuity of the VR profile. Since the latter is relatively short an assessment of paleogradient and exhumation (c. 3 km) is uncertain and seems to be overestimated compared to regional data (Botor et al., 2019a). Another discontinuity of the VR profile occurs at the base Permian unconformity, indicating a pre-Permian cooling event (Fig. 7c).
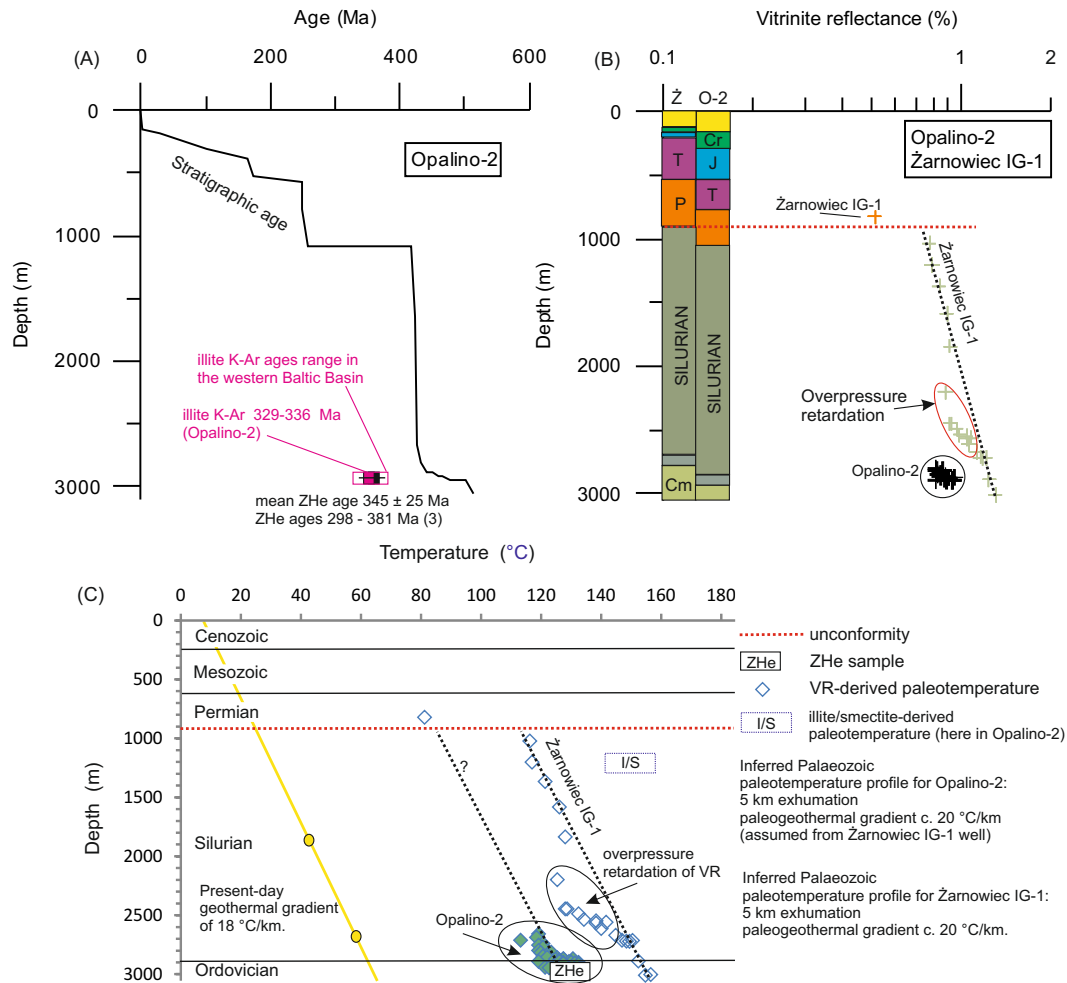
Figure 6. Opalino-2 and Żarnowiec IG-1 borehole data (western part of the Baltic Basin). (A) ZHe ages are similar to K-Ar ages (Kowalska et al., 2019) and are younger than a depositional age of sample. (B) Opalino-2 (O-2) and Żarnowiec IG-1 (Ż) vitrinite reflectance (VR) profiles. The approach by Petersen at al. (2013) was used to recalculate reflectance measurements on vitrinite-like macerals from the lower Palaeozoic strata into a standard vitrinite reflectance scale. Since the Opalino-2 VR profile is too short to define gradient and estimate exhumation the adjacent well Żarnowiec IG-1 VR was used. The profile clearly shows overpressure retardation of thermal maturity, which occurs within the mudstone/claystone dominated Silurian section in most wells in the western Baltic Basin. The Żarnowiec IG-1 VR profile shows a break across the Silurian/Permian unconformity as a VR value is 0.5% in Permian sediments. (C) Paleotemperature profile of the Opalino-2 (green diamonds) and Żarnowiec IG-1 (blue open diamonds) wells. Yellow line represents present-day geothermal gradient based on the corrected bottom hole temperature data (Górecki et al., 2006a,b). Paleogeothermal gradient of 20°C/km in the Opalino-2 well, comparable to the present-day gradient and the paleogradient in the Żarnowiec IG-1 well, gives exhumation of c. 4-5 km similar to that in the Żarnowiec IG-1 well.
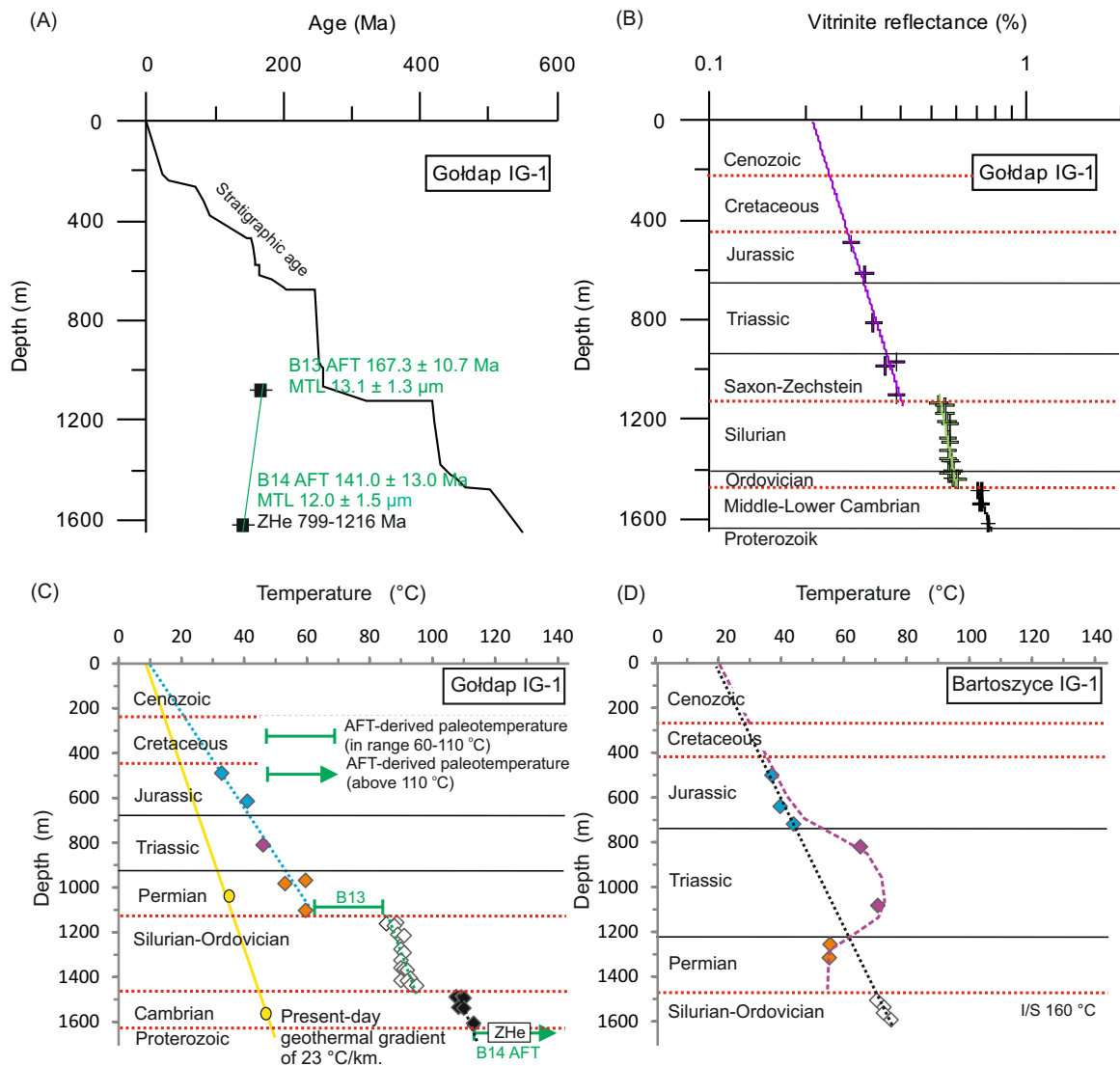
Figure 7. Gołdap IG-1 borehole data (eastern part of the Baltic Basin). (A) AFT ages vs. stratigraphic age in the Gołdap IG-1 well (ZHe age not to scale), (B) vitrinite reflectance (VR) profile in the Gołdap IG-1 well. (C) Paleotemperature profile: inferred Mesozoic profile (blue dotted line) shows no exhumation with paleogeothermal gradient of c. 45 °C/km; inferred Ordovician-Silurian profile (green dotted line) shows 3 km exhumation with paleogeothermal gradient of 33 °C/km. (D) Bartoszyce IG-1 paleotemperature profile - bell-shape curve (e.g., Ziagos and Blackwell, 1986) suggests transient fluid flow due to lateral migration in Triassic rocks. Significant discrepancy between VR$_{eq-derived}$ paleotemperature (c. 75°C) and illite/smectite-derived paleotemperature (I/S = 160°C; not to scale, data from Kowalska et al. (2019) exists in the upper part of Silurian strata that can be also caused by fluid flow. For other explanations see Figure 6

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5.3.3 Polik IG-1

The Ordovician bentonite (sample B36) from the Polik IG-1 borehole (Tab. 1) yielded a well-constrained ZHe mean age of 288 ± 6 Ma (Tab. 2; Fig. 8a). The Ordovician bentonite horizons gave also an illite K-Ar ages of 347-343 Ma (Kowalska et al., 2019), showing approximately the time of maximum paleotemperature. Therefore, the ZHe age probably represents cooling after the early Carboniferous thermal peak. A linear VR profile (Fig. 8b) shows a discontinuity across a base Permian unconformity between Ludlow (Silurian; 2.0% VR) and Rotliegend (Permian; 0.8% VR). Also, illite/smectite profile shows such a break (Kowalska et al., 2019). The paleogeothermal gradient of 35°C/km was calculated for the late Palaeozoic, the assessment implying exhumation of the Silurian-Devonian, and likely lowest Carboniferous strata by 5.6 km (Fig. 8c). This estimate agrees with the maturity modelling results by Botor et al. (2019a). The inferred Mesozoic temperature profile for the Polik IG-1 borehole suggests 700 m of exhumation experienced by the late Mesozoic strata and paleogeothermal gradient of 23°C/km, similar to the present-day gradient of 24 °C/km (Fig. 8c).

5.3.4 Tłuszcz IG-1

Two sandstone samples, B44 (Triassic) and B45 (Cambrian), were collected from the Tłuszcz IG-1 borehole (Tabs. 1–3; Fig. 9a). In both samples, AFT ages are younger than depositional ages and track lengths are shorten (Tab. 3). Sample B45 have a broader fission track distribution with 1.9 μm standard deviation that suggests re-heating or longer stay in the apatite partial annealing zone. Single ZHe ages from the B45 sample are in the range of 1088-292 Ma, indicating partial reset. Based on the difference in sensitivity of the helium system in zircons (130-220°C) and fission tracks in apatites (60-120°C), it can be concluded that sample B45 was exposed to a maximum paleotemperature of about 120-130°C. A nonlinear VR profile of the Tłłuszcz IG-1 borehole is similar to that in the Siedliska IG-1 well (Fig. 10b), suggesting a fluid flow event (e.g., Ziagos and Blackwell, 1986). There are no Carboniferous strata in the Tłuszcz IG-1 borehole, where a base Permian unconformity incises upper Silurian sediments (Fig. 9). At the bottom of the Permian-Mesozoic succession, paleotemperature was relatively high, reaching 90°C in Permian sediments. The illite/smectite data indicate 120°C at the top of the Silurian strata that is higher than the VR-derived paleotemperature, the relationship indicative of possible fluid flow influence (Fig. 9c). Alternatively, the gap between the VR- and illite/smectite-derived paleotemperatures might be the effect of contamination by detrital illite. However, in such a case the results should be highly incoherent, whereas in the Tłuszcz IG-1 and other studied boreholes the VR-derived paleotemperatures are consistently lower than those calculated based on illite/smectite data.

5.3.5 Siedliska IG-1

Bashkirian (upper Carboniferous) sandstone sample B38 (Tab. 1) yielded an AFT central age of 198 Ma (MTL 11.6 ± 1.7μm - Tab. 3). This result shows the importance of a late- to post-Carboniferous thermal event for thermal maturity of Palaeozoic succession in the area. There is a regional unconformity and stratigraphic gap between the upper Carboniferous (Moscovian) and Middle Jurassic. A non-linear VR profile below the unconformity (Fig. 10) shows that fluid flow may have contributed to maturation of organic matter (e.g., Ziagos and Blackwell, 1986). A discontinuity in the vitrinite reflectance profile exists at
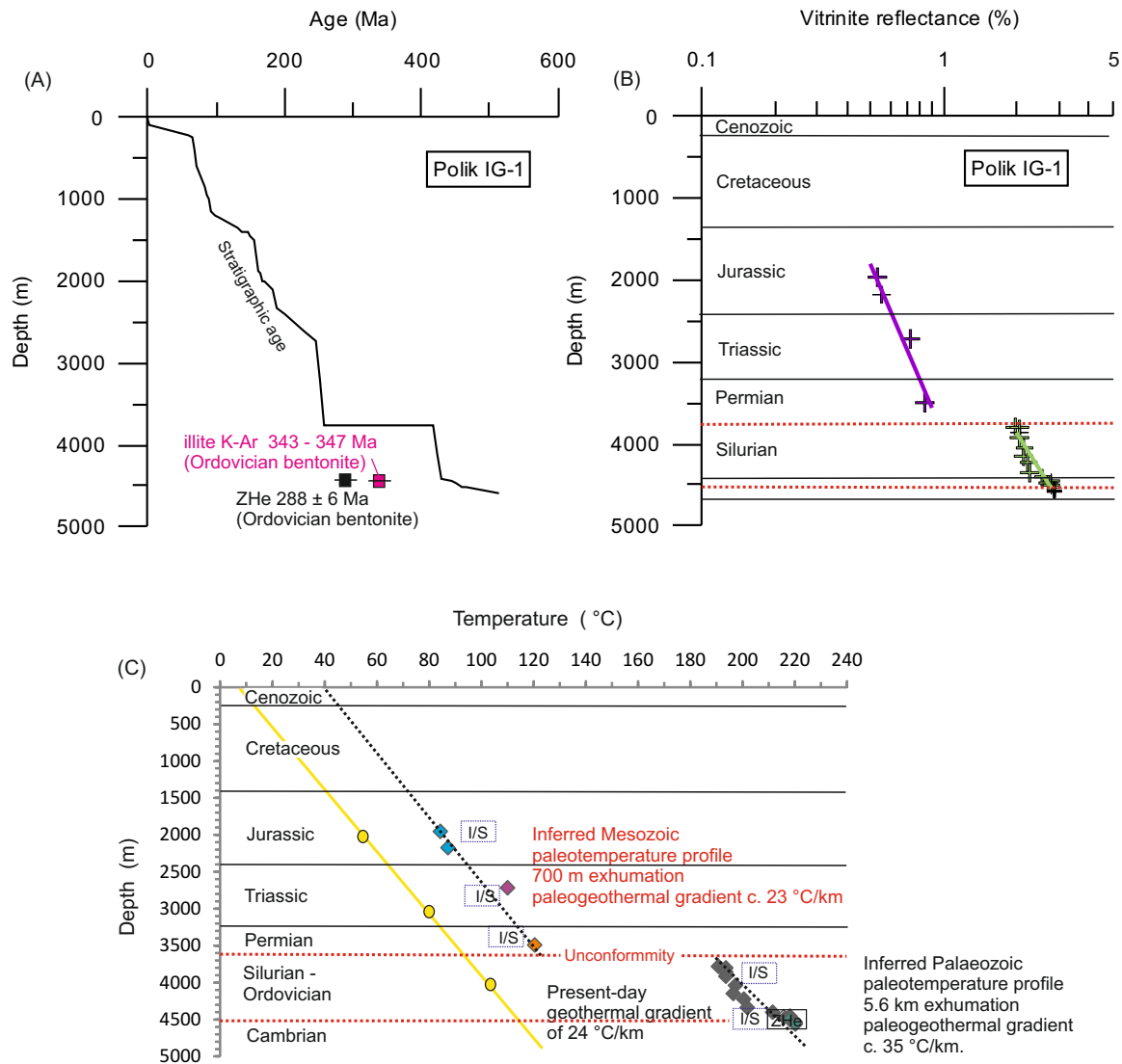
Figure 8. Polik IG-1 borehole data (SW part of the Baltic Basin). (A) Thermochronological data vs. stratigraphic age, (B) vitrinite reflectance (VR) profile, (C) paleotemperature profile. Note a thermal maturity break at the unconformity between the Permian-Mesozoic and Silurian. Other explanations as in Figures 6, 7.

the unconformity since thermal maturity of Mesozoic strata is below 0.5% VR (Fig. 10). Owing to a non-linear VR profile an assessment of post-Carboniferous exhumation might be overestimated, and it is not attempted herein.
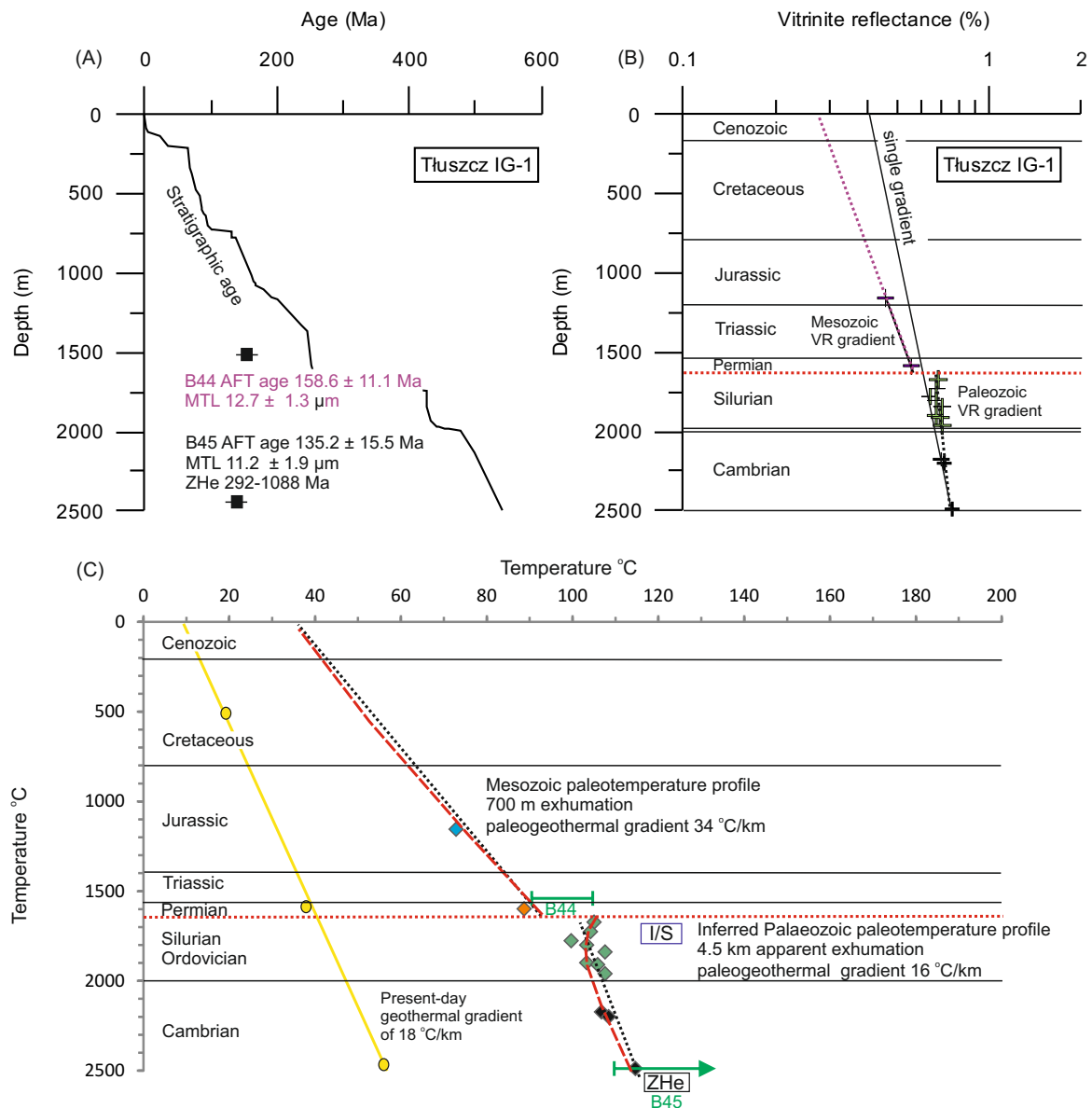
Figure 9. Tłuszcz IG-1 borehole data (Podlasie Basin). (A) Thermochronological data vs stratigraphic age, (B) VR profile, (C) paleotemperature profile. Note a thermal maturity break at the unconformity between the Permian-Mesozoic and Silurian. The Mesozoic paleotemperature profile shows 700 m of exhumation with a paleogeothermal gradient of 34°C/km. The inferred Palaeozoic paleotemperature profile shows 4.5 km of apparent exhumation with a paleogeothermal gradient of 16°C/km. A nonlinear VR profile suggests a fluid flow event. Other explanations as in Figures 6 and 7. Illite K-Ar ages from Silurian bentonites are 334-311 Ma (mid-late Carboniferous) in the adjacent Goździk OU-1 well (Kowalska et al., 2019).
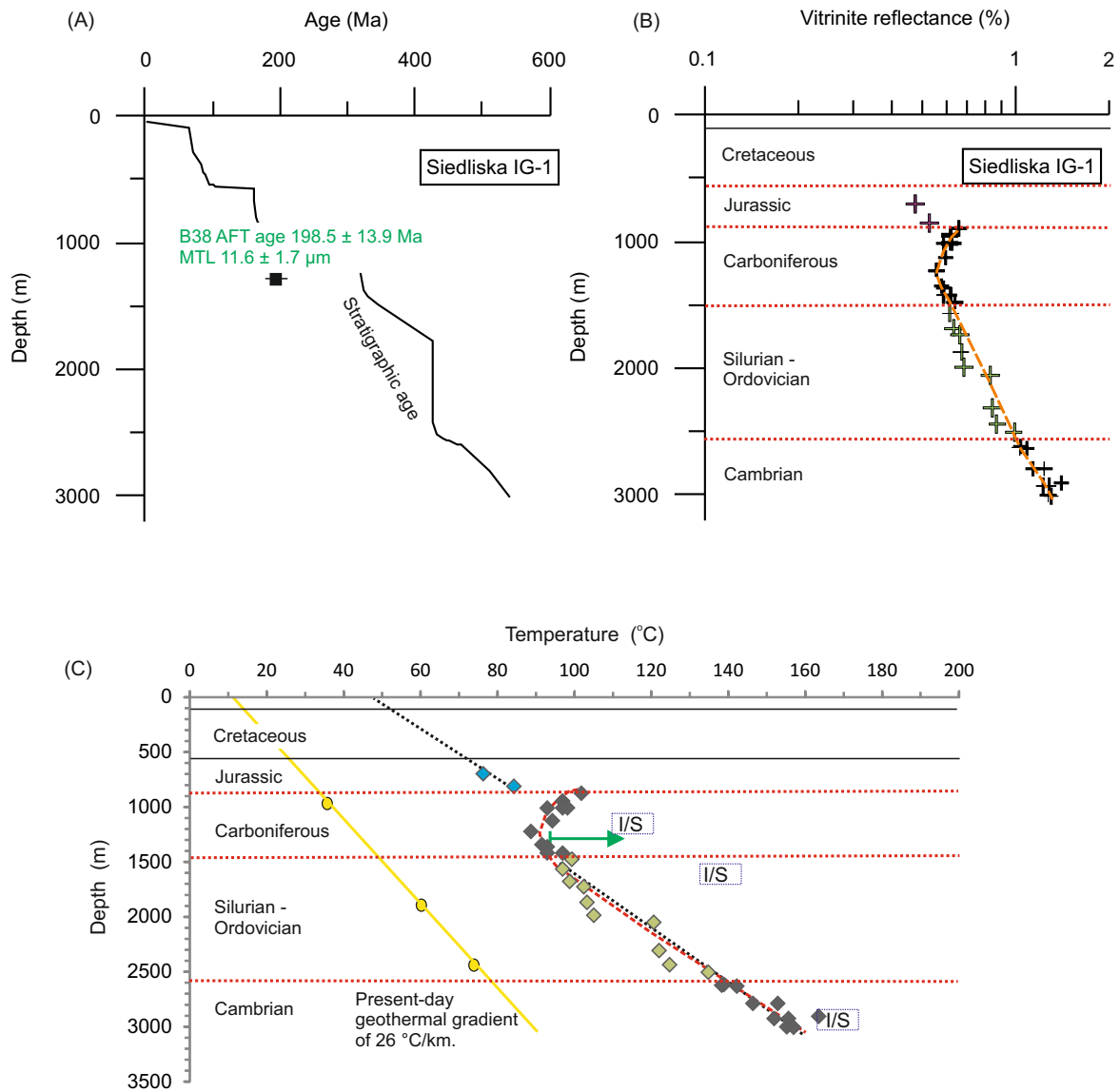
Figure 10. Siedliska IG-1 borehole data (Lublin Basin). (A) Thermochronological data vs stratigraphic age, (B) VR profile, (C) paleotemperature profile. VR and paleotemperature profiles suggest that fluid flow event is responsible for the pattern revealed. Other explanations as in Figure 6 and 7. Illite K-Ar ages from Silurian bentonites are 344-336 Ma (early Carboniferous) in the adjacent Wojcieszków-1 well (Kowalska et al., 2019).

5.3.6 Łopiennik IG-1

Sample B19 was collected from the bottom of the Viséan succession (Tab. 1, Fig. 5b) in the Łopiennik IG-1 borehole that is located in the Lublin Basin (Figs. 1, 4). Five zircons were dated in this sandstone sample by means of the ZHe method. An av-

erage ZHe age of this sample is 287 Ma after omitting the single grain age that is older than a stratigraphic age (562 Ma; Tab. 2, Fig. 11a). VR value of 1.0% for sampled depth shows a maximum paleotemperature of c. 140°C, suggesting only partial reset of the ZHe age (Fig. 11b). The Ediacaran to Carboniferous VR profile allows to estimate VR and paleogeothermal gradients for the late Palaeozoic (Fig. 11b, c). Although no VR data exist in the upper part of Carboniferous and Mesozoic sections of the profile, average VR values in the adjacent wells are c. 0.8-1.0% and 0.45-0.55% VR for the uppermost Carboniferous and Jurassic-Cretaceous strata, respectively (Botor et al., 2002, 2019a; Grotek, 2005). The Palaeozoic and Mesozoic parts of the profile are separated by an unconformity between Carboniferous (Namurian C, Bashkirian) and Middle Jurassic. The illite-smectite and VR data indicate similar paleotemperatures (Kowalska et al., 2019). A linear gradient of VR data with fitting for Palaeozoic (R^2=0.80) allows to estimate post-Namurian exhumation at 3700 m (assuming 0.2% VR as an initial value). Alternatively, it can be calculated from a paleogeothermal gradient (24°C; R^2=0.89) that a 4300 m thick pile of the sediments was removed, assuming 20°C as a surface temperature from the late Carboniferous to Early Jurassic (Fig. 11c). Both estimates are similar to that from maturity modelling based on the VR and porosity data (Botor, 2018; Botor et al., 2019a). However, assuming overpressure retardation in Silurian sediments, a paleogeothermal gradient of 23°C would suggest exhumation up to 7 km (Fig. 11c).

5.3.7 Lubycza Królewska-1

Bentonite sample LK1/1 was collected from the upper Silurian strata in the Lubycza Królewska-1 borehole in the southern part of the Lublin Basin (Tabs. 1–3, Figs. 1, 4, 12). Single grains ZHe ages are 293, 253, and 219 Ma, giving a late Permian mean ZHe age of 255 Ma with standard error of 22 Ma. There is no AFT data for this sample, but second upper Silurian sample LK1/2 gives an AFT central age of 178 ± 67.5 Ma, based on 3 apatite grains (Tab. 3, Fig. 12a). A low precision of this age is demonstrated by a large 1σ error. The timing of a maximum paleotemperature in the studied samples can be estimated based on illite K-Ar ages of 298-272 Ma (Kowalska et al., 2019). A short time gap between K-Ar ages and ZHe ages suggests a rapid decrease of temperature and can be related to late Carboniferous to early Permian rapid uplift. A pre-Middle Jurassic unconformity splits the VR profile into the matured Palaeozoic part (c. 1.7-2.0% VR) and the immature (below 0.65% VR) Mesozoic part (Fig. 12b). The Mesozoic and Palaeozoic VR profiles are shifted relative each other across the unconformity. The VR profile break suggests significant exhumation between the Silurian and Middle Jurassic. The LK1/1 sample, showing Permian ZHe ages, was heated up to c. 180°C according to the VR data (Tab. 1). However, the inferred Palaeozoic paleotemperature profiles of the Lubycza Królewska-1 and Narol IG-1/PIG-2 boreholes show unrealistic c. 11 km exhumation with a paleogeothermal gradient of c. 13°C/km. The measured VR profile is too short to confidently define a paleogeothermal gradient and, thus, the tentatively calculated 11 km exhumation is probably overestimated (Fig. 12).

5.4 Thermal modelling

Samples yielding the best quality analytical results were chosen for thermal history modelling by means of the HeFTy software (Ketcham, 2005; Ketcham et al., 2007b). For sedimentary samples, the beginning of the t-T paths was defined by an annual mean temperature estimated for the time of deposition (c. 25 ± 5°C). For sample B14 of Proterozoic granitoid, a starting point
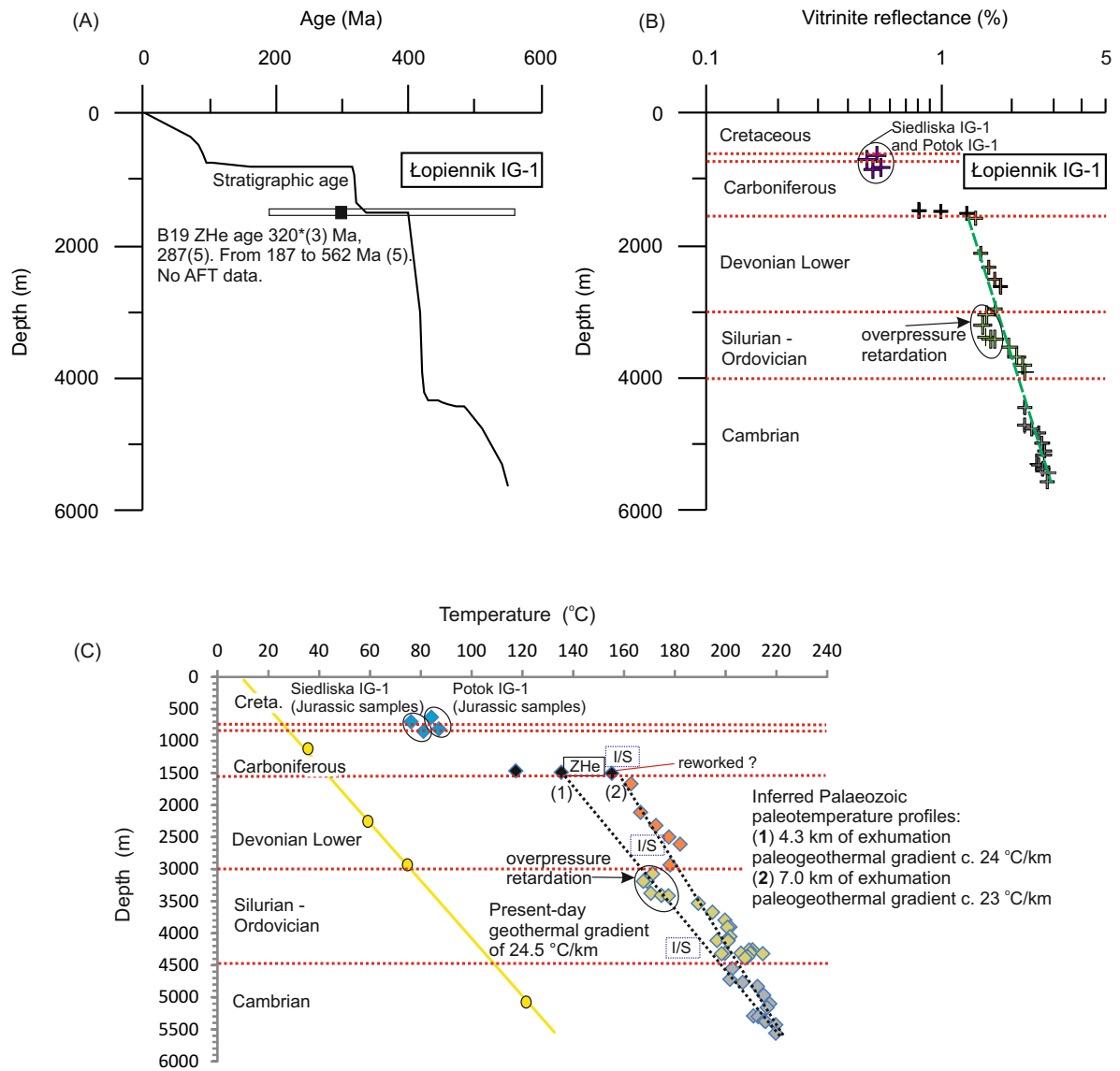
Figure 11. Łopiennik IG-1 borehole data (Lublin Basin). (A) Thermochronological data vs stratigraphic age, (B) VR profile, (C) pale­otemperature profile. Mesozoic VR data from the Siedliska IG-1 and Potok IG-1 boreholes are from the Middle-Upper Jurassic sediments (sampling depth is not in scale). Carboniferous VR data change rapidly in a short profile section and different VR and paleogeotermal gradi­ents can be established. Other explanations as in Figures 6 and 7.

of modelling was assumed at the transition from Ediacaran to Cambrian since early Cambrian sedimentary rocks, overlying basement, confirm a near surface temperature of c. 25 ± 5°C at the that time (Tab. 1). The end of t-T paths in both cases (sedimentary and crystalline samples) was defined by an average present-day temperature at a sampling depth (Tab. 1 and 4).

385   A contemporary temperature in the analysed boreholes is much lower than the sensitivity range of the thermochronological
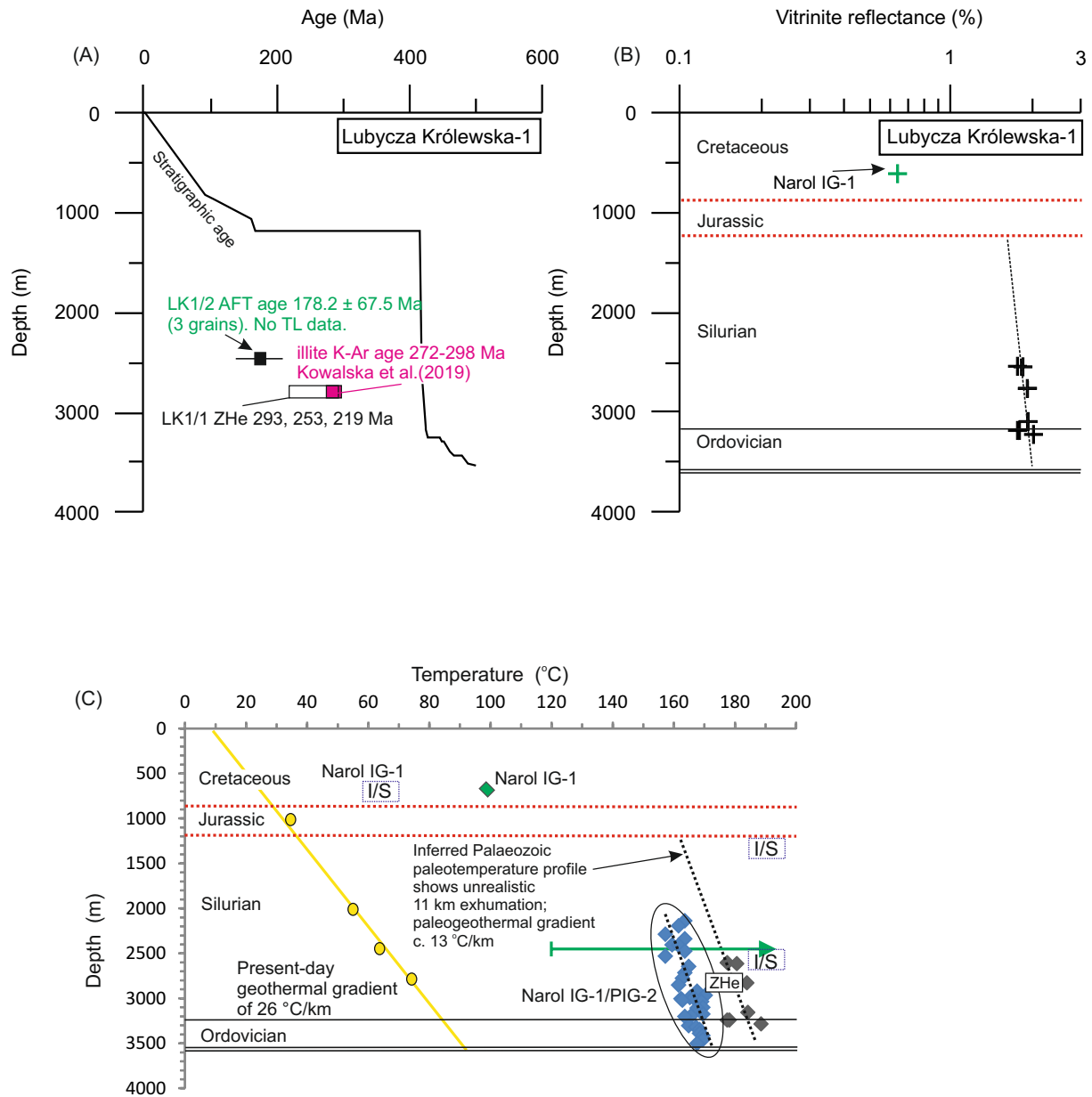
Figure 12. Lubycza Królewska-1 borehole data (Lublin Basin). (A) Thermochronological data vs stratigraphic age, (B) VR profile, (C) paleotemperature profile. Other explanations as in Figures 6 and 7. VR and illite/smectite (I/S) data from the Narol IG-1/PIG2 boreholes are not to scale. A thermal maturity break exists between the Mesozoic and Palaeozoic. VR data from the Narol IG-1/PIG2 boreholes show a similar paleogeothermal gradient to that in the Lubycza Królewska-1. The inferred Palaeozoic paleotemperature profile of the Lubycza Królewska-1 borehole shows unrealistic 11 km of exhumation with a paleogeothermal gradient of 13°C/km.

Table 4. Basic data and results of the thermal modelling.

Sample	Well	Stratigraphy	Temp. (°C)	Temp.1 (°C)	Temp.2 (°C)	Onset of cooling (Ma)	AFT age (Ma)	Range of ZHe ages (Ma)
B13	Gołdap IG-1	Permian (Saxon)	35	60	80	190	167.3 ± 10.7	n.d.
B14	Gołdap IG-1	Proterozoic	45	113	120	190	141.0 ± 13.0	799-1247
B19	Łopiennik IG-1	Carboniferous low.	50	153	130	303	n.d.	291-368
B36	Polik IG-1	Ordovician Upper	80	217	210	360	105.7 ± 21.8*	278-300
B38	Siedliska IG-1	Carboniferous upp.	35	92	100	300	198.5 ± 13.9	n.d.
B43	Stadniki IG-1	Proterozoic	45	120**	110	300	235.8 ±17.3	n.d.
B44	Tłuszcz IG-1	Triassic Lower	36	89	110	190	158.6 ± 11.1	n.d.
B45	Tłuszcz IG-1	Cambrian lower	55	116	124	190 or 330	135.2 ± 15.5	1088-292
LK1/1	LK1	Silurian upper	72	184	172	298	178.2 ± 67.5#	219-294
O2/2	Opalino-2	Ordovician Upper	60	127	145	360	n.d.	298-382

Temp. present-day temperature in the well from Tab. 1. Temp.1: % max. paleotemperature from organic maturity data from Tab. 1. Temp.2: max. paleotemperature from HeFTy modelling. * sample B9 in adjacent well - Bodzanów IG-1, ** from illite/smectite data given by (Kowalska et al., 2019). LK1: Lubycza Królewska-1 well; # sample LK1/2.

method. Thus, temperature fluctuations of ±10 °C in the boreholes from the study area that were reported in the paper by Szewczyk and Gientka (2009) do not affect the results obtained. In the HeFTy modelling, the t-T constrain boxes were based on burial history (Fig. 5; Botor et al., 2019a), calibrated by thermal maturity data and geological evidence, i.e. unconformities (Tab. 1). The temperature range of 60-200°C was established based on sensitivity of both AFT and ZHe systems. Time windows

390 for constrain boxes corresponded to subsidence periods limited by regional unconformities. Input data and results of thermal modelling are provided in Table 4. Except for the B47 sample (Jurassic sandstone, Tyniewicze IG-1) and B39 sample (Permian sandstone, Słupsk IG-1), the AFT and ZHe ages are considerably younger than the age of deposition. Therefore, we did not consider a pre-depositional thermal history of the detrital grains. The thermal modelling results are compiled in Figure 13, where the best fit results are shown. High goodness-of-fit (GOF) values (0.90 to 1.00) suggest that the resulted time-temperature

395 (t-T) paths are plausible.

5.4.1 Baltic Basin

In the western part of the Baltic Basin, thermal modelling was performed for the O2/2 sample (Upper Ordovician bentonite; Tab. 1-3, Fig. 1) from the Opalino-2 borehole using ZHe data only and based on the assumption of a positive correlation between effective uranium ($eU = U + 0.235 \cdot Th$; in ppm) and ZHe single grain ages. Constrain boxes for the O2/2 sample

400 were established to represent following regional subsidence and heating events: development of a Caledonian foreland basin (450-400 Ma), subsidence of the passive Laurussia margin (370-320 Ma), Variscan shortening and subsequent continental rifting (300-240 Ma), renewed tectonic subsidence of the Permian-Mesozoic basin (220-160 Ma) and tectonic inversion of the Permian-Mesozoic basin (100-60 Ma). The resultant model (Fig. 13a) shows a maximum temperature in earliest Carboniferous, probably owing to burial beneath a Caledonian foreland basin and Devonian passive margin cover as suggested by geologi-

405 cal evidence (e.g., Narkiewicz, 2007). Although details of the best fit curve are difficult to verify due to the missing upper
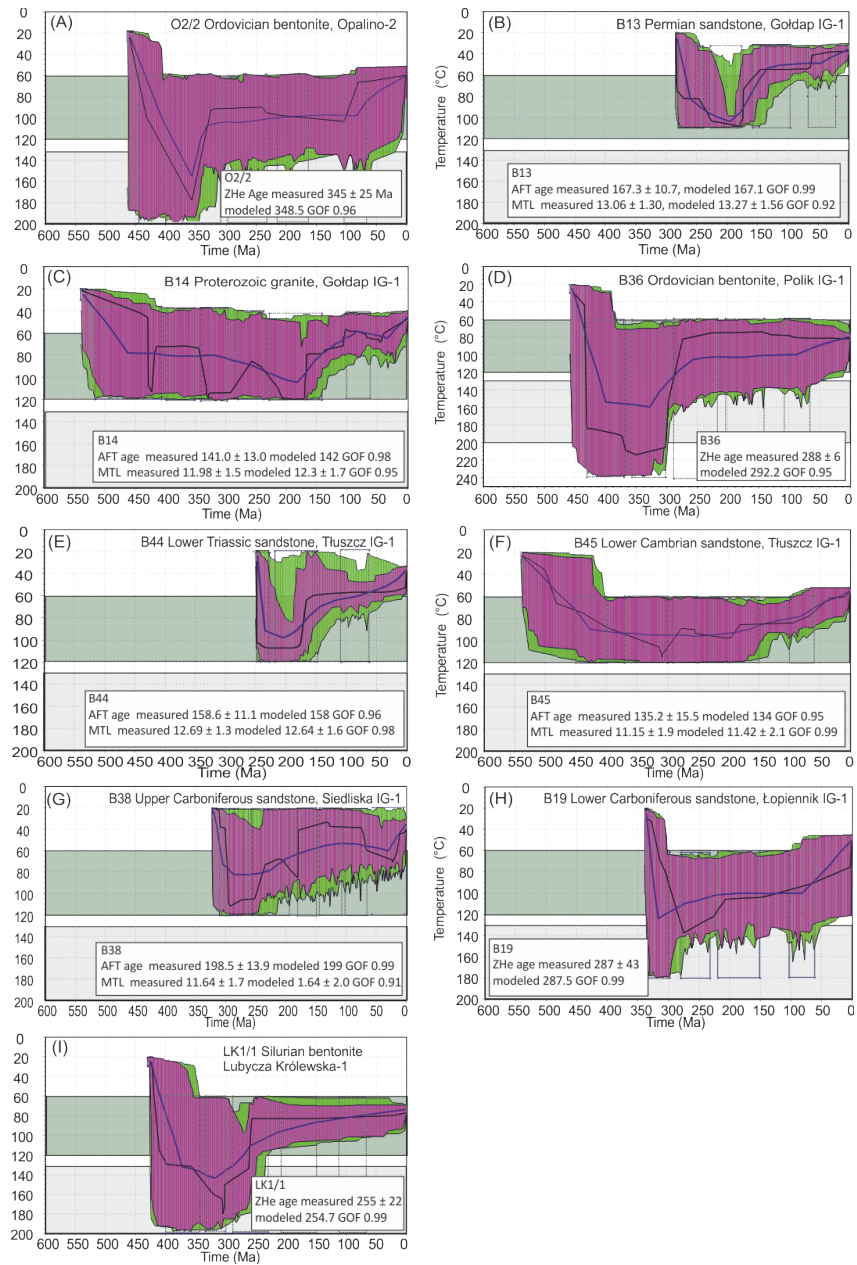
Figure 13. Thermal modelling results from HeFTy software (Ketcham, 2005). The light green range corresponds to the envelope of thermal paths with acceptable fit (goodness of fit - GOF > 0.05), the magenta range shows the envelope for thermal paths with good fit (GOF > 0.5). Bold dark blue curve shows weighted mean path, whereas thin black line is the best fit curve. Black rectangles correspond to constrain boxes. Grey rectangles show apatite partial annealing zone (60-120°C) and zircon partial retention zone (130-200°C). Further explanations in the text.

Palaeozoic sedimentary record, both the weighted mean path and best fit curve indicate an important exhumation event in the early Carboniferous (360-330 Ma). This is consistent with the illite K-Ar ages (360-322 Ma) from the same area (Kowalska et al., 2019). Subsequent Variscan shortening had a little impact on the thermal model in accord with a tectonic setting of the Opalino-2 well that is located beyond the Variscan deformation front (Fig. 1). In its late Carboniferous-Mesozoic section, the best fit curve suggests limited middle to late Mesozoic reheating. This corollary is difficult to verify owing to the lack of AFT data. Finally, the model shows acceleration of cooling coeval with the tectonic inversion of the Permian-Mesozoic basin.

In the eastern part of the Baltic Basin, the thermal maturity of lower Palaeozoic strata is one of the lowest in the Polish part of the EEP (Tab. 1; Fig. 4; Grotek, 2006, 2016; Pletsch et al., 2010). The sedimentary cover on the crystalline basement is the thinnest across the study area. Two samples were analysed in the Gołdap IG-1 borehole: Permian sandstone B13 and Proterozoic granitoid B14 (Tab. 1, Fig. 1). Since a negative correlation between eU and ZHe single grain ages in the B14 sample did not allow for helium data modelling (Green and Duddy, 2018) only AFT data were used. Constrain boxes for sample B13 (Fig. 13b) represent consecutive subsidence and thermal events: a time period following continental rifting and initiation of the Permian-Mesozoic basin (290-240 Ma), two periods of renewed Mesozoic subsidence (230-180 Ma and 160-100 Ma) and inversion of the Permian-Mesozoic basin (70-30 Ma). The model shows progressive heating up to 100°C until 200-190 Ma (earliest Jurassic). This temperature is higher than that predicted by VR and illite/smectite data (c. 60-70°C); (Kowalska et al., 2019). Therefore, we interpret early Mesozoic heating as a cumulative effect of increasing sedimentary burial and transient fluid flow in the late Permian and Triassic. The latter effect is confirmed by a bell-shaped paleotemperature profile in the nearby Bartoszyce IG-1 borehole (Fig. 7d). Consequently, the model predicts cooling after cessation of fluids influence for the remaining part of the Mesozoic and Cenozoic despite continued subsidence and sedimentation. Importantly, Mesozoic-Cenozoic sedimentary cover is thin and discontinuous. For instance, the Lower-Middle Jurassic sediments, corresponding to the time of the fastest cooling predicted by the model, are 100 m thick and their impact on a thermal state of the B13 sample was probably negligible. The total thickness of post-Triassic sediments in the Gołdap IG-1 borehole is 675 m, a sedimentary pile that generates a temperature increase of 15.5°C under the present-day geothermal gradient (23°C/km). At the same time, sample B13 from a depth of 1075 m was exposed to a present-day temperature of 35°C (Tab. 1, Fig. 7). Regardless if sample B13 was subjected to a maximum paleotemperature of 60-70°C (VR and illite/smectite data) or c. 100°C, as predicted by the thermal model (Fig. 13b), a post-Triassic period was the time of net cooling (down to 35°C) despite still ongoing sedimentation. We built a model for sample B14 (Fig. 13c) starting from the earliest Cambrian (540 Ma) with a near surface temperature of 20-40°C as the B14 sample was taken 20 m below the Cambrian unconformity (Fig. 7; Tab. 1). Constrain boxes for this model were at: 520-410 Ma (Baltica passive margin and Caledonian foreland basin), 400-350 Ma (Laurussia passive margin), 340-300 Ma (time equivalent of the Variscan orogeny), 280-240 Ma (an initial stage of the Permian-Mesozoic basin starting from continental rifting), 230-140 Ma (subsidence of the Permian-Mesozoic basin), 100-70 Ma (inversion of the Permian-Mesozoic basin). The weighted mean path of the model shows a maximum temperature in the earliest Jurassic (Fig. 13c), similarly to model B13 (Fig. 13b), and subsequent cooling. This phenomenon was already discussed above in relation to sample B13. The weighted mean path of the model predicts a maximum paleotemperature of 100°C for the earliest Jurassic, i.e., 25°C more than suggested by the Permian-Mesozoic geothermal gradient (75°C). This difference might be related to the influence of fluids.

The best fit curve reveals several details that cannot be directly verified due to incomplete sedimentary record: (1) a thermal peak in the latest Silurian potentially related to burial by a Caledonian foreland basin, (2) a thermal peak in early Carboniferous related to burial or increased heat flow (magmatism), and (3) early Mesozoic thermal peak related to burial beneath Permian-Mesozoic basin and supposed transient fluid flow. Sample B14 left the apatite partial annealing zone (141 ± Ma) after sample B13 (167.3 ± 10.7 Ma) that was located shallower in a sedimentary section.

5.4.2 Podlasie Basin

Thermal modelling for the B36 sample (Upper Ordovician bentonite, Polik IG-1) was performed based on ZHe results as no AFT data were available (Fig. 13d). All 3 zircons dated have a similar eU content up to 100 ppm and their mean ZHe age reveals low standard error (Tab. 2). A distinct break in thermal maturity profile (Fig. 8) proves that a maximum paleotemperature was achieved in the pre-Permian time. The VR-derived maximum paleotemperature calculated for sample B36 is 215°C (Tab. 1). Constrain boxes in the model were set at: 440-390 Ma (Caledonian foreland basin), 360-300 Ma (Laurussia passive margin and Variscan shortening), 290-230 Ma (initial stage of the Permian-Mesozoic basin), 200-140 Ma (further subsidence of the Permian-Mesozoic basin), and 100-60 Ma (inversion of the Permian-Mesozoic basin). The weighted mean path shows an increase of temperature until c. 420 Ma, corresponding to burial beneath a Caledonian foreland basin, and fast cooling in the Carboniferous (starting 340-330 Ma). The latter must have been related to extensive exhumation as demonstrated by a significant break in the paleotemperature profile across the base Permian unconformity (Fig. 8). The best fit curve reveals a thermal peak at 350 Ma (earliest Carboniferous), the effect that may have been related not only to burial but also increased heat flow (magmatism). A maximum paleotemperature in the early Carboniferous is consistent with the illite K-Ar ages obtained from Ordovician bentonites (347-343 Ma; Kowalska et al., 2019). The weighted mean path of the model suggests that Mesozoic was the time of thermal stability with only minor acceleration of cooling at the time of the Permian-Mesozoic basin inversion (Fig. 13d).

In the Tłuszcz IG-1 borehole two samples were analysed. Thermal modelling for sample B44 (Lower Triassic sandstone; Fig. 13e) was performed using constrain boxes for: Triassic basin subsidence (250-230 Ma), further Mesozoic subsidence (220-150 Ma), and Permian-Mesozoic basin inversion (100-60 Ma). A thermal peak of 100°C is predicted by the weighted mean path of the model for the earliest Jurassic (c. 200 Ma). This is close to a VR-based paleotemperature of 90°C in Permian sediments (Fig. 9). The post-Early Jurassic period was the time of gradual cooling despite ongoing sedimentation as already discussed above (model B13). The overlying sedimentary pile generated temperature of c. 50-55°C, using the Mesozoic geothermal gradient, that is less than 90°C noted in the Permian strata (Karnkowski, 2003). A present-day temperature at a depth of sample B44 is only 36°C (Tab. 1). Therefore, the Mesozoic was the time of net cooling that was enhanced by 700 m of exhumation calculated for the Permian-Mesozoic profile (Fig. 9).

Thermal modelling of sample B45 (Lower Cambrian sandstone) was performed using AFT data only, because the e-U trend is unclear (Tab. 2). Constrain boxes were set at: 450-400 Ma (subsidence of a Caledonian foreland basin), 370-300 Ma (subsidence and shortening of the Variscan foreland), 290-250 Ma (initial subsidence of the Permian-Mesozoic basin), 230-150 Ma (further subsidence of the Permian-Mesozoic basin), and 100-60 Ma (Permian-Mesozoic basin inversion). The

475   model predicts heating until 420 Ma (Fig. 13f), the result probably related to burial beneath a Caledonian foreland basin. Although the weighted mean path of the model does not show a clear temperature peak in the late Palaeozoic the best fit curve indicates a thermal peak of 110°C at 310 Ma (Fig. 13f). This prediction is in accord with a VR-derived paleotemperature of 116°C (Tab. 1). The thermal peak corresponds to a big erosional gap between the Silurian and Permian as well as a break in the paleotemperature profiles between the lower Palaeozoic and Permian-Mesozoic strata (Fig. 9). In the post-Carboniferous

480   period, the model shows cooling down to a present-day temperature of 55°C.

5.4.3   Lublin Basin

Thermal modelling was performed for upper Carboniferous (Bashkirian) sandstone sample B38 that was collected from the Siedliska IG-1 borehole (Figs. 1, 4). In the model for sample B38 (Fig. 13g), constrain boxes were set to represent a few succeeding subsidence and thermal events: Variscan shortening and subsequent continental rifting (320-260 Ma), subsidence

485   of the Permian-Mesozoic basin (255-195 Ma), two periods of further subsidence (180-150 and 145-110 Ma), and Permian-Mesozoic basin inversion (100-60 Ma). The model shows a temperature peak of 80°C (the weighted mean path) in the latest Carboniferous-earliest Permian. This is less than VR-derived paleotemperature in the range of 90-100°C (Fig. 10). However, the best fit curve reveals a maximum temperature of c. 110°C for roughly the same age range. The latest Carboniferous-earliest Permian thermal peak is consistent with a big erosional gap between the Carboniferous and Jurassic and a break of

490   the paleotemperature profile (Fig. 10). The thermal peak was followed by gradual cooling through the Mesozoic and Cenozoic down to the present-day temperature of 35°C.

A positive correlation between e-U and ZHe ages in sample B19 of Carboniferous sandstone from the Łopiennik IG-1 borehole favours thermal modelling despite the lack of AFT data. Modelling was carried out for a single zircon grain B-19z4 due to dispersion of the ZHe ages (Tab. 2, Fig. 13h). Four constrain boxes were used in the model: 340-300 Ma (subsidence and

495   subsequent shortening of the Variscan foreland), 280-240 Ma (initial subsidence of the Permian-Mesozoic basin), 230-150 Ma (further subsidence of the Permian-Mesozoic basin), and 100-60 Ma (Permian-Mesozoic basin inversion). The weighted mean path of the model reveals rapid heating until a thermal peak at 320 Ma. The best fit curve predicts a thermal peak in the early Permian (280 Ma). The results are consistent with a big erosional gap between the Carboniferous and Upper Jurassic as well as a major break in the paleotemperature profile across the unconformity (Fig. 11). The post-early Permian time was characterised

500   by a gradual cooling with some acceleration at the time of the Permian-Mesozoic basin inversion (Fig. 13h).

For modelling of sample LK1/1 (Silurian bentonite; Fig. 13i), the following constrain boxes were assumed: (1) post-Caledonian subsidence (400-350 Ma), (2) Variscan subsidence and shortening (340-300 Ma), (3) initial subsidence of the Permian-Mesozoic basin (290-240 Ma), and (4) further subsidence of the Permian-Mesozoic basin (230-150 Ma), and Permian-Mesozoic basin inversion (100-60 Ma). The weighted mean path indicates a thermal peak of 140°C at c. 320 Ma. The best fit

505   curve points to thermal peak of 180°C at c. 300 Ma. The latter modelled temperature is consistent with a VR-derived pale-otemperature of 175°C (Fig. 12). Furthermore, an age of c. 300 Ma is close to the illite K-Ar ages of 298-272 Ma from the same borehole (Kowalska et al., 2019). The latest Carboniferous thermal peak is consistent with a major erosional gap between Carboniferous and Jurassic in the Lubycza Królewska borehole and an important break of the paleotemperature profile across

the unconformity (Fig. 12). The thermal peak is followed by gradual cooling through Mesozoic and Cenozoic down to the present-day temperature of 65°C (Fig. 13i).

6 Discussion

Characteristic features of the SW slope of the East European Platform in Poland is an increasing thickness of sediments towards the TTZ and an erosional unconformity at the top of the lower Palaeozoic. An exception is the Lublin Basin, where the unconformity is at the top of Devonian (e.g., Narkiewicz, 2007). Another regional unconformity is located at the base of Permian to Jurassic strata. However, in the area, where Devonian and Carboniferous strata are missing (Fig. 14), there is only one major unconformity between the lower Palaeozoic and Permian. These unconformities must represent a succession of major tectonic events, but separation of their effects is impossible without thermal maturity and thermochronological data due to an incomplete sedimentary record.

6.1 Thermochronological constraints

The SW-ward increase of the total sediment thickness on the slope of the EEP has been well-known for decades based on borehole and seismic data (e.g., Młynarski, 1982; Poprawa and Pacześnia, 2002; Mikołajczak et al., 2019). Thermal maturity results are consistent with this trend (Fig. 4) since the highest paleotemperatures, based on VR data, are recorded in those boreholes that are located close to the TTZ (Tab. 1; Figs. 1, 4). Consequently, the only ZHe ages that may represent fully reset zircons are those obtained from the Polik IG-1 borehole (Tab. 2) in agreement with the VR-derived pre-Permian paleotemperature of 217°C (Fig. 8). An early Permian ZHe mean age of 288 Ma corresponds to cooling below the ZHe closure temperature of 130°C, the corollary consistent with the illite/smectite-derived paleotemperature (Fig. 8). The remaining ZHe ages represent partial reset or source ages (section 5.1). Therefore, the ZHe ages obtained, besides sample B36, cannot be directly used to constrain the time of cooling and must be supplemented by thermal modelling.

Considering significant thickness of the Permian-Mesozoic sedimentary cover on the SW slope of the EEP (Figs. 1, 4), it is not unexpected that AFT ages are dispersed in the range of 235.8 ± 17.3 to 42.1 ± 11.1 providing a record of Mesozoic to Cenozoic cooling (section 5.2). The highest frequency of the AFT ages is in the Jurassic and Early Cretaceous (Tab. 3) i.e., they are older than inversion of the German-Polish Basin (e.g., Senglaub et al., 2005; Resak et al., 2010; Łuszczak et al., 2020). Consequently, the results obtained suggest that cooling through the apatite partial annealing zone occurred in the Mesozoic before tectonic inversion in the Late Cretaceous. This might be related to the fact that tectonic inversion was significantly weaker in the part of the German-Polish Basin onlapping the EEP (Krzywiec et al., 2017b). Although inversion structures are undoubtedly recognised in this area (Krzywiec, 2009) the offset of vertical movements is probably below the sensitivity of the AFT method. Further inferences on timing and rate of Mesozoic cooling can be derived from thermal modelling. Nevertheless, the AFT data are not suitable for constraining a pre-Permian thermal history of the area.

6.2 Pre-Permian thermal history

540 Three regional unconformities at the top of crystalline basement, at the top of lower Palaeozoic and at the base of Permian or Mesozoic provide important geological constraints on thermal models limiting the number of viable solutions. Moreover, VR-derived paleotemperatures and ZHe data put important limits on the pre-Permian thermal histories modelled (section 5.4). Out of 9 models in total, 7 provides solutions for the part or entire Palaeozoic (Tab. 4, Fig. 13). Among the latter, two groups can be distinguished: (1) models predicting latest Devonian to early Carboniferous maximum paleotemperature followed by rapid

545 cooling throughout the Carboniferous and (2) models predicting a temperature peak at the end of Carboniferous (300 Ma), succeeded by cooling in the Permian-Mesozoic (Fig. 13). All 4 models in the first group were built for samples that were collected from boreholes in the Baltic and Podlasie Basins, the area missing Devonian and Carboniferous sedimentary rocks (Figs. 13, 14). Despite a regional base Permian unconformity and an important shift of the VR profile across this unconformity (Figs. 6-9), the age of the unconformity and a Palaeozoic thermal event has remained so far unconstrained, being broadly

550 limited to the interval between the end of Silurian and beginning of Permian. Therefore, our models make an important step forward predicting maximum heating at the transition from the Devonian to Carboniferous and rapid cooling soon afterwards (Fig. 13). This result is consistent with the early Carboniferous illite K-Ar ages by (Kowalska et al., 2019).

The second group of models represents samples collected from boreholes in the Lublin Basin (Figs. 13, 14). Although in the Lubycza Królewska-1 borehole, Silurian strata are directly overlain by Jurassic sediments, in the Łopiennik IG-1 and

555 Siedliska IG-1 boreholes parts of Devonian and Carboniferous sedimentary sections are preserved (Figs. 10, 11). In the latter two, no break in the VR profile is observed across the boundary between the lower Palaeozoic and Devonian or Carboniferous (Figs. 10, 11). This is consistent with the latest Carboniferous thermal peak modelled (Fig. 13) and earliest Permian illite K-Ar ages (Kowalska et al., 2019). Consequently, although a top Silurian or top Devonian unconformity exists in the Lublin Basin it is not associated with a major gap in the thermal history. In contrast to the Baltic and Podlasie Basins, a peak paleotemperature

560 was achieved there in the latest Carboniferous.

6.3 Permian-Mesozoic thermal history

Two models for Permian and Lower Triassic samples show exclusively the Mesozoic thermal history (Fig. 13). They both reveal a paleotemperature peak at the transition from the Triassic to Jurassic and subsequent cooling (Fig. 13). Since both samples (B13 and B44) come from relatively shallow part of the Permian-Mesozoic basin (c. 1000-1500 m; Figs. 7, 9) AFT

565 and VR data suggest that burial heating was strengthen by transient fluid flow in the Triassic. The Proterozoic granite sample (B14) from the Gołdap IG-1 borehole yielded the same Phanerozoic thermal history (Fig. 13c), including a peak temperature at the Triassic-Jurassic boundary and subsequent cooling.

Our data show that Mesozoic was the time of cooling despite ongoing sedimentation that might be considered a contradiction. On the other hand, we know a present-day temperature at sampling depths and a maximum paleotemperature for sampled strata

570 from VR data (Tab. 1). Furthermore, VR data from are consistent with illite/smectite data, where available. A comparison between VR data for the Permian-Mesozoic samples and present-day temperatures shows that sampled strata must have been
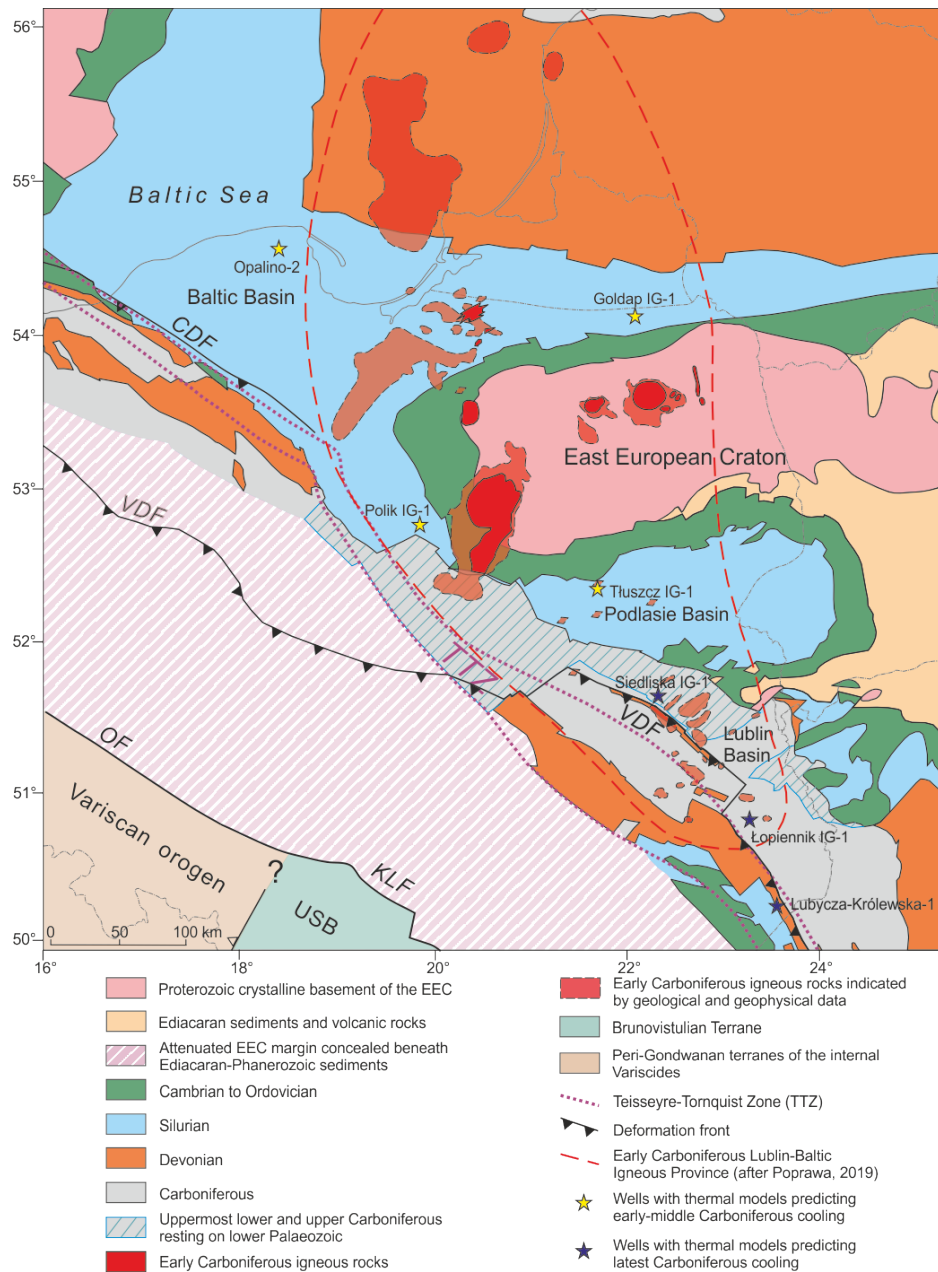
Figure 14. Pre-Permian geology map for the SW slope of the East European Platform in Poland and adjacent areas (partly based on Pożaryski and Dembowski, 1983). The extent of early Carboniferous igneous rocks and the Lublin-Baltic Igneous Province is shown after Poprawa (2019). Location of the Variscan Deformation Front is from Mazur et al. (2021). Further explanations in the text. CDF - Caledonian Deformation Front; KLF - Kraków-Lubliniec Fault; OF - Odra Fault; TTZ - Teisseyre-Tornquist Zone; USB - Upper Silesia Block; VDF - Variscan Deformation Front.

cooled during Mesozoic-Cenozoic regardless the cause is. Of course, the simplest solution would be erosional unroofing. Indeed, the Permian-Mesozoic-Cenozoic sedimentary cover is relatively thin and full of sedimentary gaps. Nevertheless, the well logs did not show any clear erosional or ungular unconformities. Furthermore, paleogeothermal gradients calculated based on VR data suggest only little to no exhumation throughout the Mesozoic (Figs. 6–12). Therefore, decreasing heat flow appears a main cause of Mesozoic cooling. Consequently, the mostly Jurassic-Early Cretaceous AFT data obtained, or earliest Jurassic thermal peaks predicted by thermal models do not represent any specific tectonic events or erosional episodes. Instead, they correspond to the time when the AFT samples left the apatite partial annealing zone. This is possible because the VR-derived paleotemperatures for the Permian-Mesozoic samples do not exceed 90°C (besides the Bodzanów IG-1 borehole with a present-day temperature of 75°C).

The models predicting an early Carboniferous thermal peak show relatively slow cooling during the Mesozoic (Fig. 13). In addition, the models built for sample O2/2 and B19 reveal acceleration of cooling at the time of Late Cretaceous basin inversion (Fig. 13a). The models characterised by a latest Carboniferous temperature maximum also demonstrate cooling throughout the Mesozoic (Fig. 13). However, this is related to sensitivity of the AFT method. The samples (B38, B45) that were at depths within the apatite partial annealing zone at the beginning of Mesozoic were unable to record early Mesozoic re-heating.

6.4 Geological implications

A deep erosional incision within the sedimentary cover of the East European Platform has been conventionally attributed to the effects of the Variscan orogeny (e.g., Żelichowski, 1987; Narkiewicz, 2007). This approach was consistent with wide-scale observations of a regional early Stephanian hiatus over much of the Variscan foreland with Stephanian and lower Permian red beds unconformably overlying truncated Westphalian series (e.g., McCann, 1996). However, our models indicate early Carboniferous exhumation of the lower Palaeozoic strata in the Baltic and Podlasie Basins of NE Poland on the SW slope of the EEP. This agrees with the presence of Carboniferous sediments resting on the deeply eroded Silurian substratum over a significant area of the Lublin and Podlasie Basins; (Fig. 14; Pożaryski and Dembowski, 1983; Żelichowski, 1987; Narkiewicz, 2007). Furthermore, early Carboniferous exhumation is consistent with data obtained in Bornholm Island (Denmark) and southern Sweden (Fig. 15; Hansen, 1995; Japsen et al., 2016; Guenthner et al., 2017).

A maximum paleotemperature achieved in the latest Devonian to earliest Carboniferous was probably related to coeval deposition of marine sediments within the Laurussia shelf basin. Thick successions of such sediments have been drilled in Pomerania (NW Poland) west of the TTZ (Żelichowski, 1987; Narkiewicz, 2007). However, based on sedimentary facies models, it is postulated that the Devonian and lower Carboniferous succession could have initially extended farther north and east than now, covering large areas of the SW slope of the EEP (Żelichowski, 1987; Matyja, 2006). This corollary is also supported by early Carboniferous illite K-Ar ages from the lower Palaeozoic sediments of the Baltic and Podlasie Basins (Kowalska et al., 2019).

The timing of the postulated early Carboniferous uplift and exhumation coincide with activity of the Lublin-Baltic Igneous Province defined by Poprawa (2019). This was a period of intra-plate alkaline magmatism in the time interval from the late Tournaisian to mid-Viséan (Pańczyk and Nawrocki, 2015). The area affected by magmatism overlaps the SW slope of the EEP
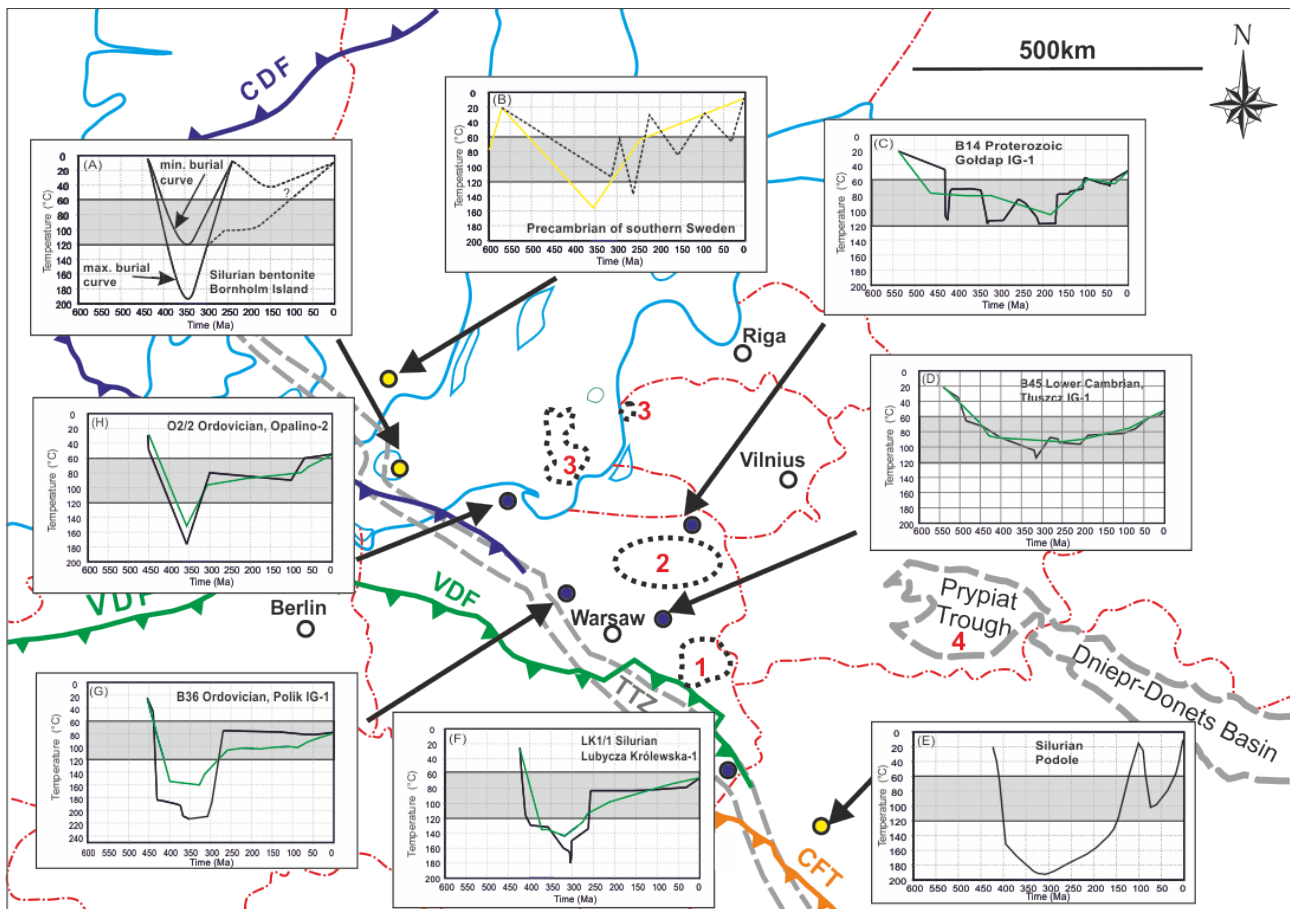
Figure 15. Simplified thermal history of different parts of the study area and adjacent regions. A. Silurian bentonite from Bornholm Island, dashed line shows possible options for t-T history in the Mesozoic (Hansen, 1995); B. Precambrian of Southern Sweden: dashed line from (Japsen et al., 2016), yellow bold curve form (Guenthner et al., 2017); C. Gołdap IG-1, well this study; D. Tłuszcz IG-1, this study, E. Silurian bentonite from Podole (Środoń et al., 2013), F. Lubycza Królewska-1, this study; G. Polik IG-1 this study; H. Opalino-2, this study. In the C-D, F-H black curve is the best-fit model, whereas green curve is average. Grey box - apatite partial annealing zone (60-120°C). Black dots show localities from this study, yellow dots - from other studies. Black dotted polygons on the map indicate the latest Devonian to early Carboniferous volcanic rocks in the study area and surroundings: 1. (Pańczyk and Nawrocki, 2015): 348 ± 0.8 Ma, alkaline basalts; 2. (Demaiffe et al., 2013): 345-354 Ma, alkaline intrusions; (Motuza et al., 2015): dolerite sills 355 Ma and a granite vein 349 Ma; 4. (Kusznir et al., 1996): 367 - 364 Ma rifting and magmatism. CDF - Caledonian Deformation Front; CFT - Carpathian Frontal Thrust; TTZ - Teisseyre-Tornquist Zone; VDF - Variscan Deformation Front.

from the Lublin to Baltic Basin (Fig. 14). The time of this event was roughly coeval with the extensional reactivation of the Dniepr-Donets-Donbas Rift in the early Viséan (Fig. 15; Stephenson et al., 2006). The formation of an igneous province is

usually a manifestation of thermal anomaly in the lithosphere. Regardless a specific tectonic setting, the anomaly promotes uplift and exhumation.

610   The early Carboniferous uplift and exhumation must have affected the entire study area as indicated by a regional erosional unconformity over the top of Silurian or Devonian. However, in SE Poland (Lublin Basin) sedimentation resumed in the late Viséan after a period of intra-plate volcanism (e.g., Narkiewicz, 2007), whereas the rest of the study area remind uplifted until the Permian. Furthermore, thermal models for the Lublin Basin predict peak temperatures for the latest Carboniferous and the VR-derived paleotemperature vertical profile reveals no break across the base Carboniferous unconformity (Figs. 13, 14).

615 Consequently, the early Carboniferous uplift and exhumation in the Lublin Basin must have been not only shorter, but also less significant than farther NW.

     Surprisingly, the Lublin Basin, the least affected by early Carboniferous exhumation in the study area, was the only area on the SW slope of the EEP, where this event has been so far recognised and correlated with "Bretonnian phase" (Narkiewicz, 2007; Krzywiec et al., 2017a). A top Devonian unconformity, equivalent to a top Silurian unconformity farther NW, was

620 accordingly termed a "Bretonnian unconformity". The "Bretonnian" tectonic event might have been also responsible for the termination of the Late Devonian subsidence in the nearby Pripyat Trough (Fig. 15; Kusznir et al., 1996). Although the earliest Carboniferous event is reflected in the Lublin Basin by block tectonics of considerable magnitude (Narkiewicz, 2007) the latter was probably insufficient to influence the low-T thermochronometers. The effects of the tectonic uplift at the transition from the Devonian to Carboniferous were not recognised farther northwest in the Podlasie and Baltic Basin because all Devonian

625 and Carboniferous sediments were removed. Paradoxically, the north-westward increasing magnitude of uplift prevented identification of its results owing to the ensuing exhumation. Therefore, thermochronology remains the best tool to recognise the scale of the early Carboniferous event, in the absence of sedimentary record.

     The Lublin Basin in SE Poland was the only part of the SW slope of the EEP that was overridden by the Variscan orogenic wedge in the late Carboniferous; (Fig. 14; Mazur et al., 2021). Therefore, we hypothesize that Carboniferous uplift and ex-

630 humation in this area was interrupted by tectonic loading associated with the adjacent orogenic wedge. Therefore, after initial uplift and exhumation at the transition from the Devonian to Carboniferous, sedimentation resumed in the late Viséan and lasted until latest Carboniferous Variscan shortening (Fig. 2c; Krzywiec et al., 2017a). Recent seismic data indicate a Variscan thin-skinned fold-and-thrust belt emplaced on SW slope of the EEP within the Lublin Basin and west of it (Krzywiec et al., 2017a,b). Hence, latest Carboniferous heating might have been caused by a combined effect of tectonic and sedimentary burial

635 (Fig. 5b; Botor et al., 2019a). A significant cover of the Variscan foreland basin was also inferred by Środoń et al. (2013) for western Ukraine (Fig. 15) based on combined clay mineralogy, K-Ar, and AFT data.

     From the beginning of Permian, the SW slope of the EEP was onlapped by marginal part of an extensive Permian-Mesozoic basin (Figs. 2, 4). Considering a present-day Permian-Mesozoic thickness (Figs. 2, 4), i.e., omitting effects of the Late Cretaceous inversion, a thermal effect of Mesozoic burial should be significant. Nevertheless, among the methods used, only the

640 AFT thermochronometer is sensitive enough to record this event. Furthermore, only models for Permian-Mesozoic samples are capable to reveal a Mesozoic thermal history. Consequently, only three models predict Mesozoic re-heating with a maximum paleotemperature at the transition from Triassic to Jurassic (Fig. 13), but this result is probably representative for the

entire study area. Comparable results were obtained by Schito et al. (2018) in the Ukrainian part of the EEP, who postulated that exhumation through the 45-120°C temperature range took place between the Late Triassic and Early Jurassic, and that no significant burial occurred afterwards.

7 Conclusions

The case of the SW slope of the EEP demonstrates a successful application of integrative approach to studying a long-term thermal history. A series of tectonic events in the area has not been yet fully resolved because of incomplete sedimentary record. Therefore, the combination of ZHe and AFT termochronometers with VR data was necessary to recognise those events that led to deep erosion and complete removal of sedimentary sequences. The same approach might be applied to craton margins elsewhere to separate superimposed effects of successive tectonic events.

Significant heating of Ediacaran to Carboniferous sedimentary successions occurred before the Permian with maximum paleotemperatures in the earliest and latest Carboniferous in the Baltic-Podlasie Basin and Lublin Basin, respectively. The results obtained suggest an important role of early Carboniferous uplift and exhumation at the SW margin of the EEP. This event was associated with a period of intra-plate magmatism in the area (Poprawa, 2019) lasting from the late Tournaisian to mid-Viséan (Pańczyk and Nawrocki, 2015). Effects of uplift and magmatism jointly suggest thermal perturbation of lithosphere. The time of this event was roughly coeval with the extensional reactivation of the Dniepr-Donets-Donbas Rift (e.g., Stephenson et al., 2006) and termination of the Late Devonian continental rifting in the Pripyat Trough (Kusznir et al., 1996).

The SW slope of the EEP was overridden in SE Poland (Lublin Basin) by the Variscan orogenic wedge (Krzywiec et al., 2017a; Mazur et al., 2021). This event interrupted Carboniferous uplift because of tectonic loading and caused resumption of sedimentation from the late Viséan. Consequently, a thermal history of the Lublin Basin is different from that in the Podlasie and Baltic Basins, but similar to other sections of the Variscan foreland. The latter is characterised by maximum burial at the end of Carboniferous and regional uplift and exhumation at the transition from the Carboniferous to Permian (e.g., McCann, 1996).

Our data are consistent with the decreasing heat flow during the Mesozoic. The elevated Permian-Triassic heat flow was probably a consequence of early Permian continental rifting. In the light of our results, the decrease of post-Permian heat flow appears an important cause of Mesozoic cooling. Our thermal models show mostly gradual cooling with little effects of the Late Cretaceous basin inversion. Although inversion structures are recognised in the part of the German-Polish Basin onlapping the EEP (Krzywiec, 2009) their vertical offset was probably below the sensitivity of the AFT method.

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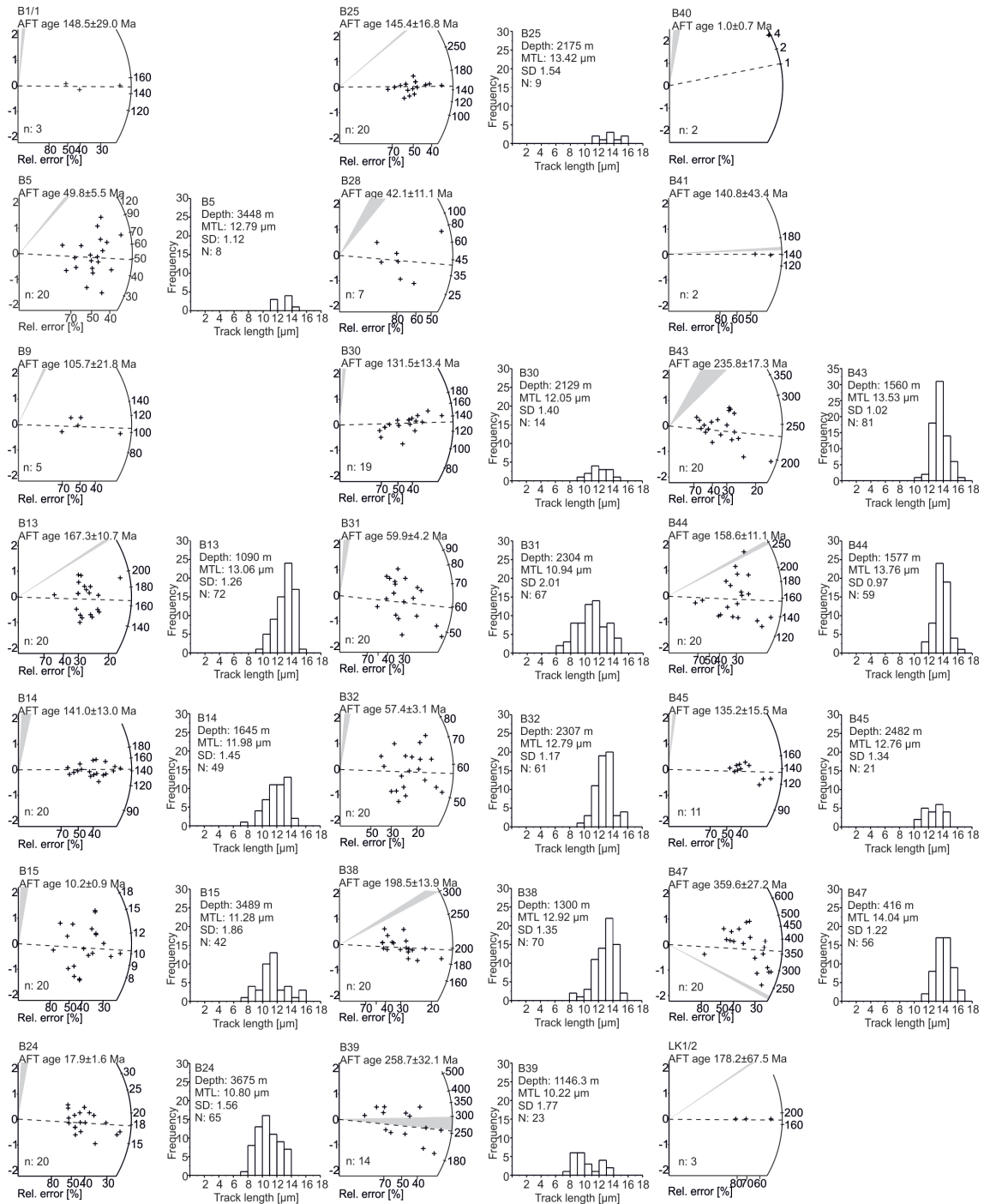
Figure A1. Radial plots of apatite fission track data and length distributions of confined fission tracks in analysed samples. Central ages of samples are defined by dashed lines, stratigraphic ages are represented by shaded areas. The position of the x-scale records the uncertainty of individual age estimates, whilst each point has the same standard error on the y-scale (illustrated as ±2σ). The age of each crystal can be determined by a line from the origin through the projection point of the crystal to the intercept on the radial age scale (Galbraith, 1990).

Competing interests. The authors declare that they have no conflict of interest.

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